

# Impacts of regional Arctic sea ice loss and the role of QBO

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Climate Dynamics Workshop, 29.06.20



European Union Horizon 2020 research and innovation program, Grant Number 727862



# Motivation

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• Arctic Sea Ice Loss  $\rightarrow$  easterly wind response

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Advanced prediction in polar regions and beyond

- Response uncertainty, Model inconsistency
- Government question (BEIS and Defra): How might response counteract climate change?

→ PAMIP: Polar Amplification Model Inter-comparison Project Led by Met Office – Doug Smith.

Rosie - PhD with Exeter University on decadal variability of the NAO and potential drivers such as Arctic sea ice.





Future Sea Ice loss



Surface temperature



- Winter (DJF): ice loss around edges
- Warming "hotspots" over ice loss regions
- Warming spreads across Arctic and to lower latitudes
- No significant Eurasian cooling

- Multi-Model International Collaboration
- Atmosphere only models
- 18 models
- 100+ ensemble members per model
- Stippled where 90% (black) or 80% (grey) of models agree on sign of response



Surface temperature

Sea level pressure





- Low pressure over warming "hotspots" → thermodynamic "heat low"
- High pressure Greenland to Siberia  $\rightarrow$  dynamical response





- Low pressure over warming "hotspots" → thermodynamic "heat low"
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- Equatorward shift of storm tracks (Atlantic and Pacific)

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- Zonal mean winds
- Robust response in troposphere
- Not robust in stratosphere
- What is the physical mechanism?
- What causes the model spread?
- What is the real world response?





























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#### Zonal wind



- Robust equatorward shift of tropospheric zonal winds
- Less robust in stratosphere  $\rightarrow$  not crucial for the mechanism

#### Temperature/winds



- Arctic warming near surface
- Upper troposphere warming not robust
- Weakening of polar (thermally direct) and Ferrel (wave driven) cells



-2.4

-1.2

0.0

### Eliassen-Palm fluxes: climatology

Upward EP flux (UEP) Northward EP flux (NEP) Pressure Level (hPa) Pressure Level (hPa) 01 2 10<sup>2</sup> 10 103 30 10 20 40 50 60 70 80 20 10 30 40 50 60 Lat. Lat.

2.4

Waves are generated at the surface by baroclinic eddies (positive upward EP flux) ٠

- Propagate upward and refracted equatorward (negative northward EP flux) •
- Flux momentum into source regions (divergence of EP flux) ٠

1.2

Drag the flow where they dissipate (convergence of EP flux)

70 80

40

-40





Correlation across models of zonal wind dipole response and divergence of NEP response

50

60 70 80

0.9

0.3

40

Lat.

- Upward EP flux reduced at 60-70°N, increased at 40-50°N
- Northward EP flux increased at 40-60°N in mid to upper troposphere
- Divergence (convergence) of northward EP flux at 30-45°N (50-70°N)
- Drives equatorward shift of zonal winds

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• Divergence of NEP response explains model spread in zonal wind response



Upward EP flux

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Correlation across models of zonal wind dipole response and refractive index response



- Arctic warming  $\rightarrow$  reduced equator to pole T gradient  $\rightarrow$  weaker storm track  $\rightarrow$  upward EP flux reduced at 60-70°N
- Increased upward EP at 40-50°N is a feedback:
  - > Storm track (wave source) shifts equatorward
  - Reinforced by refractive index response at ~40°N

Stratosphere response to sea ice loss

Correlation across models of stratospheric jet response and background zonal wind

0

Lat.

-0.3

APPLI

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τv

Pressure Level (hPa) 01 20

103

-50

-0.9



Polar jet response vs QBO (r=0.45)

- Stratospheric jet response appears to depend on background QBO
- Clear in HadGEM3 but not in IPSL

0.3

50

0.9





- Increased upward EP flux when QBO easterly
- Consistent with Labe et al 2019
- Mechanism to be explored...

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## Real world response to sea ice loss?

Zonal wind and NAO responses



- Model response is weak relative to interannual variability
- Mean NAO response ~0.1 σ
- Max NAO response ~ 0.5  $\sigma$
- But models may underestimate true response (signal-to-noise paradox)
- Need an "emergent constraint"
- · Difficult due to large uncertainties

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## Conclusions

- Robust tropospheric response to future Arctic sea ice loss
  - > Equatorward shift of jet, negative NAO
- Weakening of polar (thermally direct) and Ferrel (wave driven) cells
- Stratospheric response appears to depend on QBO
- Weak signal relative to internal variability
- Large model spread (0.1 to 0.5  $\sigma$ )
- Large uncertainties  $\rightarrow$  diagnosing real world response is challenging