

1 Article

2 Validation of RANS modelling for wave interactions 3 with sea dikes on shallow foreshores using a large- 4 scale experimental dataset

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20 **Abstract:** In this paper a RANS solver, interFoam of OpenFOAM®, is validated for wave
21 interactions with a dike, including promenade and vertical wall, on a shallow foreshore. Such a
22 coastal defence system is comprised of both an impermeable dike and a beach in front of it, forming
23 the shallow foreshore depth at the dike toe. This case necessitates the simulation of several processes
24 simultaneously: wave propagation, -breaking over the beach slope, and -interactions with the sea
25 dike, consisting of wave overtopping, bore interactions on the promenade, and bore impacts on the
26 dike-mounted vertical wall at the end of the promenade (storm wall or building). The validation is
27 done using rare large-scale experimental data. Model performance and pattern statistics are
28 employed to quantify the ability of the numerical model to reproduce the experimental data. In the
29 evaluation method, a repeated test is used to estimate the experimental uncertainty. The solver
30 interFoam is shown to generally have a very good model performance rating. A detailed analysis of
31 the complex processes preceding the impacts on the vertical wall proves that a correct reproduction
32 of the horizontal impact force and pressures is highly dependent on the accuracy of reproducing
33 the bore interactions.

34 **Keywords:** validation; wave modelling; shallow foreshore; dike-mounted vertical wall; wave
35 impact loads; OpenFOAM
36

37 1. Introduction

38 Low elevation coastal zones often have mildly to steeply-sloping sandy beaches as part of their
39 coastal defence system. For countries in north-western Europe, coastal urban areas typically have
40 high-rise buildings close to the coastline. These buildings are usually fronted by a low-crested, steep-
41 sloped and impermeable sea dike with a relatively short promenade, where the long (nourished)
42 beach in front of the dike acts as a mildly sloping shallow foreshore. This type of coastal defence
43 system therefore combines hard and soft coastal protection against flooding. Such hybrid approaches

44 are regarded by the IPCC with high agreement as a promising way forward in terms of response to
45 sea level rise [1]. Along the cross-section of this hybrid beach-dike coastal defence system, storm
46 waves undergo many transformation processes before they finally hit the buildings on top of the
47 dike. Along the shallow waters of the mildly sloping foreshore in front of the dike, sea/swell or short
48 waves (hereafter SW, $O(10^1 \text{ s})$) shoal and eventually break, transferring energy to both their super-
49 and subharmonics (or long waves: hereafter LW, $O(10^2 \text{ s})$) by nonlinear wave-wave interactions.
50 Further pre-overtopping hydrodynamic processes along the mildly sloping foreshore include: wave
51 dissipation by breaking (turbulent bore formation) and bottom friction, reflection against the
52 foreshore and dike, and wave run-up on the dike slope. Finally, waves overtop the dike crest and
53 post-overtopping processes include: bore propagation on the promenade, bore impact on a wall or
54 building, and reflection back towards the sea interacting with incoming bores on the promenade.

55 For the (structural) design of storm walls or buildings on such coastal dikes, the wave impact
56 force expected for specific design conditions needs to be estimated. Semi-empirical formulas, mostly
57 based on physical model tests, are commonly used in practice to assess wave forces and pressures on
58 coastal defences, at least in a preliminary design phase. However, semi-empirical formulas are
59 usually restricted within very specific ranges of application, currently limiting force prediction to
60 dikes with deep foreshore depths [2,3]. Such formulas do exist for dikes with very/extremely shallow
61 foreshore depths as well [4,5], but their application is also strictly limited. For the final design,
62 therefore, often detailed experimental campaigns are required [6]. Alternatively, during the last
63 decade numerical modelling of these combined processes has become feasible [7–11,3]. Numerical
64 modelling is also able to provide a detailed and accurate assessment of a specific case. Moreover,
65 numerical models can provide information on physical quantities that are difficult to measure in a
66 scaled model or in prototype (e.g. detailed velocity fields, pressure distributions, etc.).

67 To study fully two-dimensional vertical (2DV) complex fluid flows, Computational Fluid
68 Dynamics (CFD) techniques are typically applied. Relatively new mesh-free Lagrangian numerical
69 methods, such as Smoothed Particle Hydrodynamics (SPH) [12] and the particle finite element
70 method (PFEM) [13], have been recently validated and applied to several coastal engineering
71 problems [14,15,9,16,17], showing much promise. However, differently from Eulerian grid-based
72 methods, multi-phase air-fluid SPH models are still quite scarce and have a high computational cost
73 [18]. The more traditional Eulerian numerical methods are already more consolidated. For example,
74 volume-of-fluid methods (VOF) based on the Reynolds-Averaged Navier-Stokes equations (RANS)
75 have been widely employed during the last decades. Using RANS models, processes such as wave
76 transformation [19,8,20], wave overtopping [21,7,22], and wave impact on coastal structures [23–26,3]
77 have been modelled and validated, but never before at the same time (to the knowledge of the
78 authors). They are computationally very expensive to apply, but have shown their value particularly
79 for wave-structure interaction phenomena involving complex geometries. In addition, two-phase
80 water-air RANS models allow taking the effects of air entrapment on the wave impact processes into
81 account [27,28].

82 Validation of numerical models is crucial before they can be reliably applied. Even though plenty
83 of works have been published on numerical modelling and validation of individual processes
84 previously listed, there is still a lack of literature about RANS model validation for wave impacts on
85 sea dikes and dike-mounted walls in presence of a very shallow foreshore. The main goal of this
86 paper is to validate a two-phase (water-air) RANS model for this specific case. Such a modelling
87 approach is deemed necessary to fully resolve the 2DV complex fluid flows of overtopped waves and
88 bore interactions on top of the promenade. The RANS solver (interFoam) for two incompressible
89 fluids within the open source CFD toolbox OpenFOAM® is chosen because of its increasing
90 popularity for application to wave-structure interactions. Validation of this numerical model is done
91 by reproducing large-scale experiments of overtopped wave impacts on coastal dikes with a very
92 shallow foreshore from the WALOWA project [29]. The large-scale nature of these experiments
93 reduces the scale effects significantly compared to small-scale experiments, which can be particularly
94 of importance to the wave impacts on the dike-mounted vertical wall, especially in case of plunging
95 breaking bore patterns and impulsive impacts [30].

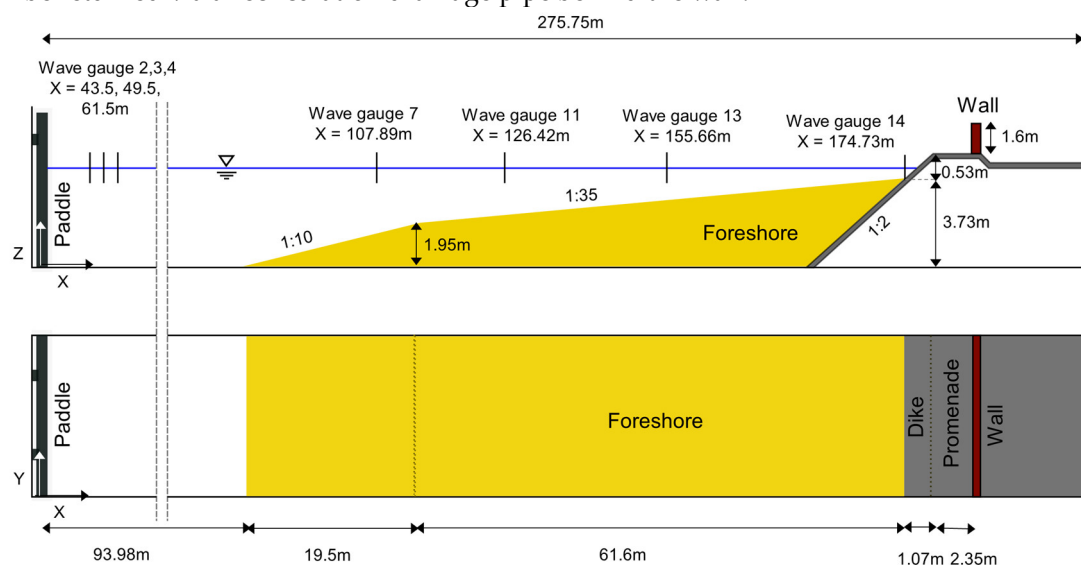
96 The paper is structured as follows. First the methods used in the paper are explained in section
 97 2, starting with the experimental model setup and a description of the tests used for the validation.
 98 This is followed by a description of the applied RANS model and the numerical model setup. Finally,
 99 the statistical model performance methods applied in this study are discussed. Next, in section 3 the
 100 results of the qualitative and quantitative numerical model validation are provided, including a
 101 comparison of model snapshots at key time instants during impacts on the vertical wall. This is finally
 102 followed by section 4 with a discussion on these results and the conclusions in section 5.

103 2. Methods

104 2.1. Large-Scale Laboratory Experiments

105 The laboratory experiments (Froude length scale 1/4.3) were done during the research project
 106 WALOWA (Wave LOads on WALLs) in the Deltares Delta Flume, which is 291 m long, 9.5 m deep
 107 and 5 m wide. This wave flume is equipped with a piston-type wave maker capable of up to second-
 108 order wave generation (in the frequency range 0.02 Hz - 1.50 Hz) and includes Active Reflection
 109 Compensation (ARC), which is an Active Wave Absorption (AWA) system to minimise reflections
 110 against the wave paddle. For a detailed description of the model setup, reference is made to Streicher
 111 et al. [29]. The WALOWA dataset is open access and is described by Kortenhaus et al. [31].

112 The model geometry consisted of a moveable sandy foreshore with a transition slope of 1:10 and
 113 a slope of 1:35 up to the toe of the dike (Figure 1). The smooth impermeable concrete dike had a front
 114 slope of 1:2, a promenade width of 2.35 m with an inclination of 1:100 in order to help drain the water
 115 in case of wave overtopping, and finally a 1.60 m high wall. The wall height was designed to be high
 116 enough to prevent wave overtopping during testing, but small amounts of overtopped water could
 117 still be returned via a recirculation drainage pipe behind the wall.



118
 119 **Figure 1.** Overview of the geometrical parameters of the wave flume and WALOWA model set-up,
 120 with indicated wave gauge locations.

121 The WALOWA dataset includes both bichromatic and irregular wave tests. For validation of the
 122 numerical model, the bichromatic wave test Bi_02_6 (EXP) and its repetition Bi_02_6_R (REXP) were
 123 selected (Table 1). The bichromatic wave tests have the advantage to be relatively short in time, while
 124 still considering the effects of wave dispersion and bound LWs, and is therefore more representative
 125 of irregular waves than monochromatic waves. In this way, even numerical models with a high
 126 computational demand are able to simulate the tests at a reasonable amount of computational time.
 127 This specific bichromatic wave test was chosen because it is the only test that was conducted shortly
 128 after a foreshore profile measurement and at the same time immediately followed by its repetition
 129 and another foreshore profile measurement [32]. Since these bichromatic wave tests are relatively

130 short in duration and only limited changes ($O(10^{-2}$ m)) were noted between the profile measurements
 131 before and after, a fixed bed is a reasonable assumption for the numerical modelling. In addition, the
 132 repeated test makes a validation of the numerical model possible relative to the experimental
 133 uncertainty.

134 **Table 1.** Hydraulic parameters for the WALOWA bichromatic wave test (EXP) and its repetition
 135 (REXP): h_o is the offshore water depth, h_t the water depth at the dike toe, $H_{m0,o}$ the incident offshore
 136 significant wave height, R_c the dike crest freeboard, f_i the SW component frequency, a_i the SW
 137 component amplitude and $\delta (= a_2/a_1)$ the modulation factor.

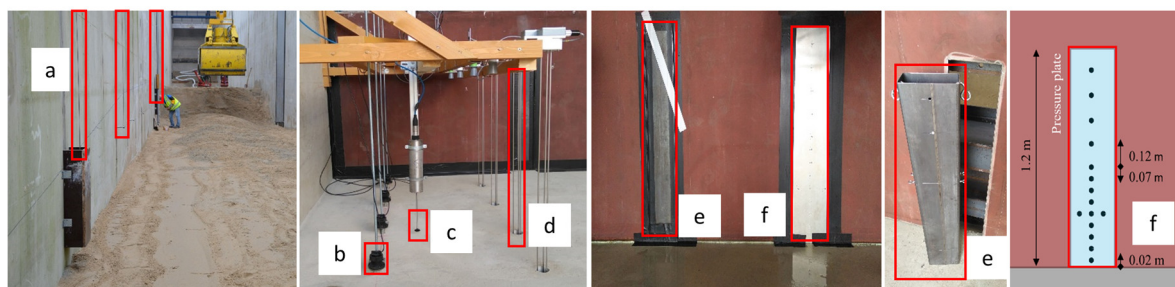
TestID	Duration	h_o	h_t	$h_t/H_{m0,o}$	R_c	f_1	a_1	f_2	a_2	δ
[-]	[s]	[m]	[m]	[-]	[m]	[Hz]	[m]	[Hz]	[m]	[-]
Bi_02_6 (EXP) & Bi_02_6_R (REXP)	209	4.14	0.43	0.33	0.117	0.19	0.45	0.155	0.428	0.951

138
 139 During these tests, three bichromatic wave groups were generated with first order wave control
 140 over 125 s, including 10 s of tapering at the beginning and end of the wave generation. Plunging
 141 breakers occurred on the 1:10 transition slope (i.e. deep water Iribarren number $\xi_0 = \tan\alpha/(H/L_0)^{1/2}$
 142 with α the foreshore slope angle, H the wave height and L_0 the deep water wave length [33]: $0.5 < \xi_0$
 143 $\approx 0.7 < 3.3$) and spilling breakers on the 1:35 foreshore slope ($\xi_0 \approx 0.2 < 0.5$). Considering this was a test
 144 of a dike with a very shallow foreshore depth (Table 1: $0.3 < h_t/H_{m0,o} < 1.0$ [34]), the wave energy at the
 145 toe of the dike was dominated by LW energy.

