

Carbon cycle response to temperature overshoot beyond 2 °C – an analysis of CMIP6 models

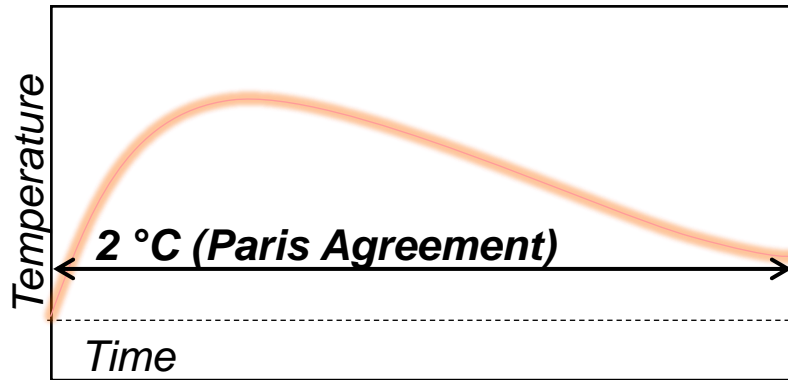
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Motivations for Achieving the Paris Agreement Temperature Targets after Overshoot (PRATO) project



Gap

2030

- **Science (IPCC SR15):** Net anthropogenic CO₂ emissions need to decline by 2030 up to 45% (25%) from 2010 levels for 1.5 °C (2 °C) target
- **National policies:** NDCs indicate a slight increase in CO₂ emissions from 2020 levels (emission pledges need to be more than halved (Höhne et al., 2020))

Converging?

2050

- **Science (IPCC SR15):** Net zero CO₂ emissions required to be on track for 1.5 °C (or by 2070 for 2 °C)
- **National policies:** Major economy countries adopted "net zero" targets by 2050-2060 (December 2020, Climate Action Tracker)
- **Hausfather & Peters, 2020 :** global warming is on course to 3 °C by the end of this century.

End of the 21st century

- **Paris Agreement**
- To limit warming below 2 °C (1.5 °C)

Overshooting the temperature target is becoming more likely

Background studies

Study	Model	Design	Results
Boucher <i>et al.</i> , 2012, Zickfield <i>et al.</i> , 2016	HadGEM2-ES, Uvic ESCM	572-1144 ppm CO ₂ , 2.0 to 5.5 °C warming (<i>idealized</i>)	Ocean becomes a source in a few decades, land is sink for long time
Tokarska <i>et al.</i> , 2019	UVic ESCM	480-620 ppm CO ₂ , 1.5 to 2.5 °C warming (<i>idealized</i>)	Ocean and land are decreasing sinks after ramp-down, more carbon is stored in the ocean
Palter <i>et al.</i> , 2018	GFDL-ESM2M	>580 ppm, ≈2 °C warming	Land and ocean become a carbon source nearly two to three decades after the peaks of CO ₂ concentration and temperature
Schwinger & Tjiputra, 2018	NorESM	>800 ppm, ≈4 °C warming (<i>idealized</i>)	Carbon source only after a strong reduction in atmospheric CO ₂

