

# Ocean-atmosphere coupling over mid-latitude oceans

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SST has been viewed as the surface signature of ocean's response to atmospheric forcing or the oceanic driver of atmospheric circulation. Local and remote connections with atmosphere have long been studied

- Local influence
- Teleconnection

## Limited Extra-tropical Coupling

- **Ocean has slow and small-scale, but atmosphere has fast and large-scale processes.**

**Coupling in mid-latitudes at long time period is controversial.**

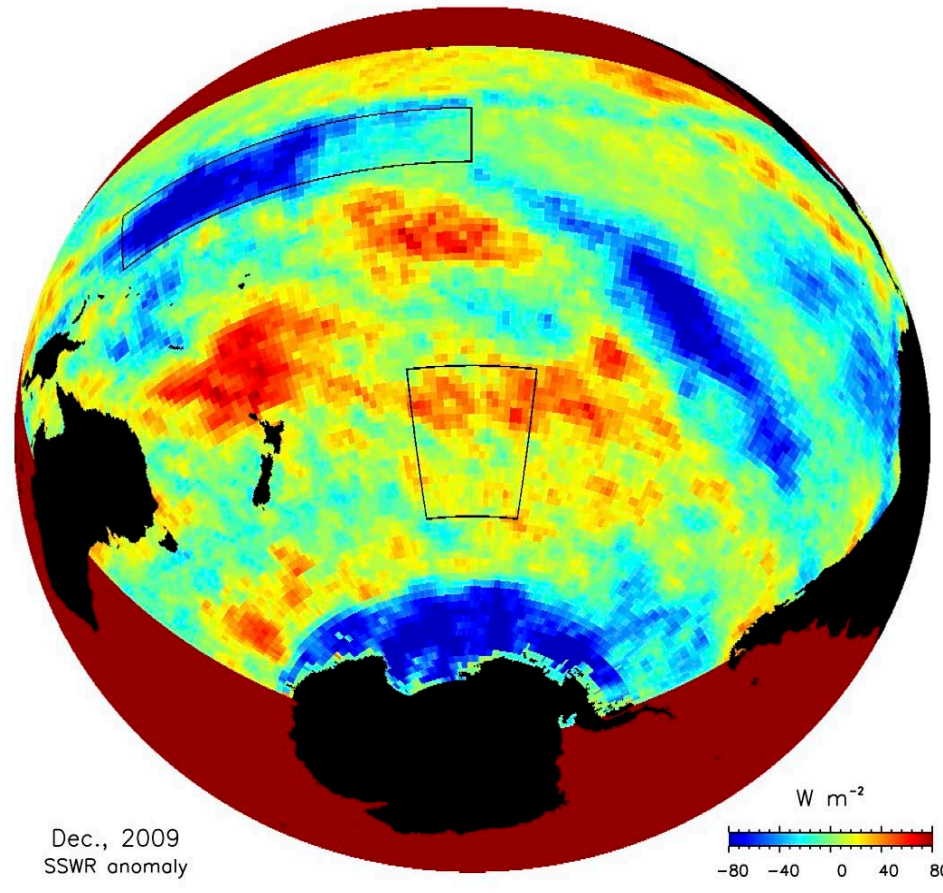
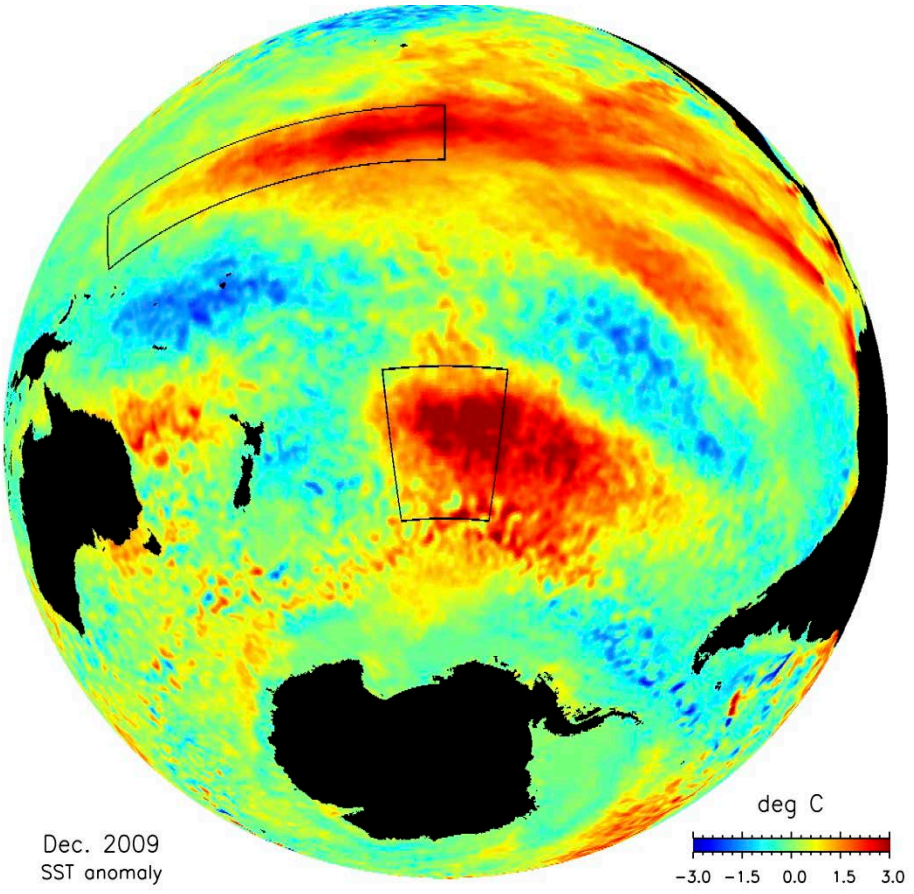
**GCM fails to generate systematic response to prescribed midlatitude SST anomalies above boundary layer**