146 The measurement setup consisted of instruments to measure the water surface elevation along
 147 the flume and on the promenade, the velocity of the overtopped flow on the promenade and the
 148 impact pressure and force on the vertical wall (Figure 2). All measurements were sampled at 1000 Hz
 149 frequency and synchronized in time.

150 The water surface elevation η (with the vertical origin at $z = h_o$) was measured with resistance
 151 type wave gauges (WG) deployed at seven different locations along the Delta Flume side wall (Figure
 152 1 and Figure 2a). WG02-WG04 were installed over the flat bottom part of the flume close to the wave
 153 paddle. These wave gauges were positioned to allow a reflection analysis following the method of
 154 Mansard and Funke [35]. WG07 was installed along the transition slope; WG11 and WG13 along the
 155 foreshore slope. WG14 was installed close (~ 0.35 m) to the dike toe. The data of WG11 is not
 156 considered further in the present analysis, because of faulty data. Furthermore, to remove unwanted
 157 noise in the η signals measured by the other WG's from the wave paddle up to the dike toe, a low-
 158 pass 3rd order Butterworth filter with a cut-off frequency of 1.50 Hz was applied. This frequency is
 159 well above the frequencies of the super-harmonics of the primary waves and frequency components
 160 due to triad interactions between the primary components and the difference frequency, which gain
 161 energy in the shoaling and surf zone [36].

162 Flow layer level measurements η on the promenade were obtained by four resistance type Water
 163 Level Distance Meters (WLDM01 – WLDM04, Figure 2d). Flow velocity measurements on the
 164 promenade were obtained by four Paddle Wheels (PW01 – PW04, Figure 2b), measuring the
 165 horizontal flow velocity U_x in one direction (i.e. towards the wall) 0.026 m above the promenade.
 166 Additionally, a bidirectional Electromagnetic Current Meter (ECM, Figure 2c) was installed at the
 167 same cross-shore location as WLDM02 and PW02 to get directional information of the incoming or
 168 reflected flow. The ECM disc was positioned 0.03 m above the promenade and sampled the
 169 horizontal velocity at 16 Hz. Further detailed information on the sensor setup on the promenade and
 170 the post-processing of the η and U_x data measured on top of the promenade was provided by
 171 Cappiotti et al. [37]. During return flow, positive U_x values were possibly incorrectly measured by
 172 the PWs, indicated by the ECM that measured negative U_x values during return flow (compared to
 173 the measurements of the co-located PW02). This will be further discussed when comparing with the
 174 numerical model result (section 3.1). However, no such co-located measurements are available for
 175 other paddle wheels than PW02, so that no correction of the PW measurements during return flows
 176 was attempted.



177

178 **Figure 2.** (a) WGs deployed along the flume side wall to measure η ; (b) PWs; (c) ECM to measure U_x ;
 179 (d) WLDMs installed on the promenade to measure η ; (e) Hollow steel profile attached to two LCs
 180 and (f) aluminium plate equipped with pressure sensors (PS) to measure F_x and p respectively.

181 The overtopped wave impacts on the wall were measured by horizontal force F_x and pressure p
 182 measurement systems integrated into the wall. The horizontal impact force was measured by two
 183 compression type load cells (LC) connecting the same hollow steel profile to the very stiff supporting
 184 structure (Figure 2e). Impact pressures were measured by 15 pressure sensors (PS). The first 13 PSs
 185 were spaced vertically over a metal plate flush mounted in the middle section of the steel wall, with
 186 PS14 and PS15 placed horizontally next to PS05 or the fifth PS from the bottom (Figure 2f). The initial
 187 post-processing of the F_x and p signals, including baseline correction and filtering, is discussed by
 188 Streicher [38]. Additional filtering is applied to remove the high frequency oscillations caused by
 189 stochastic processes during dynamic or impulsive impacts, so that the signal can be reproduced by a
 190 deterministic numerical model [39]. To achieve this, an additional 3rd order Butterworth low-pass
 191 filter with a cut-off frequency of 6.22 Hz was necessary. This corresponds to a cut-off frequency of 3.0
 192 Hz at prototype scale, which is still well above the natural frequency of about 1.0 Hz for typical
 193 buildings found along e.g. the Belgian coast [40]. Furthermore, local spatial variability over the width
 194 of the flume of the resultant F_x (i.e. derived from the LCs and pressure integrated) and p (i.e. PS05,
 195 PS14 and PS15) time series was found to be low (not shown). This spatial variability over the width
 196 of the experimental flume is therefore further neglected in the quantitative numerical model
 197 validation: for F_x the LC-derived signal is used and for p the PS05 signal is used.

198 2.2. Numerical model

199 2.2.1. Model description

200 In this work OpenFOAM v6 [41] is applied and validated, or more specifically interFoam, a
 201 solver of the Reynolds-Averaged Navier-Stokes (RANS) equations, where the advection and
 202 sharpness of the water-air interface is handled by an algebraic VOF method [42] based on MULES
 203 [43–45]. InterFoam with MULES has already been successfully applied before for wave propagation
 204 [44], wave breaking [20,46–49], wave run-up [20,49], wave overtopping [50,51] and bore impact on a
 205 vertical wall [26].

206 Several open source contributions of boundary conditions for wave generation and absorption
 207 exist for interFoam, of which the main developments are: IHFOAM [52], olaFlow [53] and
 208 waves2Foam [54]. In the present study, olaFlow was chosen, which was found to be the most
 209 computational efficient [52,55,56] and feature complete package at the time of the simulations
 210 presented in this paper.

211 The turbulence is modelled by the $k-\omega$ SST turbulence closure model [57], which has been shown
 212 to be one of the most proficient in modelling wave breaking [46]. Two-equation turbulence closure
 213 models are known to cause over-predicted turbulence levels beneath computed surface waves,
 214 leading to unphysical wave decay for wave propagation over constant water depth and long distance
 215 [58,59,48]. Turbulence modelling was therefore stabilized in nearly potential flow regions by Larsen

216 and Fuhrman [48], with their default parameter values [60]. Hereafter, the OpenFOAM numerical
217 model as presented here is simply referred to as OF.

218 2.2.2. Computational domain and mesh

219 Wave breaking is an inherently three-dimensional (3D) process due to the formation of 3D
220 vortices extending obliquely downward in the inner surf zone [61]. Even so, many examples exist
221 where the wave kinematics during wave breaking could be approximated well by vertical two-
222 dimensional (2DV) RANS modelling [62,19,8,63,46,49,47,48]. To reduce the computational time as
223 much as possible, OF is therefore applied in a 2DV configuration (i.e. cross-shore section of the wave
224 flume).

225 The OF model domain (Figure 3) starts at the wave paddle zero position ($x = 0.00$ m) and ends
226 on top of the vertical wall ($x = 178.80$ m). The bottom boundary is at its lowest point ($z = 0.00$ m) along
227 the flume bottom between the wave paddle and the foreshore toe, and extends up to $z = 7.20$ m, well
228 above the maximum measured surface elevations along the flume. The bottom is further defined by
229 the measured foreshore and dike geometry as described in section 2.1. The vertical wall is included
230 up to its height of 1.60 m including the top which was given a slight inclination towards the model
231 boundary to allow overtopped water (limited to mainly spray in this case) to exit the model domain.

232 The computational domain is discretised into a structured grid. To optimise the computational
233 time, a variable grid resolution is applied, where a higher resolution is defined only where it is
234 necessary. This is mostly the areas of the model domain where the water-air interface is expected to
235 pass [45,55]. The expected location of the free surface along the flume during the entire test was
236 estimated first by a fast preliminary one-layer depth-averaged SWASH calculation (not shown: see
237 [64] for the SWASH model setup description). The minimum and maximum η along the flume and
238 over the complete test duration were used from the SWASH model result to define areas in which
239 mesh refinement should be done. These locations are delineated by the dotted lines in Figure 3,
240 defining several areas around the still water level (SWL). In front of the wave paddle, the refinement
241 area is slightly higher to accommodate the stabilisation of the newly generated waves, after which
242 the refinement zone can decrease in height when the waves have fully developed. Then the
243 refinement area is increased in height again to allow room for wave shoaling and incipient wave
244 breaking on the foreshore. The upper limit can subsequently be lowered again due to wave breaking,
245 but the lower limit is extended to include the bottom boundary. This is to resolve properly the
246 entrained air pockets that have been shown to travel towards the bottom during the breaking process
247 in the inner surf zone [65]. The height of the refinement zone on the dike was defined based on the
248 maximum measured water level in the experiment by the WLDM's on the promenade and extended
249 to the upper model boundary along the vertical wall to resolve the run-up and splashing against the
250 vertical wall.

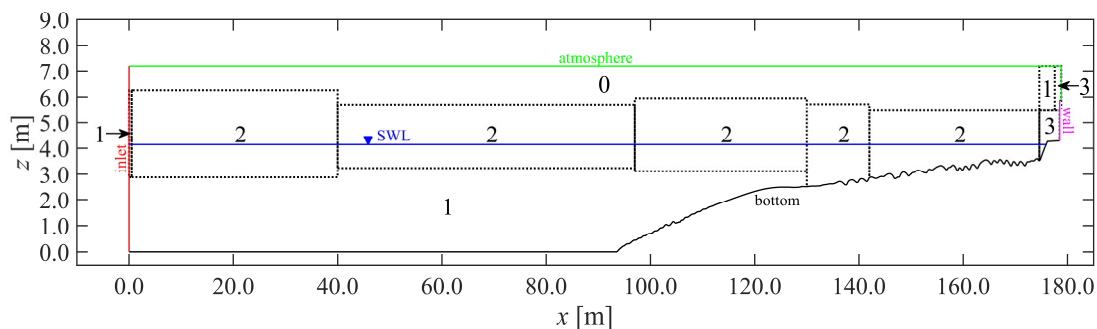
251 In terms of the grid cell size in these refinement zones, about 20 cells are typically recommended
252 over the wave height H of a regular wave (i.e. $H/\Delta z = 20$, with Δz being the vertical cell size) [45,56].
253 Applied to the wave heights of the primary wave components of the bichromatic wave in Table 1, a
254 minimal vertical cell size of $\Delta z = 0.045$ m to 0.043 m is obtained. Smaller wave heights in the
255 bichromatic wave group are less resolved with this choice, but this is deemed acceptable because of
256 their relatively low steepness. A value of $\Delta z = 0.045$ m was chosen, because the water depth at the
257 wave paddle h_0 is divisible by it (i.e. $h_0/\Delta z = 4.14/0.045 = 92$), meaning that the SWL can lie perfectly
258 along cell boundaries. Or in other words, α -values between 0 and 1 are thereby minimised at the start
259 of the simulation, which simplifies the initialisation of the SWL and is beneficial for an effectively still
260 SWL at the start of the simulation.

261 The mesh maintains an aspect ratio $\Delta x/\Delta z$ of 1 (with Δx being the horizontal cell size) throughout
262 the entire computational domain, which has been shown necessary for accuracy [54,65,45] and
263 numerical stability in this study. One exception is a higher aspect ratio along the bottom and wall,
264 where layers were locally added to the mesh to resolve the boundary layer. Six layers were added
265 over the vertical cell size along those boundaries, with a growth rate of 1.2, leading to a maximum
266 aspect ratio of 18.

267 Outside the refinement zones, in the air and water phases, the mesh can be coarser [45,56]. The
 268 structured mesh was given a base grid resolution of 0.18 m. This base resolution is multiplied by a
 269 refinement ratio r , here defined as:

$$r = \frac{1}{2^\beta}, \quad (1)$$

270 in which β signifies the refinement level. Each refinement level effectively refines every cell into four
 271 new cells. The applied refinement levels are provided for each mesh subdomain in Figure 3. For the
 272 air in the model domain the base resolution was assumed ($\beta=0$), except for a small area over the dike
 273 ($\beta=1$). In the water phase, refinement level 1 was assumed ($\Delta x = \Delta z = 0.09$ m) and was further refined
 274 in the zone of the surface elevation up to the dike toe (level 2 or $\Delta x = \Delta z = 0.045$ m). Close to the inlet
 275 boundary, however, a lower refinement level was necessary for numerical stability ($\beta=1$) over a very
 276 short distance ($0 \text{ m} < x < 0.50$ m) where locally high water velocities (i.e. low Courant numbers and
 277 low time steps) at the interface can occur due to the wave generation. On the dike up to the wall, the
 278 mesh was refined even more (level 3 or $\Delta x = \Delta z = 0.0225$ m) to resolve thin layer flows, the complex
 279 flows of bore interactions, and impacts on the vertical wall. In addition, a refinement level 3 was
 280 necessary to resolve the experimental pressure sensor locations along the vertical wall.



