

# InitMIP-Antarctica experiments with the ice sheet model SICOPOLIS



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# InitMIP-Antarctica: Antarctic ice sheet model initialisation experiments

Early  
initiative within



- Initialise ice sheet to present day with method of choice (init)
- Run three forward experiments (100 a)
  - unforced control run (ctrl)
  - prescribed schematic SMB anomaly (asmb)
  - prescribed sub-ice-shelf melting anomaly (abmb)

(Seroussi et al., in prep.)

# Spin-up with SICOPOLIS

([www.sicopolis.net](http://www.sicopolis.net))

Model time:  $t = -135 \text{ ka} \dots 0 \text{ ka}$  (one glacial cycle).

Grid spacing:  $\Delta x = 16 \text{ km}$ .

(1) Without ice-stream dynamics / (2) with ice-stream dynamics.

Fixed topography (except first 100 a and last 500 a).

Surface temperature anomaly  $\Delta T$  from Vostok  $\delta D$  record.

(Petit et al., 1999)

Precip = precip\_today x fct( $\Delta T$ ).

(Precip\_today: Arthern et al., 2006; Le Brocq et al., 2010)

Runoff: PDD.

Parameterization of ice-shelf basal melting (next slides).

# Parameterisation of ice-shelf basal melting

Basal melting rate  $a_b$  = function of ocean temperature  $T_{oc}$ :

(modified after  
Beckmann and Goosse, 2003)

$$a_b = \frac{\rho_w c_w \gamma_t}{\rho L} \Omega \left( \frac{d}{d_0} \right)^\alpha (T_{oc} - T_{m,b})$$

$$T_{m,b} = -\beta_{sw} d - 1.85^\circ \text{C}$$

$\Omega$ : sensitivity parameter for each sector.

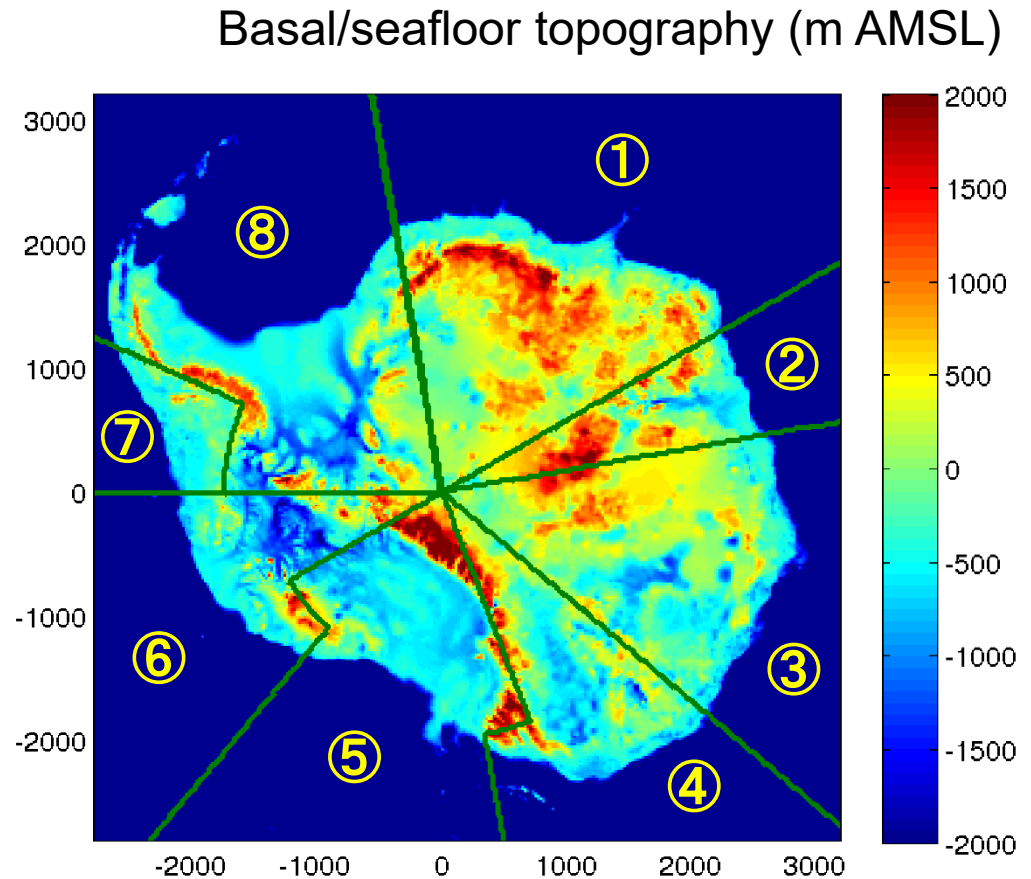
$d$ : ice-shelf draft (reference value  $d_0 = 200$  m).

$\alpha$ : non-linearity parameter.

$T_{m,b}$ : melting temperature at ice-shelf base, depends on draft  $d$ .

# Definition of eight sectors

- ① Western East Antarctica.
- ② Amery/Prydz Bay.
- ③ Sabrina Coast/Aurora subglacial basin.
- ④ George V Coast/Wilkes subglacial basin.
- ⑤ Ross Sea.
- ⑥ Amundsen Sea.
- ⑦ Bellingshausen Sea
- ⑧ Weddell Sea