**Lapse rate is believe to be too weak to generate deep convection, as over the tropical oceans.**

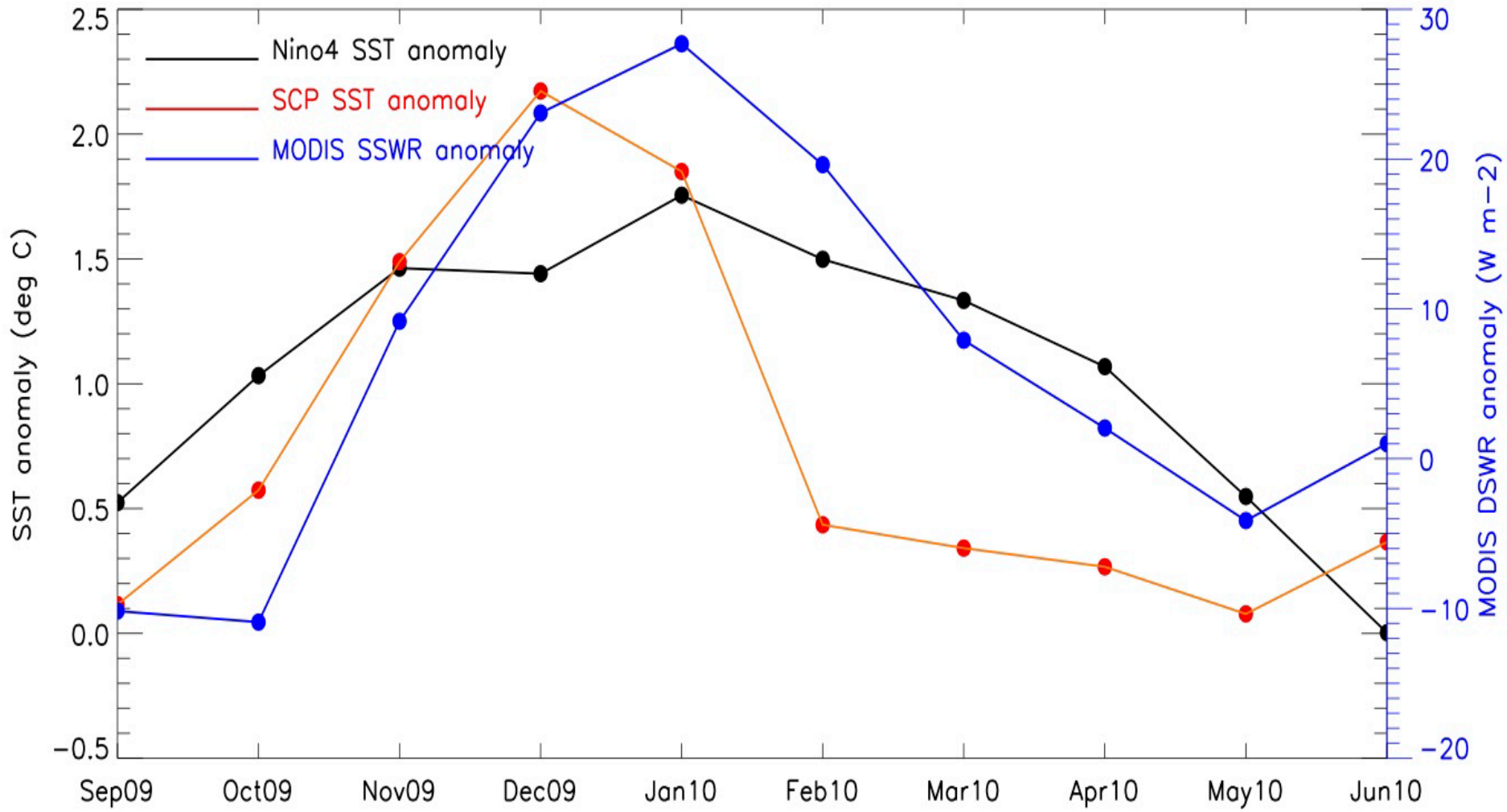
## Atmospheric Bridges Hypothesis (e.g., Kushnir et al. 2002; Alexander et al. 2002)

- In the tropical Pacific, SST anomalies during El Niño cause deep convection that affects large-scale atmospheric circulation, which alters the near-surface air temperature, humidity, and wind far from the equatorial Pacific.
- The resulting variation in heat, water, and momentum flux drives local SST anomalies at extra-tropical oceans.

- A record warming-event occurred in the south-central Pacific, starting in September and peaking in December of 2009.
- Lee et al. [2010] attribute the warming to reduced evaporative cooling and Ekman advection caused by persistent anticyclone.
- Boening et al. [2011] attributed to mass convergence in the ocean resulted from wind stress curl associated with the anticyclone.
- Solar heating was deemed unimportant because NCEP reanalysis show reduced solar heating.



Liu et al., 2013, IJRS, I npress

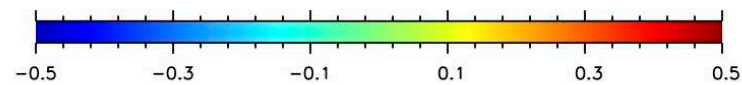
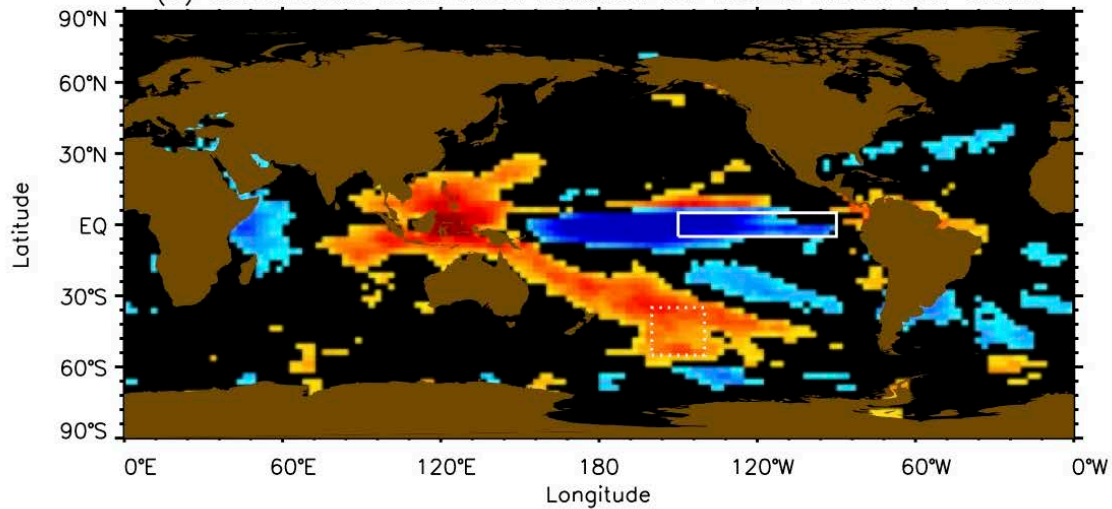