281
 282 **Figure 3.** Definition of the OF 2DV computational domain, with coloured indication of the model
 283 boundary types. The still water level (SWL) is indicated in blue ($z = 4.14$ m). The number in each of
 284 the mesh subdomains of the model domain (demarcated by black dotted lines) is the refinement level
 285 β applied in each subdomain (for $\beta = 0, 1, 2$ and 3 : $\Delta x = \Delta z = 0.18$ m, 0.09 m, 0.045 m and 0.0225 m).
 286 Note: the axes are in a distorted scale.

287 The mesh was generated by applying the *cartesian2DMesh* algorithm of cfMesh [66], which resulted
 288 in a mesh with 318,381 cells, for the refinement levels indicated in Figure 3.

289 The adaptive time stepping is controlled by a predefined maximum Courant number $maxCo$ ($Co = \Delta t |U| / \Delta X$, where Δt is the time step, $|U|$ is the magnitude of the velocity through that cell and ΔX
 290 is the cell size in the direction of the velocity [67]) and a maximum Courant number in the interface
 291 cells $maxAlphaCo$. Generally $maxCo = maxAlphaCo$ is chosen, as well as in this paper. Larsen et al. [44]
 292 have shown that a relatively low $maxCo$ (~ 0.05) is necessary to obtain a stable wave profile over more
 293 than five wave periods propagation duration. Here, however, a $maxCo$ of 0.25 is used to balance the
 294 accuracy and computational costs. Since the primary waves of the bichromatic wave group only
 295 propagate over about three wave lengths up to the mean breaking point location ($x_b = \sim 120$ m), this is
 296 considered an acceptable assumption. Both the refinement level in the refinement zones around the
 297 surface elevation zones (β_{sez}) and the $maxCo$ were verified in a convergence analysis (Appendix A).
 298

299 2.2.3. Boundary conditions

300 Since the model domain represents a 2DV simulation, no solution is necessary in the y -direction
 301 and the lateral boundaries of numerical wave flume were assigned an “empty” boundary condition.
 302 Non-empty boundary conditions were defined for the remaining boundaries in the xz -plane (Figure
 303 3).

304 The bichromatic waves from Table 1 were generated at the inlet by applying a Dirichlet-type
305 boundary condition: the experimental wave paddle velocity was imposed. The paddle displacement
306 time series is used by olaFlow to calculate the wave paddle velocity by a first order forward derivative
307 [68]. Since the reflection in the numerical wave flume is expected to behave close to, but not exactly
308 the same as in the experiment, the theoretical paddle displacement without ARC was selected and
309 the AWA by olaFlow was activated instead. In addition to the paddle displacement, the surface
310 elevation at the wave paddle is provided, which allows olaFlow to trigger the AWA with fewer
311 assumptions [68]. The AWA implementation in olaFlow is most effective for shallow water waves.
312 The primary components of the bichromatic wave group are intermediate waves for the water depth
313 at the wave paddle, but their reflection is expected to be low, since most of their wave energy
314 dissipates over the foreshore in the surf zone. However, reflected free long (infragravity) waves are
315 expected to be non-negligible (section 3.2). They are shallow water waves and are by definition
316 absorbed well by the AWA system in olaFlow, preventing their re-reflection and therefore replicating
317 the behaviour of the ARC in the experiment.

318 Both the bottom and wall boundaries are fixed boundaries, including the sandy foreshore
319 (section 2.1), along which the velocity vector field U has a Dirichlet-type boundary condition ($U = (0,$
320 $0, 0)$ m/s), while the pressure p and α are given a Neumann boundary condition. Along the foreshore,
321 dike and wall, no-slip boundary conditions are assumed and a continuous scalable wall function
322 based on Spalding's law (Spalding, 1961) is implemented. The six boundary layers that were
323 previously added in the mesh along these no-slip fixed boundaries make sure that the scalable wall
324 function criterion for the dimensionless wall distance z^+ (i.e. $1 < z^+ < 300$) is complied. For the
325 remaining boundary conditions, initial conditions and solver settings, the same settings were chosen
326 as those reported by Devolder et al. [47].

327 The OF simulations were run in parallel on a 24-core Intel Xeon E5-2680 @ 2500 MHz computer
328 with 128 GB of RAM. The scotch decomposition algorithm was used to divide the mesh into equal
329 amounts of cells for each processor, while minimising the number of processor boundaries [41]. The
330 cells along the inlet patch were forced onto the same processor, which benefits the computational
331 efficiency. On this setup, the simulation required a CPU time of about 85h.

332 2.2.4. Data sampling and processing

333 The same data was sampled in OF at the same cross-shore locations as in the experiment (section
334 2.1). Applying the same sampling frequency of 1000 Hz in OF, however, would increase the
335 calculation time to unpractical levels because it affects the time stepping. Instead a sampling
336 frequency of 80 Hz was maintained throughout, which is a compromise between temporal resolution
337 of the output data and calculation time.

338 To obtain η in OF, α was recorded at a fixed interval over a vertical line at each wave gauge
339 location. In post-processing, η was then obtained by vertical integration of α , thereby excluding air
340 inclusions produced in the surf zone, but taking into account all water volumes (i.e. even air-borne
341 water, e.g. in case of plunging waves, spray,...). This corresponds best to how η in the experiment
342 was measured: resistive wave gauges give a response proportional to the wire wet length [69],
343 thereby similarly excluding air pockets. However, it is acknowledged that still some uncertainty
344 remains on how resistive type wave gauges measure the free surface in the presence of air-water
345 mixtures along the gauge. This could lead to discrepancies in the numerical-experimental model
346 comparisons in de surf zone and on top of the promenade [70].

347 The resulting numerical time series were filtered in the same way as the experimental data
348 (section 2.1) and were synchronised to the experimental time reference. The synchronisation was
349 done based on the η time series at the three most offshore located wave gauges (i.e. WG02-03-04) by
350 means of a cross-correlation. The obtained numerical-experimental time lags for each of these WG
351 locations were subsequently averaged and rounded to the nearest multiple of the time series time
352 step. This time lag was then used to synchronise all numerical time series to the experimental time
353 reference. This makes sure that numerical errors (such as phase lag), which are important for model
354 validation, were retained.

355 Furthermore, to investigate the model performance for the SW and LW components separately,
 356 the η time series were separated into η_{SW} and η_{LW} by applying a 3rd order Butterworth high- and low-
 357 pass filter respectively. A separation frequency of 0.09 Hz was employed, which is in between the
 358 bound long wave frequency ($f_1 - f_2 = 0.035$ Hz) and the lowest frequency of the primary wave
 359 components ($f_2 = 0.155$ Hz).

360 2.3. Validation method

361 The validation of the numerical model OF to the large scale experiment EXP is done both
 362 qualitatively and quantitatively. The qualitative validation entails a comparison of time series of the
 363 main measured parameters. However, it is recommended to apply model performance statistics as
 364 well for a more quantified and objective validation [71]. Therefore, general numerical model
 365 performance will be evaluated by applying a skill score or dimensionless measure of average error,
 366 such as Willmott's refined index of agreement d_r [72]:

$$d_r = \begin{cases} 1 - \frac{MAE}{cMAD}, & MAE \leq cMAD \\ \frac{cMAD}{MAE} - 1, & MAE > cMAD \end{cases} \quad (2)$$

367 where c is a scaling factor and is taken equal to 2, to obtain a balance between the number of
 368 deviations evaluated within the numerator and within the denominator of the fractional part of d_r ,
 369 MAE is the mean-absolute-error defined by:

$$MAE = \frac{1}{N} \sum_{i=1}^N |P_i - O_i|, \quad (3)$$

370 with N the number of samples in the time series, and P the predicted time series together with the
 371 pair-wise-matched observed time series O (for $i = 1, 2, \dots, n$), and MAD is the mean-absolute deviation:

$$MAD = \frac{1}{N} \sum_{i=1}^N |O_i - \bar{O}|, \quad (4)$$

372 where the overbar represents the mean of the time series. This model performance index d_r is
 373 bounded by [-1.0, 1.0] and, in general, more rationally related to model accuracy than other existing
 374 model performance indices or skill scores. For the purposes in this paper, d_r is used as a general
 375 measure of the model performance and a d_r value of 0.5 is already considered to be a poor model
 376 performance. Since it is a single measure of model performance, it can be more easily used to evaluate
 377 for example the spatial model performance over the length of the wave flume.

378 Because a repetition of the selected experimental test is available (REXP), d_r can be evaluated
 379 between REXP and EXP as well. This can serve as a limit above which a d_r value of the numerical
 380 model signifies that the numerical model performance cannot be improved beyond the experimental
 381 model uncertainty due to model effects, etc. Therefore, similar to the relative errors as defined by van
 382 Rijn et al. [73], a *relative* refined index of agreement d'_r is proposed here which provides the
 383 performance of the numerical model relative to the experimental model uncertainty:

$$d'_r = \begin{cases} 1 - \frac{MAE_{num} - MAE_{rexp}}{cMAD} = 1 - (d_{r,num} - d_{r,rexp}), & MAE_{num} - MAE_{rexp} \leq cMAD \\ \frac{cMAD}{MAE_{num} - MAE_{rexp}} - 1 = (d_{r,num} - d_{r,rexp}) - 1, & MAE_{num} - MAE_{rexp} > cMAD \end{cases} \quad (5)$$

384 where the subscripts *num* and *rexp* indicate that the statistic is evaluated respectively for the
 385 numerical and repeated experimental data, and c is again taken equal to 2. When the numerator
 386 $MAE_{num} - MAE_{rexp}$ is negative (i.e. < 0), the numerical error compared to the experiment is smaller than
 387 the experimental uncertainty, which means that the numerical model performance cannot be
 388 improved. In that case $MAE_{num} - MAE_{rexp} = 0$ is forced, so that $d'_r = 1$. A classification of model
 389 performance based on ranges of d'_r values and corresponding rating terminology is proposed in Table
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Table 2. Proposed classification of the relative refined index of agreement d'_r and corresponding rating.

d'_r classification [-]	Rating
0.90 – 1.00	<i>Excellent</i>
0.80 – 0.90	<i>Very Good</i>
0.70 – 0.80	<i>Good</i>
0.50 – 0.70	<i>Reasonable/Fair</i>
0.30 – 0.50	<i>Poor</i>
(-1.00) – 0.30	<i>Bad</i>

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To obtain more insight into where the error of the model originates from, pattern statistical parameters are considered as well. They are here explained in terms of what they represent for a time series of η . The first additional statistical parameter is the standard deviation σ , which is a measure of the wave energy or wave height of a η time series. The normalised standard deviation is given by:

$$\sigma^* = \frac{\sigma_p}{\sigma_o}, \tag{6}$$

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where σ_p and σ_o are the standard deviations of the predicted and observed time series, respectively. Another important statistical parameter is the bias B , given by:

$$B = \bar{P} - \bar{O}, \tag{7}$$

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The bias indicates whether the model under- or over-predicts the observation, but provides no further assurances on the accuracy of the model result. The bias represents the difference in wave setup between two η time series. It is normalised by the standard deviation of the observed time series:

$$B^* = \frac{B}{\sigma_o}, \tag{8}$$

405

And finally the correlation coefficient R , defined by:

$$R = \frac{\frac{1}{N} \sum_{i=1}^N (P_i - \bar{P})(O_i - \bar{O})}{\sigma_p \sigma_o}, \tag{9}$$

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which is a measure of the phase similarity between two time series and the wave periods in case of η time series.

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The length of the time series used for the analysis is based on the duration of the generated bichromatic waves including tapering (i.e. 125 s), beginning at the first time step when the baseline is first significantly exceeded (i.e. indicating arrival of the first wave). Since the experimental and numerical time series have different sampling frequencies, the time series with the highest sampling frequency was interpolated to the time steps of the time series with the lowest sampling frequency.

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For some locations where wetting and drying occurs (i.e. on the dike, promenade and vertical wall), the measurement regularly returned to the baseline or zero-line (Figure 4g-j, Figure 6 and Figure 7), meaning that as a bore passed by, reflected against the wall and ran back down the dike slope, intervals were created in the time series of (near-) zero values. Including these “non-event” times in the statistical analyses would bias the statistics by:

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- unnecessarily penalising the numerical model performance for an experimental measurement error. For example, in the experimentally measured and processed time series of p and F_x , often some residual instrumental noise or oscillations persisted during such non-event (or “dry”) times;
- unnecessarily rewarding the model performance towards (almost) perfect agreement. For example, during the time between impacts no water reaches the wall and model performance would be perfect during such times (disregarding measurement noise).