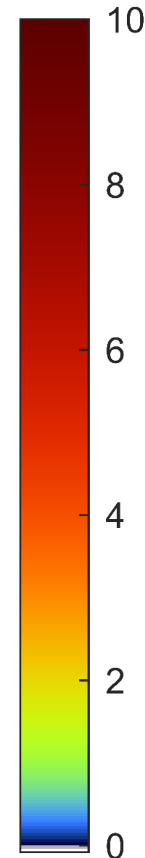
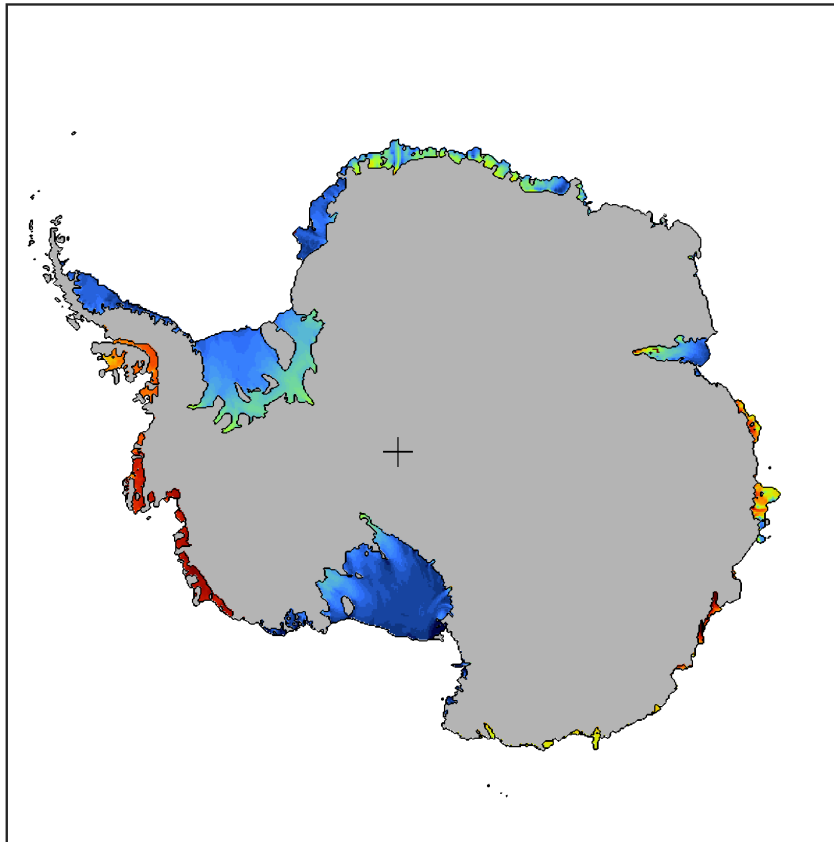


**Tuning by “observed” melting rates...**

# Distribution of basal melting

(on a 10-km grid)

$a_b$  (m/a)



Enhanced melting near deep grounding lines reproduced well.

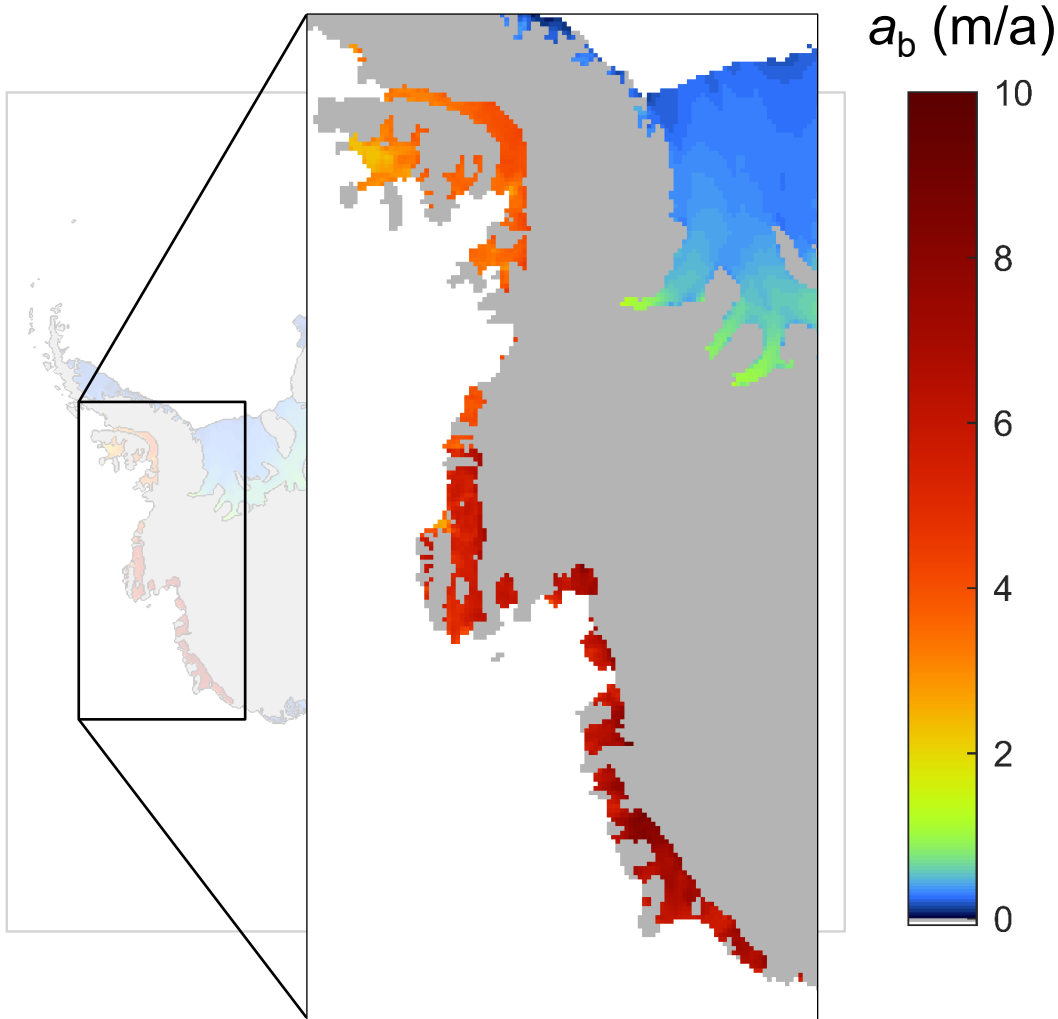
Due to different physics:

No enhanced melting near calving fronts.

No accretion of frazil ice.

# Distribution of basal melting

(on a 10-km grid)



Enhanced melting near deep grounding lines reproduced well.