## Solar contribution to SCP warming

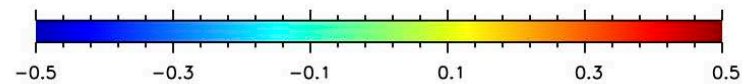
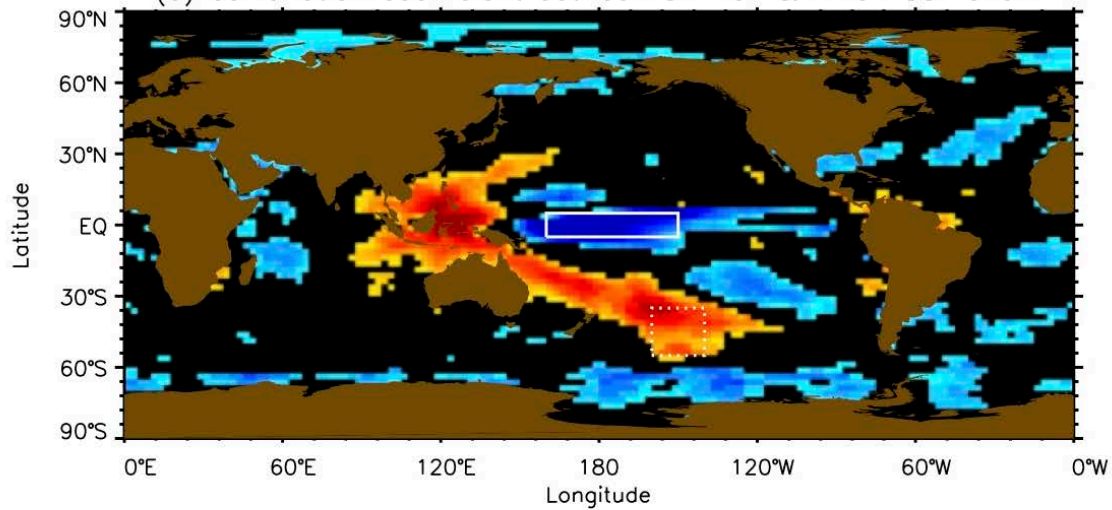
	<b>(a) SSWR anomalies (W/m<sup>2</sup>)</b>	<b>(b) Ta (°C)</b>	<b>(c) SSTa (°C)</b>	<b>(d) Ta / SSTa (%)</b>
<b>Sept. 2009</b>	<b>-10.17</b>	<b>-0.1280</b>	<b>0.1153</b>	<b>-</b>
<b>Oct. 2009</b>	<b>-10.92</b>	<b>-0.1373</b>	<b>0.5736</b>	<b>-</b>
<b>Nov. 2009</b>	<b>9.17</b>	<b>0.1154</b>	<b>1.4890</b>	<b>7.75</b>
<b>Dec. 2009</b>	<b>23.07</b>	<b>0.2903</b>	<b>2.1725</b>	<b>13.36</b>
<b>Jan. 2010</b>	<b>27.70</b>	<b>0.3485</b>	<b>1.8502</b>	<b>18.84</b>
<b>Feb. 2010</b>	<b>19.62</b>	<b>0.2469</b>	<b>0.4350</b>	<b>56.76</b>
<b>Mar. 2010</b>	<b>7.91</b>	<b>0.0995</b>	<b>0.3415</b>	<b>29.14</b>
<b>Apr. 2010</b>	<b>2.04</b>	<b>0.0257</b>	<b>0.2666</b>	<b>9.64</b>
<b>May 2010</b>	<b>-4.14</b>	<b>-0.0521</b>	<b>0.0780</b>	<b>-</b>
<b>Jun. 2010</b>	<b>1.00</b>	<b>-0.0126</b>	<b>0.3661</b>	<b>-</b>



(a) correlation coefficient between SW flux & nino3 SST anom

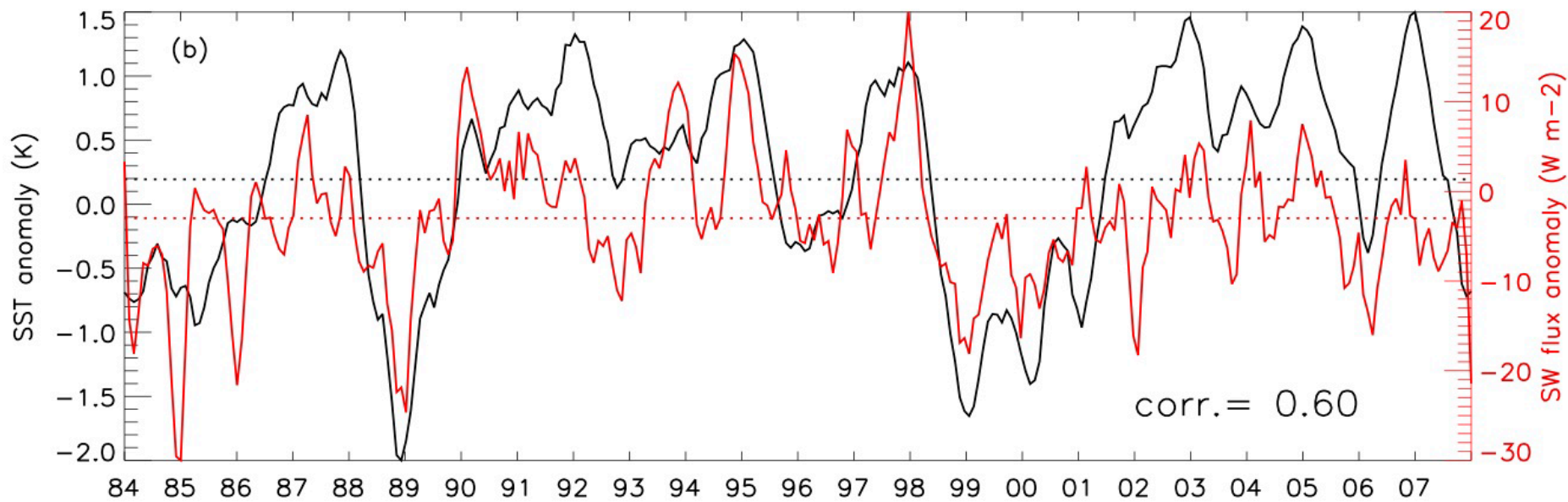
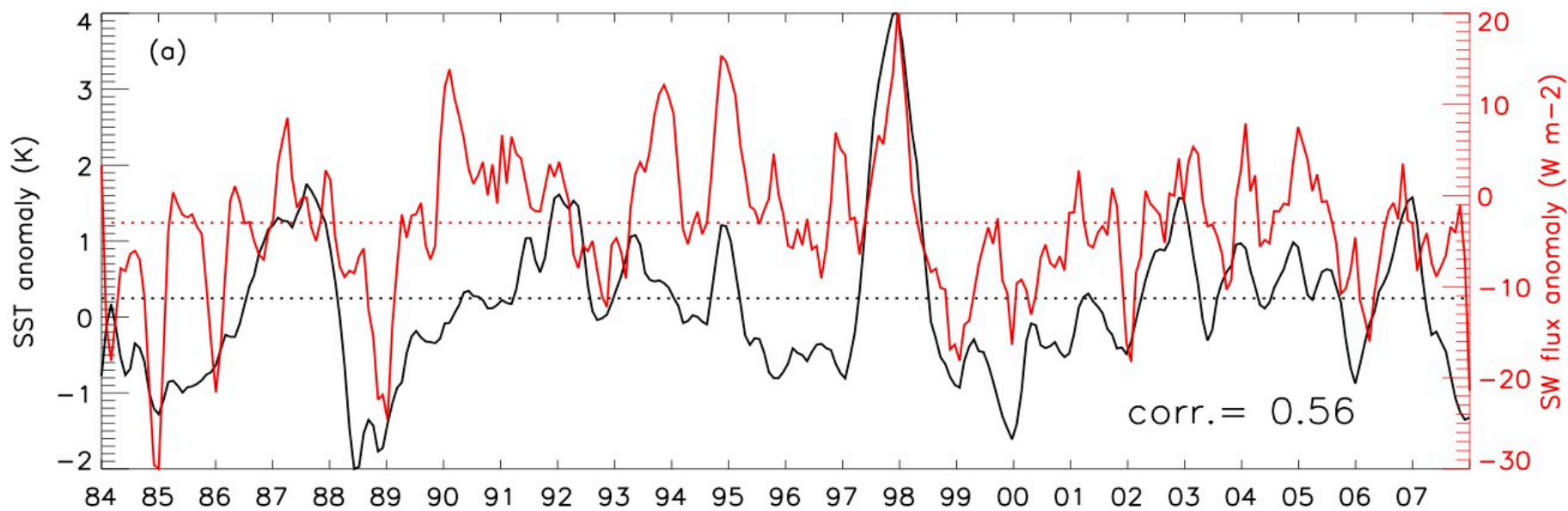