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It is therefore decided to focus the analysis on the event instances when the values of the time series (either experimental or numerical, to penalise phase differences or impacts not modelled by the

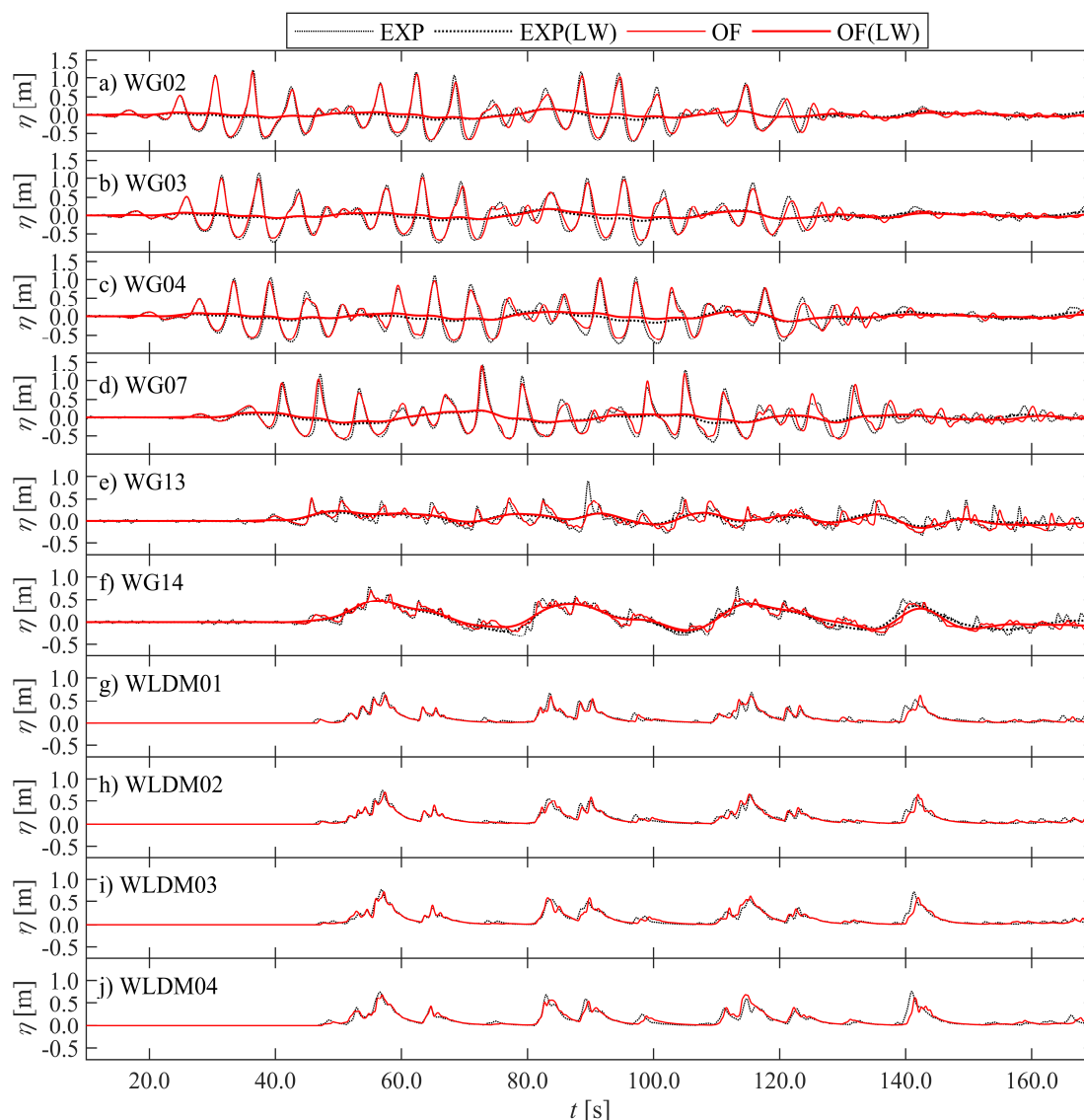
427 numerical model) is larger than a certain threshold above the baseline. The threshold for each such
 428 time series is chosen to be as low as possible, but higher than the residual noise in the experiment.

429 **3. Results**

430 *3.1. Time series*

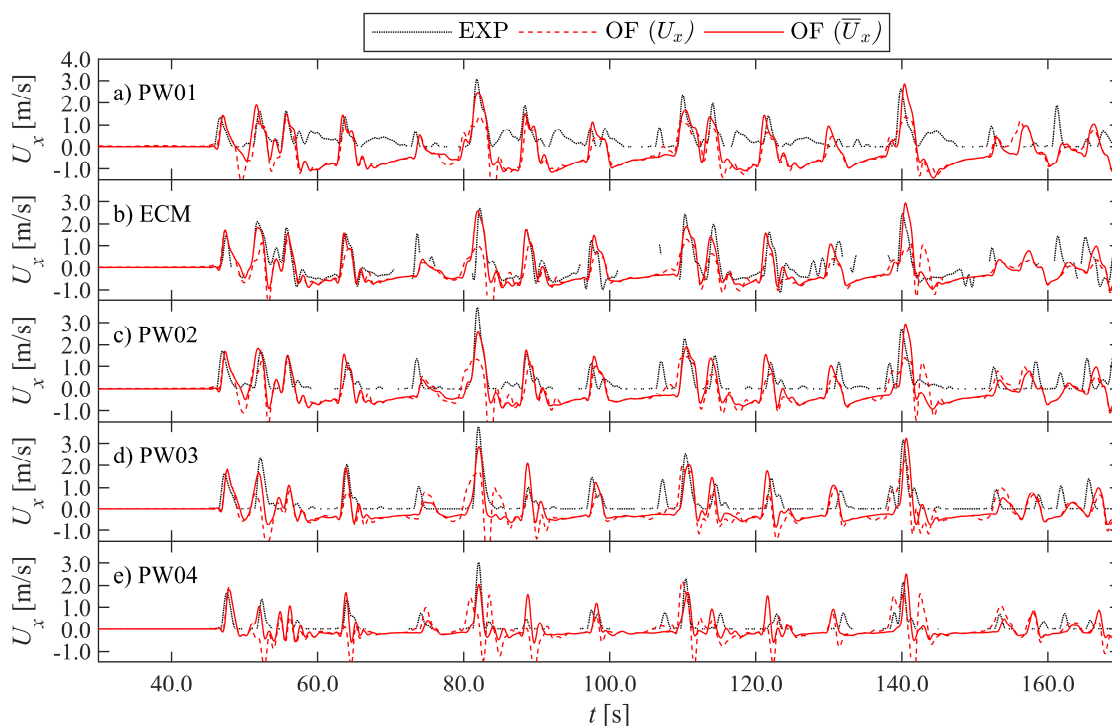
431 The numerical model results are first compared qualitatively in the time domain to the
 432 experimental measurements of test EXP. The surface elevations η are compared in Figure 6, the
 433 horizontal velocity U_x on the promenade in Figure 5, and the total horizontal force F_x and pressures
 434 p on the vertical wall in Figure 6 and Figure 7, respectively.

435 The η time series compare very well between OF and EXP (Figure 4), especially at the beginning
 436 of the simulation, but more discrepancies start to show over time and further along the flume.
 437 Overall, frequency dispersion, the non-linear wave transformation processes (i.e. SW shoaling
 438 (Figure 4d), breaking (Figure 4e-f), energy transfer to the subharmonic bound LW (Figure 4d-f),
 439 overtopping (Figure 4g), bore interactions and reflection processes (Figure 4g-j) seem to be well-
 440 represented by OF.



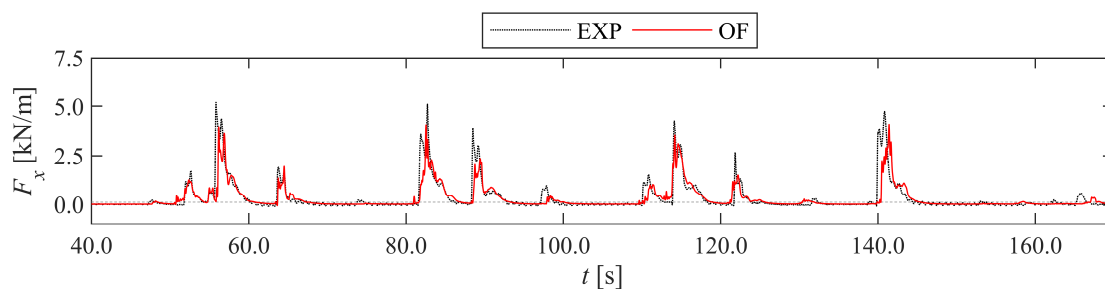
441
 442 **Figure 4.** Comparison of the η time series at all sensor locations, including η_{LW} in (a) – (f) (bold lines).
 443 The zero-reference is the SWL for (a) – (d) and the promenade bottom at the sensor location for (e) –
 444 (h).

445 The simulated U_x on top of the promenade appears to significantly underestimate the
 446 experimental measurements (Figure 5). This underestimation mostly disappears when using the OF
 447 depth-averaged velocity \overline{U}_x instead, which is done for the remainder of the validation. In addition,
 448 OF shows much better correspondence to the ECM than the PWs during return flow of a reflected
 449 bore ($U_x < 0$). This confirms that the PWs did not measure correct velocities during those instances
 450 (e.g. $t = [57, 63]$ s in Figure 5b-c).

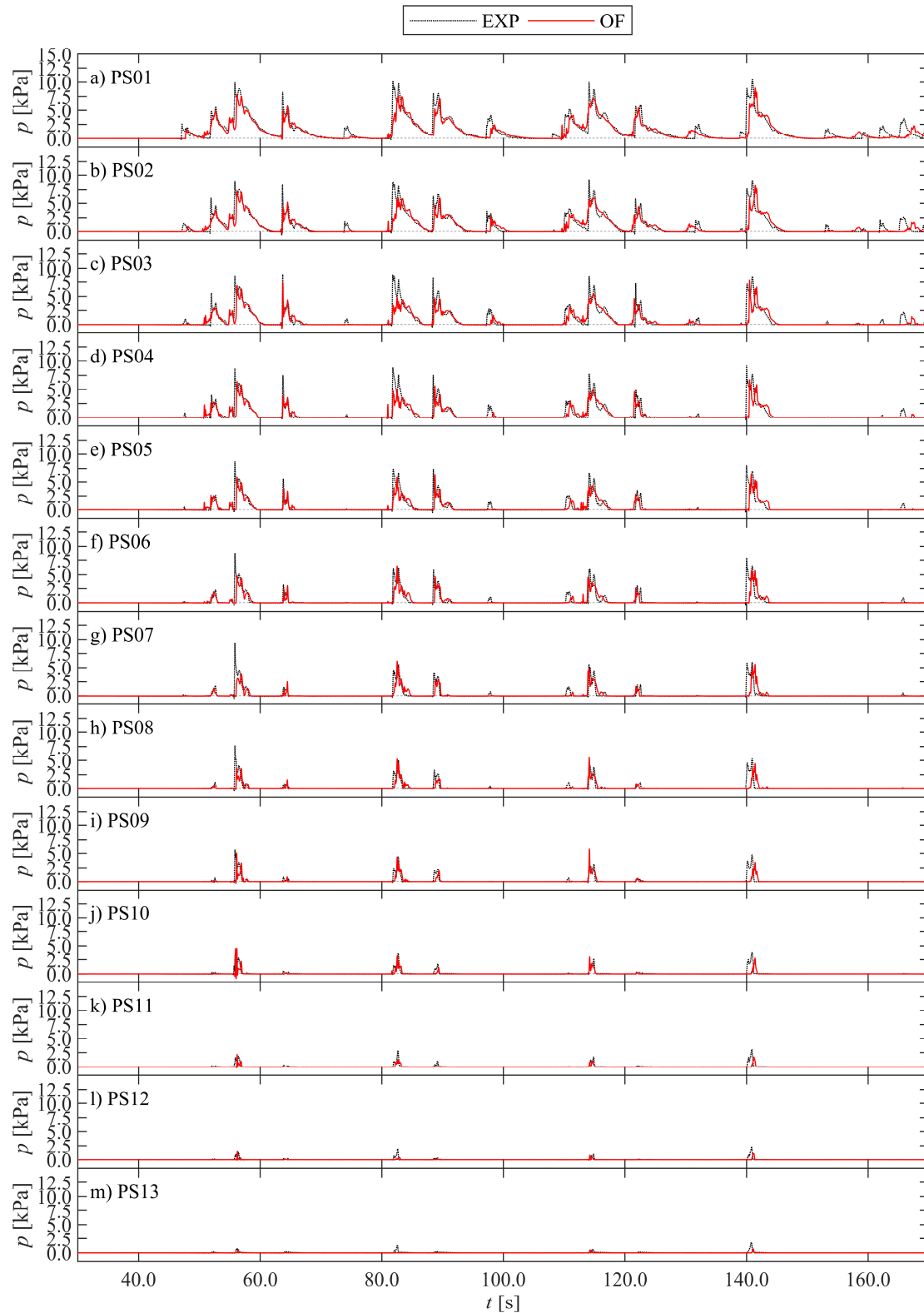


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 452 **Figure 5.** Comparison of U_x time series at all sensor locations. The zero-reference is the promenade
 453 bottom at the sensor locations. For OF both U_x at the measured height above the promenade and the
 454 depth-averaged \overline{U}_x time series are shown.

455 In terms of F_x and p on the vertical wall, OF generally reproduces the timing of the impact events,
 456 including the evolution over time (Figure 6 and Figure 7). However, the EXP time series peak values
 457 appear to be underestimated by OF for both F_x and p , and for a few impacts the first dynamic impact
 458 peak is not entirely captured either (e.g. $t = 82$ s and 140 s). In the experiment, the lowest PSs were
 459 loaded more often than the PSs positioned higher up the vertical wall, because of different bore
 460 impact run-up heights. The lowest PSs also registered the highest values, indicating a mostly
 461 hydrostatic pressure distribution along the vertical wall [74]. Both these observations are reproduced
 462 by OF. Validation of the pressure distribution along the vertical wall is further investigated in section
 463 3.4.



464
 465 **Figure 6.** Comparison of F_x time series at the vertical wall. The experiment is the load cell force
 466 measurement.



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Figure 7. Comparison of p time series at all vertical pressure sensor locations, PS01 being the bottom PS and PS13 the top most PS.