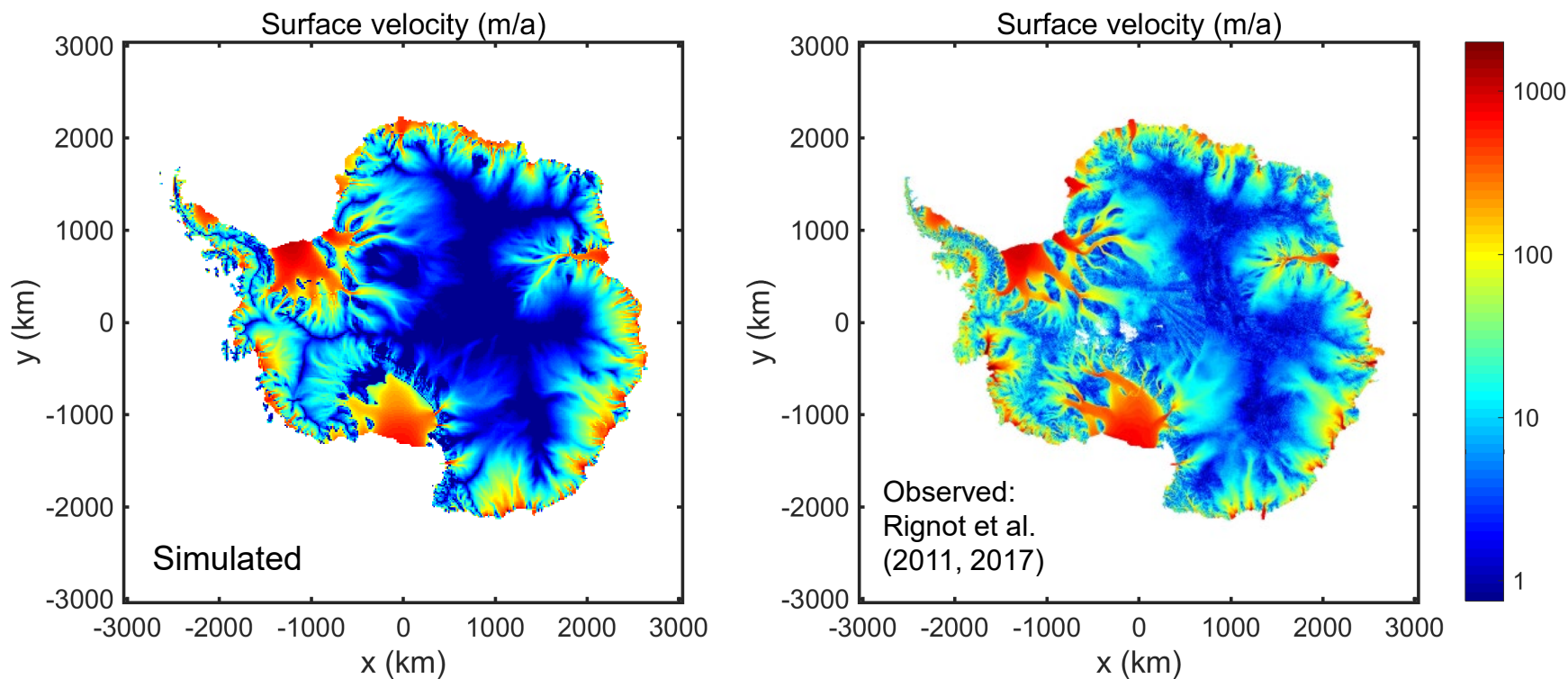
Due to different physics:

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# Spin-up: simulated vs. observed surface velocity

Without ice-stream dynamics:

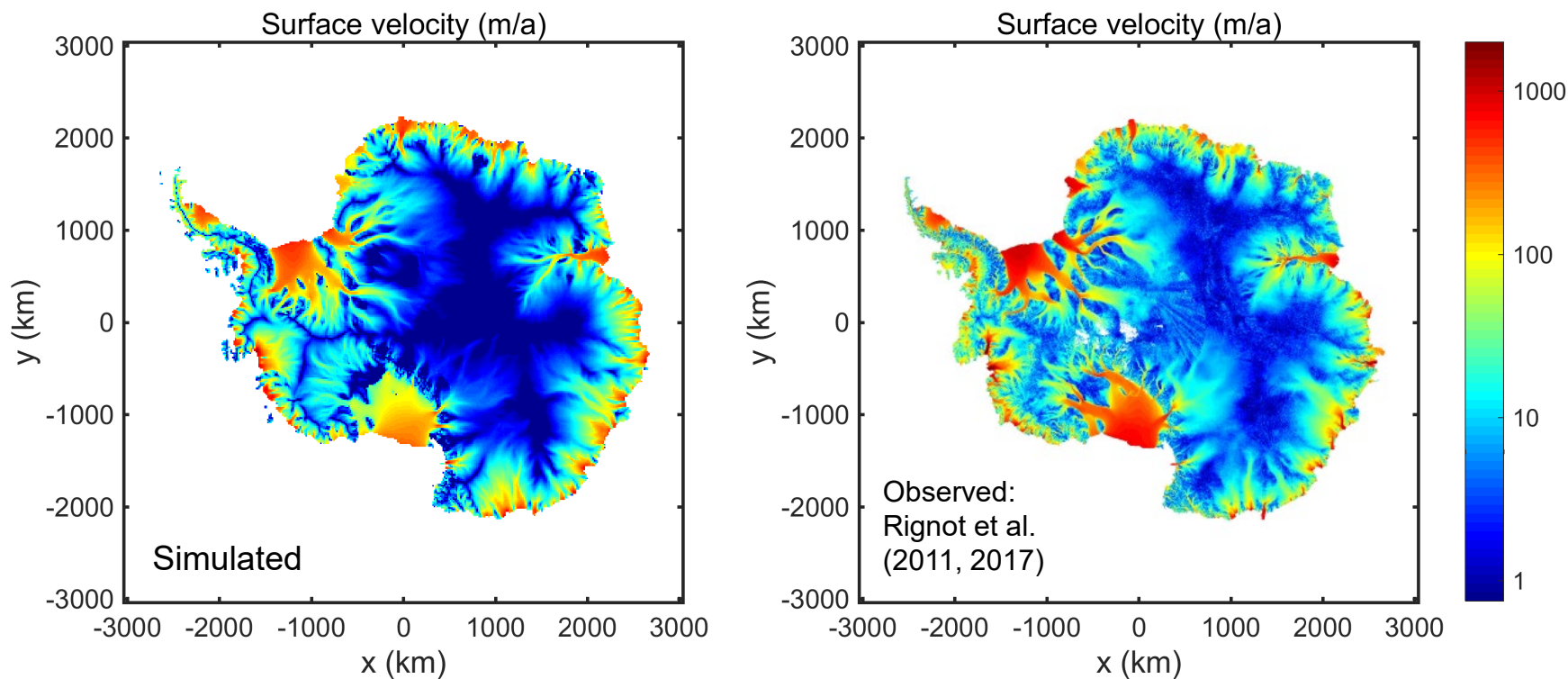


Very sharp transitions between sheet and shelves.