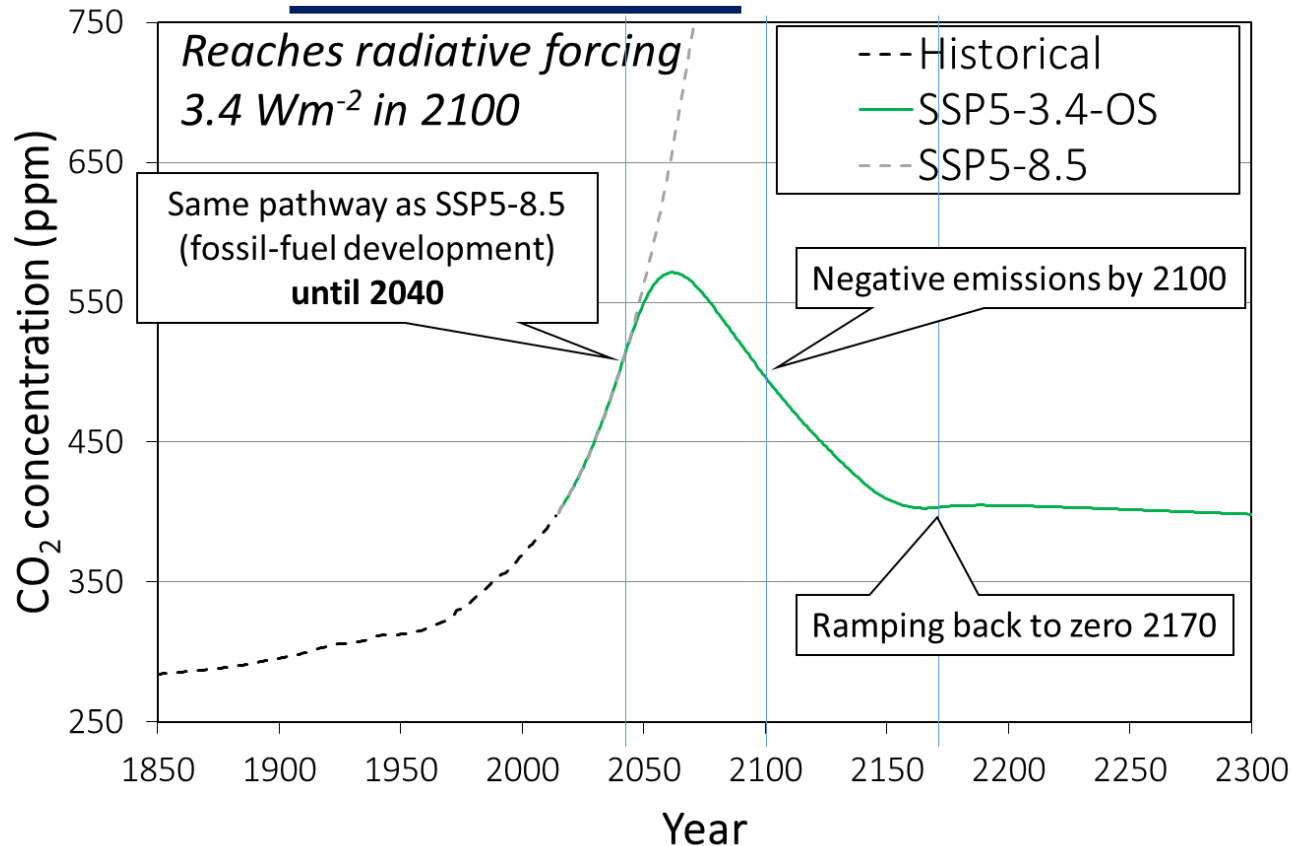
- Different, highly idealized scenarios
- Limited number of models, sometimes of intermediate complexity

What controls the response of the carbon cycle to the CO₂ and temperature overshoots is not well-known

Study purpose and methods

To explore the carbon cycle feedbacks over land and ocean under overshoot scenario

