

PROBLEM N°17

Balancing Pebble



Ecole polytechnique

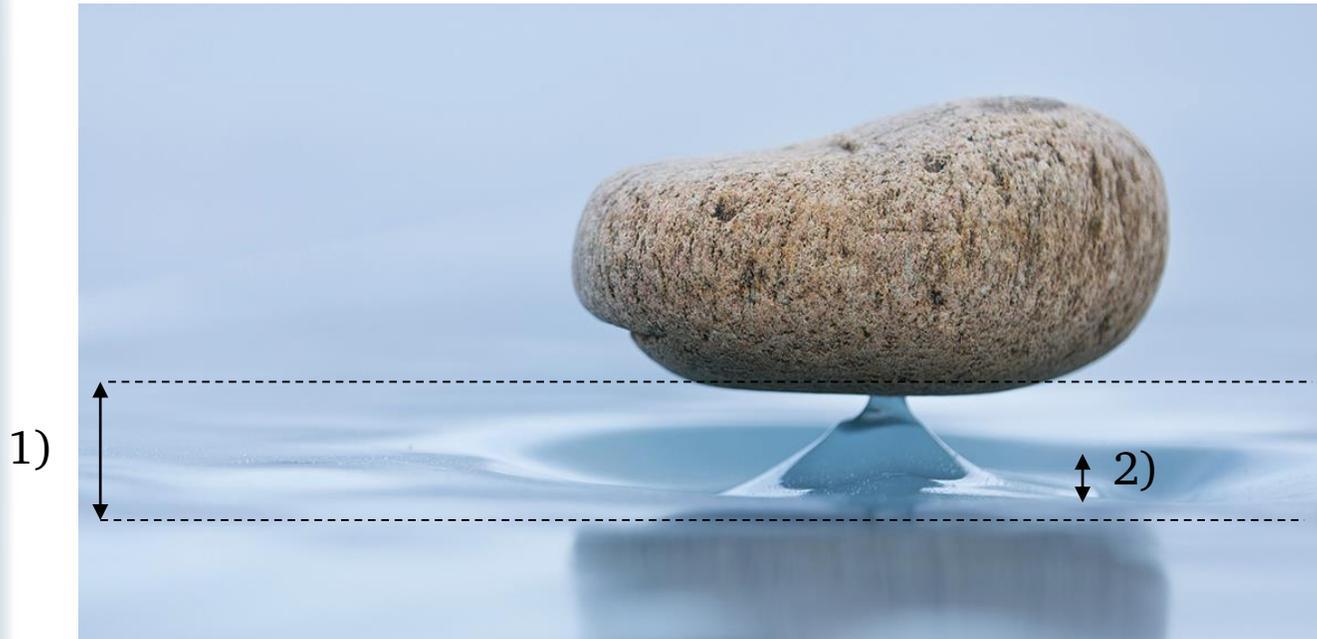
Balancing Pebble



Stones which are taken by wind on the ice of Baikal Lake can be found after some time staying on a thin «stand».

- **Reproduce** and **explain** this «stand» phenomenon
- Estimate **the curve** of the stand depending on the important parameters.

What characterizes a “Balancing Pebble” ?

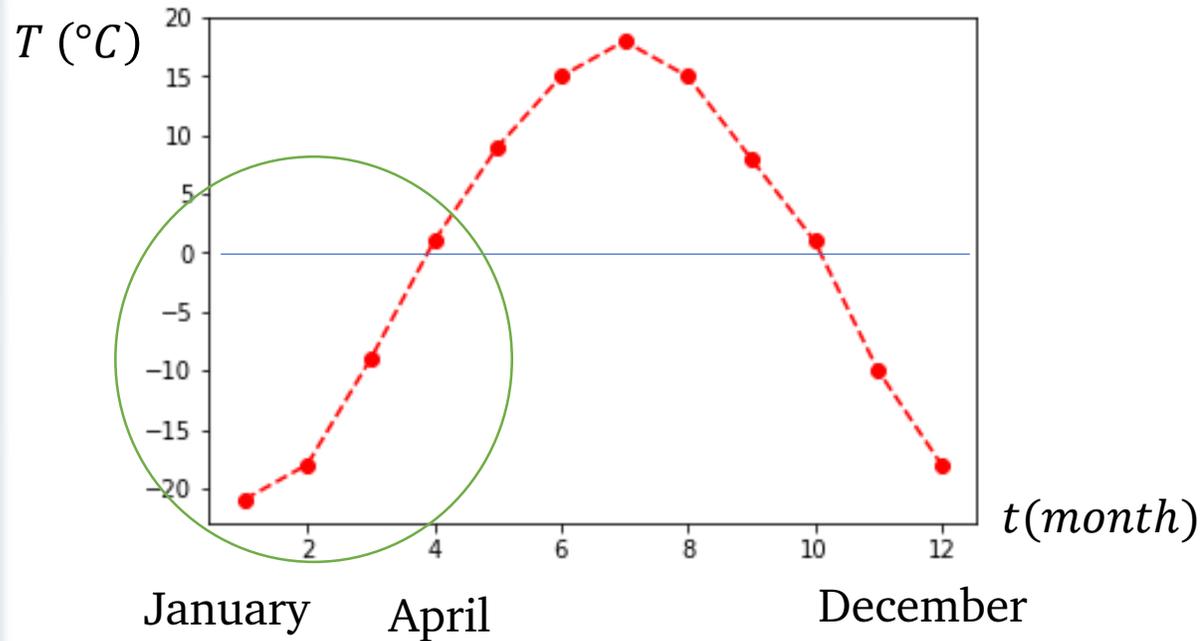


- 1) Stone stands above the lake
- 2) Ice lower just around the stone
- 3) Asymmetry possible

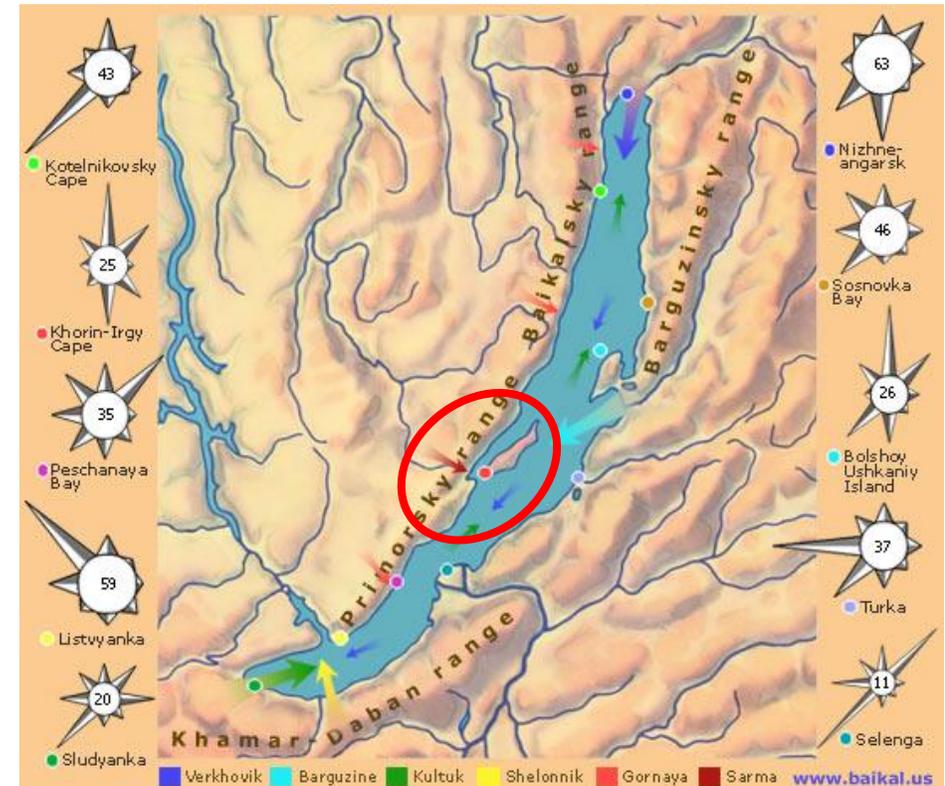


Context on Lake Baikal, Siberia

Temperature

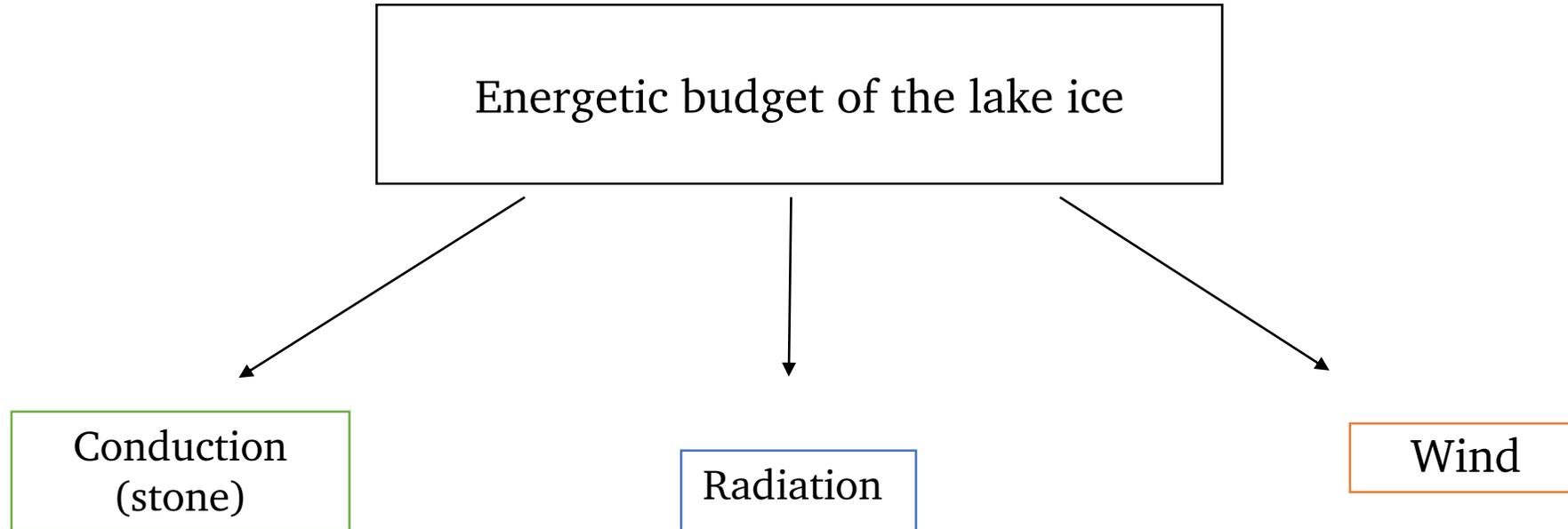


Wind





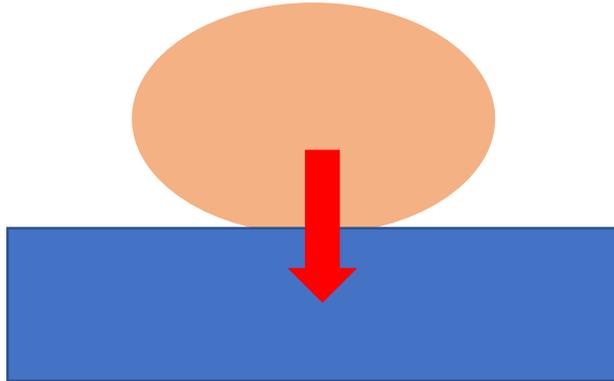
Sources of energy



Formation of a stand ?



1st experiment : Conduction



- 1) Stone stands above the lake X
- 2) Ice lower just around the stone X
- 3) Asymmetry possible X



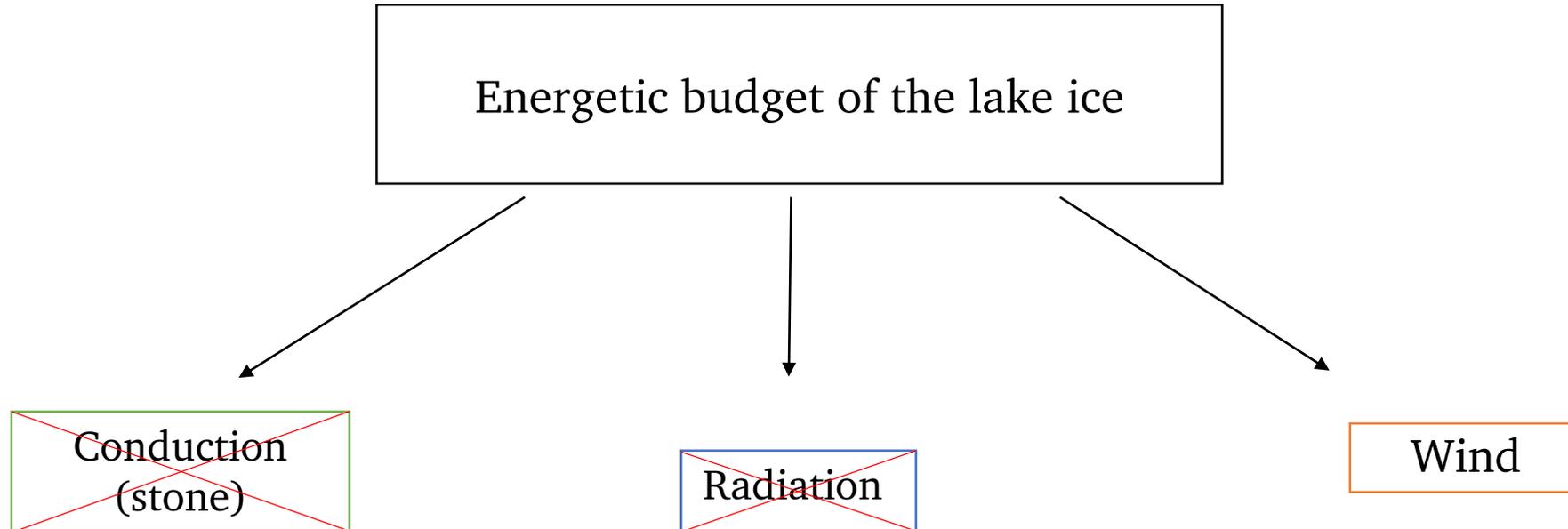
Radiation ?



Not needed! Balancing pebbles found in a cave

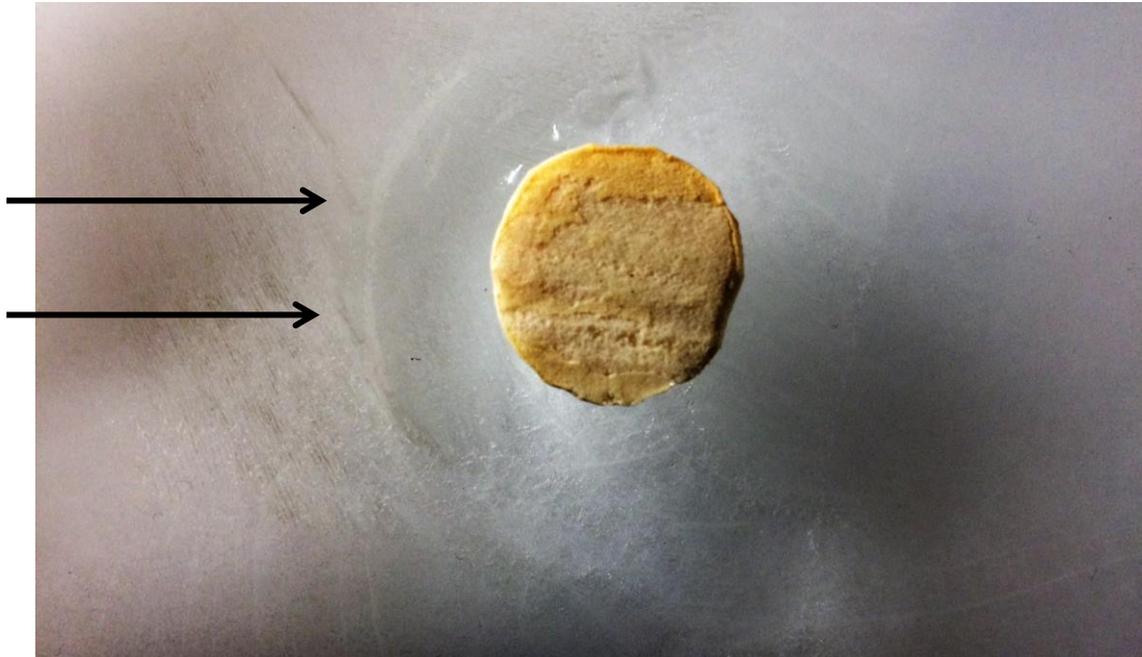


Why does ice melt ?





Experiment : blowing hot wind



A stand
Asymmetric pattern



Experiment : with turning wind !