# Spin-up: simulated vs. observed surface velocity

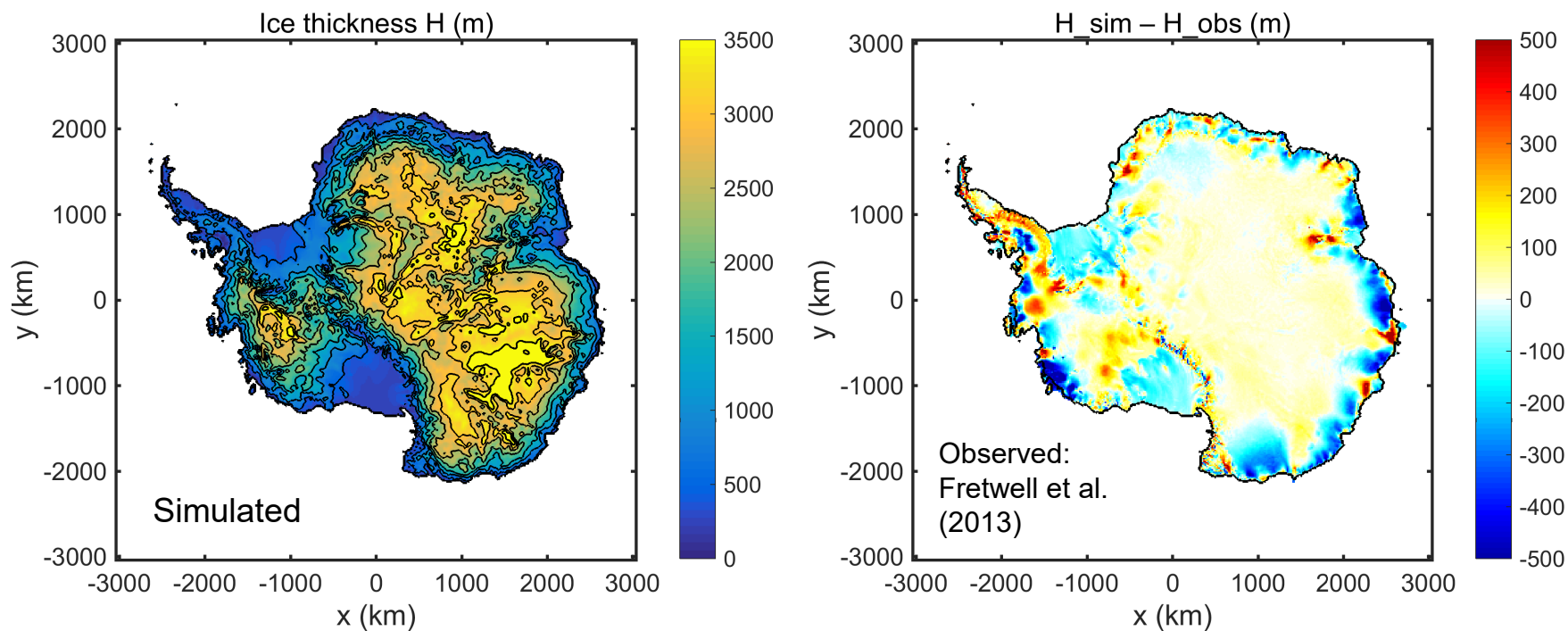
With ice-stream dynamics:



Smother transitions between sheet and shelves.

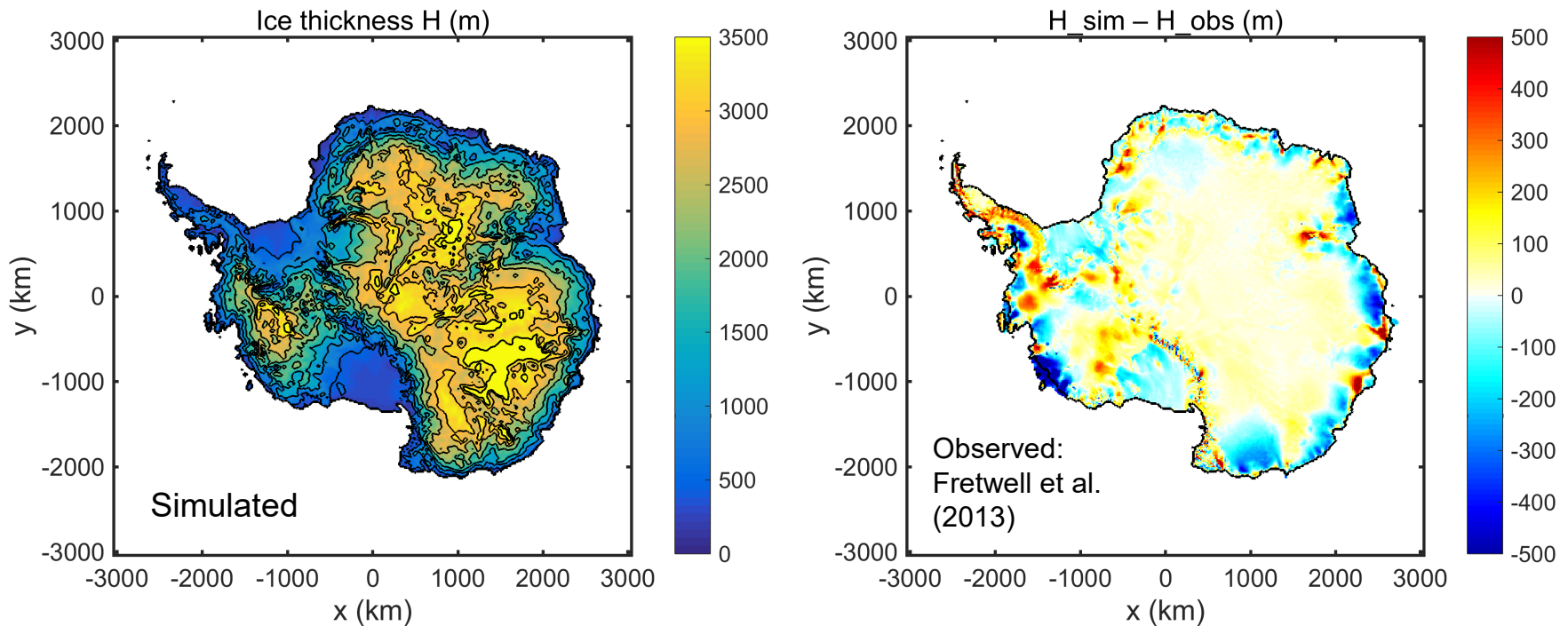
# Spin-up: simulated vs. observed ice thickness

Without ice-stream dynamics:

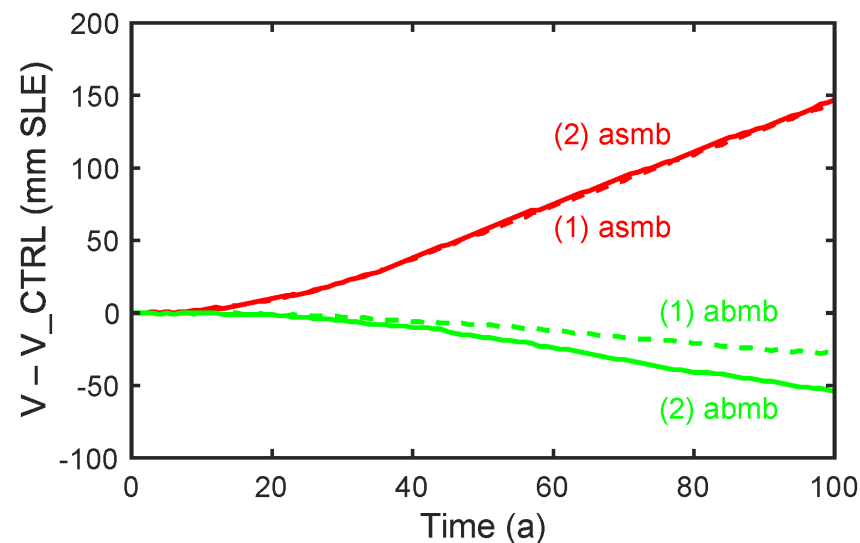
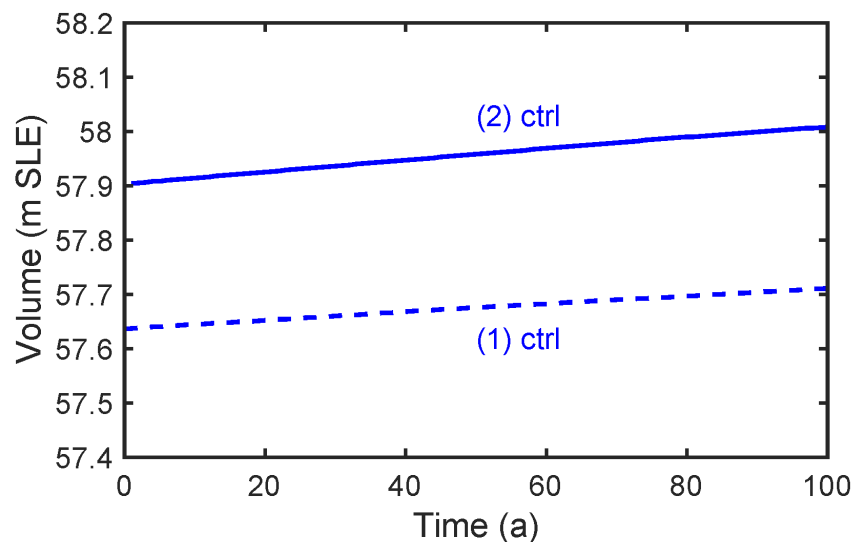


# Spin-up: simulated vs. observed ice thickness

With ice-stream dynamics:



# Simulated ice volume for future climate runs ctrl, asmb and abmb



- (1) Without ice-stream dynamics.
- (2) With ice-stream dynamics.

(1) vs. (2): Sensitivity of abmb notably different.

# Outlook

- Work in progress...
- Grid spacing: 16 km → 8 km.
- Generally too cold & too slow  
→ larger geothermal heat flux, tuning of basal sliding?
- InitMIP-Antarctica final submission deadline 31 July 2017  
→ ensemble of model results.
- Longer-term goal:  
Realistic climate change scenarios, input for IPCC AR6.



**Thank you**

Funding acknowledgement:

JSPS: ProGrIS project (KAKENHI Kiban A, No. 16H02224)