SSP5-3.4-OS



CMIP6 ESMs

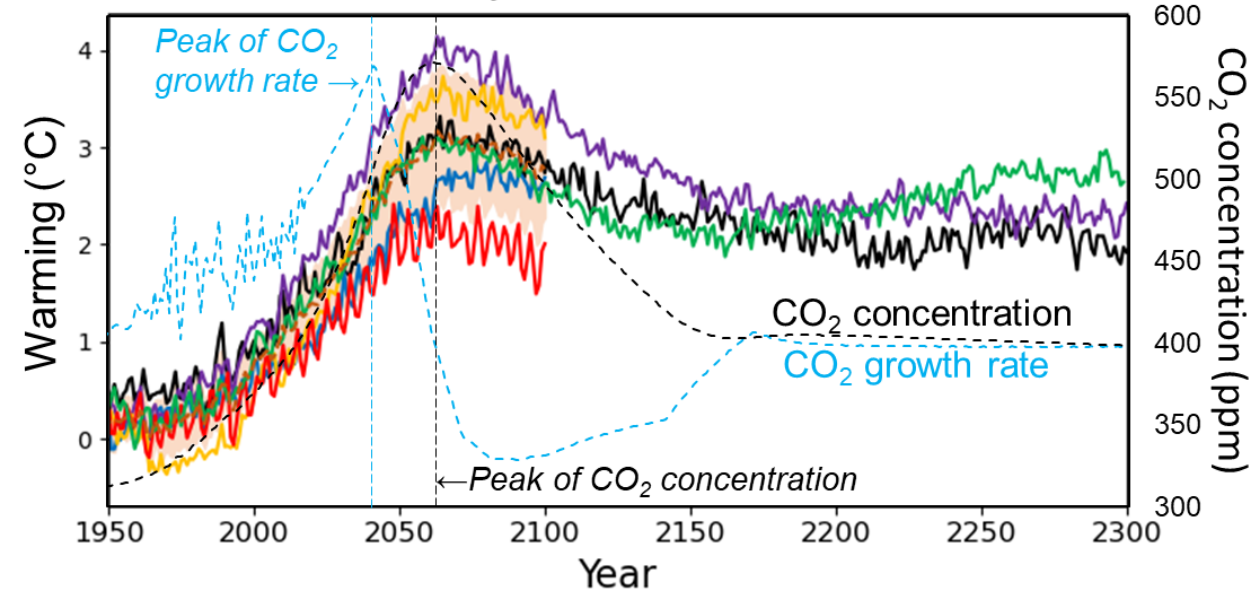
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- 1) IPSL-CM6A-LR
- 2) CNRM-ESM2-1
- 3) CanESM5
- 4) UKESM1-0-LL
- 5) MIROC-ES2L
- 6) CESM2-WACCM

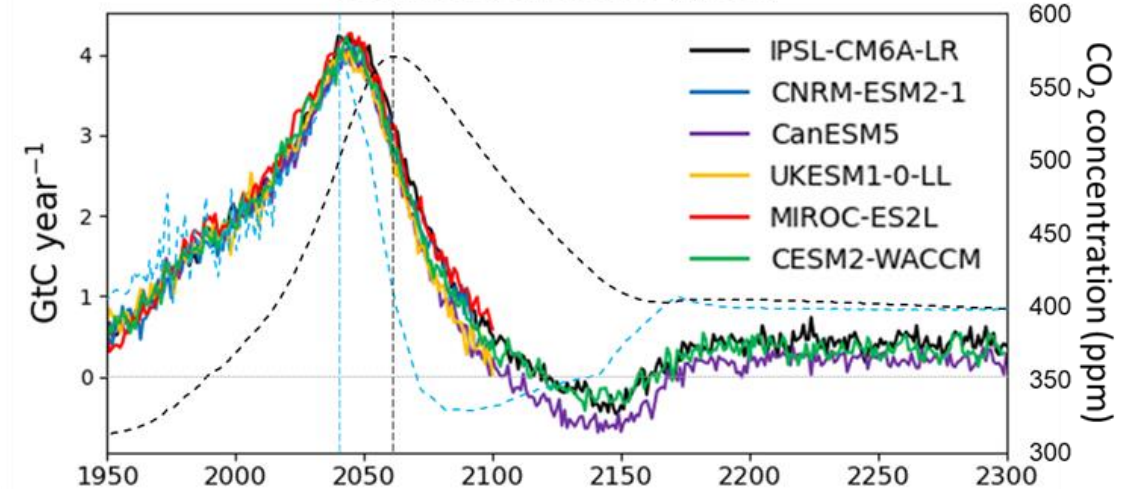
- We investigate temporal changes in carbon fluxes
- We apply **the carbon cycle feedback framework** to quantify the carbon cycle feedbacks