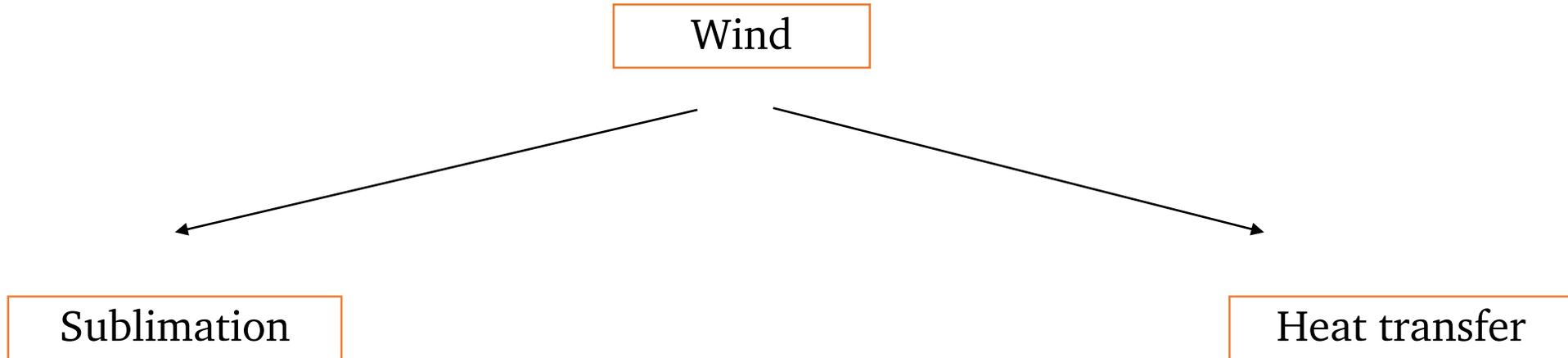


- 1) Stone stands above the lake
- 2) Ice lower just around the stone
- 3) Asymmetry possible





Why does ice melt ?





Why does ice melt ?

Wind

Sublimation

Heat transfer

Origin: Vapor pressure gradient

Origin: Temperature gradient

$$\Phi_{sublimation} = h_m(\mathbf{u}) (n_{sat}(T) - n_{wind}) \quad \sim$$

n_{sat}, n_{wind} : number of water molecules per m^3 of air

$$\Phi_{heat} = h(\mathbf{u}) (T_{wind} - T_{ice})$$

h_m : mass transfer coefficient

h : heat transfer coefficient

Conditions of temperature and humidity for Baikal lake:

Ablation rate of $2 \cdot 10^{-5} \text{ kg/s}$ for sublimation \gg Ablation rate of $2 \cdot 10^{-8} \text{ kg/s}$ for melting through heat transfer



How does ice melt ?

Modelling transfer phenomena

Newtons law of convective heating:

$$\phi_{heat} = h(\mathbf{u})(T_{wind} - T)$$

Heat transfer equation:

$$\rho c \partial_t T = h(\mathbf{u})(T_{wind} - T)$$

$$T = T_i + (T_{wind} - T_i) \left(1 - e^{-\frac{h(\mathbf{u})}{\rho c} t}\right)$$

Convective mass transfer:

$$\phi_{sub} = h_m(\mathbf{u})(n_{wind} - n)$$

Convective mass transfer equation:

$$\partial_t n = h_m(\mathbf{u})(n_{wind} - n)$$

$$n = n_i + (n_{wind} - n_i) \left(1 - e^{-h_m(\mathbf{u}) t}\right)$$

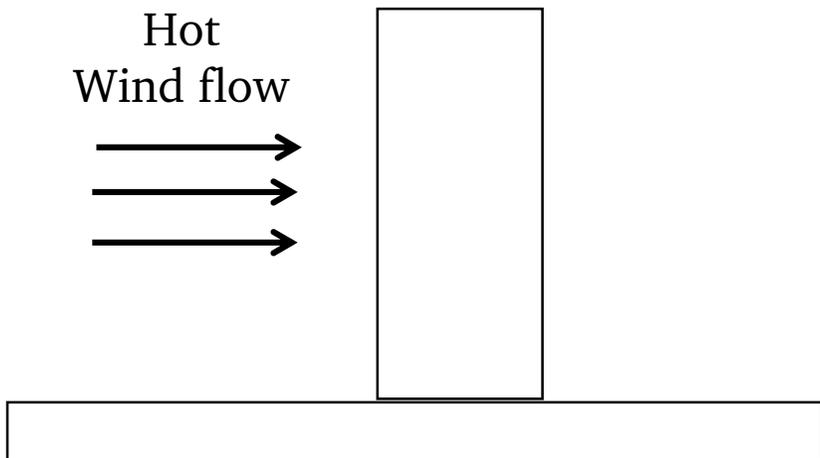
Same exponential laws



How does ice melt ?

Modelling the heat transfer $h(u)$

$$T = T_i + (T_f - T_i) \left(1 - e^{-\frac{h(u)}{\rho c} t}\right)$$





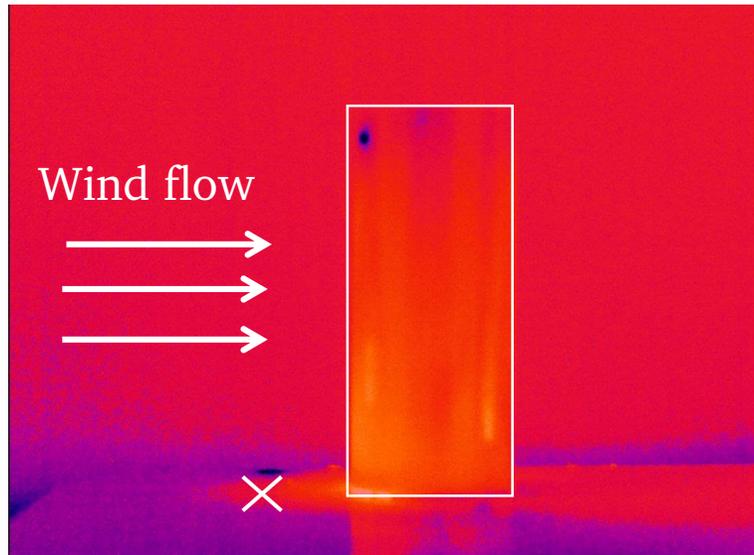
How does ice melt ?

Modelling the heat transfer $h(u)$

Newton's law gives :

$$T = T_i + (T_f - T_i) \left(1 - e^{-\frac{h(u)}{\rho c} t} \right)$$

Thermal camera from 20°C to 35°C





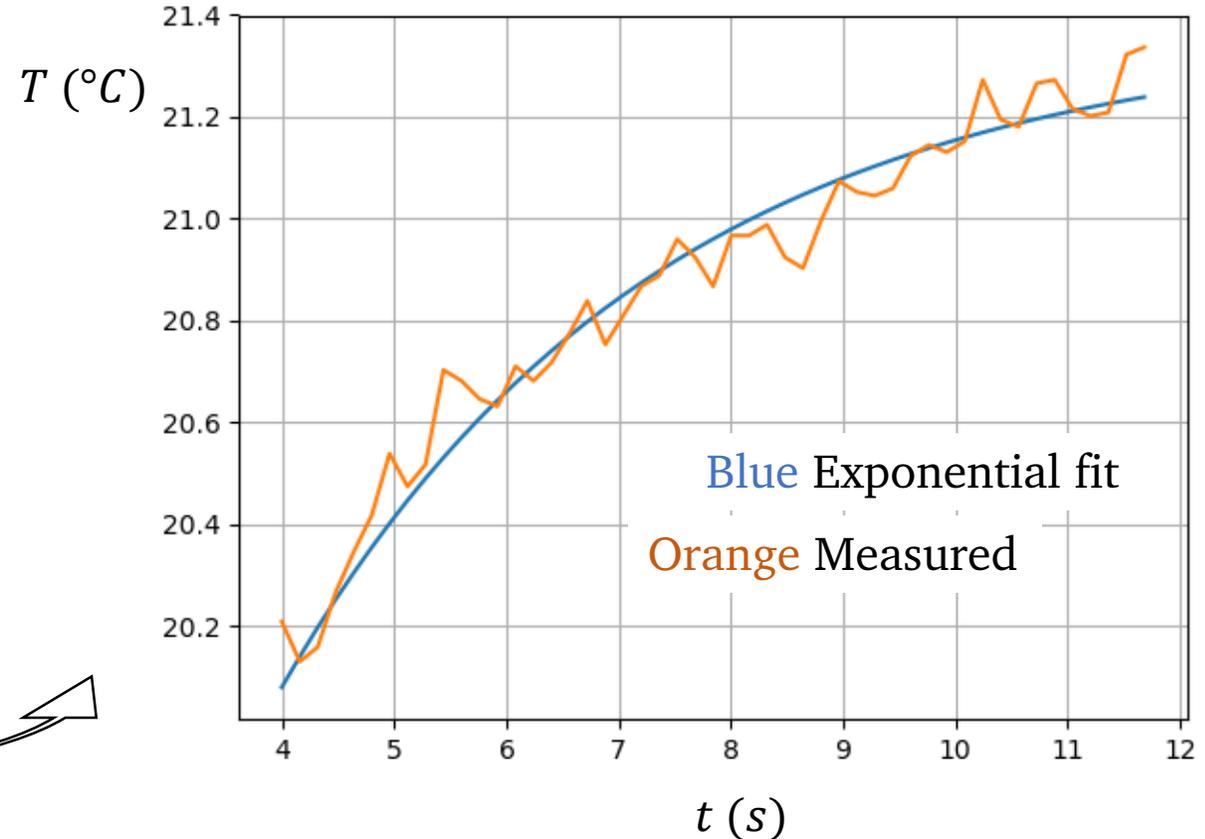
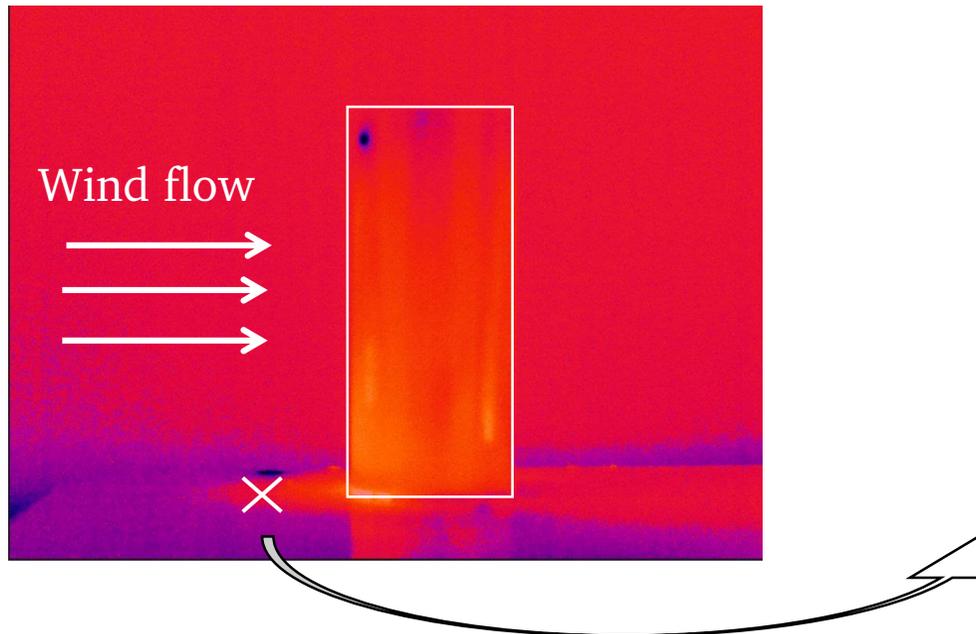
How does ice melt ?

Modelling the heat transfer $h(u)$

Newton's law gives :

$$T = T_i + (T_f - T_i)(1 - e^{-\frac{h(u)}{\rho c}t})$$

Thermal camera from 18°C to 25°C





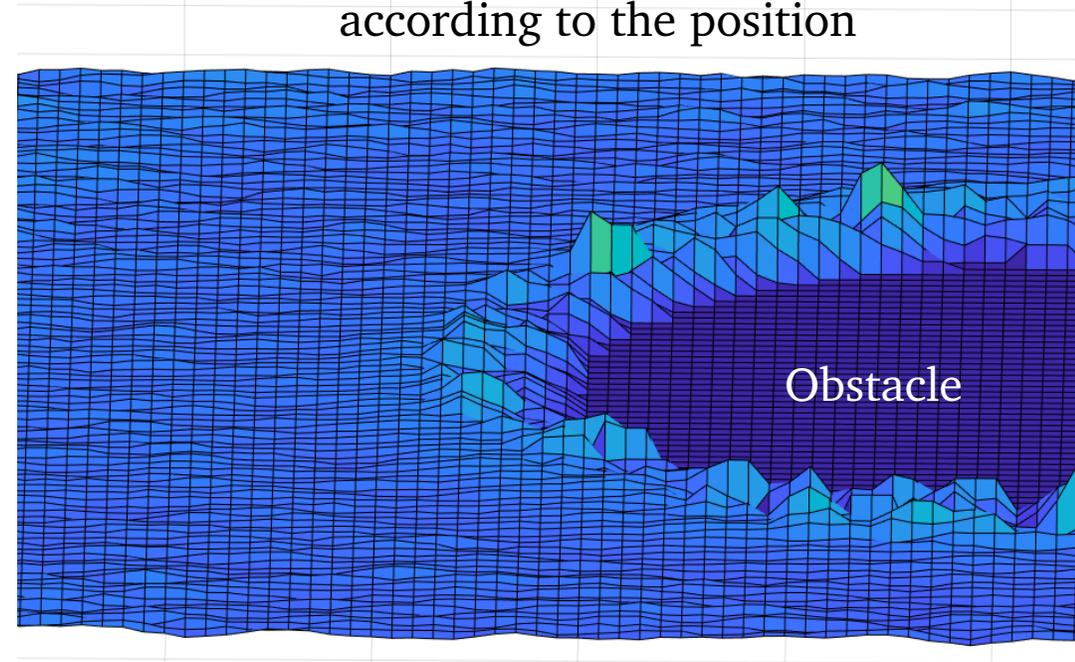
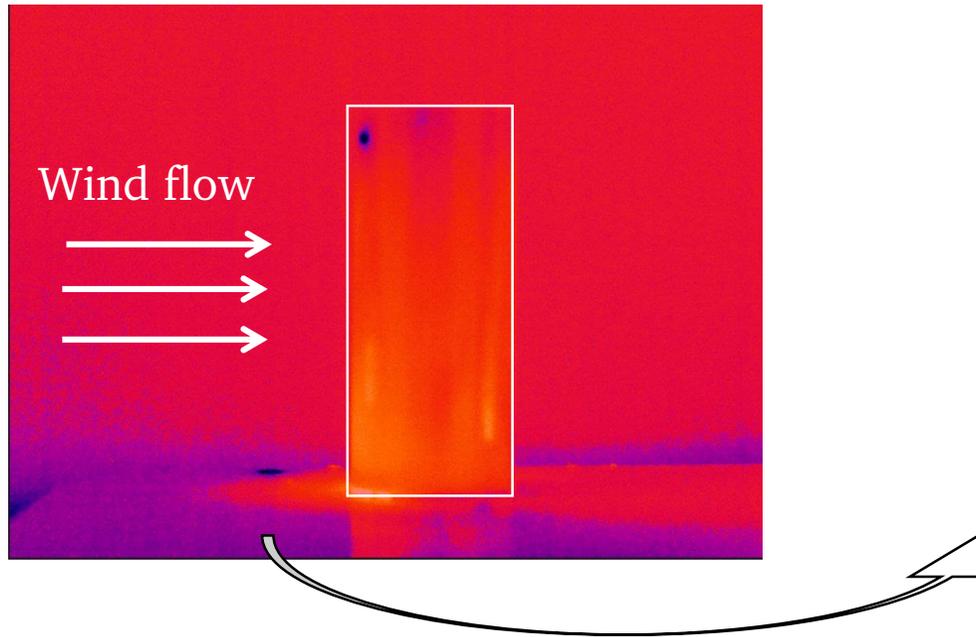
How does ice melt ?