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3.2. Wave characteristics

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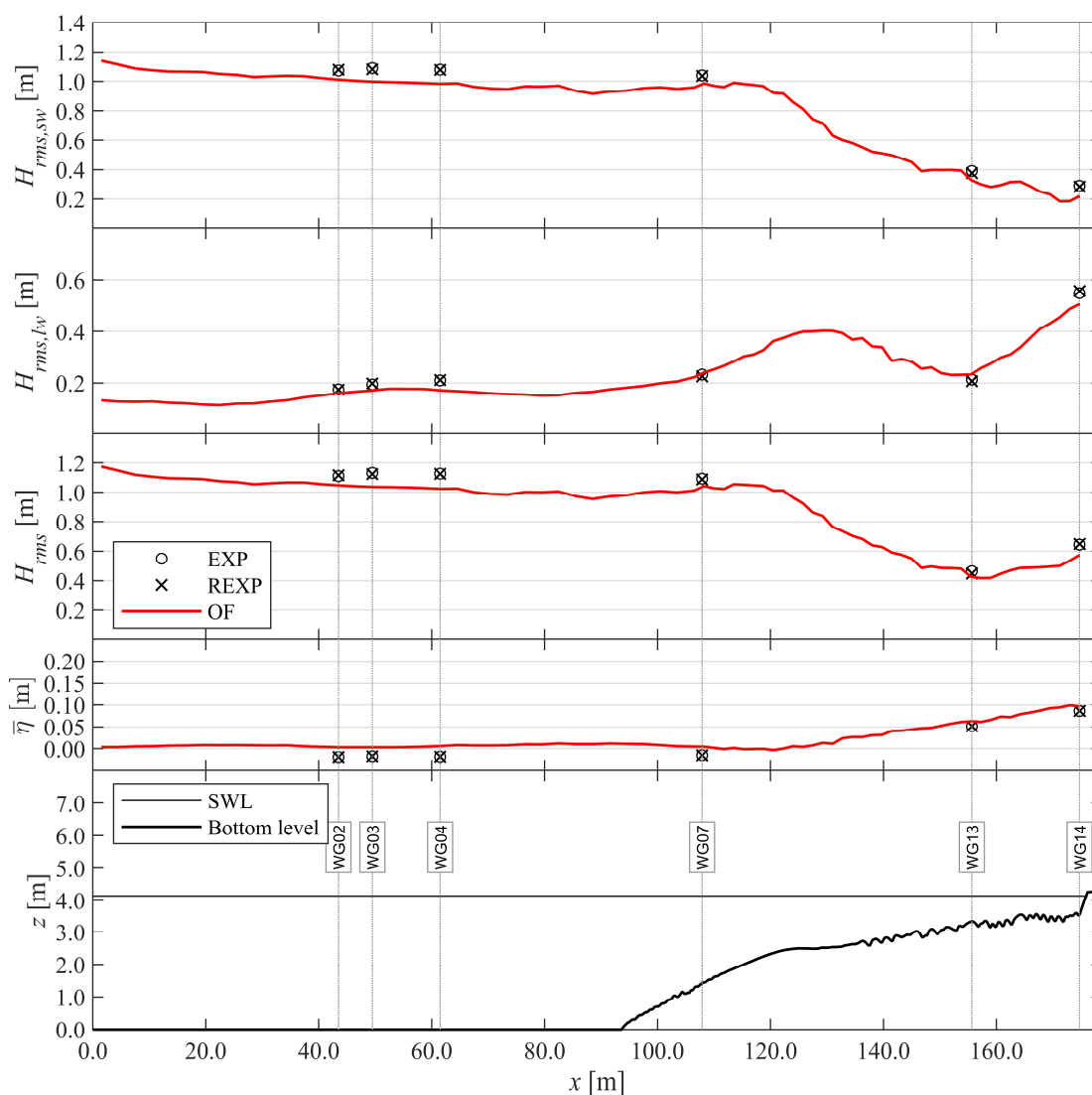
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Based on the η time series the root-mean-square wave height H_{rms} is calculated in the time domain and represents a characteristic wave height and measure of the wave energy. The evolution of H_{rms} , the short- and long-wave components (i.e. $H_{rms,sw}$ and $H_{rms,lw}$), and the mean surface elevation

474 $\bar{\eta}$ or wave setup over the wave flume up to the toe of the dike are displayed in Figure 8. The
 475 experimental repeatability of H_{rms} appears to be near-perfect, since the EXP and REXP data points are
 476 almost indistinguishable. The OF results for these wave characteristics are available along the
 477 complete distance from the wave paddle till the toe of the dike location. The numerical results seem
 478 to follow the experiments very well, although some discrepancies can be seen. The total and SW wave
 479 heights (respectively H_{rms} and $H_{rms,sw}$ in Figure 8) decrease in the OF result from the wave paddle up to
 480 to the toe of the foreshore and underestimate the EXP wave height along this distance. Over the
 481 foreshore, the SWs start to shoal until their steepness becomes too high and, according to OF, start to
 482 break about 11 m from WG07 towards the dike. The location of incipient wave breaking (or decrease
 483 in H_{rms}), x_b , cannot be validated with the experiment, because of insufficient wave gauges in the wave
 484 breaking zone. In any case, the EXP wave height increase due to shoaling (WG07) and decrease due
 485 to breaking (WG13-14) is reproduced well by OF. However, also over the foreshore OF slightly
 486 underestimates the wave amplitude. The experimental LW wave height ($H_{rms,lw}$ in Figure 8) is slightly
 487 underestimated by OF in front of the wave paddle (WG02 – WG04), and at the dike toe (WG14).

488 In terms of the wave setup $\bar{\eta}$, the wave set-down observed in the experiment offshore from the
 489 foreshore toe is not reproduced by OF ($\bar{\eta}_{OF}$ remains close to zero). Further along the flume in the surf
 490 zone, however, $\bar{\eta}$ is better predicted by OF, showing a smaller overestimation.



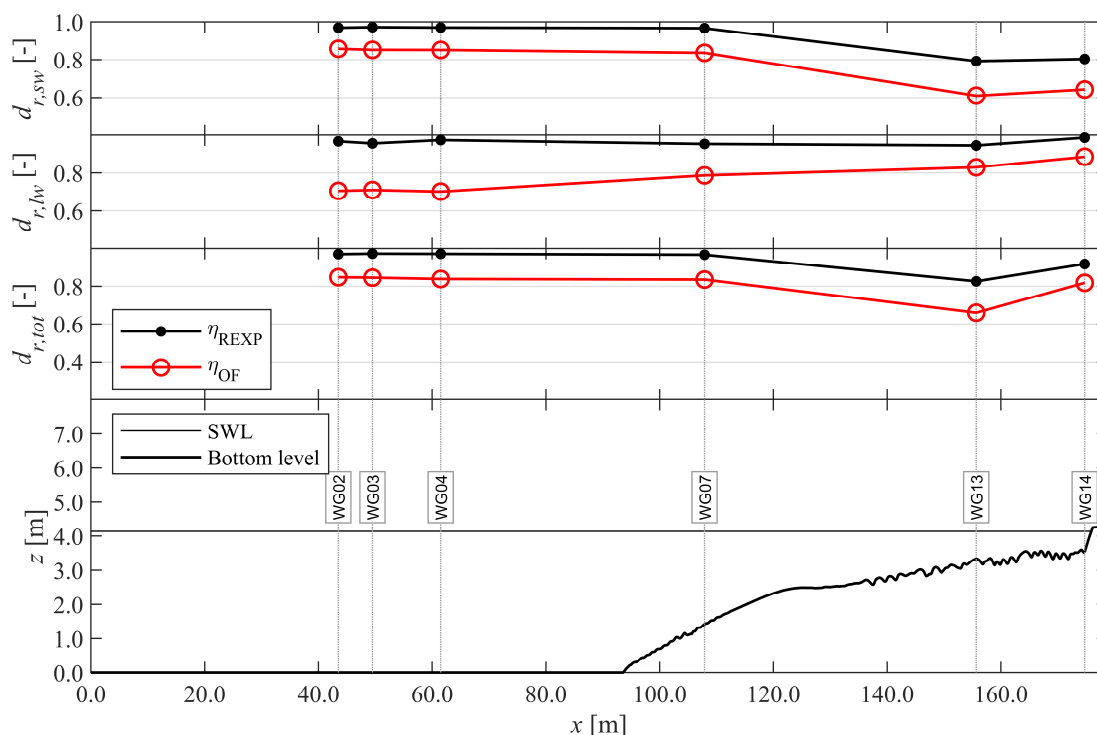
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Figure 8. Comparison of H_{rms} between OF and (R)EXP up to the dike toe. From top to bottom: $H_{rms,sw}$ for the SW components, $H_{rms,lw}$ for the LW components, H_{rms} for the total η , the wave setup $\bar{\eta}$ and finally an overview of the sensor locations, SWL and bottom profile.

495 3.3. Model performance and pattern statistics

496 In this section, the model performance and pattern statics introduced in section 2.3 are applied
 497 to obtain a quantitative numerical model performance evaluation as well. Tables 3 and 4 provide the
 498 pattern and model performance statistics for all sensor locations along the flume up to the vertical
 499 wall. The evolution of d_r at the WG locations along the wave flume up to the toe of the dike is
 500 visualised in Figure 9 for η_{SW} ($d_{r,sw}$), η_{LW} ($d_{r,lw}$) and η ($d_{r,tot}$), and in Figure 11 for η and U_x on the
 501 promenade.

502 The evolution of $d_{r,tot}$ along the flume is very similar for both REXP and OF (Figure 9 and Table
 503 3): it remains constant till the shoaling zone (WG02-WG07), decreases over the surf zone (WG07-13),
 504 and increases back up to the dike toe (WG13-14). This indicates that the decreased experimental
 505 model repeatability of the surface elevation in the surf zone is at least part of the cause of the
 506 decreased numerical model performance. The relative model performance d'_r for η is consequently
 507 fairly constant, corresponding to a model performance rating of *very good*, which remains consistently
 508 so up to the last sensor location in front of the vertical wall. Considering η_{SW} and η_{LW} separately,
 509 reveals that $d_{r,sw}$ mostly follows the same trend as $d_{r,tot}$, and that $d_{r,lw,OF}$ clearly has a different behaviour:
 510 $d_{r,lw,OF}$ is not as high as $d_{r,sw,OF}$ in front of the wave paddle (i.e. $d_{r,lw,OF} = \sim 0.70$ and $d_{r,sw,OF} = \sim 0.85$ at WG02-
 511 WG04), but steadily increases towards the dike toe, while $d_{r,lw,exp}$ remains relatively constant, causing
 512 d'_r to slightly increase as well.



513

514 **Figure 9.** Refined index of agreement d_r of REXP and OF with EXP up to the dike toe. From top to
 515 bottom: $d_{r,sw}$ for η_{SW} , $d_{r,lw}$ for η_{LW} , $d_{r,tot}$ for η , and finally an overview of the sensor locations, SWL and
 516 bottom profile.

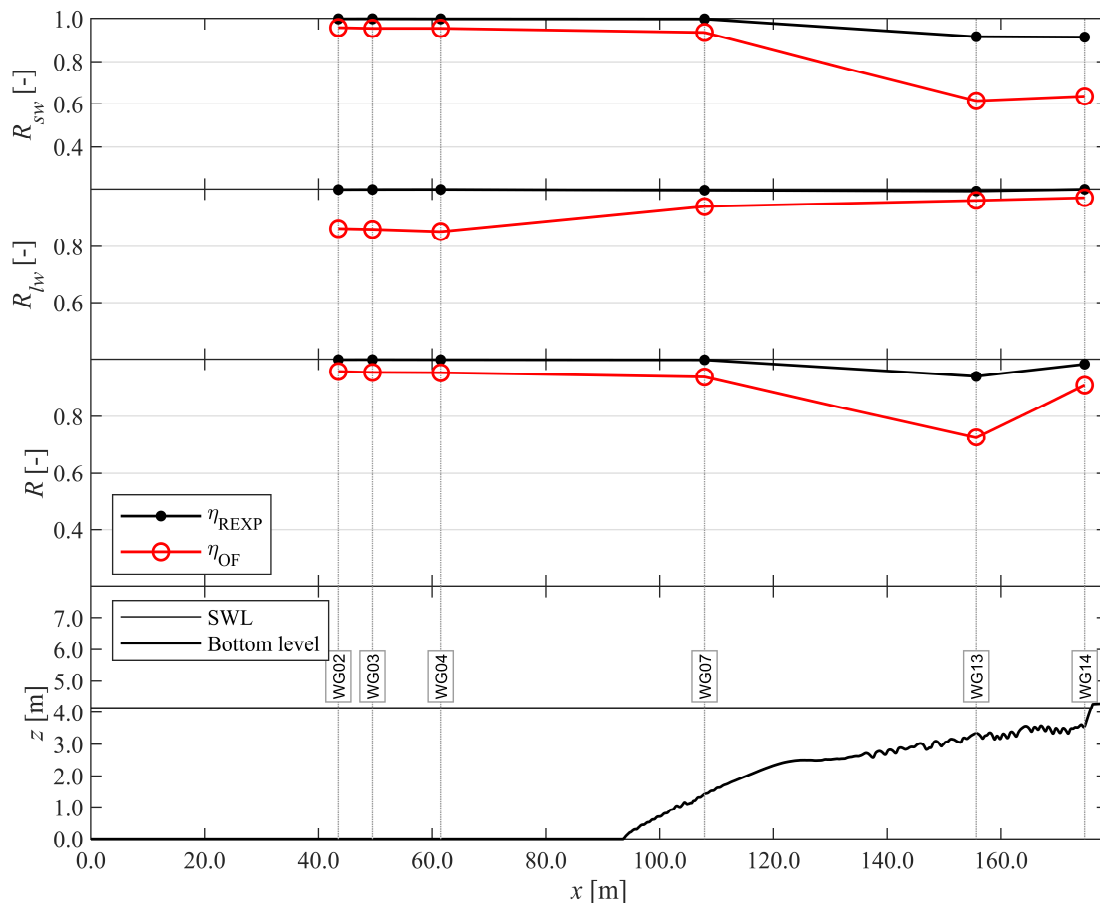
517 The pattern statistics B^* and σ^* represent respectively the accuracy of the wave setup and wave
 518 height from offshore till the dike toe, and confirm the qualitative observations made in section 3.2.
 519 However, spatial information about the accuracy of the numerical wave phase modelling was not
 520 included previously, and is shown separately here in Figure 10. The SW phase accuracy of OF
 521 decreases significantly over the surf zone ($R = \sim 0.90$ to ~ 0.60), while it increases for the LWs ($R = \sim 0.85$
 522 to ~ 0.97). The total wave phase prediction accuracy of OF decreases at WG13 because it is located at
 523 a node of the standing long waves in front of the dike (Figure 8), thus R_{sw} has a higher weight in R

524 there. Conversely, the dike toe (WG14) is located at an antinode, and therefore R_{tw} has higher weight
 525 in R than R_{sw} , leading to an increase of R again at the dike toe.