Land and ocean carbon uptakes

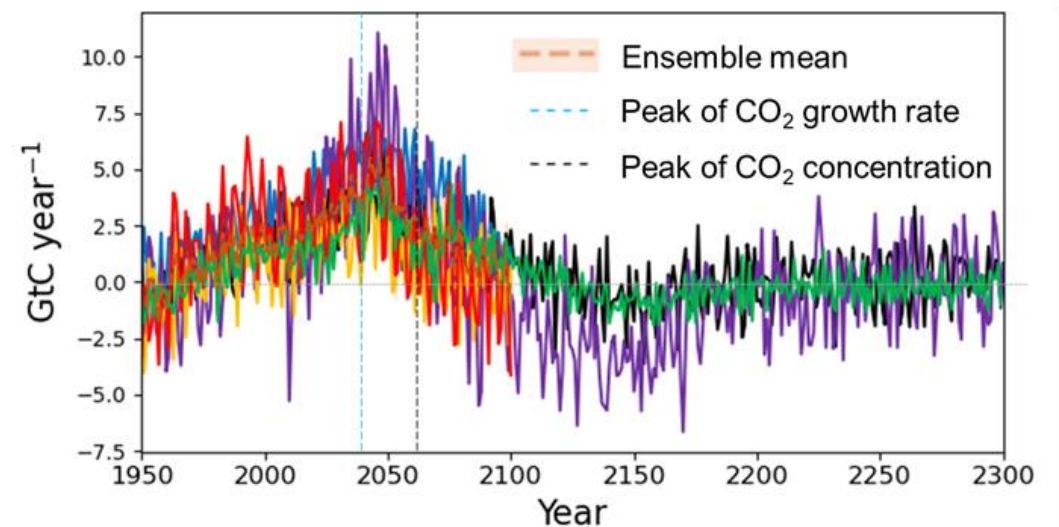
GSAT anomaly relative to 1850-1899



Ocean carbon uptake (fgco2)

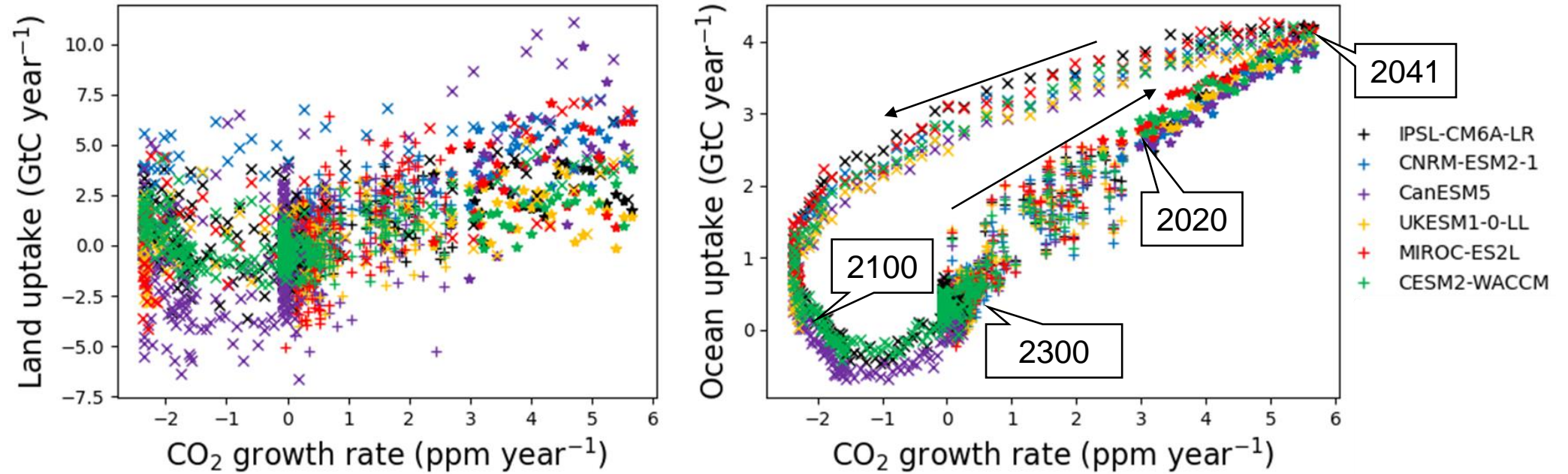


Land carbon uptake (NBP)