Modelling the heat transfer $h(u)$

$$T = T_i + (T_f - T_i) \left(1 - e^{-\frac{h(u)}{\rho c} t} \right)$$

Experimental data for $h(u)$ in $\text{W m}^{-2} \text{K}^{-1}$ according to the position

Thermal camera from 20°C to 35°C

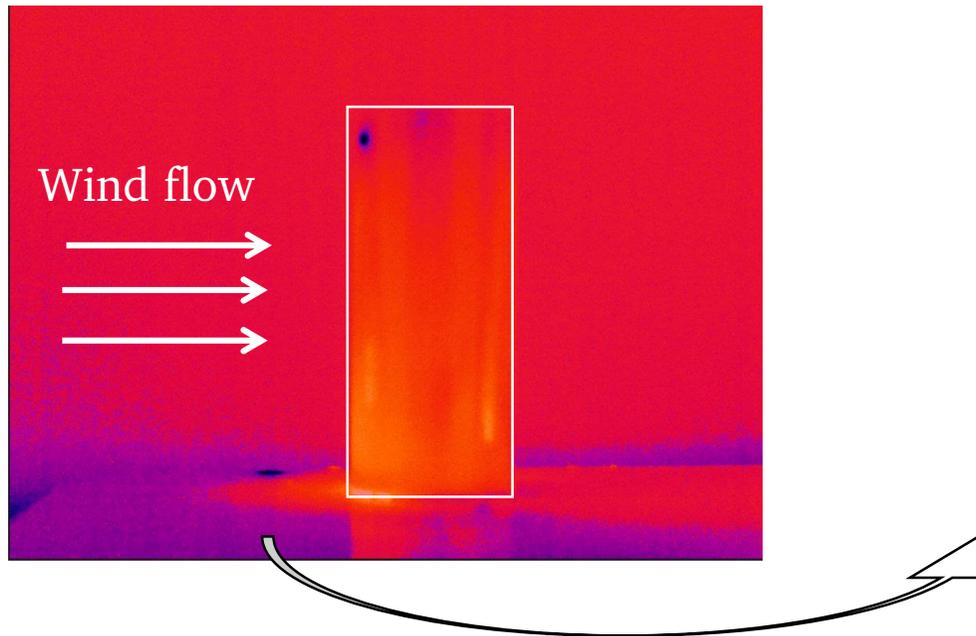




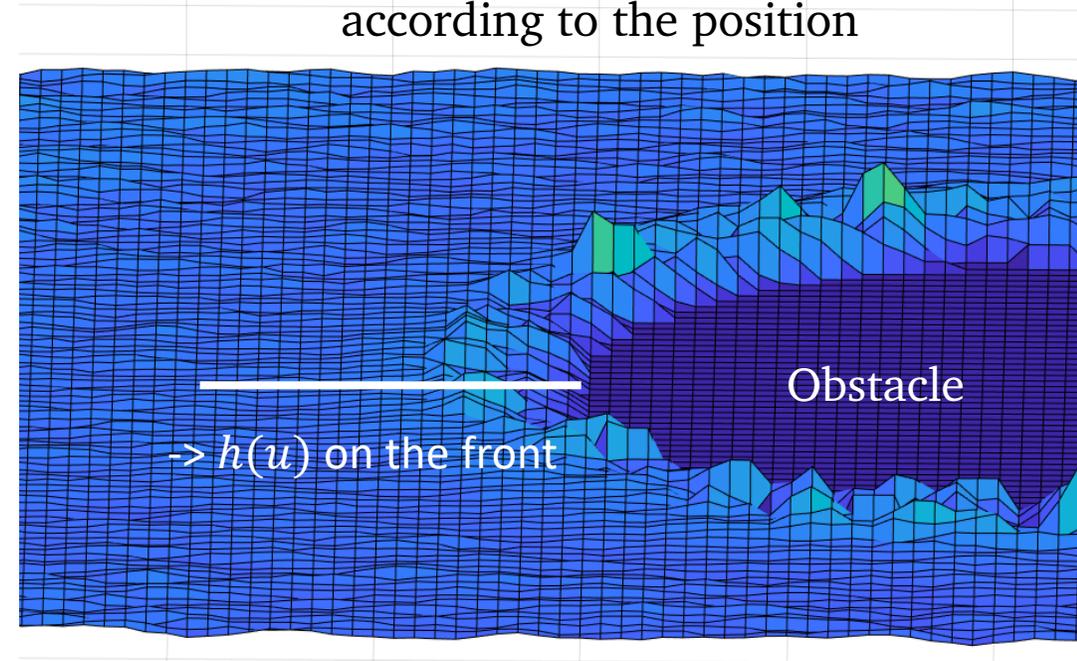
How does ice melt ?

Modelling the heat transfer $h(u)$

Thermal camera from 20°C to 35°C



Experimental data for $h(u)$ in $\text{W m}^{-2}\text{K}^{-1}$
according to the position

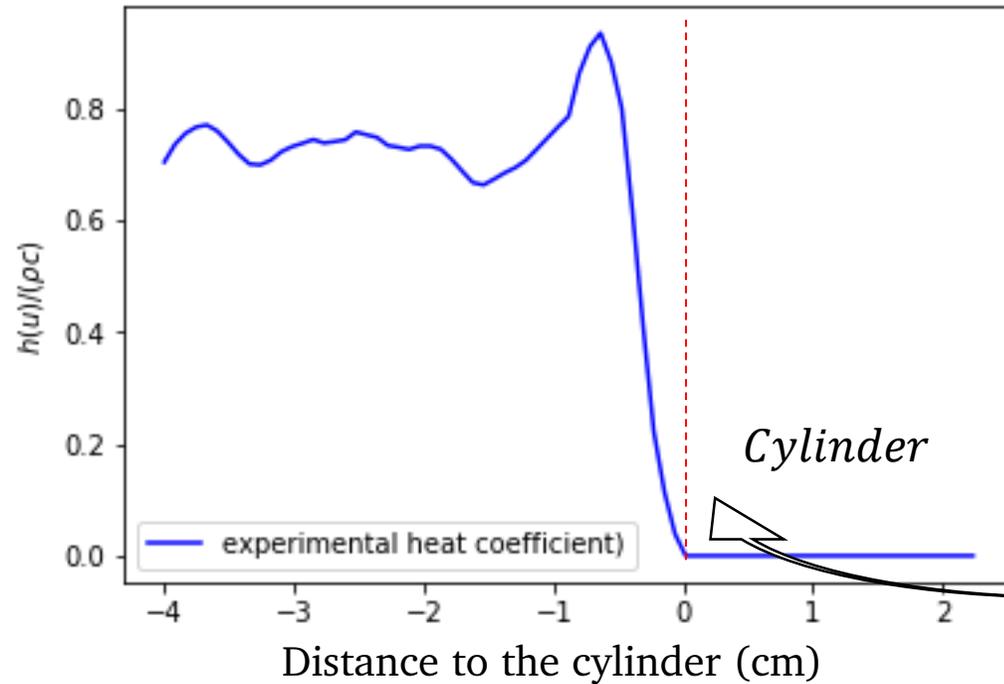




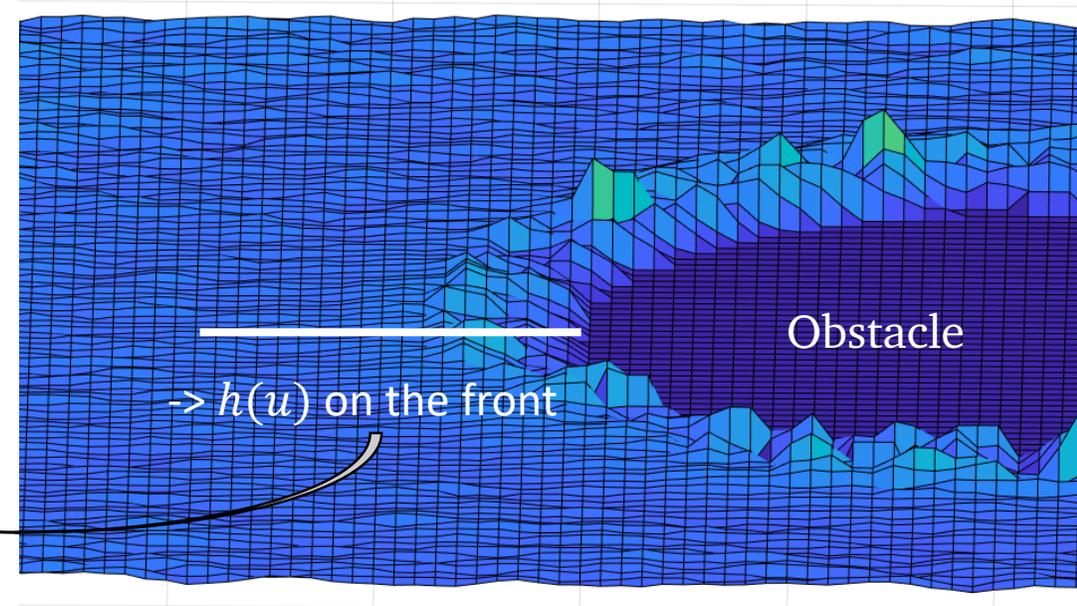
How does ice melt ?

Modelling the heat transfer $h(u)$

$h(u)$ on the front part



Experimental data for $h(u)$ in $W m^{-2} K^{-1}$ according to the position

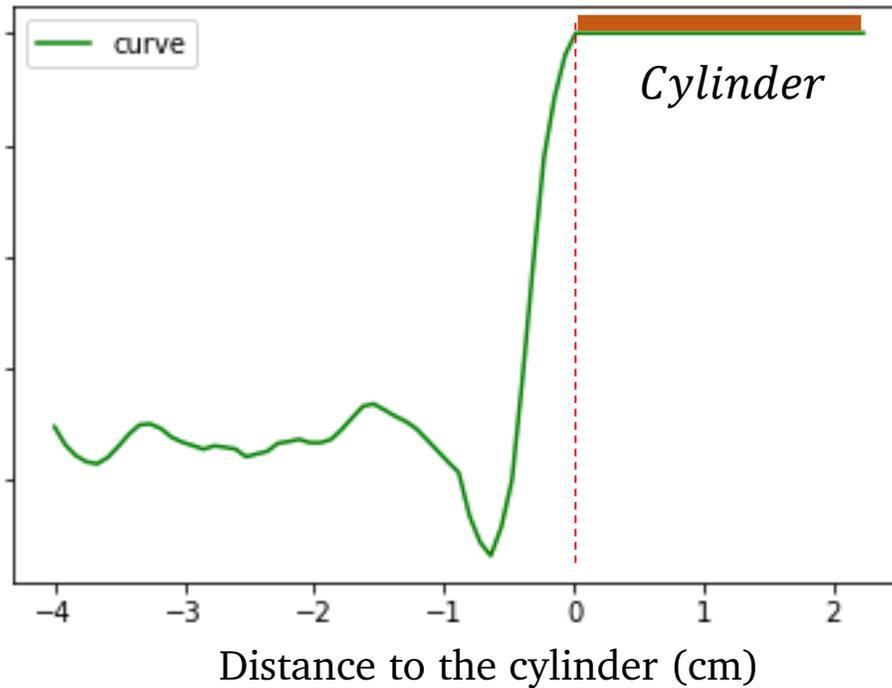




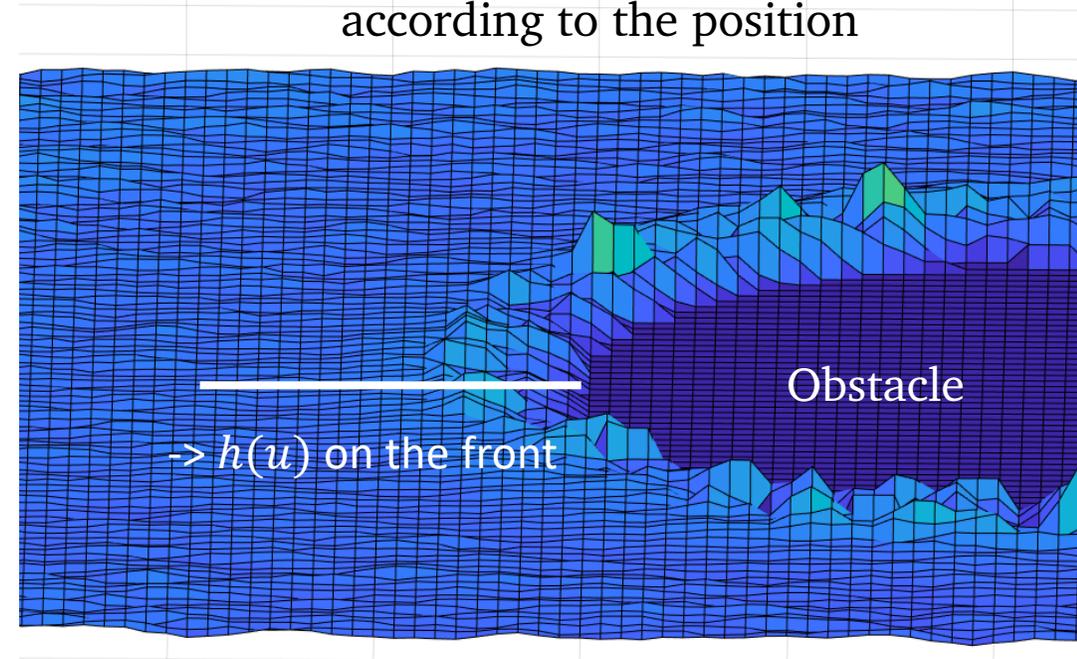
How does ice melt ?

Modelling the heat transfer $h(u)$

Estimated shape of the melting ice
At $t = 0+$



Experimental data for $h(u)$ in $\text{W m}^{-2}\text{K}^{-1}$
according to the position