526 **Table 3.** Pattern and model performance statistics for all surface elevation η sensor locations.

Location	REXP				OF					
	B^* [-]	σ^* [-]	R [-]	d_r [-]	B^* [-]	σ^* [-]	R [-]	d_r [-]	d'_r [-]	Rating [-]
WG02	-0.01	1.01	1.00	0.97	0.06	0.94	0.96	0.85	0.88	Very Good
WG03	-0.01	0.99	1.00	0.97	0.05	0.92	0.95	0.85	0.87	Very Good
WG04	-0.01	1.00	1.00	0.97	0.06	0.91	0.95	0.84	0.87	Very Good
WG07	0.01	1.00	1.00	0.97	0.06	0.94	0.94	0.84	0.87	Very Good
WG13	0.00	0.97	0.94	0.83	0.04	0.95	0.73	0.66	0.83	Very Good
WG14	0.00	1.00	0.98	0.92	0.05	0.89	0.91	0.82	0.90	Very Good
WLDM01	-0.02	0.99	0.99	0.92	-0.08	1.00	0.89	0.80	0.88	Very Good
WLDM02	-0.02	1.01	0.99	0.92	-0.05	1.01	0.91	0.82	0.89	Very Good
WLDM03	0.00	0.98	0.99	0.92	-0.03	0.98	0.90	0.82	0.90	Very Good
WLDM04	0.01	0.97	0.98	0.92	-0.00	1.00	0.87	0.79	0.87	Very Good

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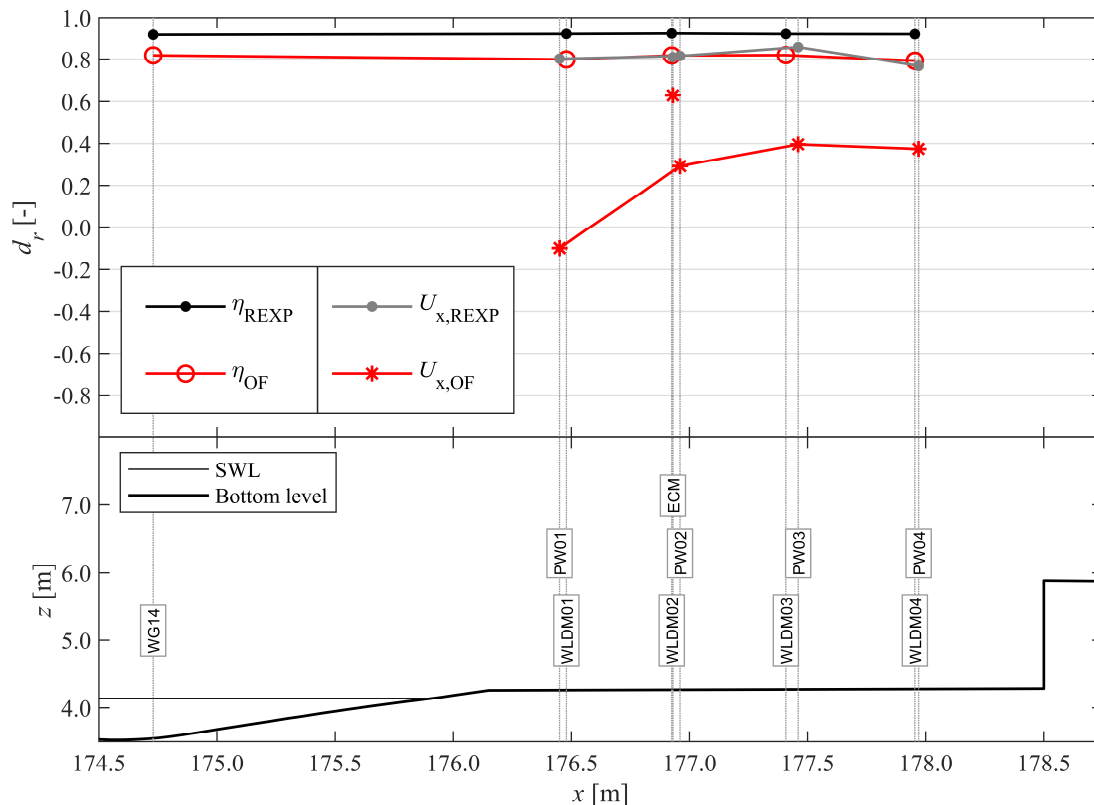
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529 **Figure 10.** Comparison of R for η of REXP and OF with EXP up to the dike toe. From top to bottom:
 530 R_{sw} for η_{sw} , R_{tw} for η_{tw} , R for η , and finally an overview of the sensor locations, SWL and bottom
 531 profile.

532 Along the promenade, the d_r for η and U_x is shown in Figure 11 and – on first sight – seems to
 533 indicate that the OF model performance for U_x is much worse than for η , primarily for comparisons
 534 to the PW measurements, but also for the ECM measurement. Taking into account the experimental
 535 uncertainty, however, the model performance rating for U_x of ECM is actually *very good* ($d'_{r,ECM}$ in

536 Table 4), which is the same as the OF model performance rating for η on the promenade ($d'_{r,WLDM01-04}$
 537 in Table 3). For the PW measurements, the OF rating for U_x is still worse (*reasonable/fair to bad*), but
 538 was explained before by the fact that the PW's had faulty positive U_x measurements during return
 539 flow (section 3.1).

540 Although the wave setup at the dike toe is overestimated by OF ($B^*_{WG14} > 0$), η on the promenade
 541 is on average underestimated ($B^*_{WLDM01-04} < 0$) and U_x as well ($B^* < 0$). Conversely, the bore wave height
 542 is well-represented on the promenade ($\sigma^*_{WLDM01-04} \approx 1.00$), while the wave height is underestimated
 543 at the dike toe ($\sigma^*_{WG14} = 0.89$). The surface elevation phase difference between OF and EXP observed
 544 at the dike toe ($R_{WG14} = 0.91$) is carried over on the promenade ($R_{WLDM01-04} \approx 0.90$), but higher phase
 545 differences are detected for U_x ($R_{ECM} = 0.73$).



546

547 **Figure 11.** Refined index of agreement d_r of REXP and OF with EXP from the dike toe up to the vertical
 548 wall. From top to bottom: d_r for η and U_x , and finally an overview of the sensor locations, SWL and
 549 bottom profile.

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Table 4. Pattern and model performance statistics for U_x on the promenade.

Location	REXP				OF					Rating
	B^* [-]	σ^* [-]	R [-]	d_r [-]	B^* [-]	σ^* [-]	R [-]	d_r [-]	d'_r [-]	
PW01	0.02	0.96	0.91	0.80	-1.24	1.55	0.58	-0.10	0.10	Bad
ECM	-0.02	1.05	0.87	0.81	-0.25	0.94	0.73	0.63	0.82	Very Good
PW02	-0.05	0.99	0.88	0.82	-0.66	1.22	0.65	0.29	0.48	Poor
PW03	-0.02	1.00	0.92	0.86	-0.57	1.06	0.68	0.40	0.54	Reasonable/Fair
PW04	-0.03	1.02	0.88	0.77	-0.42	0.88	0.58	0.37	0.61	Reasonable/Fair

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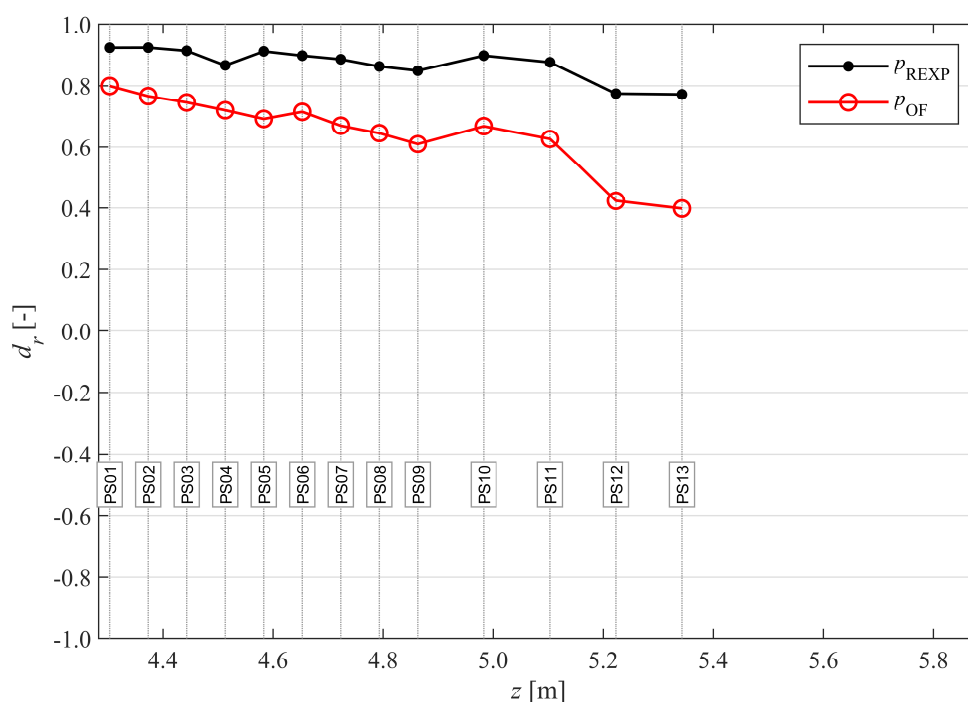
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Finally, the model performance in terms of p and F_x are evaluated at the vertical wall (Figure 12 and Table 5). Both REXP and OF show the highest model performance at the lowest pressure sensor location and a more or less linear decreasing model performance at PS locations higher along the

555 vertical wall. The relative difference between the d_r of REXP and OF increases as well higher along
 556 the vertical wall, leading to a numerical model performance rating from *very good* for PS01-PS06, to
 557 *good* for PS05-PS11 and finally to *reasonable/fair* at the highest PS locations (PS12-PS13) (Table 5).
 558 Considering that the bottom PSs registered the highest p values and are therefore the most
 559 determinative in the calculation of F_x , it follows that the numerical model performance for F_x is rated
 560 *very good* as well. The pattern statistics in Table 5 reveal the remaining numerical errors to be that p
 561 and F_x are generally underestimated by OF (i.e. $B^* < 0.00$ and $\sigma^* < 1.00$) and that the impact events still
 562 slightly mismatch in time between OF and EXP ($R < 1.00$).



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Figure 12. Refined index of agreement d_r of REXP and OF with EXP for p at the vertical wall (horizontal axis).

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Table 5. Pattern and model performance statistics for all surface elevation sensor locations.