- CO₂ concentration peaks **two decades** after CO₂ growth rate
- ESMS differ in the response of **temperature** to the prescribed changes in CO₂ concentration, but all agree on a **delayed response**
- After the peak of CO₂ growth rate, **land and ocean take up carbon for at least 50 years** but at a declining rate
- In the 23rd century, **ocean is a weak carbon sink**, the response of **land** has a larger range of **uncertainties**
- Land and ocean carbon uptakes **peak before the CO₂ concentration peak**

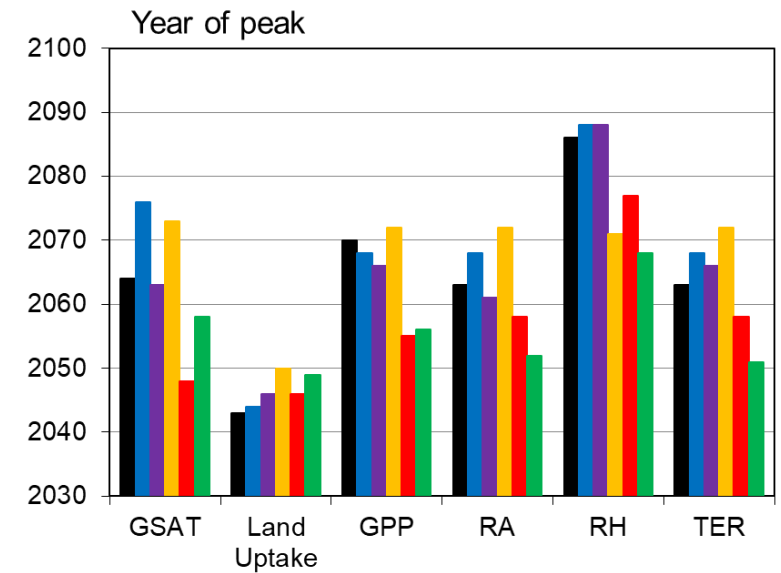
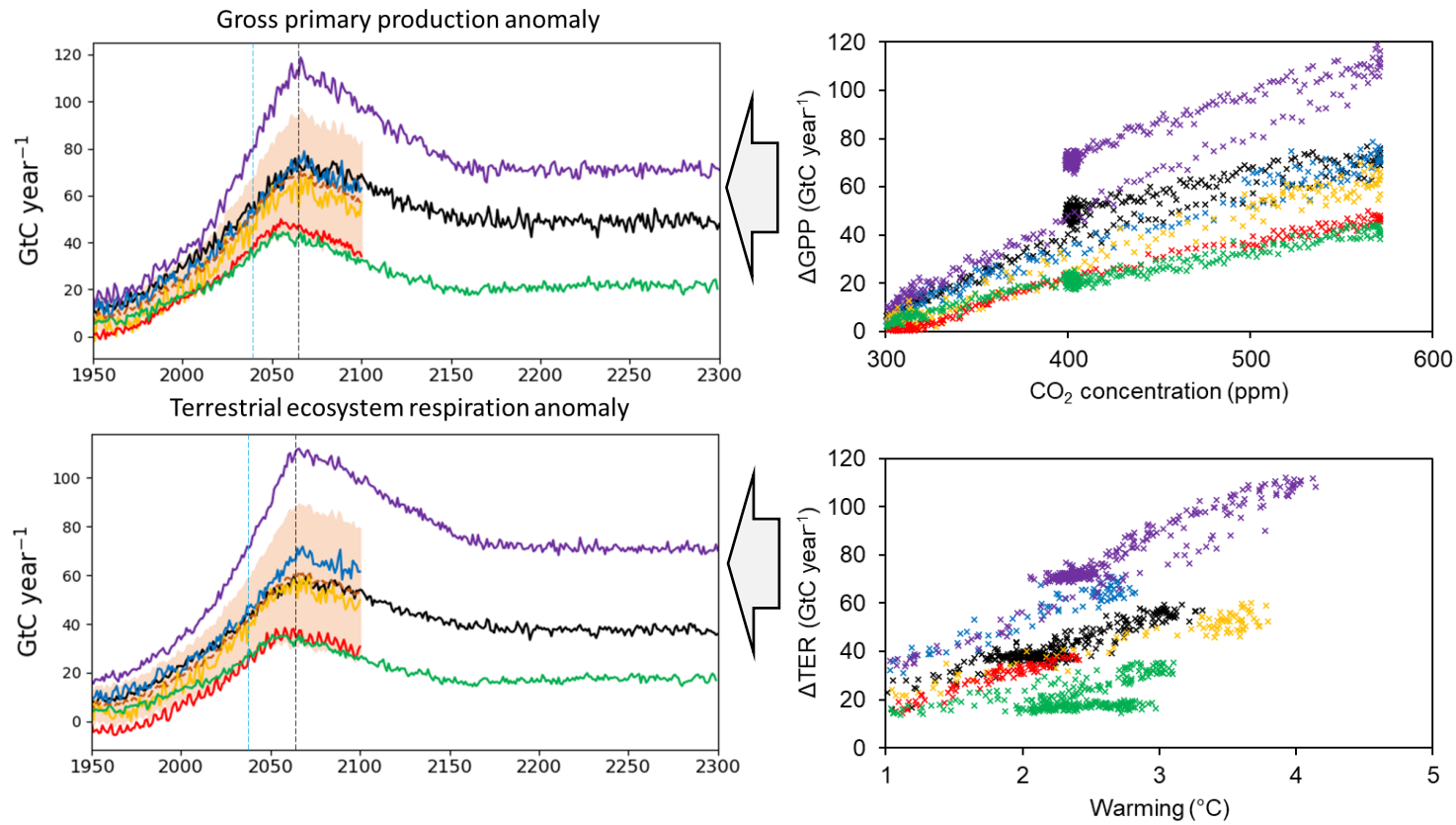
Peak of ocean carbon uptake