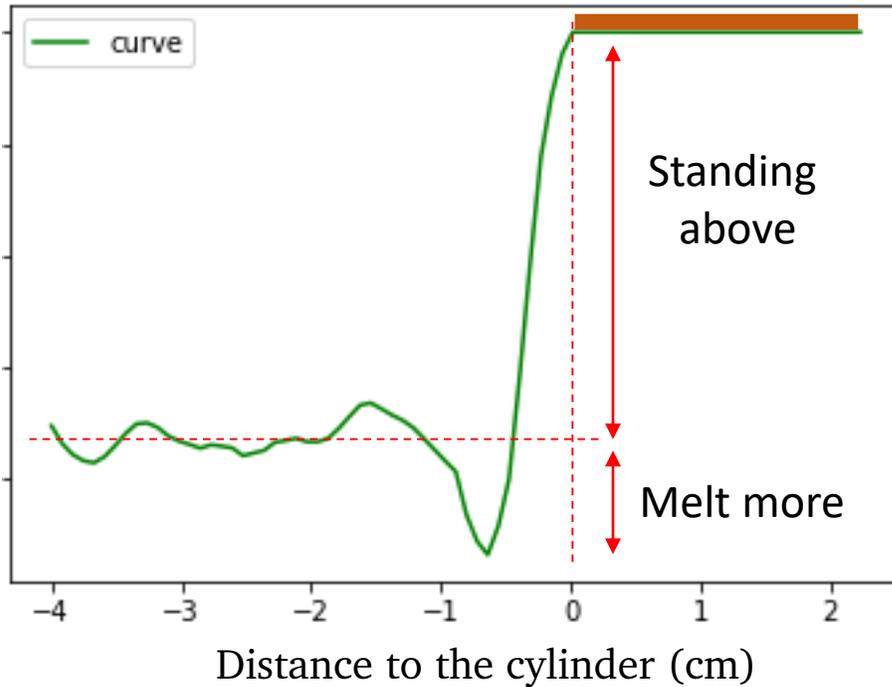




Explaining the balancing pebble

Wind creates the stand

Estimated shape of the melting ice
At $t = 0+$



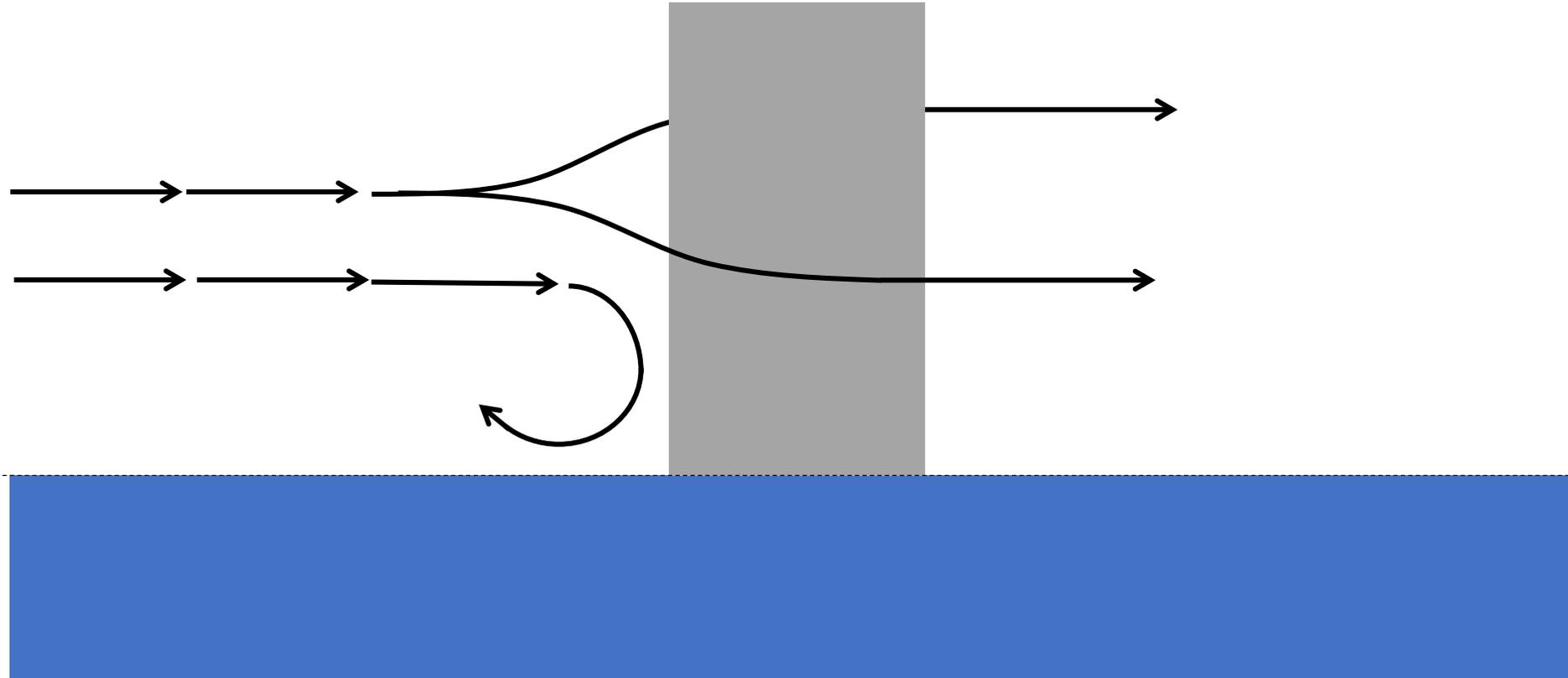
Curve at $t = 10$ min





Influence of an obstacle

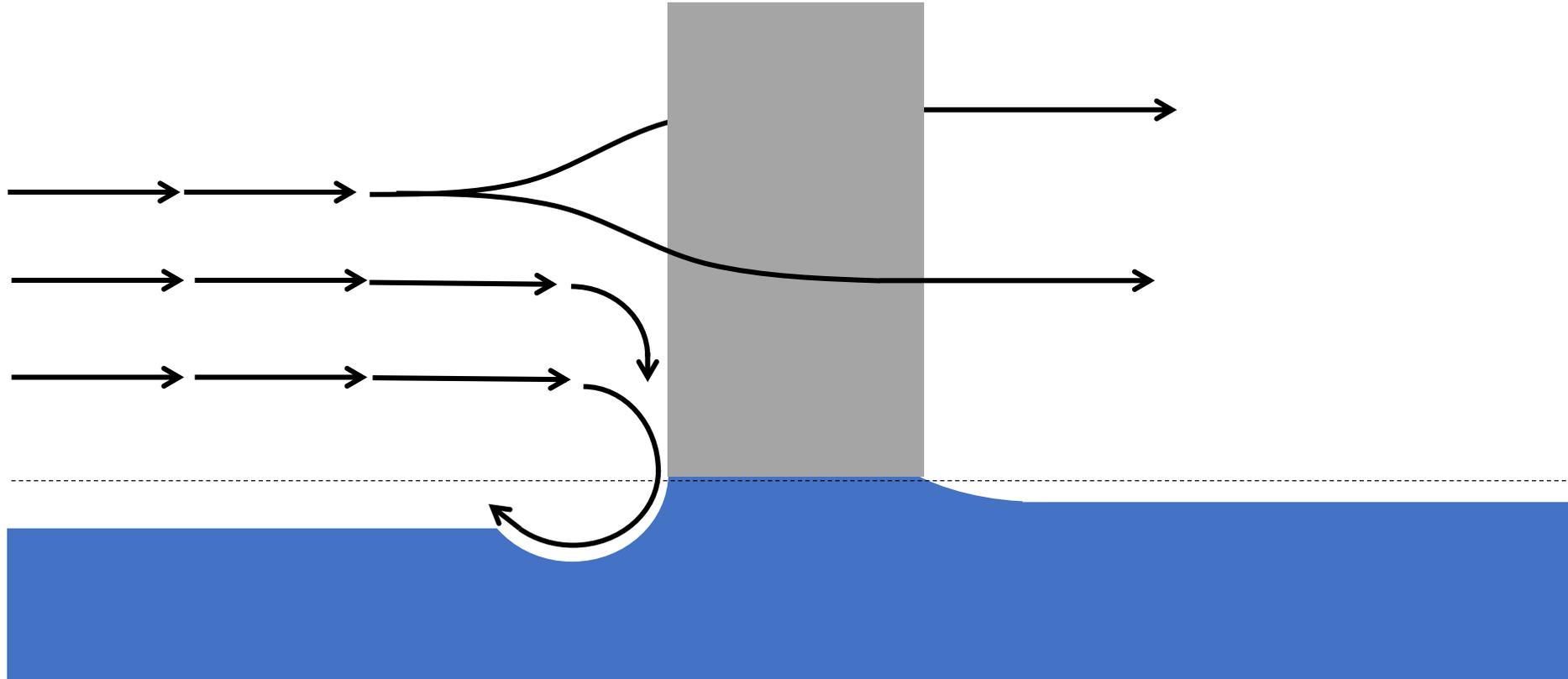
Air flow ?





Influence of an obstacle

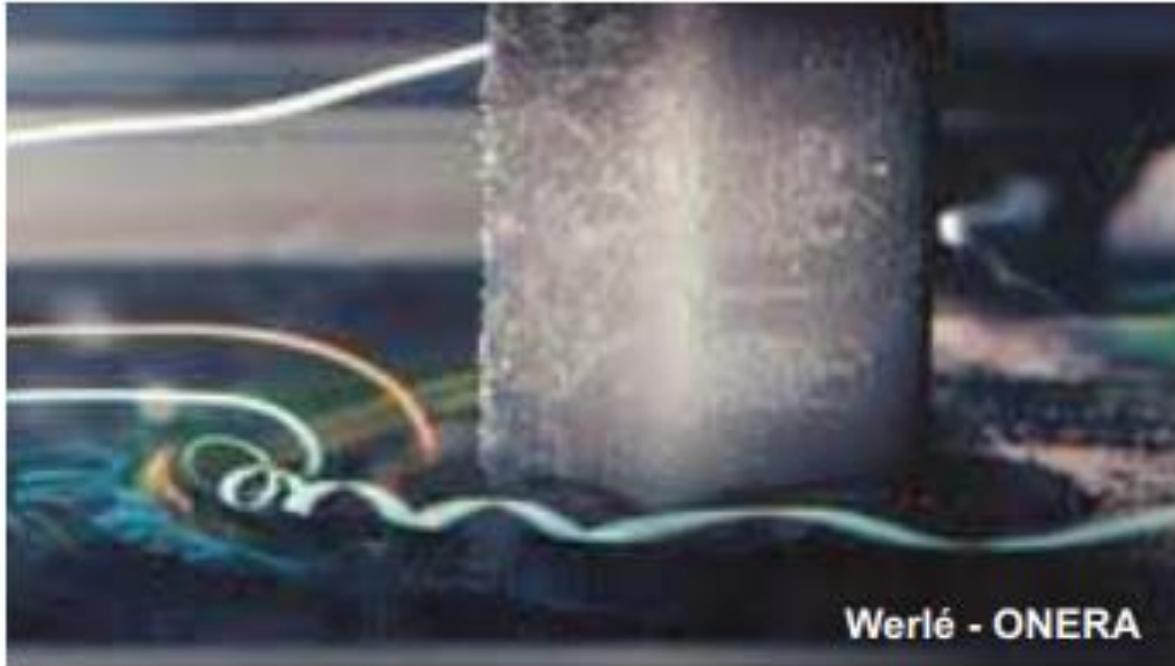
Turbulent vortex





Influence of an obstacle

Turbulent vortex

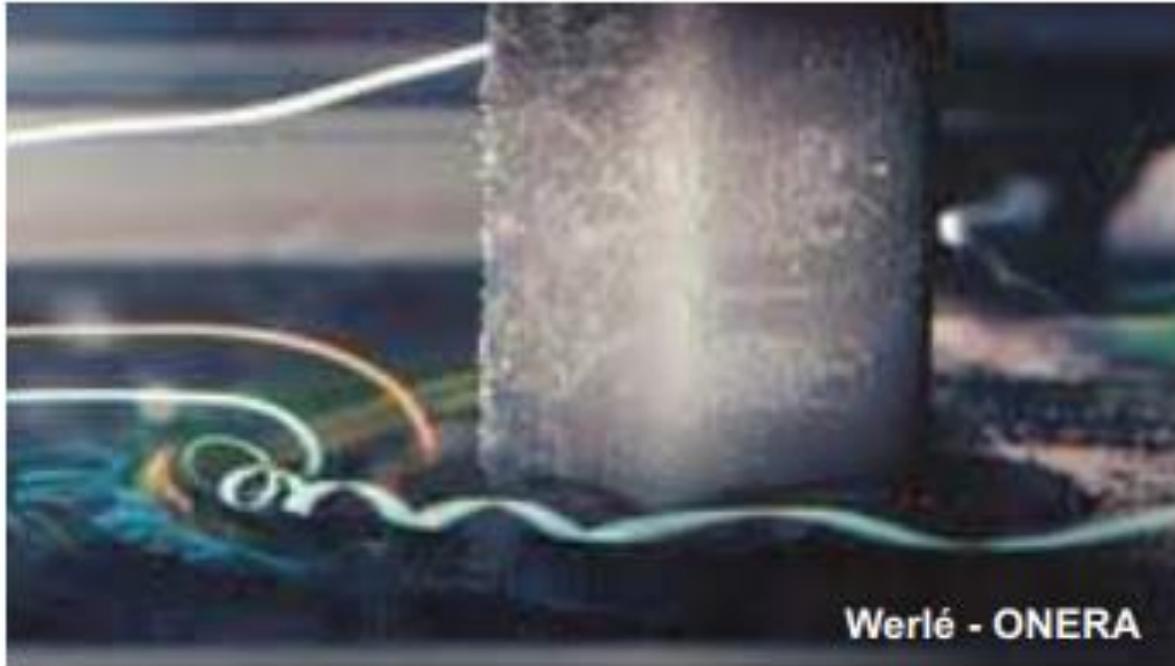


$$Re = \frac{\rho_{air} u D}{\eta_{air}} \approx 10^4$$



Influence of an obstacle

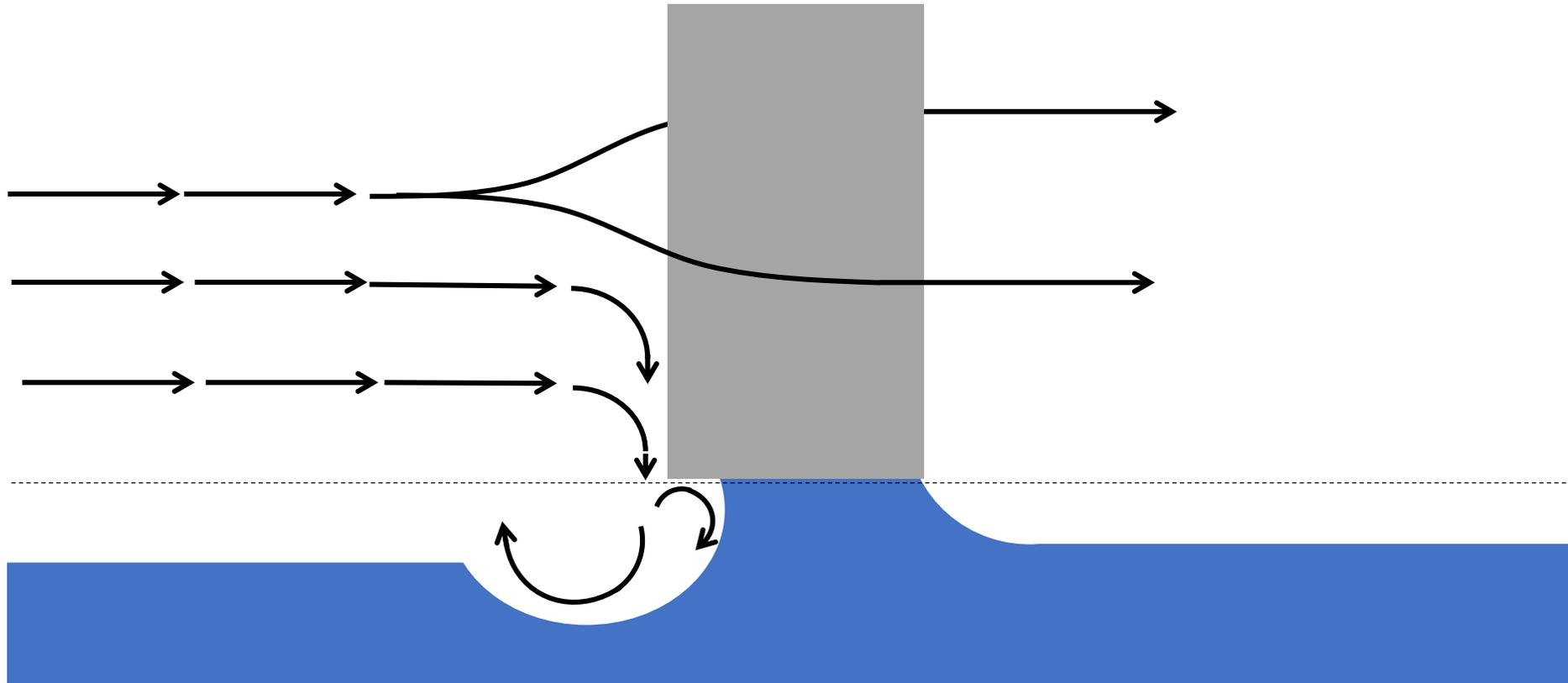
Turbulent vortex





Influence of an obstacle

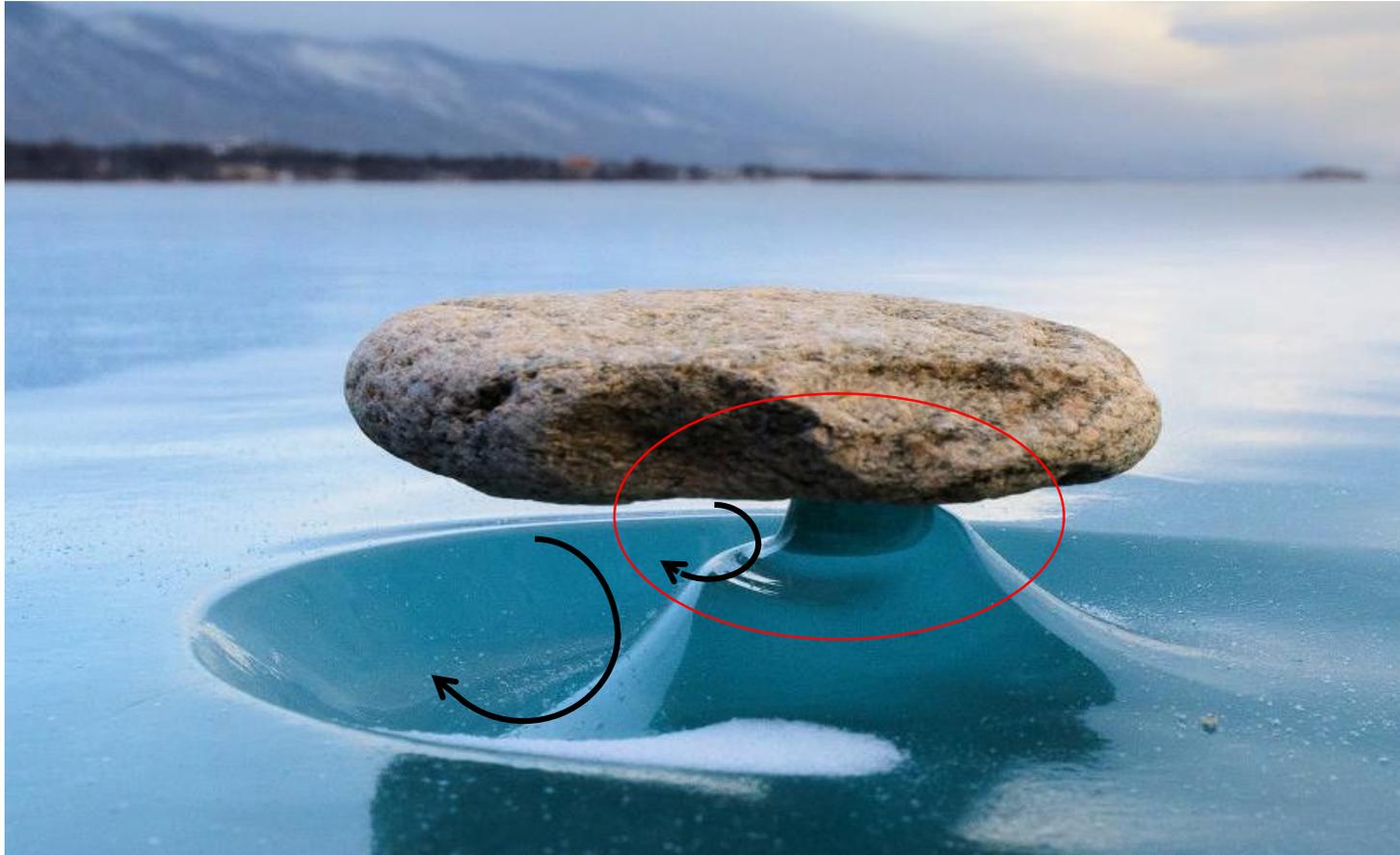
Turbulent vortex





A double stand

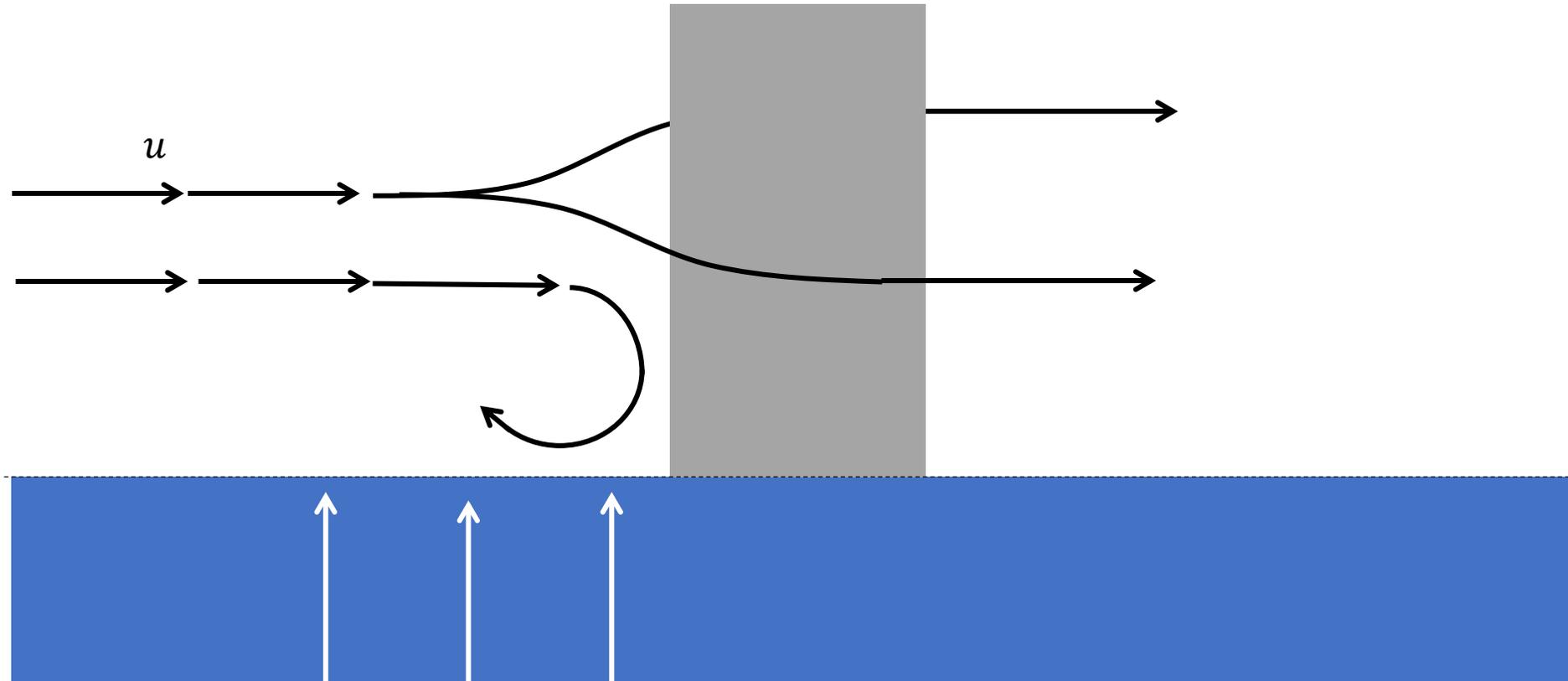
Turbulent vortex : 2nd order effect ?