(b) correlation coefficient between SW flux & nino4 SST anom

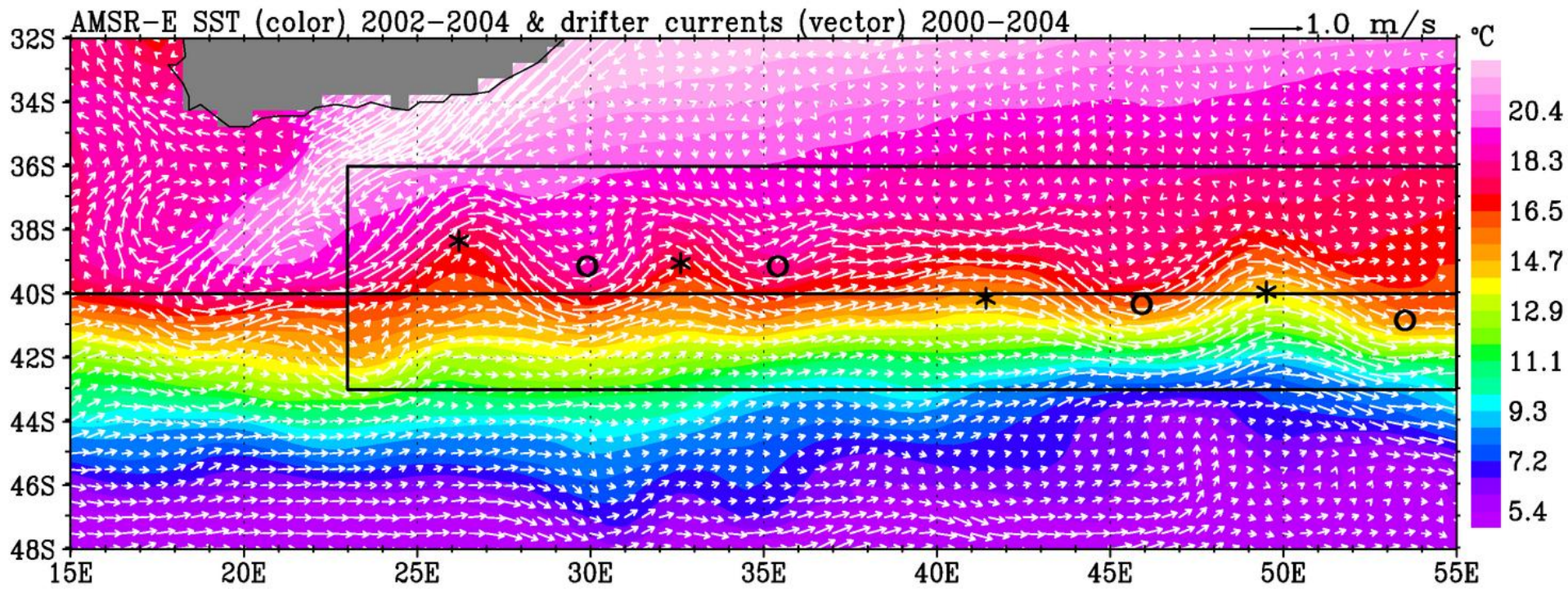
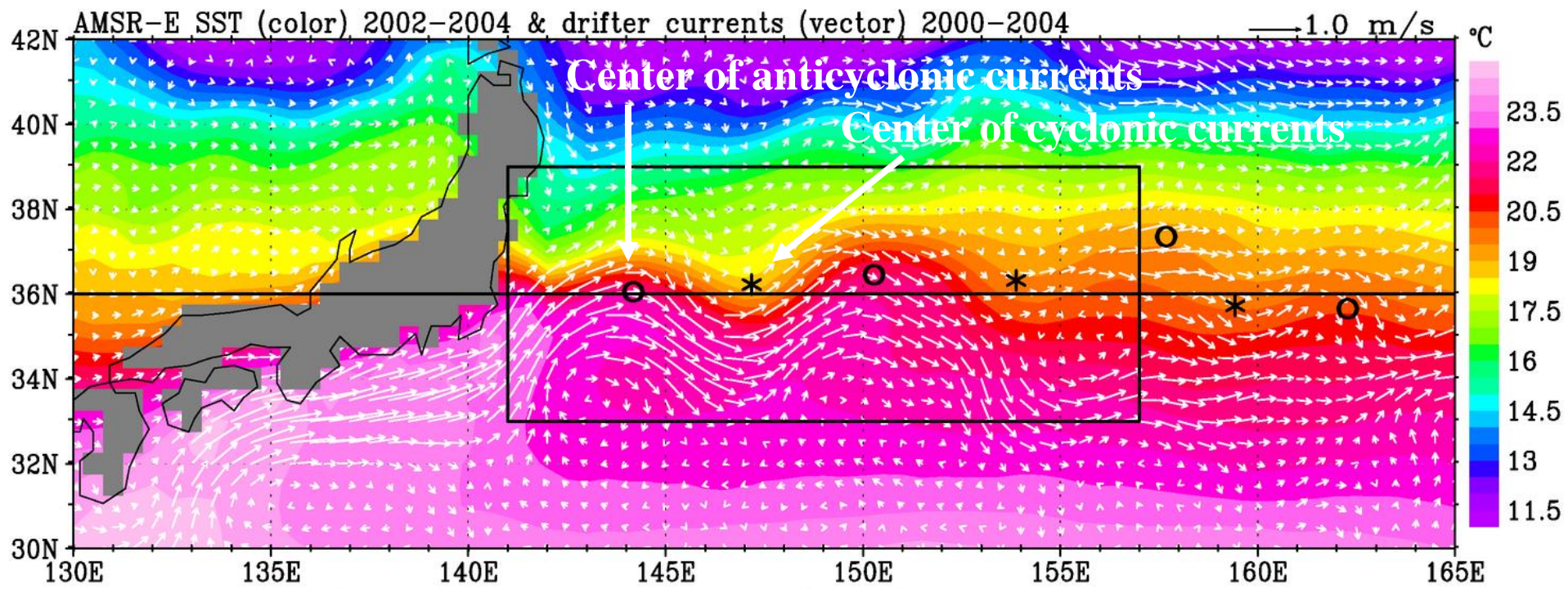