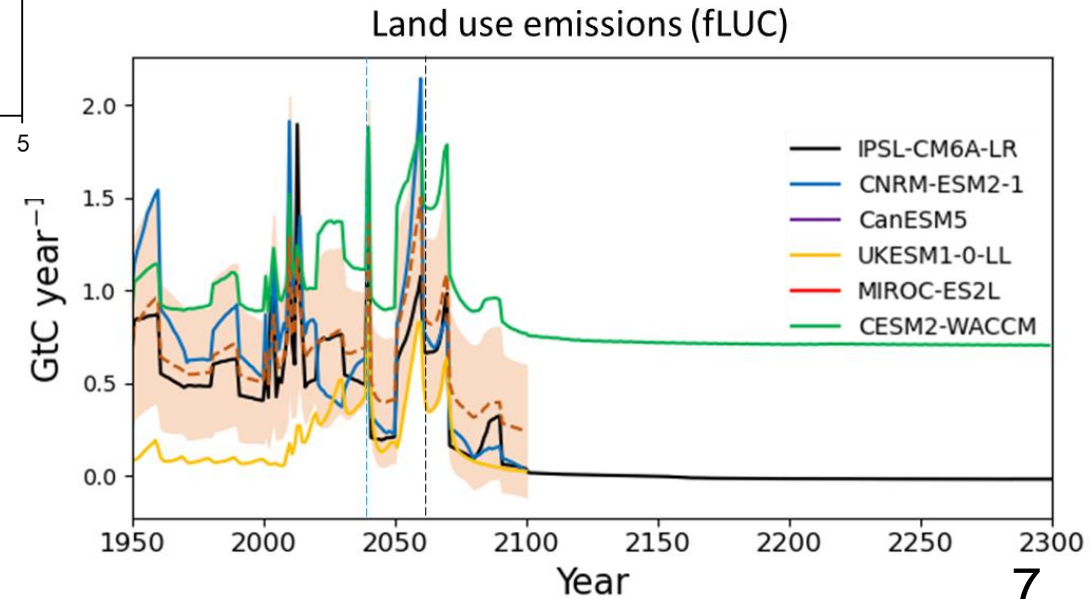
Global ocean carbon uptake has a nearly linear dependency on the atmospheric CO₂ growth rate with a hysteresis after the CO₂ growth rate peak