Location	REXP				OF					Rating
	B^* [-]	σ^* [-]	R [-]	d_r [-]	B^* [-]	σ^* [-]	R [-]	d_r [-]	d'_r [-]	
p (PS01)	0.00	1.00	0.98	0.92	-0.14	0.84	0.84	0.80	0.88	Very Good
p (PS02)	-0.01	0.99	0.97	0.92	-0.10	0.82	0.77	0.77	0.84	Very Good
p (PS03)	0.00	1.00	0.96	0.91	-0.13	0.75	0.71	0.75	0.83	Very Good
p (PS04)	0.02	0.99	0.94	0.87	-0.13	0.74	0.66	0.72	0.85	Very Good
p (PS05)	0.01	1.00	0.96	0.91	-0.11	0.75	0.61	0.69	0.78	Good
p (PS06)	-0.01	0.97	0.96	0.90	-0.13	0.78	0.61	0.72	0.82	Very Good
p (PS07)	-0.01	0.93	0.95	0.89	-0.17	0.76	0.53	0.67	0.78	Good
p (PS08)	-0.05	0.86	0.94	0.86	-0.20	0.74	0.46	0.65	0.78	Good
p (PS09)	-0.07	0.88	0.93	0.85	-0.25	0.78	0.39	0.61	0.76	Good
p (PS10)	-0.04	0.93	0.94	0.90	-0.24	0.77	0.48	0.67	0.77	Good
p (PS11)	-0.04	0.91	0.94	0.88	-0.33	0.57	0.37	0.63	0.75	Good
p (PS12)	-0.20	0.79	0.89	0.78	-0.55	0.53	-0.05	0.42	0.65	Reasonable/Fair
p (PS13)	-0.15	0.57	0.92	0.77	-0.59	0.33	0.12	0.40	0.63	Reasonable/Fair
F_x (LC)	0.00	0.97	0.90	0.90	-0.12	0.74	0.73	0.76	0.85	Very Good

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568 3.4. Bore interactions and impact

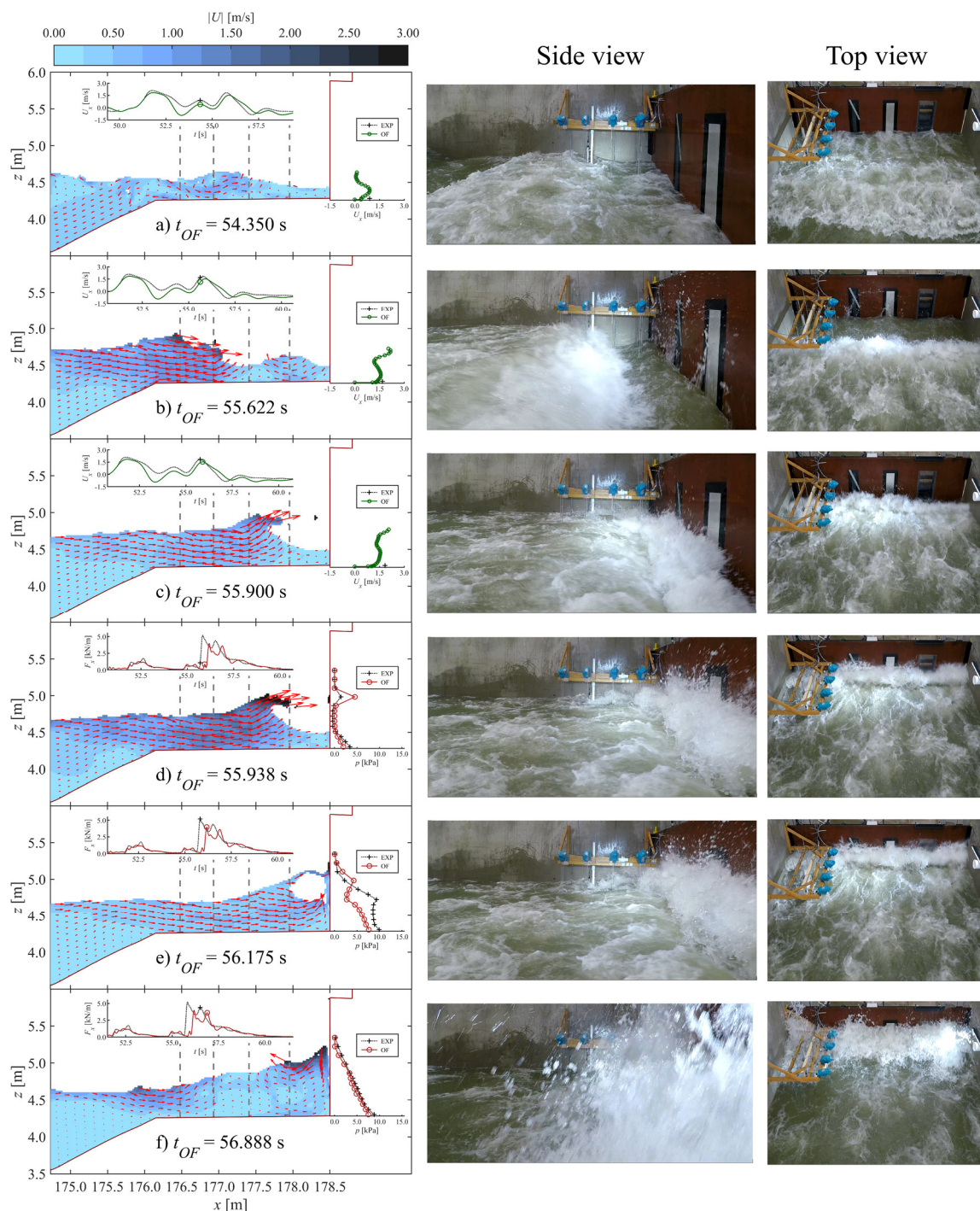
569 To explain some of the numerical successes and failures encountered in the reproduction of the
 570 experimental bore impacts on the vertical wall, a detailed analysis is done of a selection of individual
 571 impact events and the bore interactions leading up to them. The analysis is based on an investigation
 572 of snapshots at important time instants during the first two largest impact events in the modelled
 573 time series (Figure 7). The first ($t \sim 56$ s) and second ($t \sim 82$ s) main impact events are chosen because
 574 they are good examples of respectively a successful and less successful numerical reproduction of the
 575 experimental impacts.

576 Numerical snapshots of the flow on the dike, including the velocity distribution along the
 577 vertical cross-section at the ECM location or the pressure distribution along the vertical wall are
 578 compared in Figure 13 and Figure 14 to the equivalent experimental data and snapshots based on
 579 side and top view video images. Key time instants of overtopped bore behaviour are selected during
 580 these two main impacts and are listed chronologically in Table 6. Some of the key time instants occur
 581 at slightly different times in each model (due to slight wave phase differences). In those cases, the key
 582 time instants were selected from each model result based on identifiable features in the bore
 583 interaction images, the U_x time series or the F_x time series (e.g. peaks, troughs,...), making sure a
 584 relevant comparison is made of the bore interaction and the velocity or pressure profile.

585 **Table 6.** Description of the snapshots shown in Figures 13 and 14.

<i>Time instant number</i>	<i>Description</i>	<i>Figure</i>
<i>Main impact 1</i>		
1a	Pre-impact of small overtopped wave.	Figure 13a
1b	Pre-collision of large overtopped bore and small wave reflected from vertical wall.	Figure 13b
1c	Collision of large overtopped bore and reflected small wave.	Figure 13c
1d	Impact on vertical wall of high velocity spray from overturned bore.	Figure 13d
1e	Dynamic impact of overturned bore on vertical wall.	Figure 13e
1f	Quasi-static impact of overturned bore on vertical wall.	Figure 13f
<i>Main impact 2</i>		
2a	Very small overtopped bore.	Figure 14a
2b	Impact of small overtopped bore on vertical wall.	Figure 14b
2c	Impact of large overtopped bore on vertical wall.	Figure 14c
2d	Impact of large overtopped bore on vertical wall, continued.	Figure 14d
2e	Impact of large overtopped bore on vertical wall, continued.	Figure 14e
2f	Return flow of large bore reflected from vertical wall.	Figure 14f

586 The first series of impacts mainly occurred while the LWs overtopped and reflected on the dike-
 587 wall structure for the first time. A good indication of this time period is when η at the dike toe (Figure
 588 4f) was larger than the freeboard (i.e. $t = [47, 70]$ s). During the LW overtopping/reflection several
 589 SWs propagated on top of the LW crest, overtopped the dike and impacted the vertical wall along
 590 with the LWs: after a very small first overtopped bore ($t \sim 48$ s in Figure 6), a second larger bore
 591 impacted and reflected on the vertical wall ($t \sim 52.5$ s). While the reflected second bore returned
 592 seawards, a third small wave overtopped and headed towards the vertical wall (Figure 13a, termed
 593 *sequential overtopping bore pattern* by Streicher et al. [74]). This small wave then reflected against the
 594 vertical wall, while a very large turbulent bore was overtopping the dike crest (Figure 13b). At that
 595 moment the small wave and large bore were propagating in opposite directions on the promenade.
 596 Eventually they collided, and the larger incident turbulent bore was forced to overturn (Figure 13c).
 597 This collision also caused spray to be ejected at a high velocity from the overturning wave tongue
 598 (see $[x, z] = [178.3 \text{ m}, 4.9 \text{ m}]$ in Figure 13c). This airborne water volume hit the vertical wall first and
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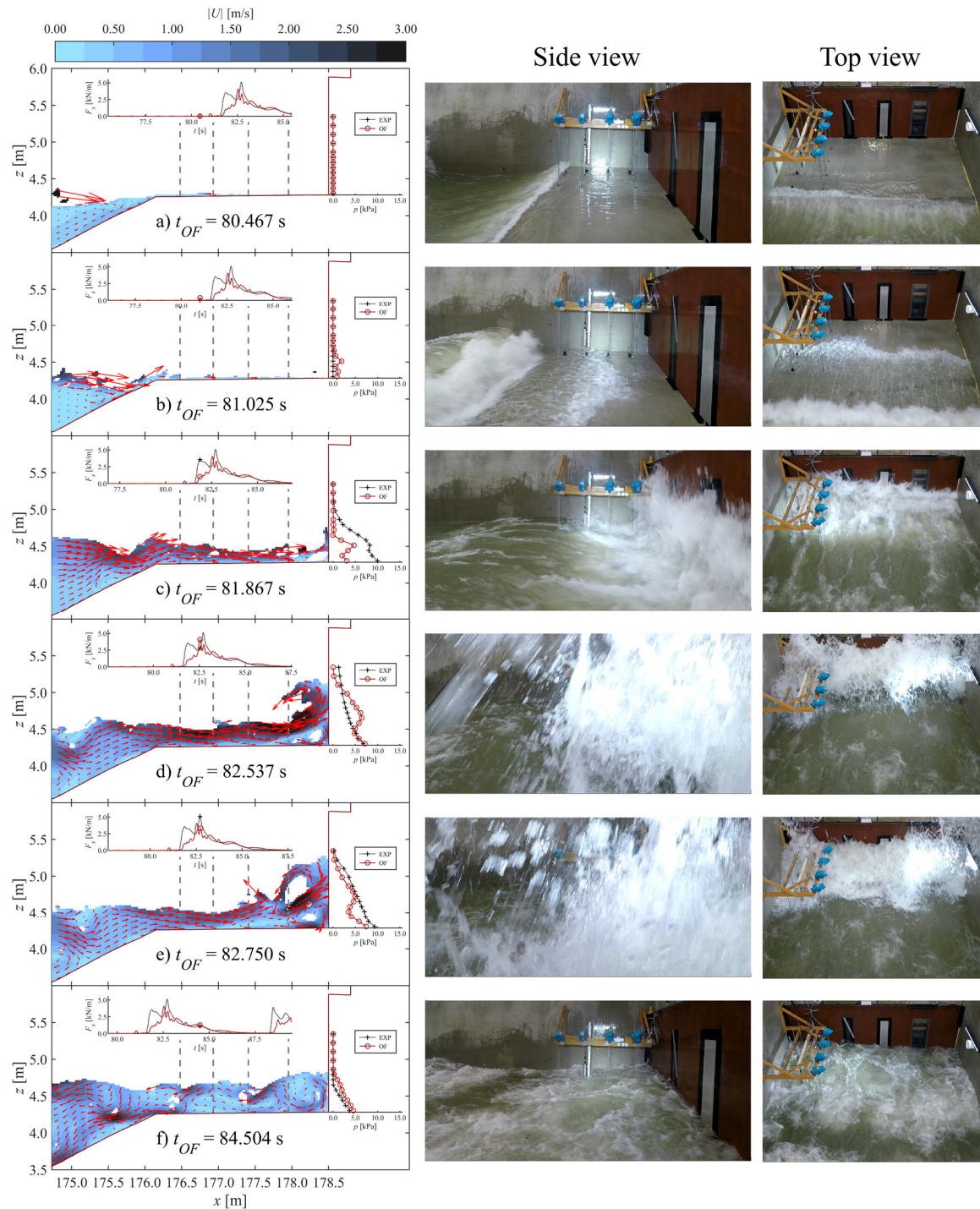
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Figure 13. Snapshots of selected key time instants chronologically over the first main impact (a-f). The OF snapshot (left) is compared to the equivalent EXP snapshot from the side view (centre) and top view (right) cameras. In the OF snapshots, the colours of the water flow indicate the velocity magnitude $|U|$ according to the colour scale shown at the top. The red arrows are the velocity vectors, which are scaled for a clear visualisation. Each OF snapshot has two inset graphs: at the top is a time series plot of U_x (for EXP and \bar{U}_x for OF) (a-c) or F_x (d-f), in which a circle marker (o) and a plus marker (+) indicate the time instant of the numerical and experimental snapshot respectively. Along the vertical wall U_x (a-c) or p (d-f) is plotted at respectively the ECM sensor location or each PS location (the vertical axis is z [m]). Along the promenade four vertical grey dashed lines indicate the sensor locations on the promenade, of which the WLDM gauges are also visible in the experimental snapshots (topped by blue plastic bags). The location of the ECM is at the second vertical grey dashed line from the left. The time instant of the numerical snapshot is provided by t_{OF} .



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Figure 14. Snapshots of selected key time instants chronologically over the second main impact (a-f). See caption of Figure 13 for further descriptions.