Estimation of the curve

Size of the vortex



Measuring kinetic pressure



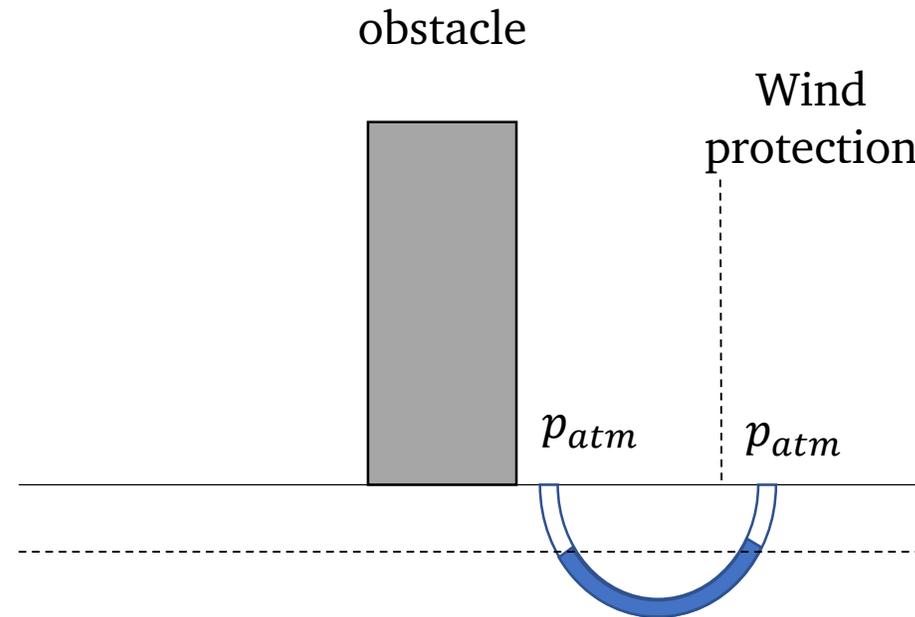
Estimation of the curve

Size of the vortex

Estimation of the characteristic size of the vortex



Set up side view



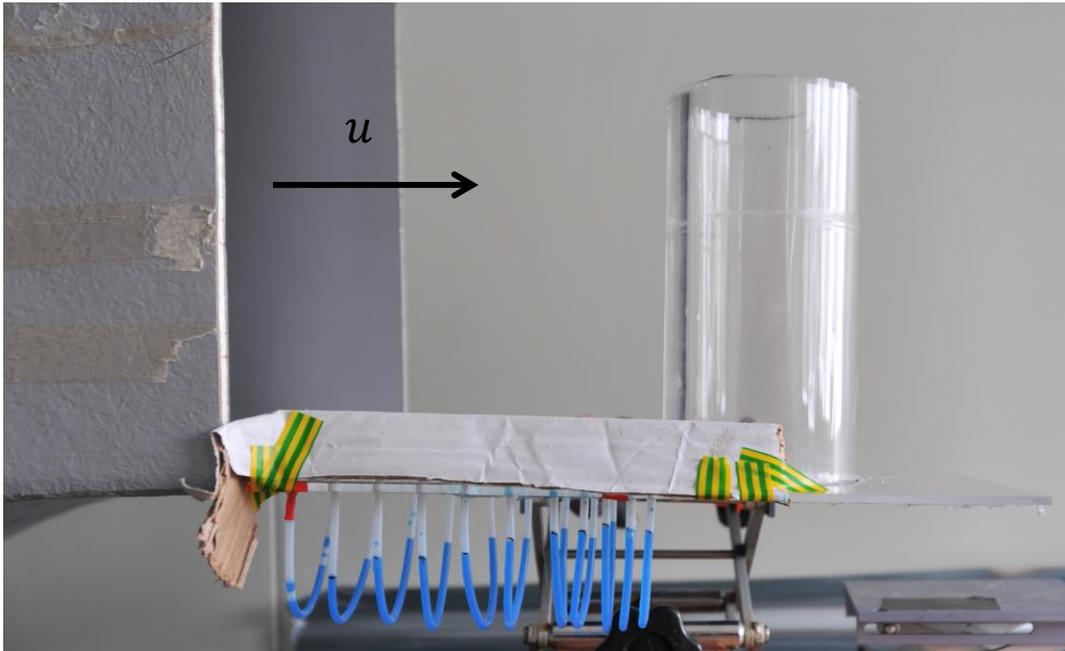
Set up front view



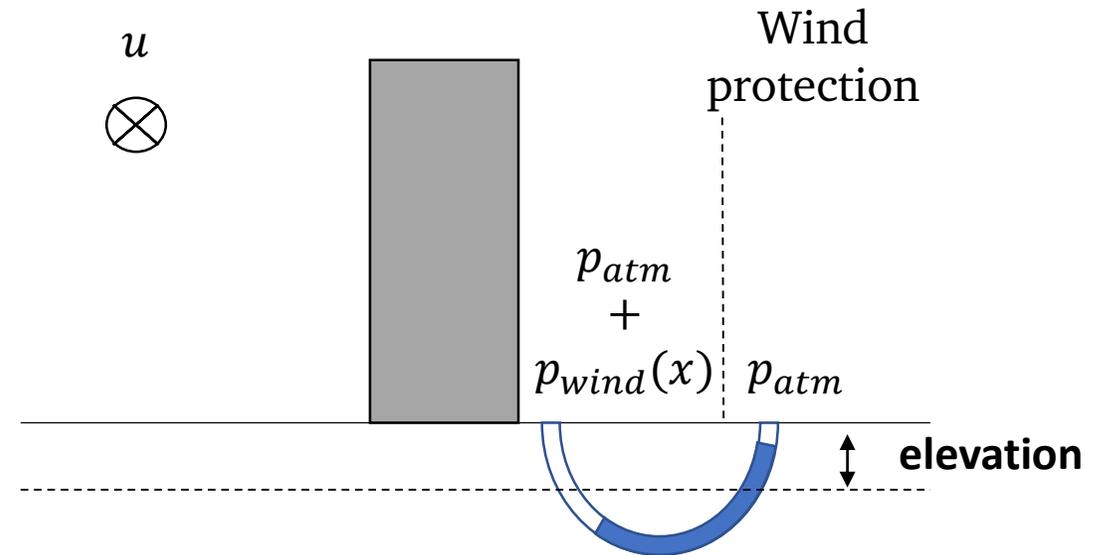
Estimation of the curve

Size of the vortex

Estimation of the characteristic size of the vortex



Set up side view



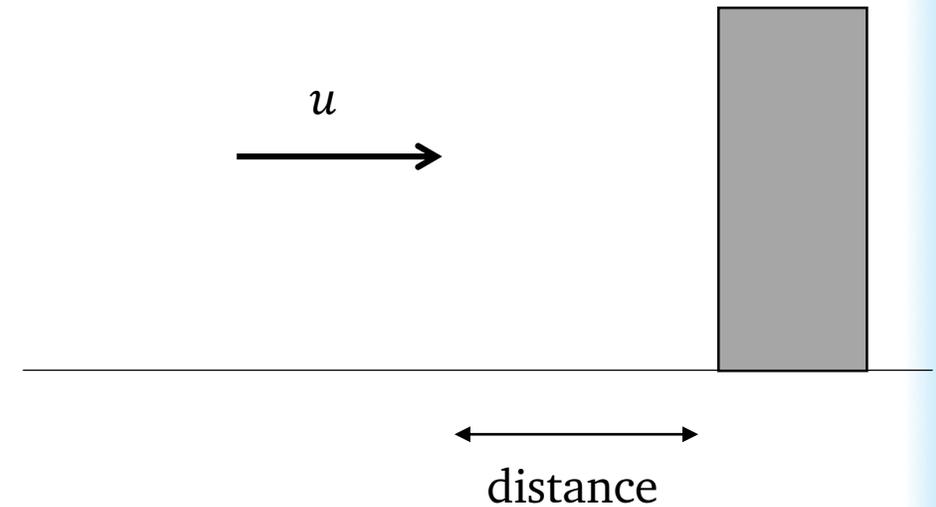
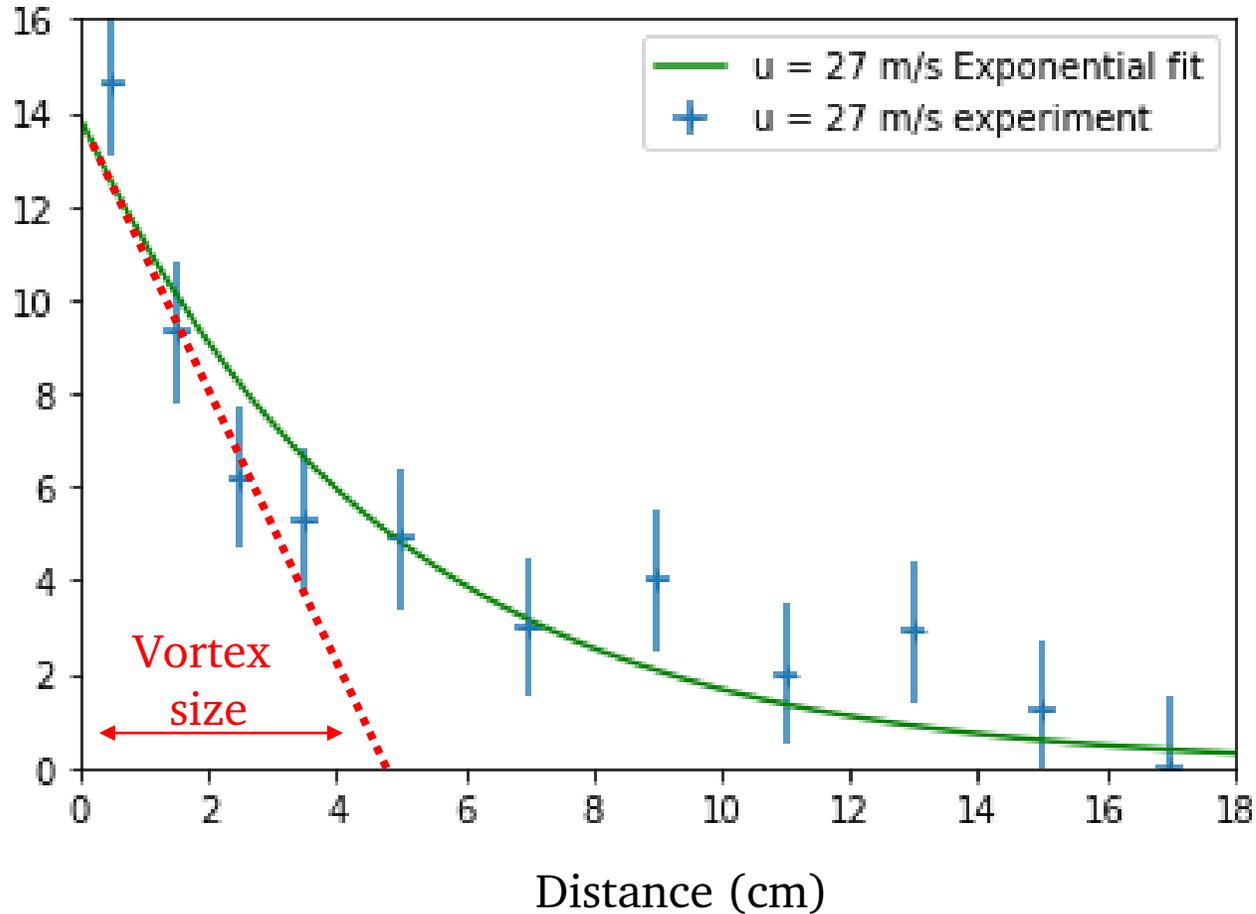
Set up front view



Estimation of the curve

Size of the vortex

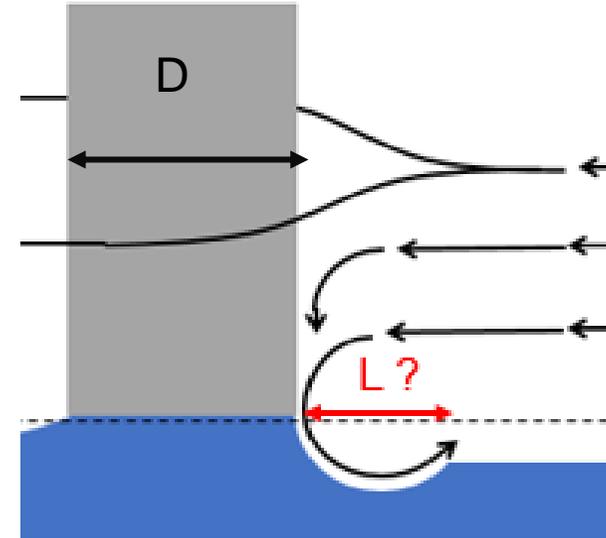
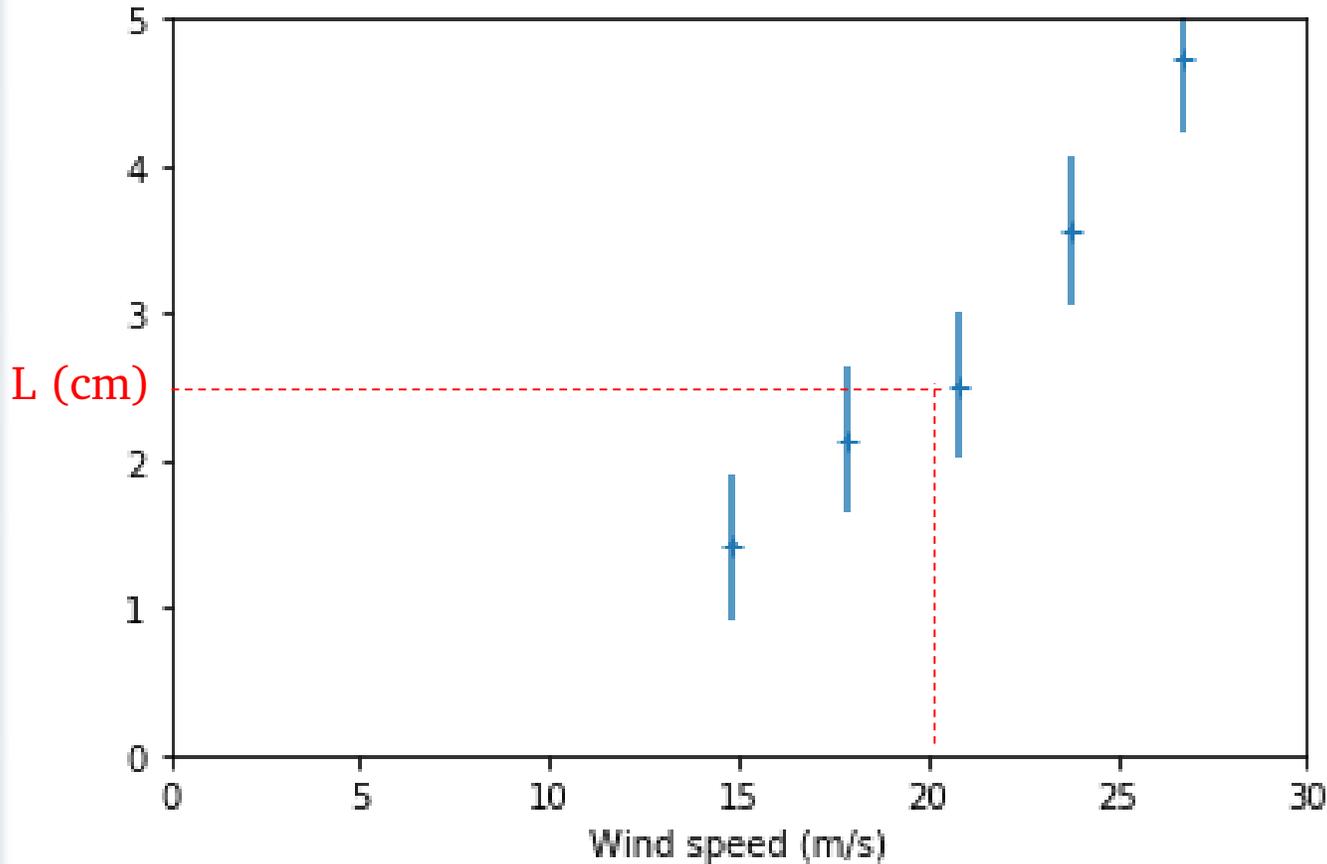
Elevation in Pito tube (mm)





Estimation of the curve

Size of the vortex



For $u = 20 \text{ m/s}$, $D = 7 \text{ cm}$:
 $\Rightarrow L \approx 2.5 \text{ cm}$

Curve = a circle or radius $\approx L$



Conclusion

✓ **Reproduce** this «stand» phenomenon



✓ **Explain** this «stand» phenomenon

Sublimation – Heat transfer

Same equation

Sublimation quicker on Baikal Lake

✓ Estimate **the curve** of the stand depending on the important parameters.