- McKinley et al. (2020) and Schwinger and Tjiputra (2018): the CO₂ growth rate dominates the variability in the global ocean on year-to-year timescales
- Additional analysis on the underlying reasons is needed
- A weaker linear dependency of the land carbon uptake on the CO₂ growth rate

Peak of land carbon uptake



The land sink decrease is due to the **larger growth of respiration than photosynthetic production and concurrent land-use change emissions**



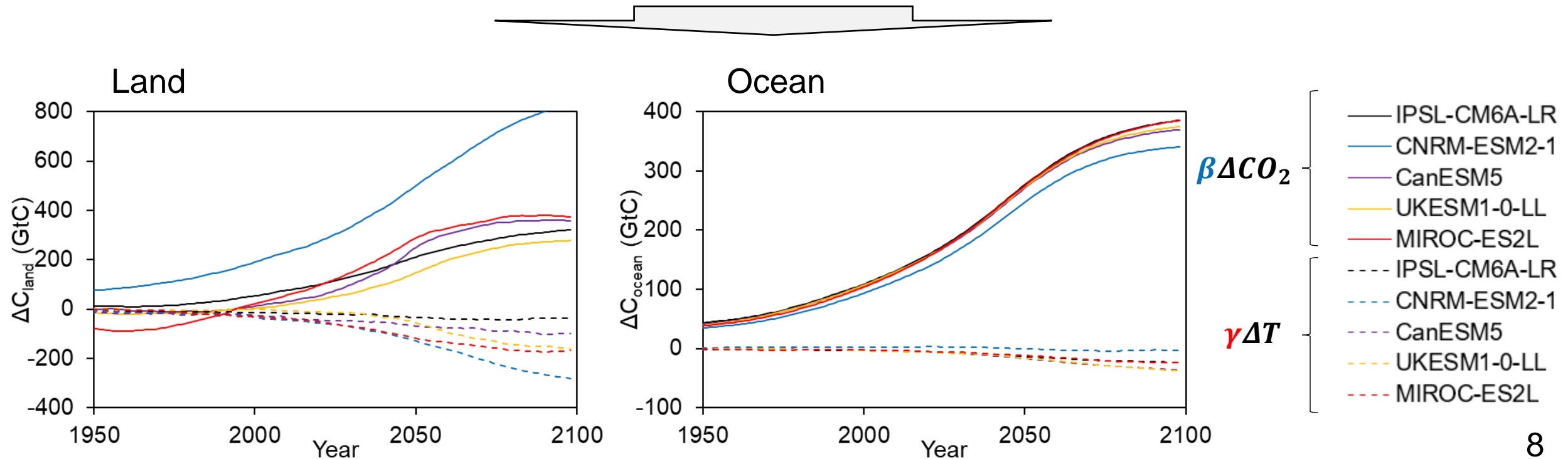
Carbon cycle feedback framework (SSP5-3.4-OS)