# **Appendix A: SICOPOLIS**

# Ice sheet model SICOPOLIS

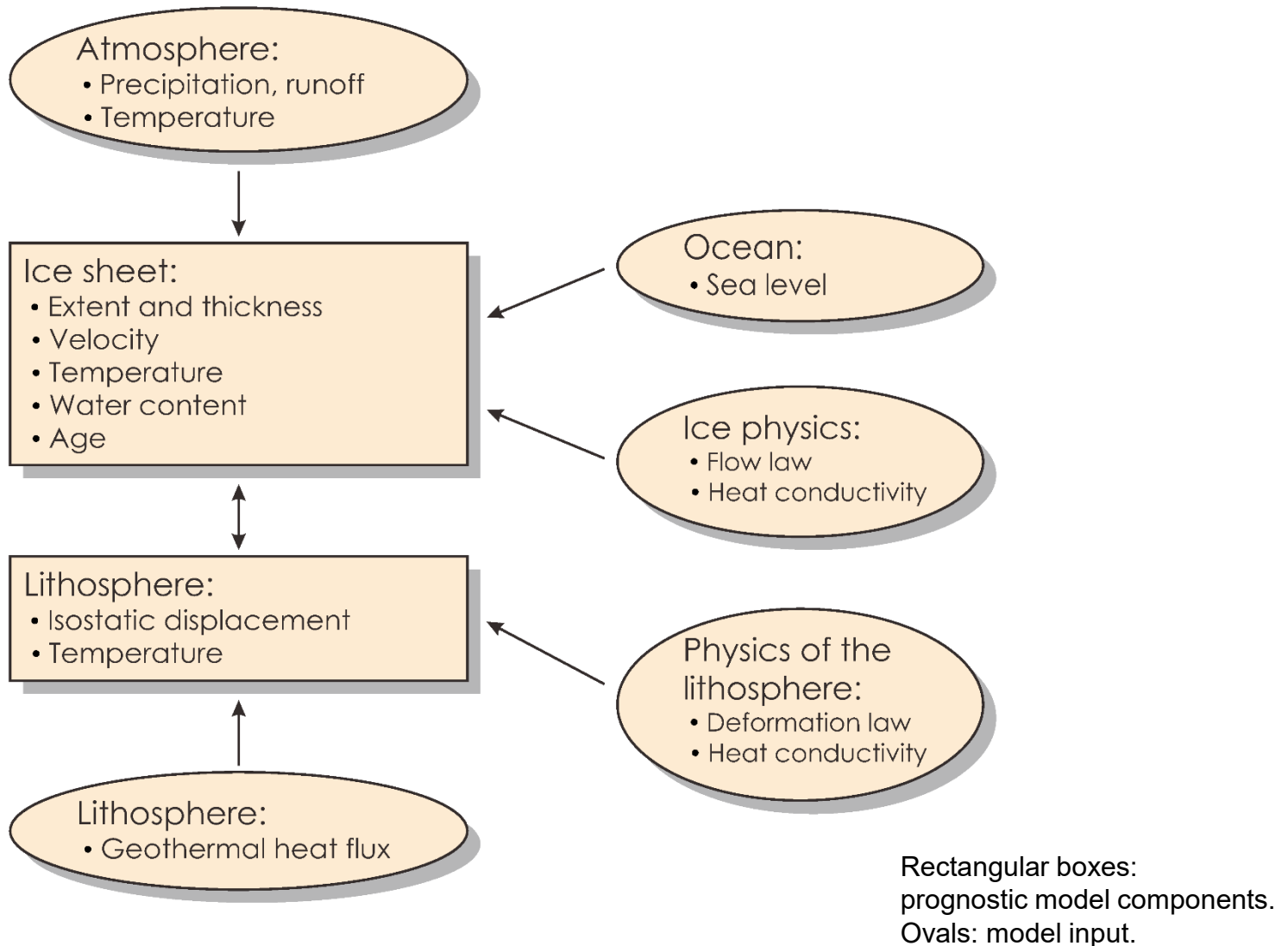
```
...
kc=kGMAX
kce=0
lgs_a0(kr) = ccb2
lgs_a1(kr) = -(ccb1+ccb2)
lgs_a2(kr) = ccb1
lgs_b(kr) = ccb3+ccb4
do kc=1, KGMAX-1
  lgs_a0(KGMAX+kc) = -0.5d0*(ct1(kc)-ct2(kc)-ct3(kc)-ct4(kc)) &
    -ct5(kc)*ct6(kc-1)
  lgs_a1(KGMAX+kc) = 1.0d0*(ct1(kc)*ct6(kc-1)
    -ct5(kc)*ct6(kc)) &
    -ct5(kc)*ct6(kc)
  lgs_a2(KGMAX+kc) = 0.5d0*(ct1(kc)-ct2(kc)-ct3(kc)-ct4(kc)) &
    -ct5(kc)*ct6(kc)
  !if ADV_HOB==1
    lgs_b(KGMAX+kc) = temp_c(kc,j,i) + ct7(kc) &
      -dti_2dxi* &
      ( vx_c(kc,j,i)-abs(vx_c(kc,j,i))) &
      *(temp_g11_gg(kc,j,i)
      +insq_g11_gg(kc,j,i)
      *(vx_c(kc,j,i-1)+abs(vx_c(kc,j,i-1))) &
      *(temp_c(kc,j,i)-temp_c(kc,j,i-1)) &
      *insq_g11_gg(kc,j,i-1) ) &
      -dti_2deiva* &
      ( vy_c(kc,j,i)-abs(vy_c(kc,j,i))) &
      *(temp_g22_gg(kc,j,i)
      +insq_g22_gg(kc,j,i)
      *(vy_c(kc,j,i-1)+abs(vy_c(kc,j,i-1))) &
      *(temp_c(kc,j,i)-temp_c(kc,j,i-1)) &
      *insq_g22_gg(kc,j,i-1) )
  !endif ADV_HOB==2
  lgs_b(KGMAX+kc) = temp_c(kc,j,i) + ct7(kc) &
    -dti_dxi*(ctx_c_r(kc)-ftx_c_l(kc)) &
    -dti_deva*(fvy_c_r(kc)-ftvy_c_l(kc))
  !endif
end do
kc=kGMAX
lgs_a0(KGMAX+kc) = 0.0d0
lgs_a1(KGMAX+kc) = 1.0d0
lgs_b(KGMAX+kc) = temp_s(j,i)
...
```

## “Simulation COde for POLythermal Ice Sheets”

- Open-source model, mainly developed at ILTS ([www.sicopolis.net](http://www.sicopolis.net)).
- Coded in Fortran.
- Shallow ice + shallow shelf approximations.
- Finite difference method.