99% significant, radiation flux positive into ocean







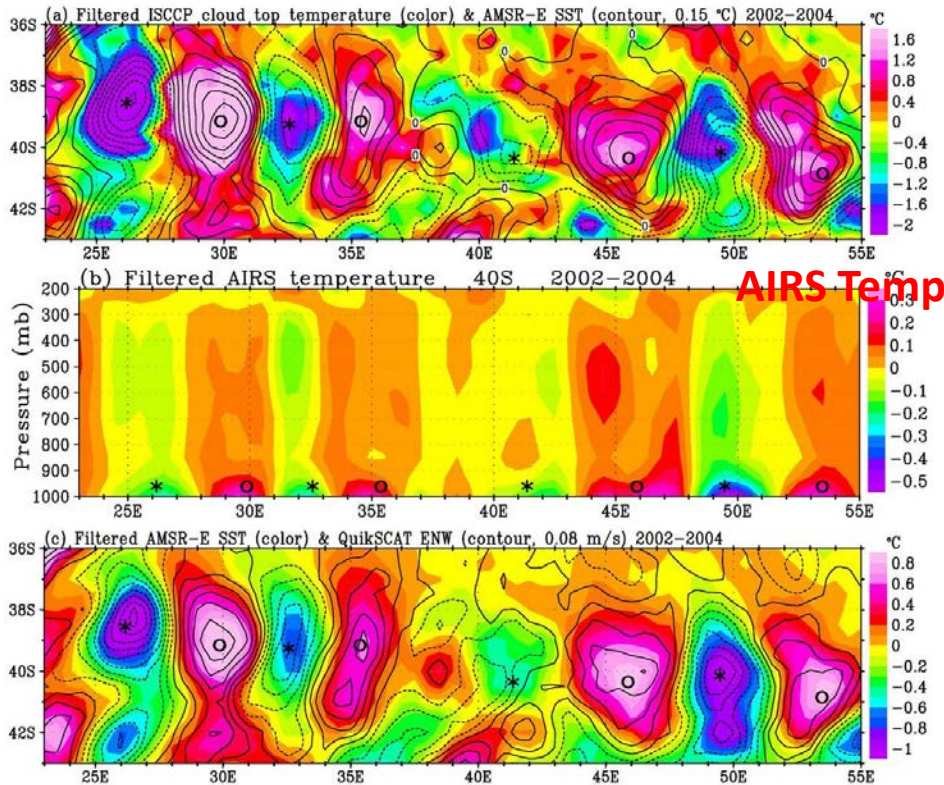




# From Ocean Surface to Tropopause

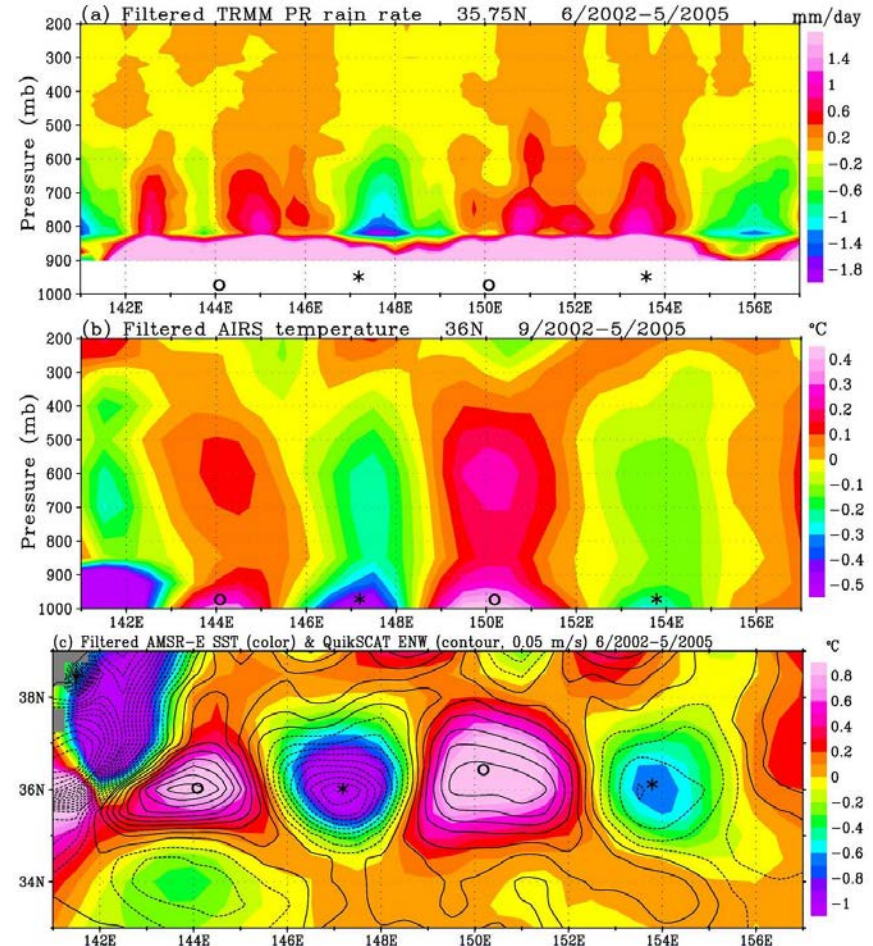
Over mid-latitude ocean, lapse rate is too weak to generate deep convection and transfer ocean effects beyond boundary layer?