$$\Delta C = \beta \Delta CO_2 + \gamma \Delta T$$

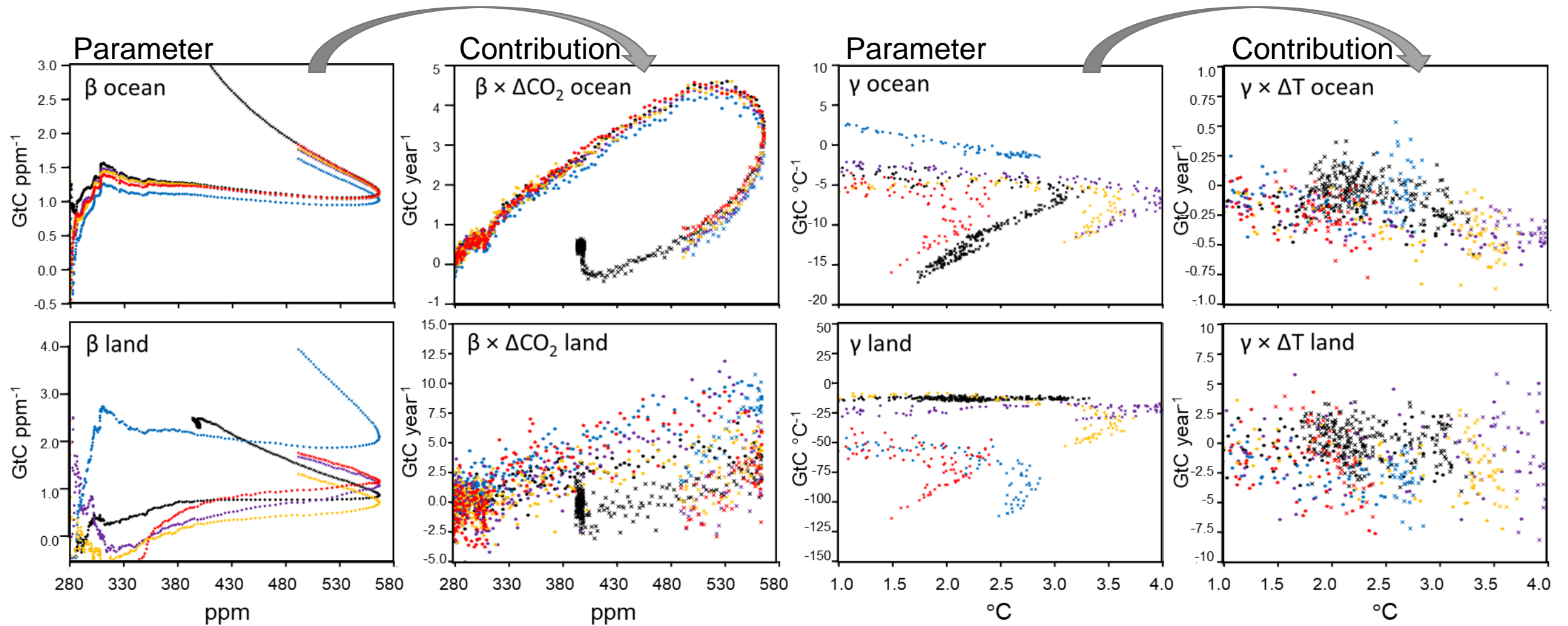
Beta, β : carbon-concentration feedback ($\Delta C \text{ ppm}^{-1}$)
Gamma, γ : carbon-climate feedback: ($\Delta C \text{ }^\circ\text{C}^{-1}$)

ΔC , ΔCO_2 , ΔT are the changes of carbon mass (GtC), CO_2 concentration (ppm) and air temperature ($^\circ C$) relative to their unperturbed preindustrial states

Gregory *et al.* 2009, Jones *et al.*, 2016



Carbon cycle feedbacks (SSP5-3.4-OS)



- Under declining emissions, both land and ocean continue to take up (via β) and lose carbon (via γ) at an asymmetrically larger rate
- The changes in the carbon cycle are dominated by β rather than γ till the end of the 21st century.

Conclusions

- In SSP5-3.4-OS, the land and ocean continue to remove carbon from the atmosphere at least for 50 years after the peak of the CO₂ growth rate. They turn to a source in the first half of the 22nd century for a short period and become a weak sink later, i.e., reach a new steady-state
- The ocean uptake decreases due to its dependency on the CO₂ growth rate
- The land uptake decreases due to a larger growth of respiration than GPP and concurrent LUC emissions
- The inertia in β and γ after their peaks causes amplification of carbon cycle feedback parameters.
- The total β carbon gain $\beta \times \Delta\text{CO}_2$ is larger than γ loss $\gamma \times \Delta T$, and, thus, land and ocean continue to uptake carbon during the CO₂ and temperature ramp-down periods.