Ablation Parameters :

Vapor pressure, Temperature

Absolute wind speed

Shape parameters :

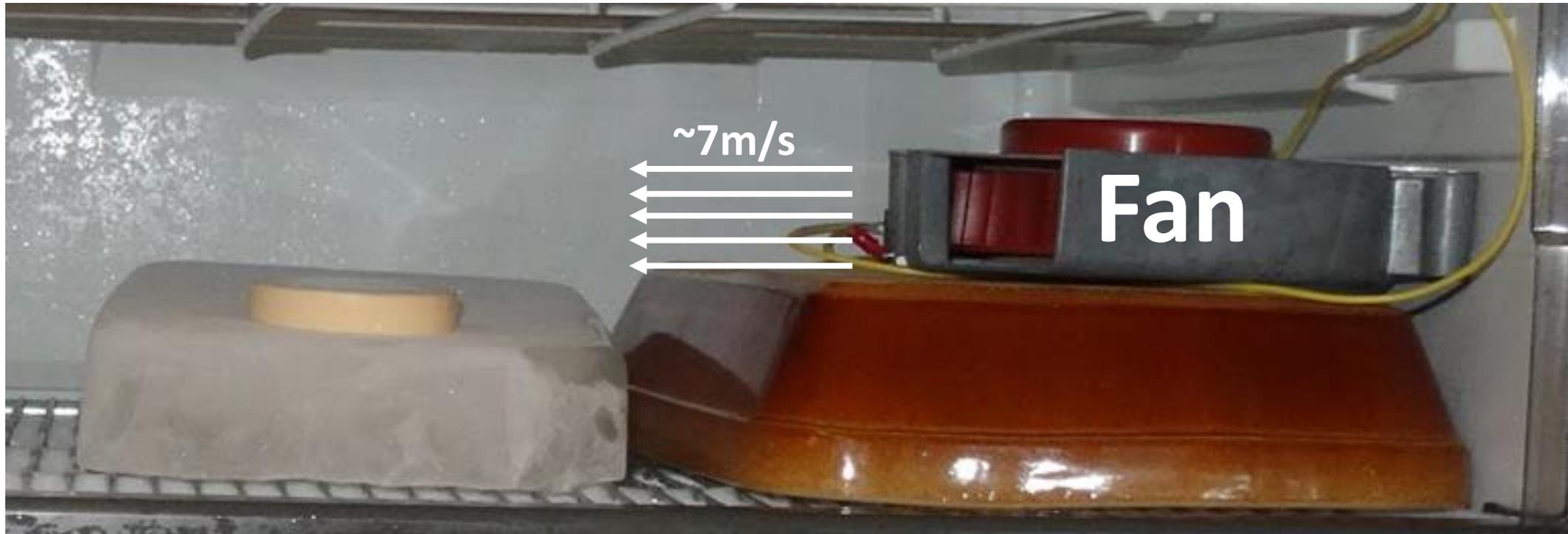
Turbulent flow

Radius of curvature $\sim L$, typical size of the vortex depends on the wind speed



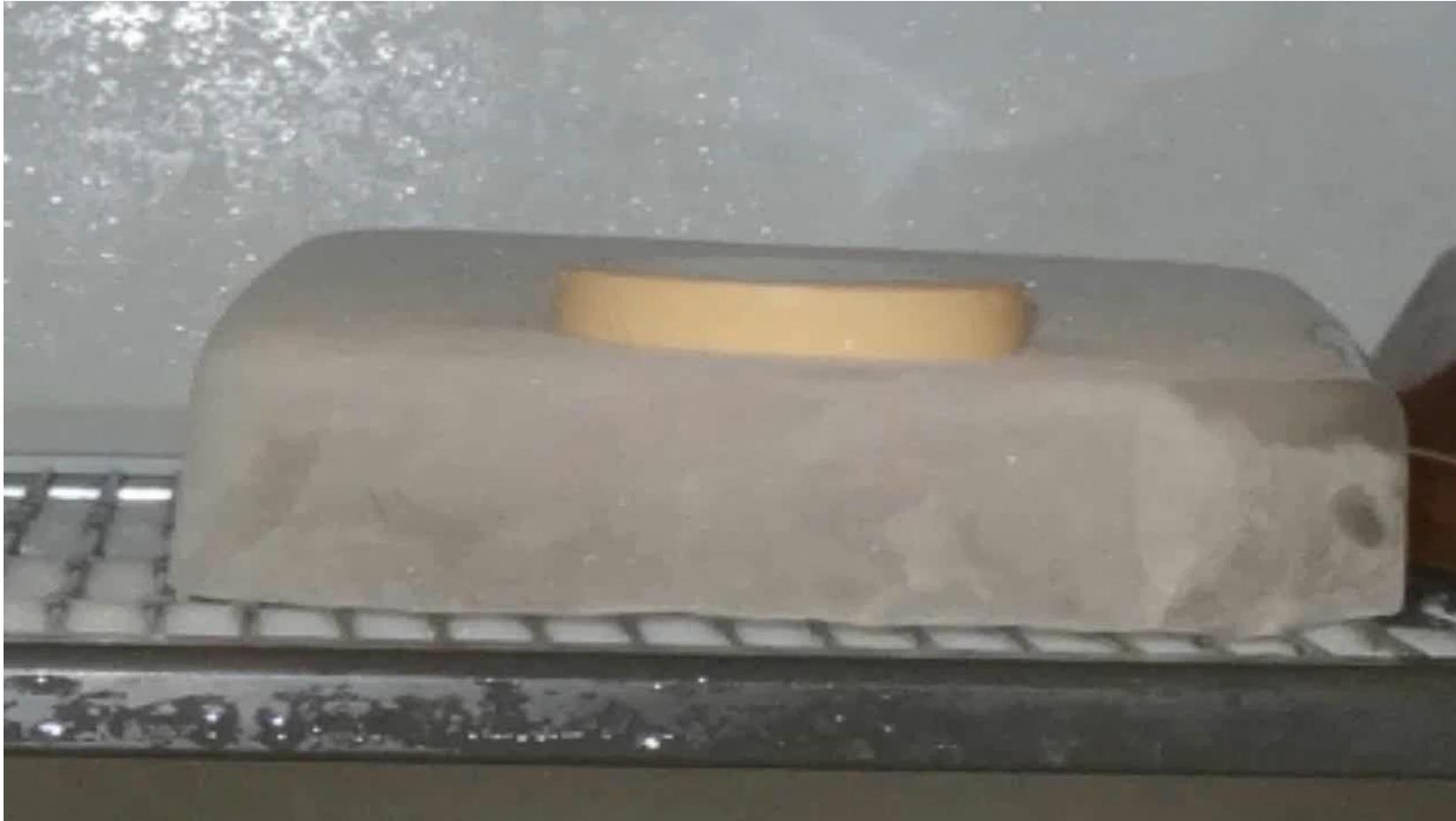
Sublimation

$T = -22^{\circ}\text{C}$





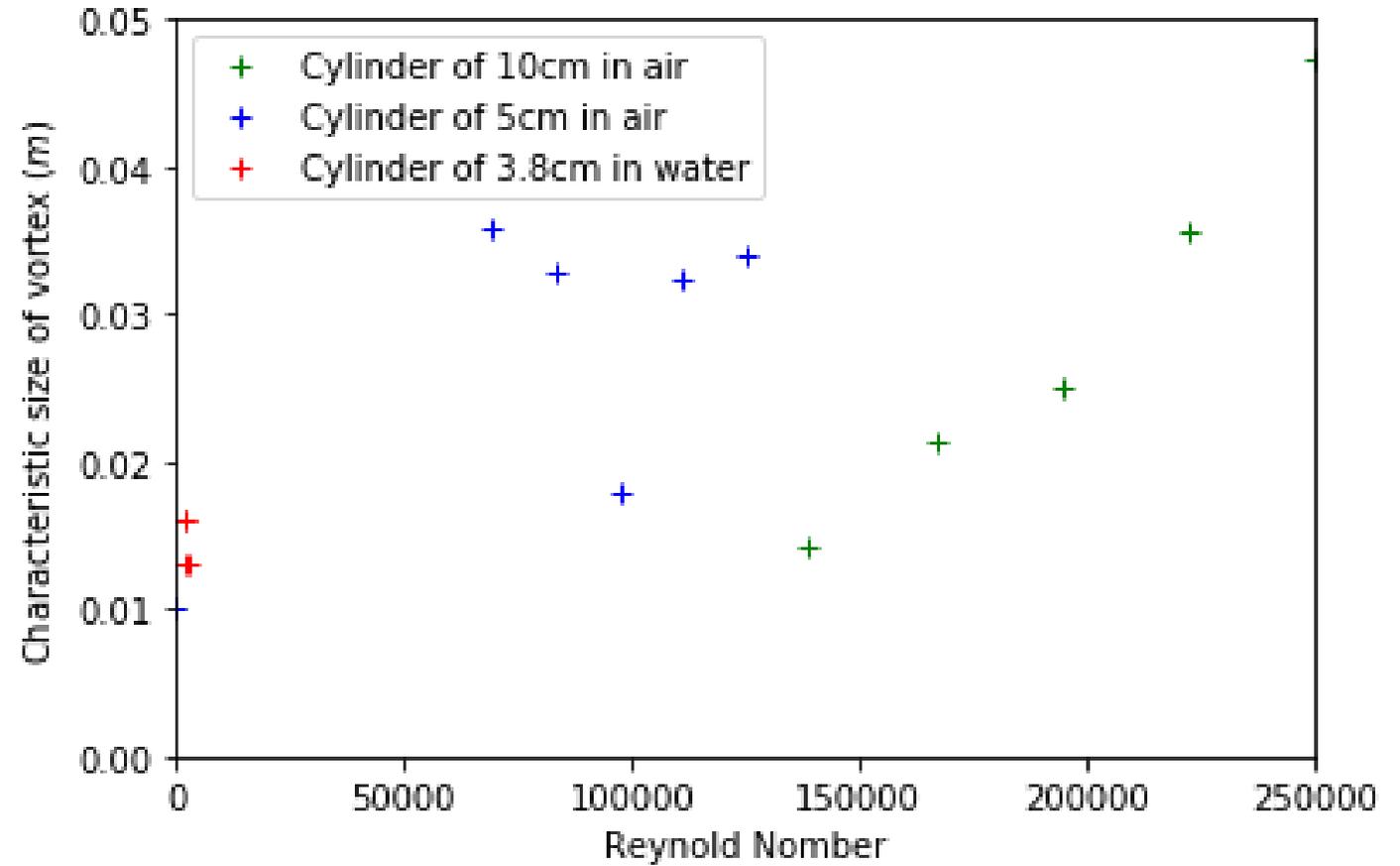
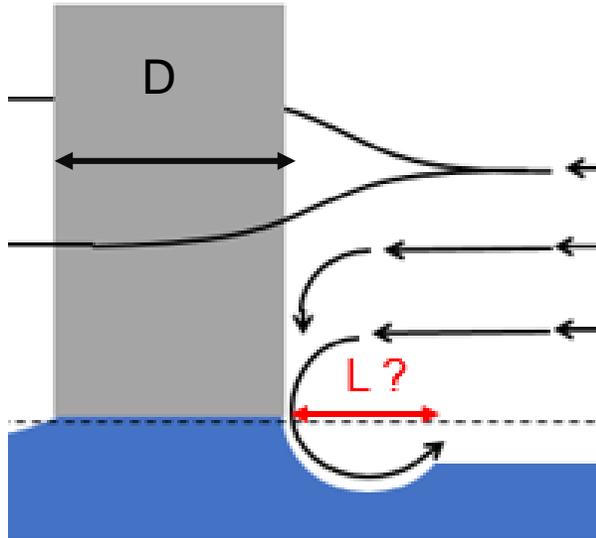
Sublimation



In 20 days !



Size of vortex VS Reynolds Number





How does ice melt ?

Heat and mass transfer

Analogous causes:

- ✓ Mass: concentration gradient
- ✓ Heat: temperature gradient

Chilton-Colburn J-analogy:

Similar formulations of basic equations:

Conduction:

- ✓ Mass: $\partial_t C = D_m \Delta C$
- ✓ Heat: $\partial_t T = D_h \Delta T$

D_m = Mass diffusivity ; D_h = Heat diffusivity

Convection:

Mass: Sherwood number

$$Sh = h_m$$

$$Sh = Ct. \times Sc^{\frac{1}{3}} \times Re$$



Why does ice melt ?

Wind

Sublimation

Vapor pressure

$$\Phi_{sublimation} = \rho_a C_E L_{sub} \mathbf{u} (q_{sat}(T) - q_{wind})$$

$(q_{sat}, q_{wind} : \text{mass of water per } m^3 \text{ of air})$

Heat transfer

From ice to water then vapor

$$\Phi_{heat} = \rho_a c_a C_H \mathbf{u} (T_{wind} - T_{ice})$$

~



Why does ice melt ?

Sublimation

Bulk aerodynamic approach :

$$\Phi_{\text{humidity}} = \rho_a C_E L_{\text{sub}} u (q_{\text{sat}}(T) - q_{\text{wind}})$$

$$\Phi_{\text{humidity}} \approx 25 \text{ W/m}^2 \text{ for } u = 5 \text{ m/s}$$

with

$$\left\{ \begin{array}{l} \rho_a: \text{density of air} \\ L_{\text{sub}}: \text{sublimation latent heat} \\ q_{\text{sat}}, q_{\text{wind}} : \text{mass of water in } 1\text{m}^3 \\ C_E: \text{a constant} \end{array} \right.$$

(Ref : *The Physics of glaciers, by K.M Cuffier and W.S.B Paterson*)

Characteristic time :

$$\Delta e = 10 \text{ cm}$$

$$q_{\text{wind}} = 0.8 q_{\text{sat}}$$

Δe 

$$\Delta t \approx 4 \text{ months}$$

On Lake : possible

In lab : long and difficult to reproduce



Why does ice melt ?