## Agulhas, 40S, ISCCP CTT & AMSR-E SST



## AMSR-E SST & QuikSCAT $U_N$

## Kuroshio, 36N, TRMM PR rain



From tropical to mid-latitude oceans,

GPM is needed to monitor western boundary currents



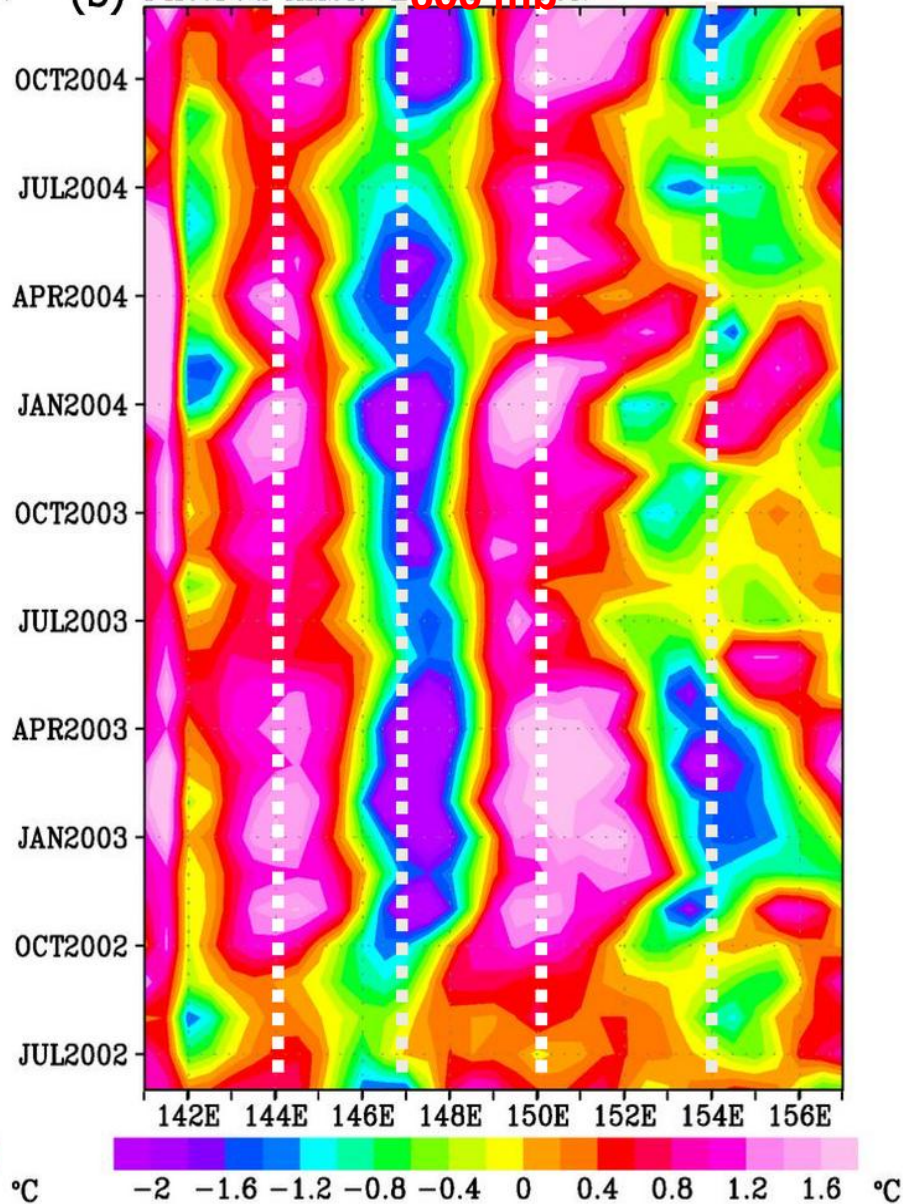
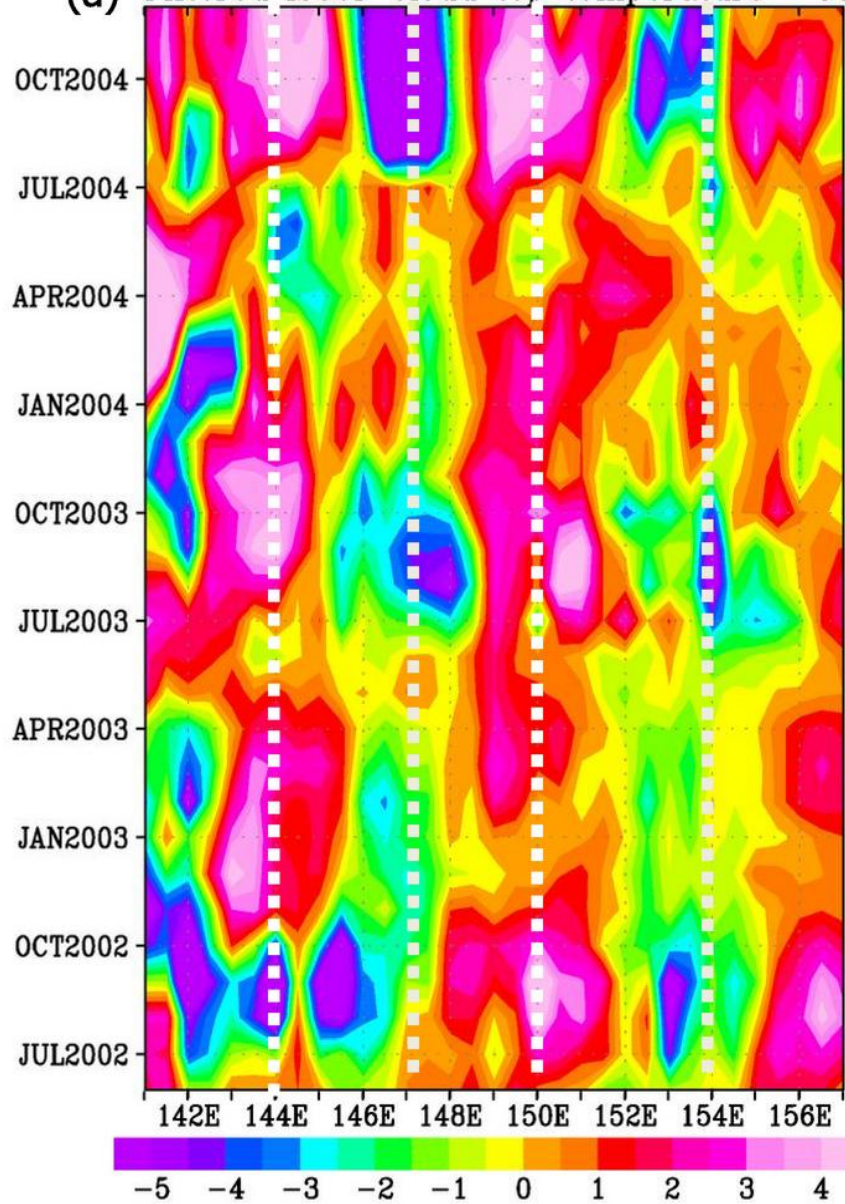
# ISCCP cloud top temperature