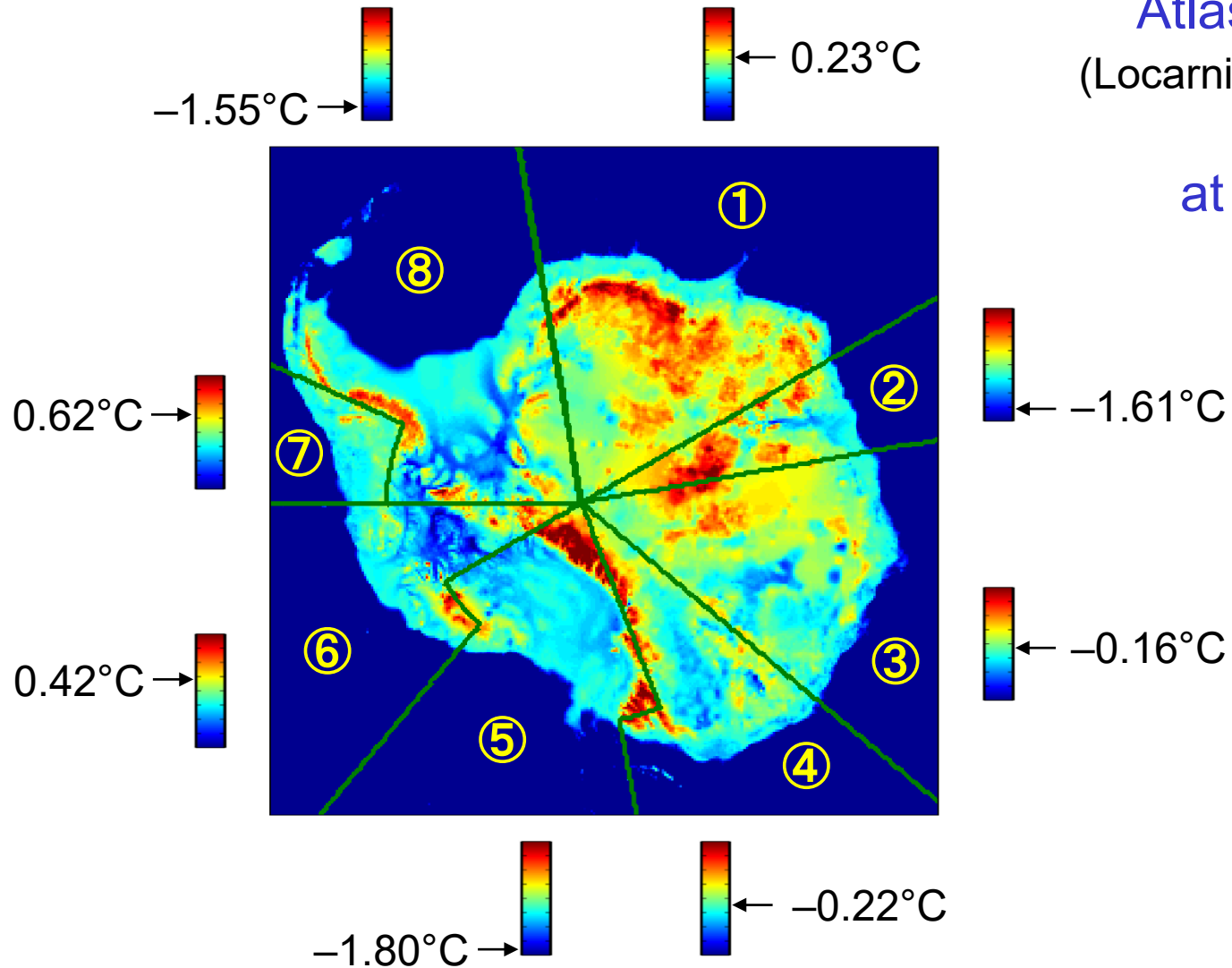
# Ice sheet model SICOPOLIS



## **Appendix B: Ice-shelf basal melting**

# Sector-averaged ocean temperature

[World Ocean Atlas 2009  
(Locarnini et al., 2010)  
at 500 m depth]



# Tuning strategy (for each sector)

(1) Target basal melting rate:

“observed”, from  $a_b = -\text{div}(H\mathbf{v}) + a_s$ .

(2) Assume  $\alpha$  (non-linearity parameter).

(3) Compute point-wise  $\Omega$  (sensitivity parameter) with nearest-neighbor  $T_{oc}$  (ocean temperature).

(4) Set up the parameterisation for  $a_b$  with sector-averaged  $\Omega$  and  $T_{oc}$ .

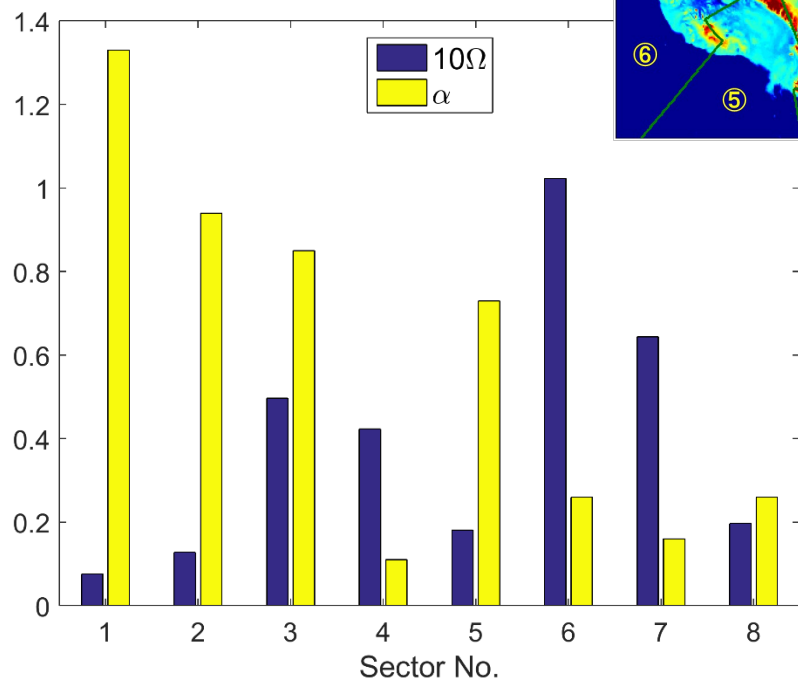
(5) Compute misfit to “observed”  $a_b$ .

(6) Repeat with different  $\alpha$ 's  $\rightarrow$  minimize misfit.

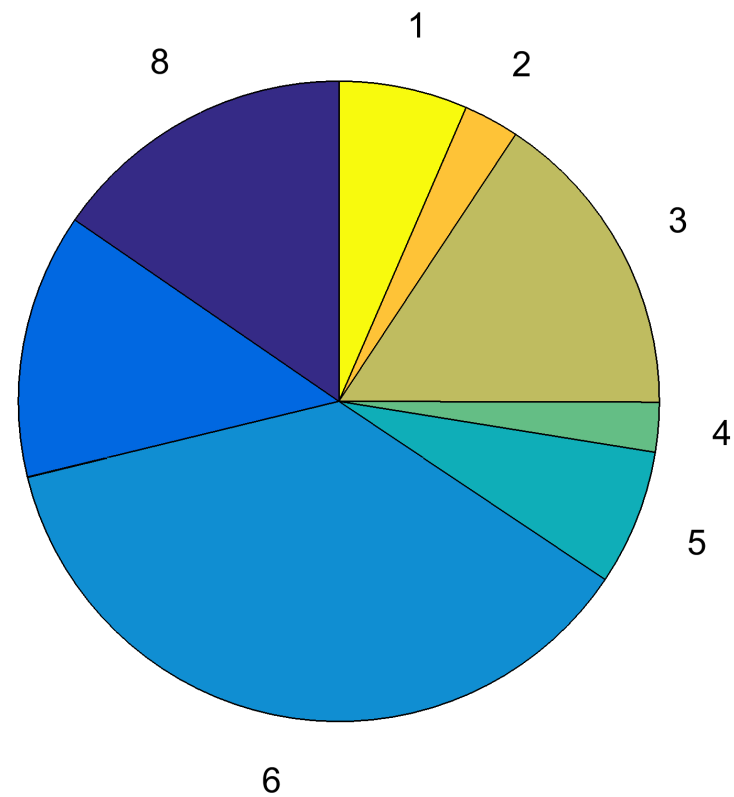
(7) Scale  $\Omega$ 's  $\rightarrow$  sectorial and total melt by Rignot et al. (2013)