Heat transfer

Bulk aerodynamic approach :

$$\Phi_{heat} = \rho_a c_a C_H \mathbf{u} (T_{wind} - T_{ice})$$

$$\Phi_{heat} \approx 50 \text{ W/m}^2 \text{ for } u = 5 \text{ m/s}$$

with $\begin{cases} \rho_a : \text{density of air} \\ c_a : \text{thermal capacity} \\ C_H : \text{a constant} \end{cases}$

(Ref : *The Physics of glaciers*, by K.M Cuffier and W.S.B Paterson)

Characteristic time :

$$\Delta e = 10 \text{ cm}$$

$$T_{wind} - T_{ice} = 5^\circ\text{C}$$

Δe 

$$\Delta t \approx 2 \text{ months}$$

On Lake : possible

In lab : possible by increasing $T_{wind} - T_{ice}$

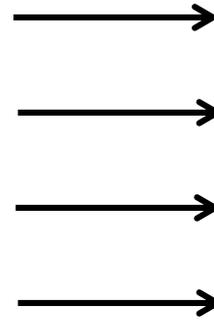
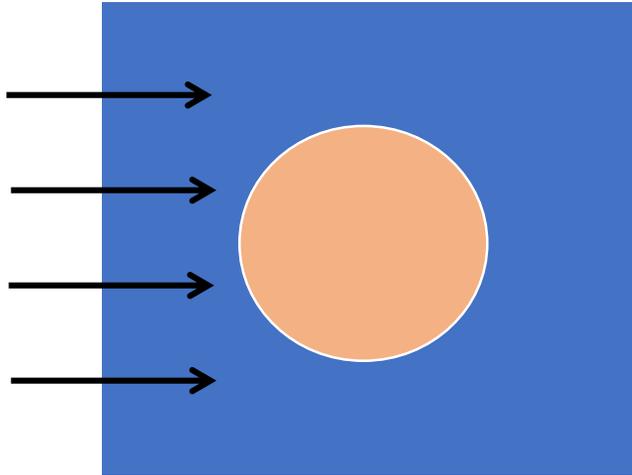


Hoodoos





4th experiment : Erosion on sand

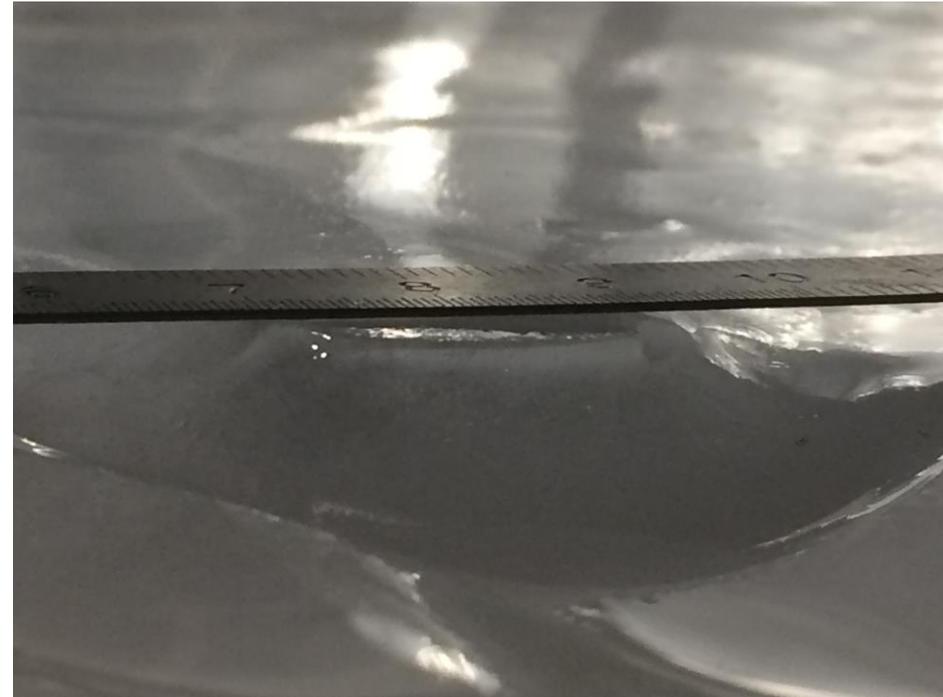
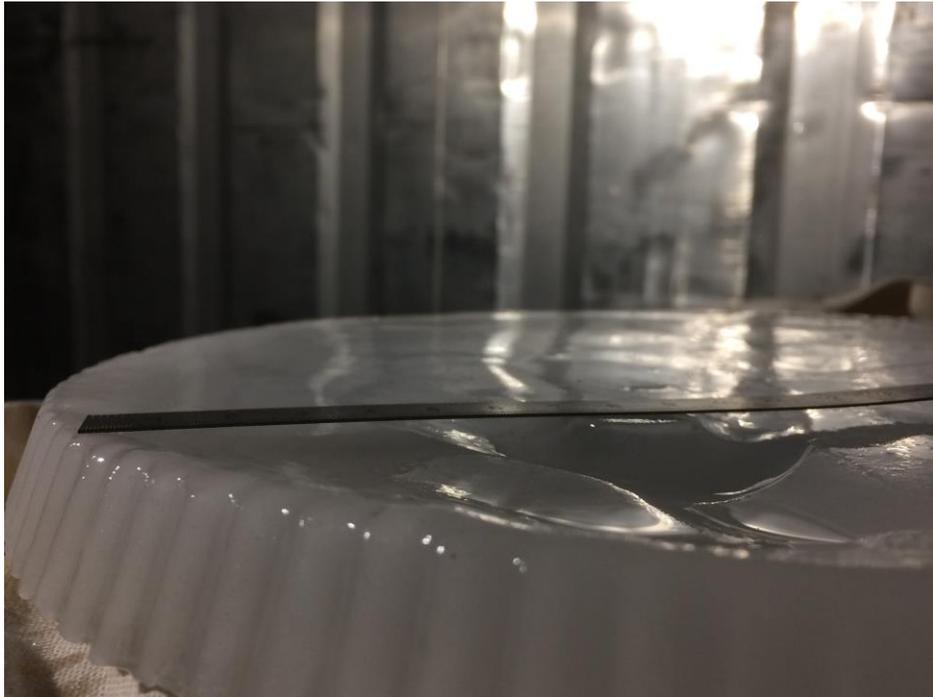


- 1) Stone stands above the lake
- 2) Ice lower just around the stone
- 3) Symmetry or Axisymmetry





No elevation : only melting !



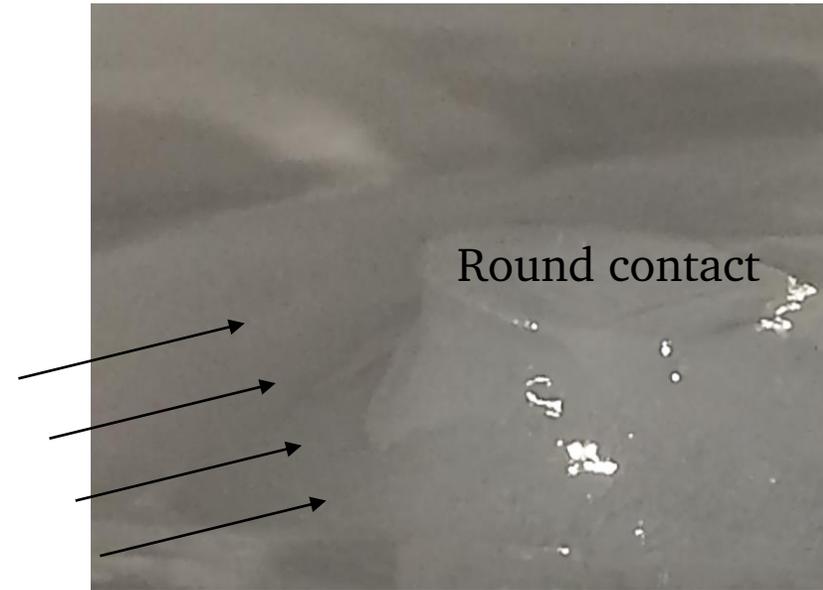
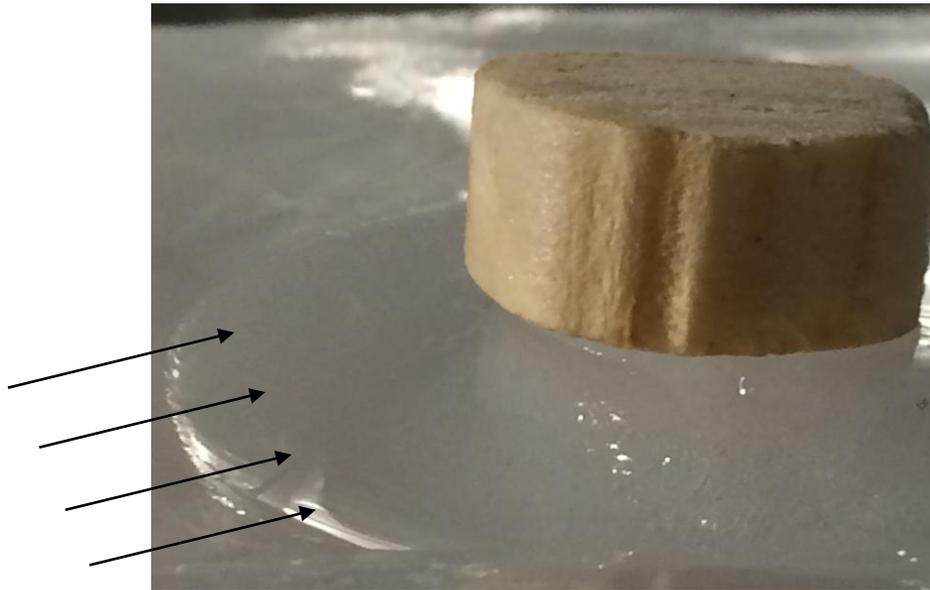


Dependence of the form of contact





Dependence of the form of contact



If no contact : the wind can blow on ice, more and more.
If contact with a solid : the ice is protected from the wind.



3rd experiment : Wind





Understanding the phenomenon

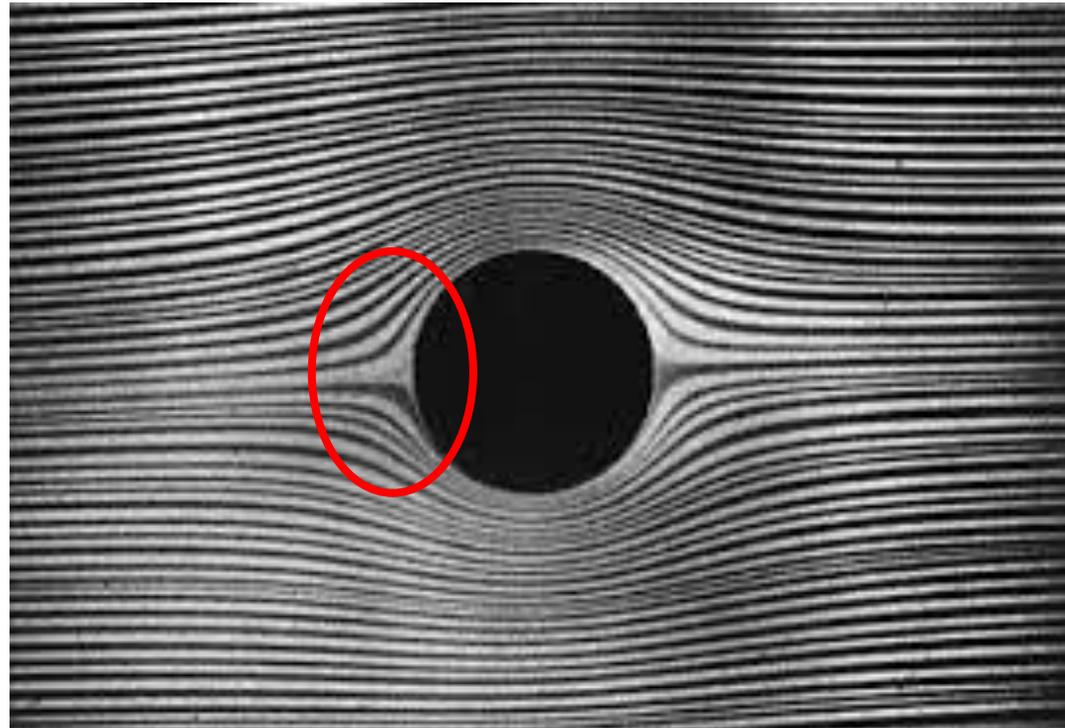
Aerodynamic flow

A simple model : Potential flow

$$\Phi_{heat} = h(u) (T_{wind} - T)$$

Should melt less
at the front

$$Re = \frac{\rho_{air} u D}{\eta_{air}} \approx 10^4$$

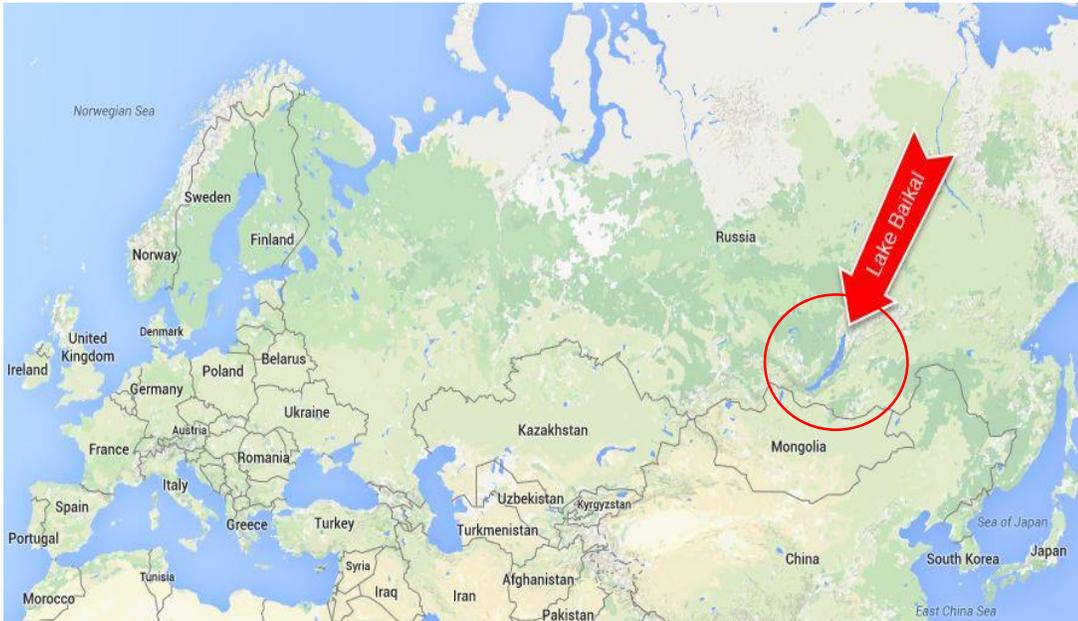


There must be something else... Turbulence !



Context of the picture

Lake Baikal, Siberia



Location of Lake Baikal

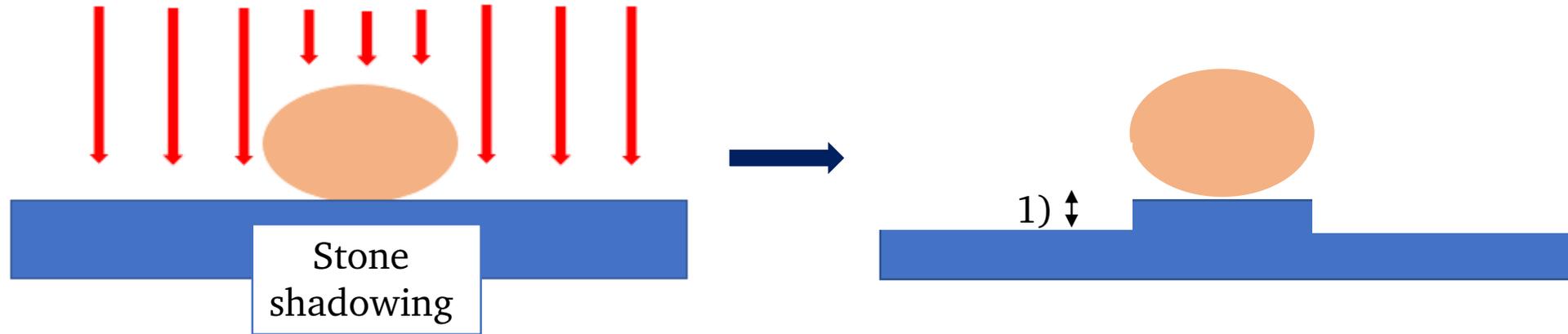


Picture of Lake Baikal

Structure formation : from December to April



2nd experiment : Radiation



- 1) Stone stands above the lake
- 2) Ice lower just around the stone
- 3) Asymmetry possible

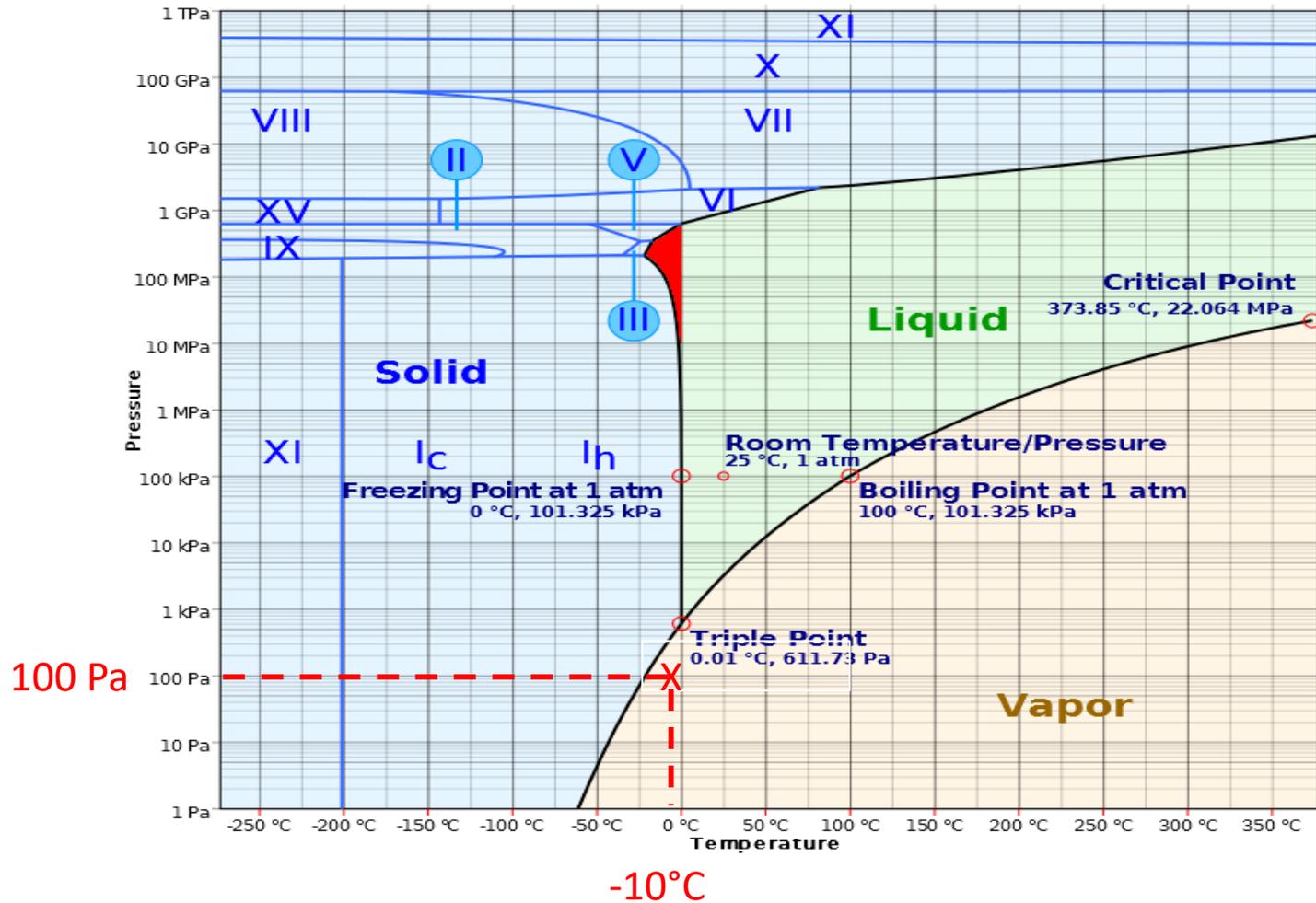
✓

X

X



Why does ice melt ?