# AMSR sea surface temperature

## Kuroshio Extension, 36N,

(a) Filtered ISCCP cloud top temperature 36N

(b) Filtered AMSR-E SST 36N



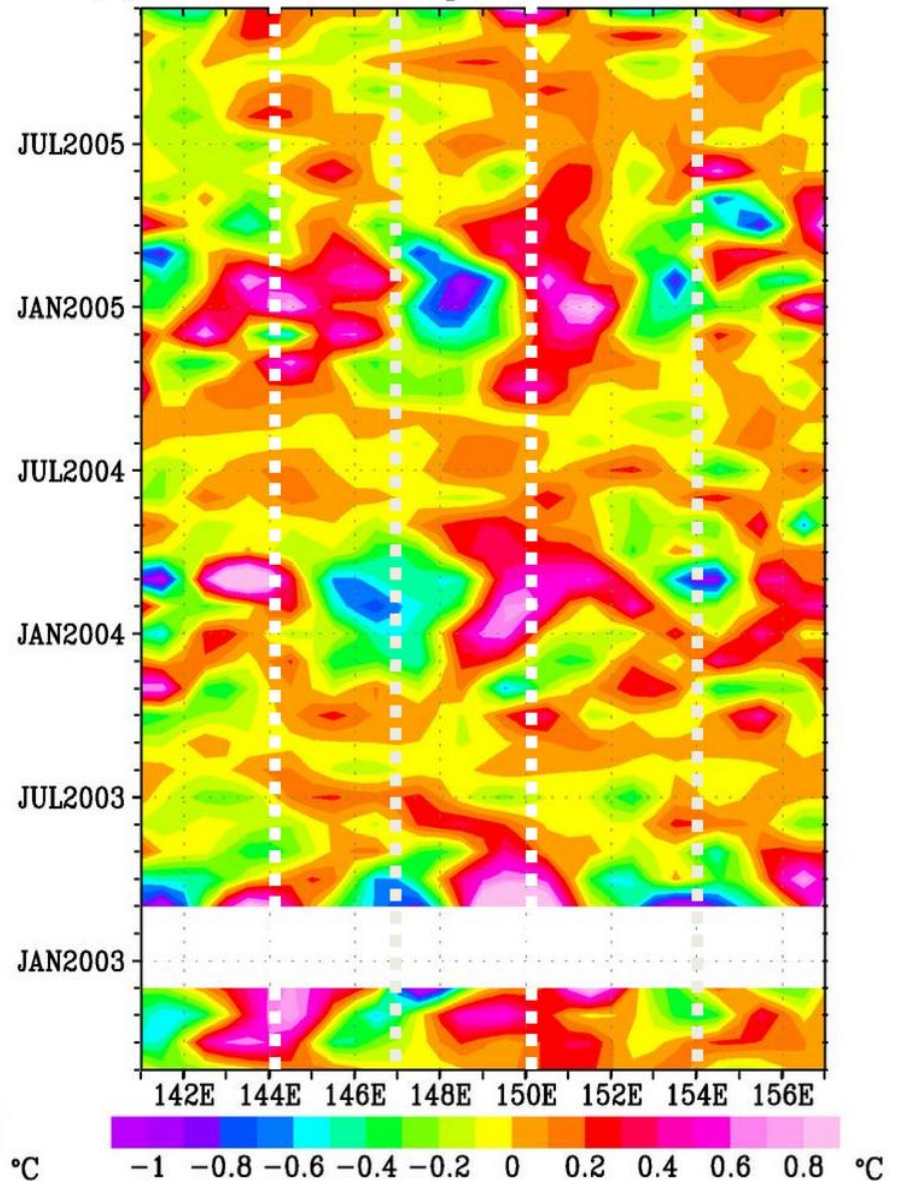
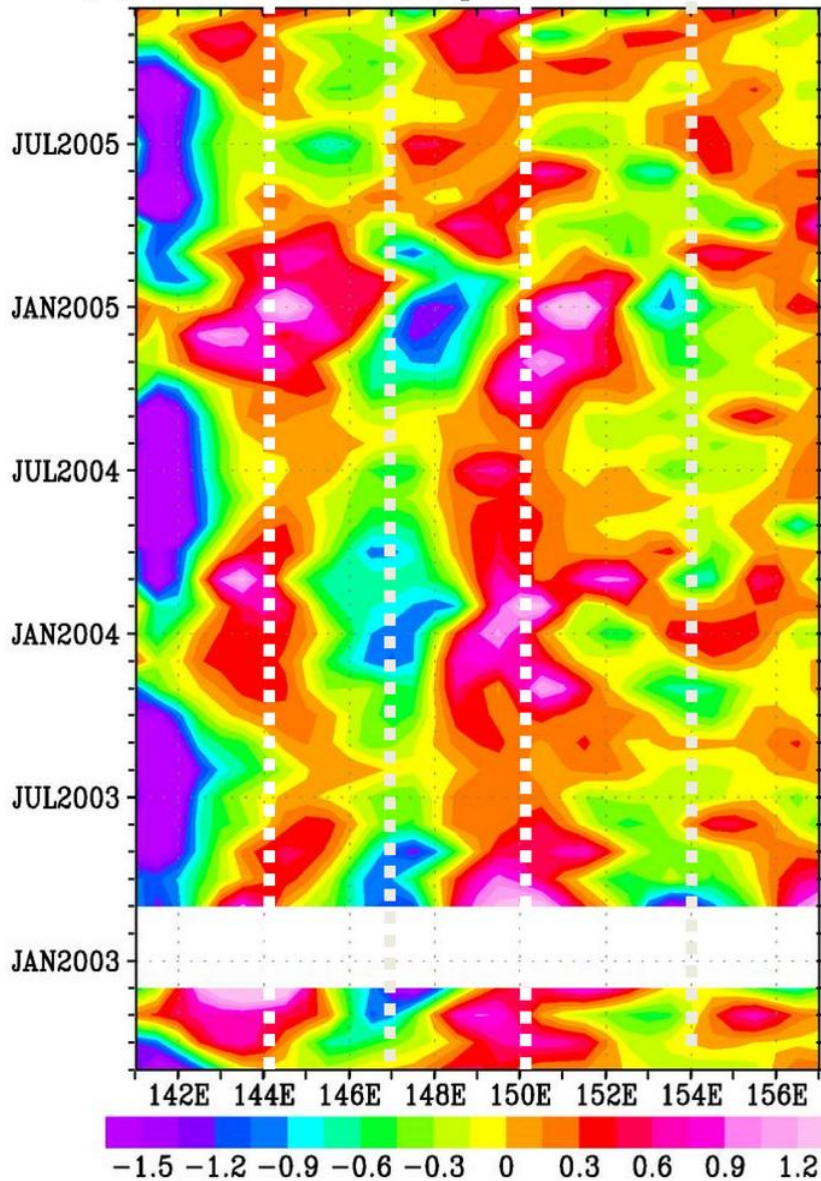