# $\Omega$ 's, $\alpha$ 's and melt rates

$\Omega$ 's,  $\alpha$ 's



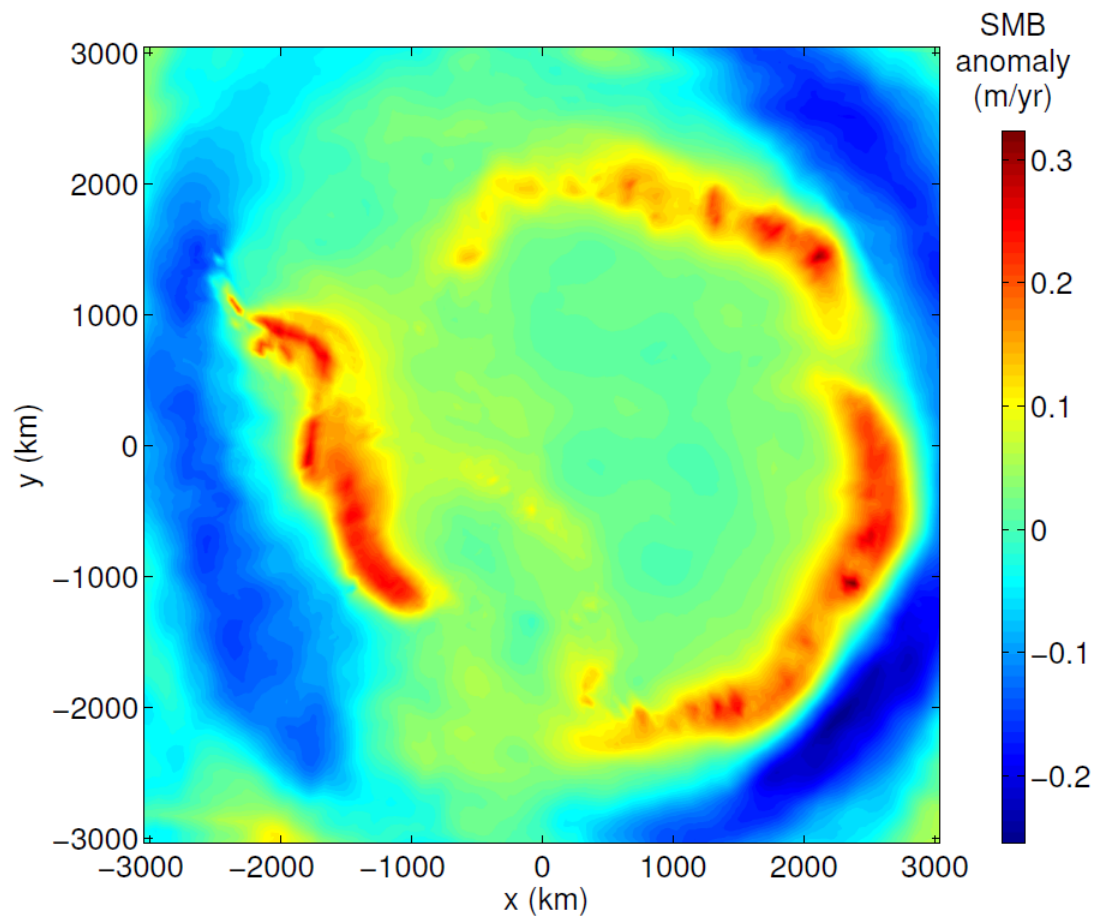
Sectorial melt rates



$$Q_{\text{tot}} = 1323.9 \text{ Gt/a}$$

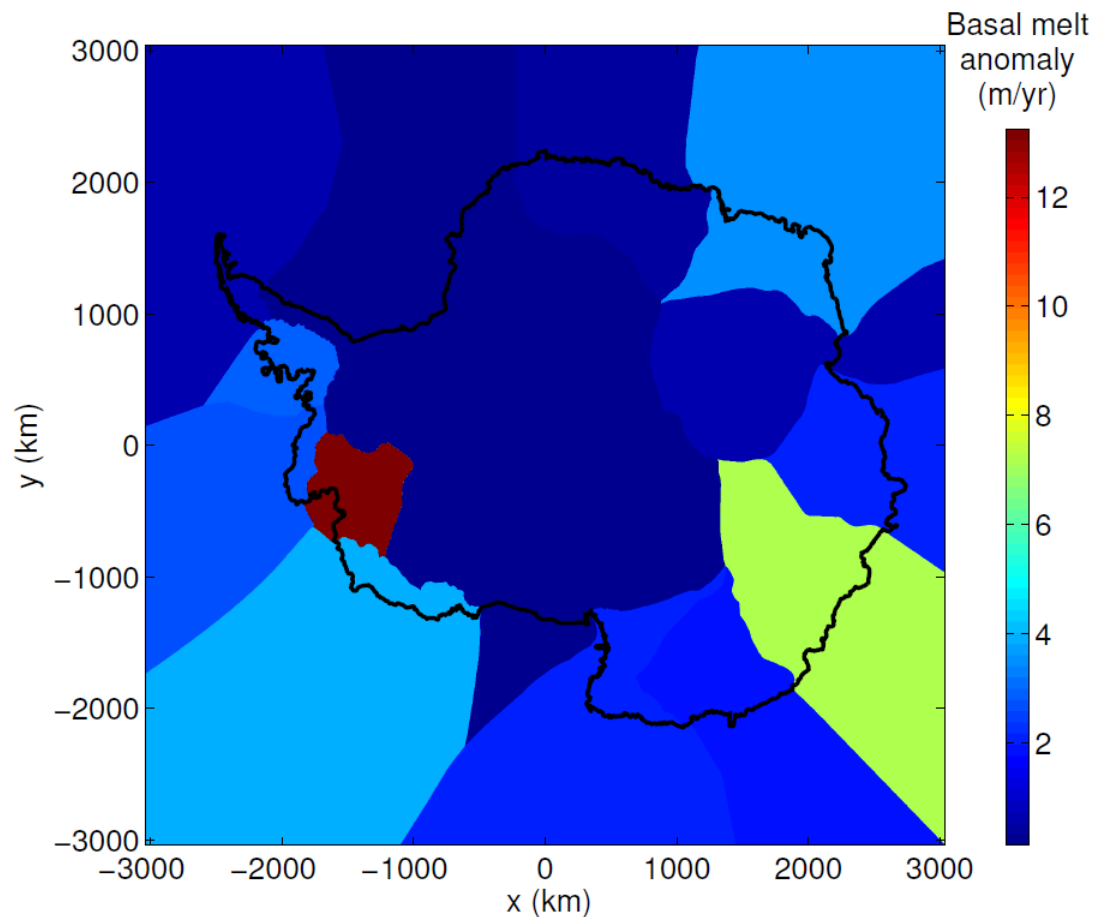
## **Appendix C: InitMIP anomalies**

# Schematic SMB anomaly (asmb)



(Seroussi et al., in prep.)

# Schematic basal melt anomaly (abmb)



(Seroussi et al., in prep.)



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