- Most stable phase: vapor
⇒ sublimation
- Endothermic reaction
⇒ energy needed



How does ice melt ?

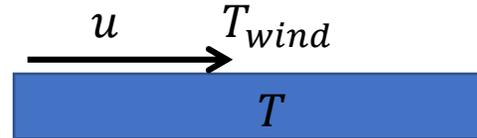
Modelling the heat transfer

Newton's law

$$\Phi_{heat} = h(u) (T_{wind} - T)$$

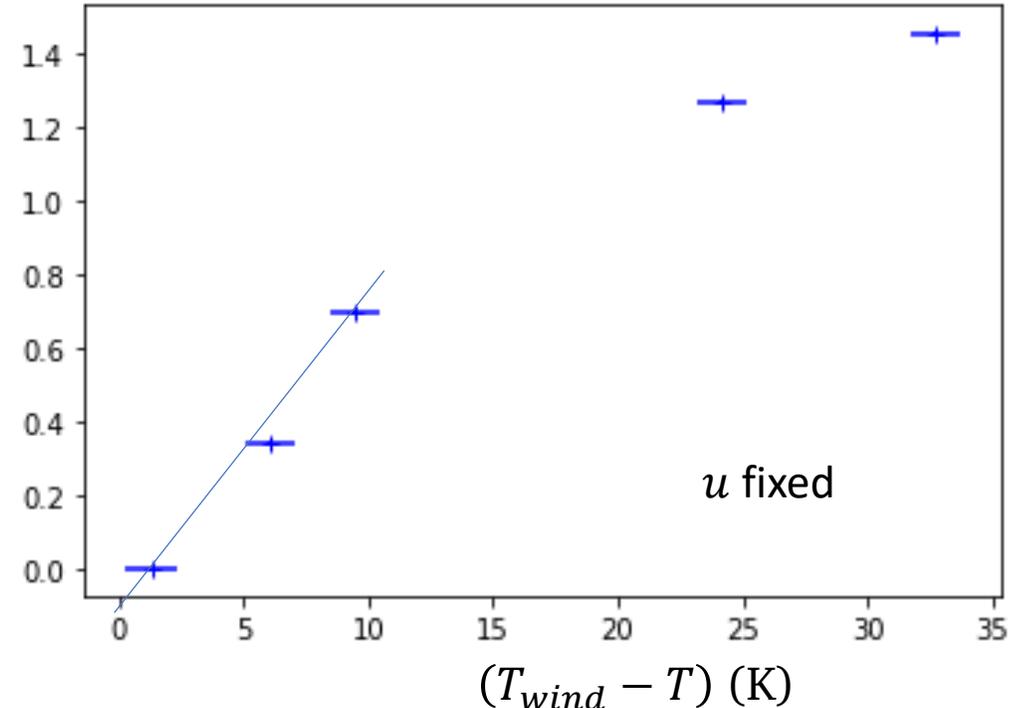
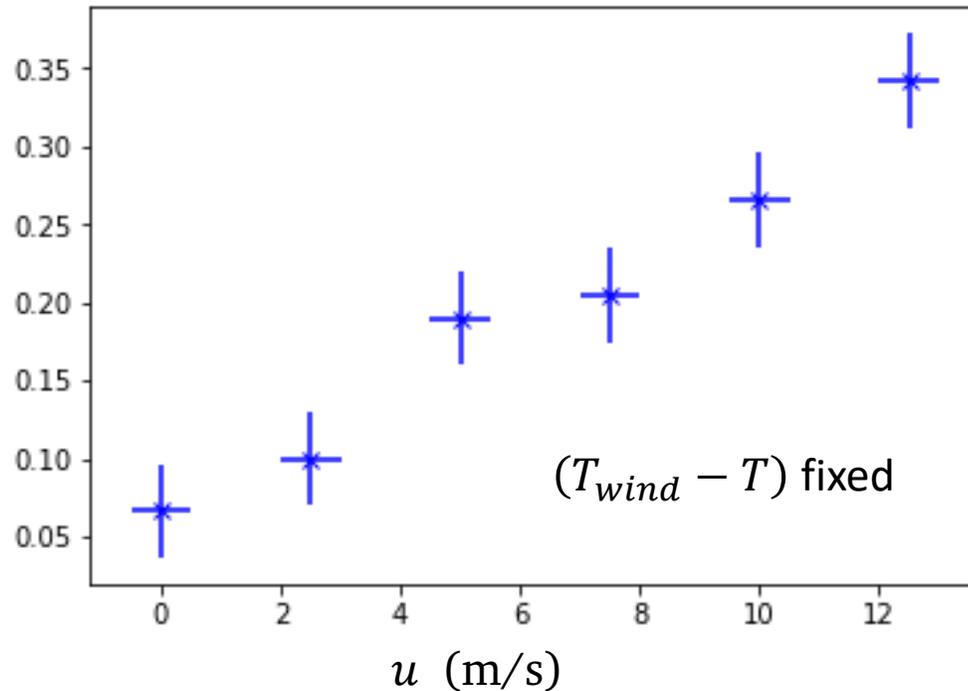


On a simple case :



Melting Ratio (%)

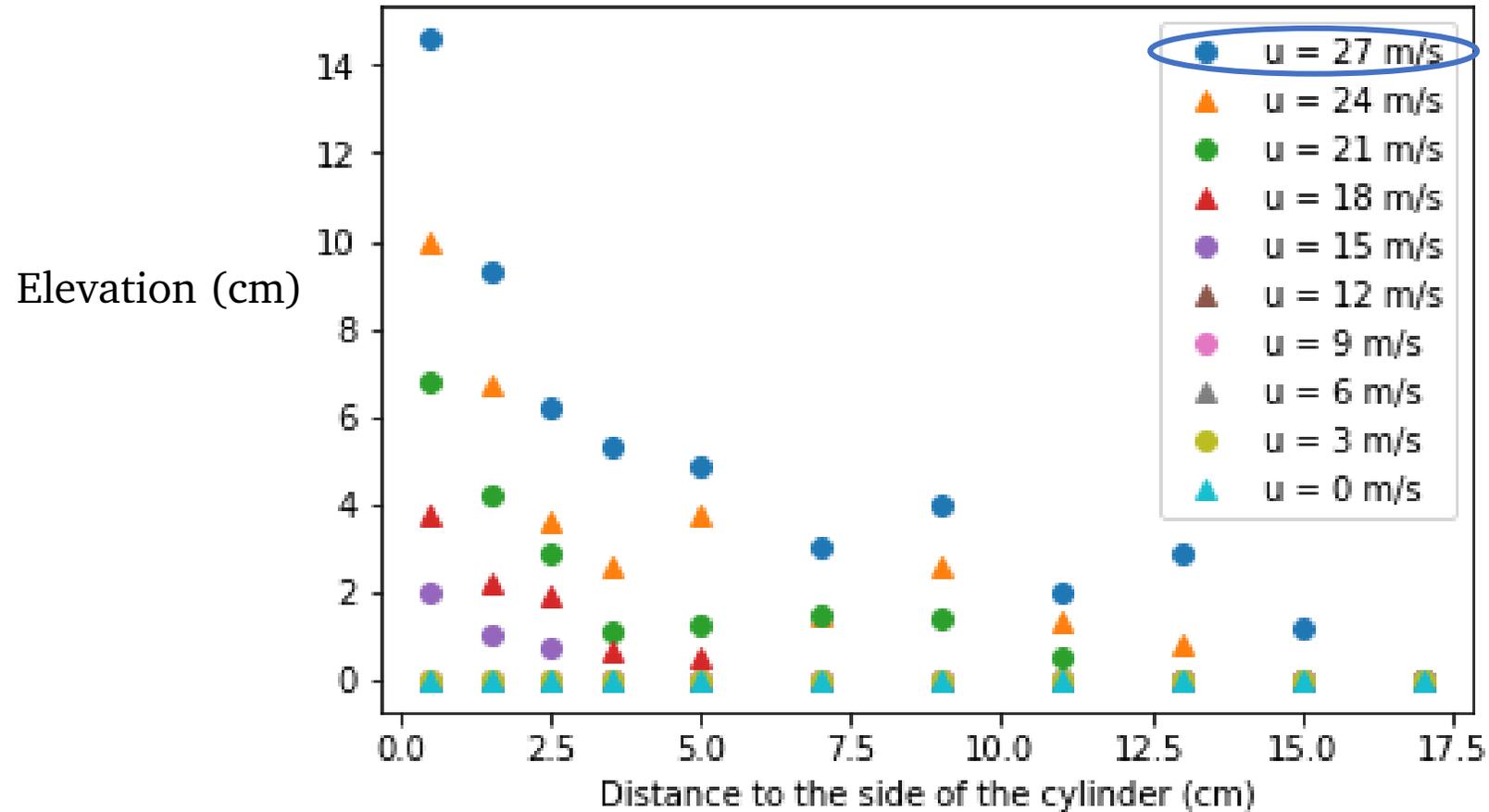
Melting Ratio (%)





Estimation of the curve

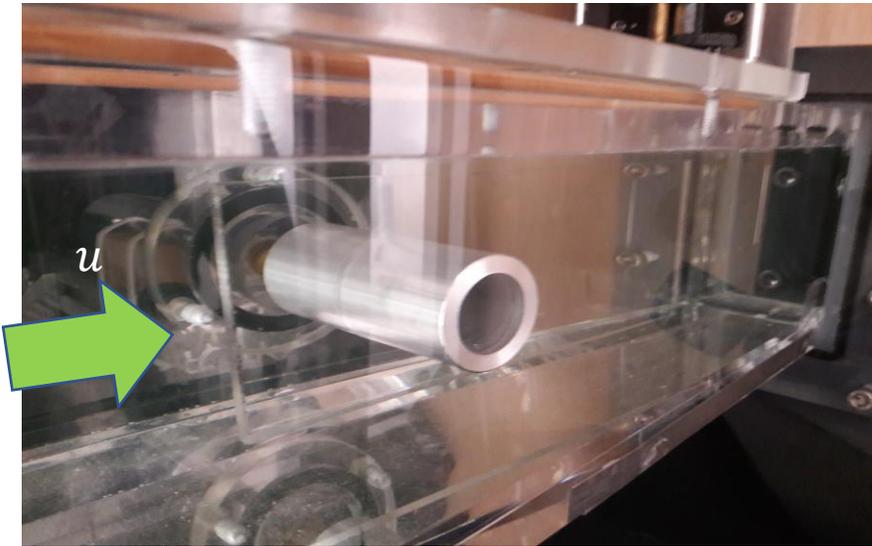
Size of the vortex





Estimation of the curve

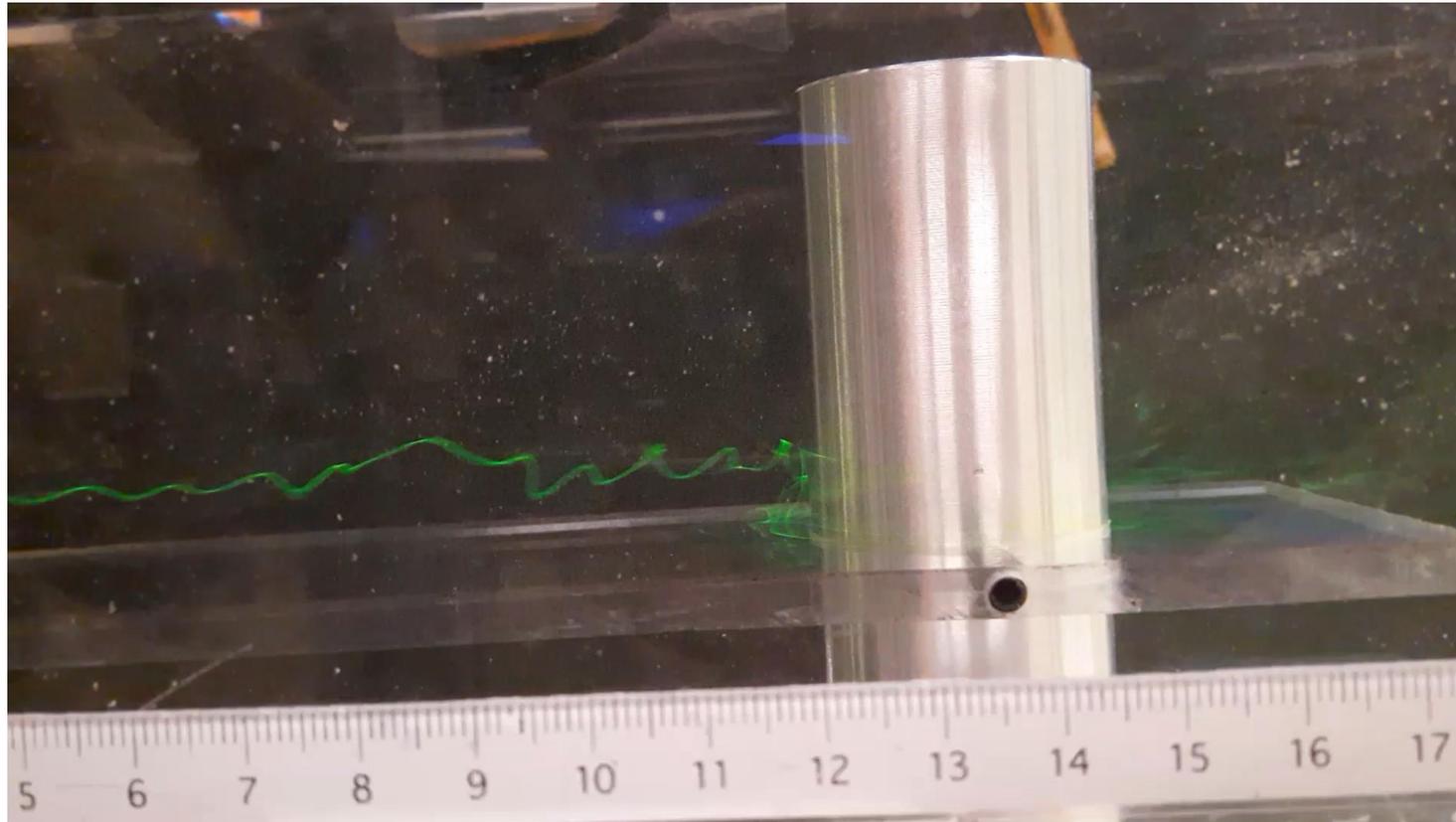
Visualizing the vortex





Estimation of the curve

Visualizing the vortex





Why does ice melt ?

Wind

Sublimation

Heat transfer

Origin: Vapor pressure gradient

$$\Phi_{sublimation} = h_m(\mathbf{u}) (q_{sat}(T) - q_{wind})$$

(q_{sat}, q_{wind} : mass of water per m^3 of air)

Sherwood number:

$$Sh = \frac{h_m L}{D_m} = 0.03 \times Re^{4/5} \times Sc^{1/3}$$

h_m : mass transfer coefficient

~

Origin: Temperature gradient

$$\Phi_{heat} = h(\mathbf{u}) (T_{wind} - T_{ice})$$

Nusselt number:

$$Nu = \frac{hL}{D} = 0.002 \times Re^{4/5} \times Sc^{1/3}$$

h : heat transfer coefficient

$$\frac{\Phi_{sublimation}}{\Phi_{heat}} \gg 1$$