**AIRS temperature, Kuroshio Extension, 36N,  
1000 mb**

**600 mb**

**(a)** AIRS 1000 mb temperature 36N

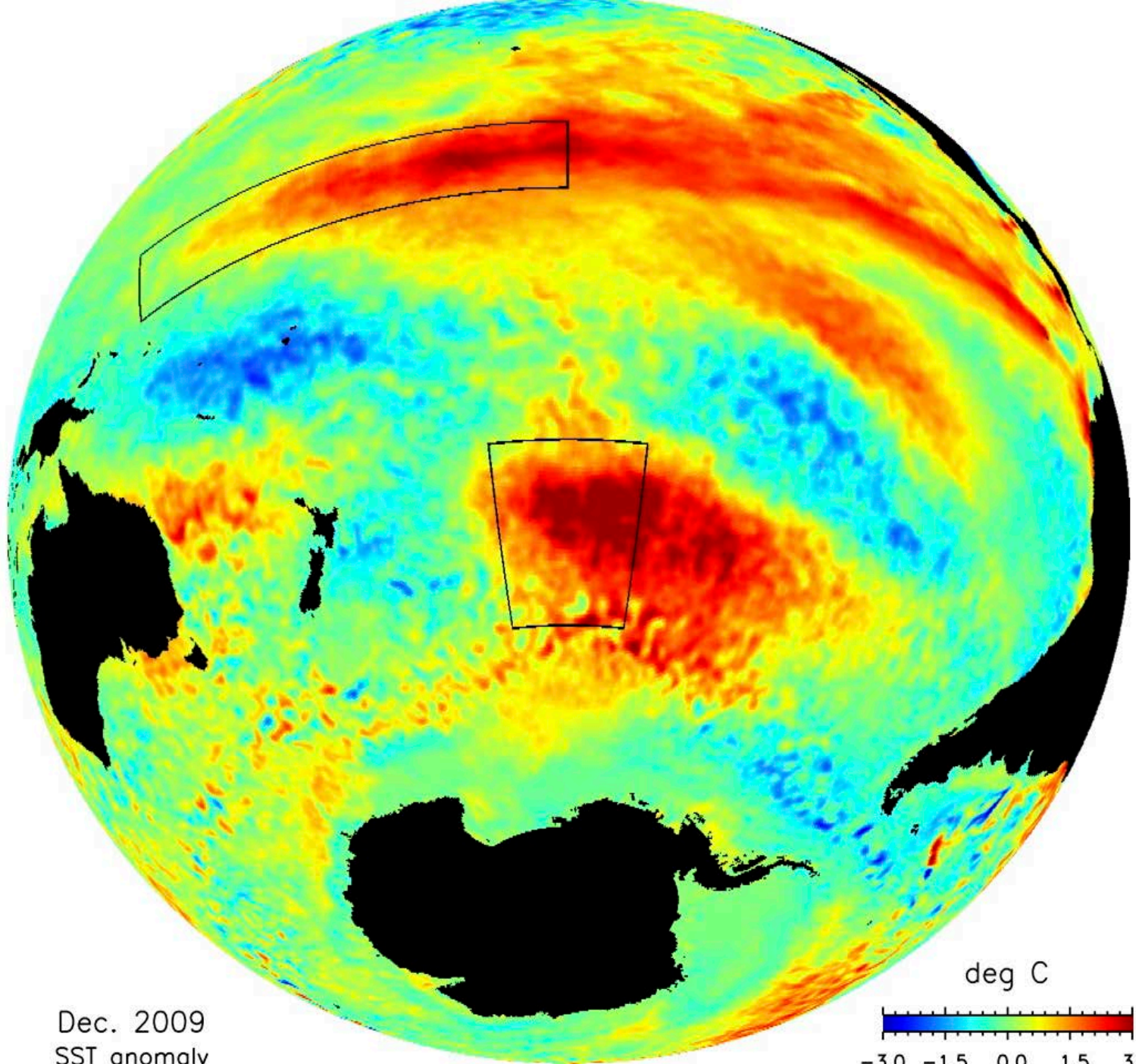
**(b)** AIRS 600 mb temperature 36N



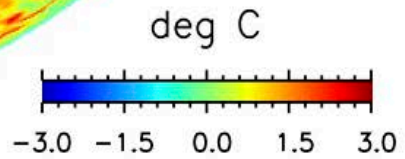
# Backup

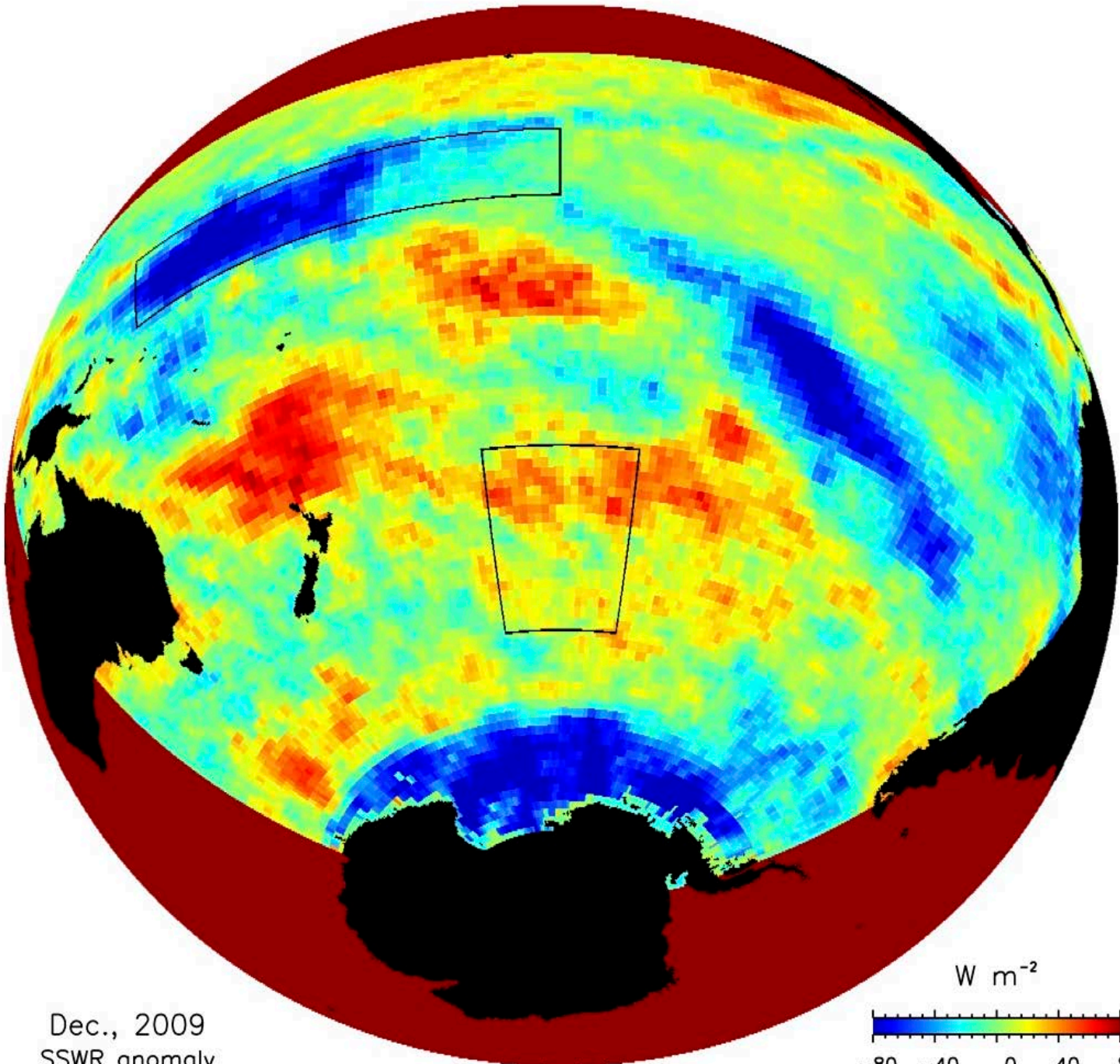




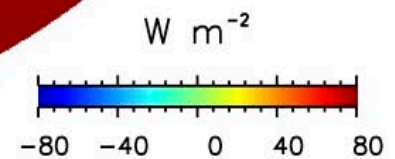


Dec. 2009  
SST anomaly





Dec., 2009  
SSWR anomaly





- **AIRS, TRMM, and ISCCP reveal signature of SST way above the atmospheric boundary layer at low frequencies.**
- **Present GCM or operational NWP data, do not show vertical propagation of ocean SST effect beyond atmospheric boundary layer.**
- **The observations present a challenge to understand the transition from random turbulence to organized convection and to improve **spatial versus temporal scales parameterization in GCM****

- The correlation coefficients of the geographic variations (585 pairs of data at a  $0.5^\circ$  grid in the region of  $42^\circ\text{--}38^\circ\text{S}$ ,  $23^\circ\text{--}55^\circ\text{E}$ ) between SST and CTT for 3-yr, 1-yr, and 3-m, 1-m periods are nearly constant, over 0.7