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separately from the main overturning wave tongue (see $[x, z] = [178.5 \text{ m}, 4.95 \text{ m}]$ in Figure 13d), causing a local pressure peak at the location of PS10 (see the p -profile in Figure 13d). Subsequently, the main overturning wave hit the wall, causing a dynamic force peak $F_{x,1}$ (Figure 13e), and ran vertically up the wall temporarily reducing F_x during maximum run-up (not shown). The following run-down and reflection from the wall corresponds to a second force peak $F_{x,2}$, this time of quasi-static nature (Figure 13f). This type of bore interaction was called a *plunging breaking bore pattern* by Streicher et al. [74], which – in this case – caused a *quasi-static impact* ($F_{x,1}/F_{x,2} < 1.20$, according to Streicher et al. [74]). This is valid for both the experiment and the numerical model result, indicating

624 that OF was able to reproduce these processes leading to a very similar shape of the pressure
 625 distribution along the vertical wall (see pressure profiles in Figure 13d-f) and time evolution of F_x (see
 626 time series graph insets in Figure 13d-f). Comparing $U_{x,ECM}$ from EXP with the velocity profile from
 627 OF at the ECM location (see velocity profiles in Figure 13a-c) reveals that OF locally, but consistently
 628 underestimated U_x at the vertical measurement position of the ECM, which was also observed in
 629 Figure 5b.

630 The second series of impacts occurred during the second LW overtopping and reflection event
 631 (Figure 4f-j: $t = [74, 100]$ s). Again, SWs propagated on top of the LW crest, bringing bore interactions
 632 to the promenade. This time, however, the bore interaction pattern modelled by OF that caused the
 633 main impact was different than the pattern observed in EXP. First a very small bore overtopped the
 634 dike crest and was immediately followed by a much larger bore. In EXP, the smaller bore was
 635 overtaken by the larger bore (Figure 14b-c, termed *catch-up bore pattern* by Streicher et al. [74]), leading
 636 to a *quasi-static impact*. In the result from OF, however, the very small wave overtopped sooner (Figure
 637 14a), so that it had time to reflect against the wall (Figure 14b) before colliding with the incoming
 638 larger bore (not shown). OF therefore modelled a *collision bore pattern* instead of a *catch-up bore pattern*,
 639 greatly reducing the first impact force peak of the main impact (by ~65% compared to EXP, Figure
 640 14c). This also clearly affected the pressure profiles along the vertical wall: during the first F_x peak, p
 641 is severely underestimated, but the distribution is still similar, with a local peak at PS04. The p -profiles
 642 differentiate more at the F_x peak of the OF result (Figure 14d) and at the quasi-static F_x peak in the
 643 EXP result (Figure 14e). In the experiment a quasi-hydrostatic pressure profile was measured, at both
 644 those time instants. In the OF result, however, a pressure peak is found at PS06, caused by a vortex
 645 formed at the foot of the vertical wall upon which a strong flow impinged on the wall at that location.
 646 After reflection of the bore, both models correspond again, showing a hydrostatic pressure profile
 647 along the wall (Figure 14f).

648 4. Discussion

649 4.1. Wave transformation processes till the dike toe

650 In sections 3.1 and 3.2 it was already established that OF is capable of reproducing the wave
 651 shoaling and breaking processes in terms of evolutions in η and H_{rms} . This section discusses the
 652 processes related to the LW transformations over the foreshore as modelled by OF and their
 653 correspondence to observations in EXP.

654 The modulation factor δ of the SWs is high for the considered bichromatic wave conditions
 655 (Table 1), indicating that the incident bound LW amplitude was relatively high as well. Furthermore,
 656 the normalised bed slope parameter β_b can be calculated [36]:

$$\beta_b = \frac{h_x}{\omega} \sqrt{\frac{g}{h_b}}, \quad (10)$$

657 where h_x is the foreshore slope (= 1:35), ω is the radial frequency of the bound LW ($= 2\pi(f_1 - f_2)$), g the
 658 gravitational acceleration and h_b a characteristic breaking depth (= 2.12 m at $x_b = 115$ m). A value of
 659 0.28 is obtained, which means that the bound LW shoaling had a mild slope regime ($\beta_b < 0.3$), so that
 660 the growth rate of the incoming LWs was much higher than given by Green's Law (conservative
 661 shoaling), indicating significant energy transfer from the primary SWs to the bound LW [75].
 662 Additionally, in a mild-slope regime, LW shoreline dissipation and shoreline reflection are high and
 663 low respectively [36]. However, the beach considered here is not a beach by itself, but acts as a
 664 foreshore to a steep-sloped dike. Consequently, no such expected decrease in LW energy towards the
 665 shoreline is observed (i.e. $H_{rms,lw}$ in Figure 8). Indeed, the dike was positioned in the shoaling zone of
 666 the long waves, thereby preventing the LWs to break. Instead, LWs reflected against the dike,
 667 indicated by the oscillations of $H_{rms,lw}$ towards the dike in the OF result, which implies the presence
 668 of a (partial) standing wave system. Wave gauges WG13 in the inner surf zone and WG14 at the dike
 669 toe were positioned at a node and anti-node of this standing wave system. This is also clearly visible
 670 in the η time series plot, where η_{LW} is much closer to zero at WG13 (Figure 4e) than at WG14 (Figure
 671 4f). In the surf zone the LW previously bound to the wave group became a free wave, traveling at its

own wave celerity. Due to first order wave generation at the boundary, other spurious free LWs were generated as well at the wavemaker and propagated as free waves towards the dike [76]. During a standing LW crest at the dike toe, the LWs themselves overtopped the dike (i.e. when $\eta >$ freeboard $R_c = 0.117$ m, Figure 4f) thereby temporarily aiding several breaking SWs to overtop the crest of the dike (the wave length of the free LWs was more than five times longer than the primary SW components in the inner surf zone). These results have illustrated OF's ability to reproduce the wave energy transfer to the subharmonics and LW transformations over the foreshore till the dike toe. All these observations also confirm that the contribution of LWs to the processes on the dike, including the wave impact loading on the vertical wall, is very important in the case that is considered here.

4.2. Importance of differences in wave generation methods

Although the overall OF model performance was rated to be *very good*, a few differences between the OF and EXP results remain to be explained. One of the largest OF inaccuracies was an underestimation of the wave height, primarily observed at the offshore WG locations (WG02-WG04, see Figure 8 and Table 3), suggesting an underestimation of the incident wave energy and/or numerical diffusion. The underestimation was likely caused by differences between the numerical wave generation method with static boundary in OF and the physically moving wave paddle in the EXP [68]. The wave boundary condition by olaFlow allows for a tuning factor to be applied to U_x and η at the boundary, to overcome a possible underestimation of the incident wave height. Such a calibration of the OF model (with a tuning factor of 1.13) was found to solve the underestimation of the wave height (not shown), but introduced or exacerbated other errors, finally leading to lower values of d_r and decreased model performance ratings for $U_{x,ECM}$ and F_x .

Another remaining discrepancy between OF and EXP is found in $\bar{\eta}$, which was primarily overestimated by OF in the offshore region (Figure 8). Also, after calibration of the incident wave height to EXP, H_{rms} – and consequently $\bar{\eta}$ – increased in the surf zone, exacerbating the $\bar{\eta}$ overestimation there (not shown). The root cause of this difference is likewise related to the different wave generation methods applied in EXP and OF. In the experimental wave flume, the finite body of water and conservation of mass caused water mass to be redistributed from offshore to the surf zone during build-up of the wave setup, thereby causing a lowering of the mean water level in the offshore region. This process developed differently in OF because of the static boundary condition including AWA. The AWA assures a constant mean water level at the boundary [8,52], meaning that a net water mass is added to the computational domain until a quasi-steady state is achieved when wave setup is fully developed [54]. In this case, OF's method is closer to the field condition, where generally a large enough body of water is available to supply water mass for the wave setup to develop without noticeably lowering the offshore mean water level. Nevertheless, in the context of the validation, this difference in $\bar{\eta}$ is at the cause of many of the remaining inaccuracies in the OF result compared to EXP, because the waves propagated in slightly different mean water depths, which affected the non-linear wave-wave interactions and wave phases in the surf zone. Consequently, it is believed to be the root cause of the strong decrease of R_{sw} observed in the surf zone (i.e. locations WG13-14 in Figure 10).

These two remaining inaccuracies in the OF results compared to EXP (i.e. underestimation of H_{rms} and overestimation of $\bar{\eta}$), are both attributable to the differences in wave generation methods applied. Although still an overall *very good* model performance rating was achieved by OF, it is expected that even better results can be obtained by applying a closed dynamic wave boundary condition in OF, which mimics the EXP wave paddle movement. However, application of the dynamic boundary condition of olaFlow proved to be highly unstable for the present case, and no result was achieved to confirm this hypothesis.

4.3. OF model performance for impacts on a dike-mounted vertical wall

The accuracy of a numerical wave model to reproduce wave overtopping over a dike with a very shallow foreshore depends on the quality of the incident waves at the dike toe location [10]. The same should therefore hold true for impacts on a dike-mounted vertical wall by such overtopped waves.

722 The overall *very good* model performance of OF in terms of p and F_x at the vertical wall can be
723 explained by a generally correct reproduction of bore interactions over the promenade of the dike.
724 Conversely, discrepancies – even small ones – in bore interactions between OF and EXP can lead to
725 significant differences in the impact type on the vertical wall, and consequently in p and F_x (section
726 3.4). In addition, the much lower values of B^*_{OF} and R_{OF} compared to B^*_{REXP} and R_{REXP} for $U_{x,ECM}$ (i.e.
727 $B^*_{REXP} = -0.02$ and $R_{REXP} = 0.87$, $B^*_{OF} = -0.25$ and $R_{OF} = 0.73$ in Table 4) indicate an important contribution
728 of the underestimation of U_x and of phase differences in U_x between OF and EXP to the remaining
729 errors in the impact prediction by OF. The bore interactions on their part depend on the wave
730 conditions at the dike toe location. This is illustrated by the calibrated OF model results, which was
731 found to improve the wave height reproduction at the dike toe compared to the OF model (section
732 4.2), while errors increased for the wave setup and wave phases at the dike toe location, leading to a
733 lower model performance for the processes on the dike (not shown).

734 Even when the incident wave conditions at the dike toe would be perfectly reproduced, other
735 model limitations would still contribute to residual errors in the numerical results for the wave
736 impacts on the vertical wall:

- 737 • 3D effects in EXP (i.e. irregular and oblique wave fronts, wave breaking-induced 3D vortex
738 formation), which are unreproducible by a 2DV RANS model;
- 739 • Water-air mixing in bores and air pressure fluctuations in entrained air pockets by overturning
740 wave impacts on the wall, which are both processes not resolved by a multiphase numerical
741 model of two incompressible and immiscible fluids.
- 742 • Douglas and Nistor [77] have shown that – compared to a dry-bed condition – a bore
743 propagating on over a thin layer of water on the bed (i.e. wet-bed condition) can substantially
744 increase the steepness and depth of the bore-front and consequently affect the impact of the bore
745 on the wall. The near-bed resolution of the OF grid along the promenade might not have been
746 able to reproduce correctly wet-bed bore propagation in cases of a very thin layer of water,
747 possibly even modelling a dry-bed bore propagation instead.
- 748 • Differences between OF and EXP in the treatment of friction on the bed of the promenade. The
749 no-slip boundary condition and applied wall function in OF modelled a boundary layer, which
750 lowered U_x close to the bed more than was measured in EXP. On average, U_x has been
751 underestimated by OF at the measurement locations of the PWs and ECM close to the
752 promenade bed (Figure 5, B^* in Table 4 and Figure 13a-c).

753 Errors in the reproduction of the impact type and the first two model limitations listed above are also
754 apparent in the numerical reproduction of the pressure distribution along the vertical wall: higher
755 up the wall a decreasing OF model performance rating of p was observed (Figure 12, Table 5). The
756 highest PS locations are the most sensitive to errors in the impact and run-up patterns along the
757 vertical wall and to overly simplified water-air mixture modelling.

758 5. Conclusions

759 A RANS multiphase solver for two incompressible and immiscible fluids (water and air),
760 interFoam of OpenFOAM® with olaFlow wave boundary conditions (OF), was applied in 2DV for
761 bichromatic wave transformations over a cross-section of a hybrid beach-dike coastal defence system,
762 consisting of a steep-sloped dike with a mildly-sloped and very shallow foreshore, and finally wave
763 impact on a vertical wall. OF was not validated before in this context, where – prior to impact – waves
764 undergo many nonlinear transformations and interact with a dike slope and promenade. A large-
765 scale experiment of bichromatic waves and its repetition were selected for this validation. The
766 repeated test allowed to assess the accuracy of the measurements, uncertainty due to model effects
767 and variability due to stochastic processes in the experiment.

768 The validation consisted of both qualitative and quantitative comparisons. Pattern and model
769 performance statistics were employed for the quantitative validation. Based on Willmott's refined
770 index of agreement d_r , calculated for OF and the repeated test REXP with reference to the first test
771 EXP, a relative refined index of agreement d'_r was proposed, which takes the experimental

772 uncertainty, derived from REXP, into account in the numerical model performance evaluation. Based
773 on value ranges of d_r' , a classification into model performance ratings was proposed as well.

774 After a convergence analysis of the most important numerical parameters (i.e. grid resolution
775 and CFL number), and without calibration of the numerical model, a model performance rating of
776 *very good* was achieved by OF compared to the experiment for all relevant design parameters (i.e. η ,
777 U_x , p and F_x), which demonstrates OF's applicability for the design of such hybrid coastal defence
778 systems. Remaining discrepancies were found to be mainly caused by the different wave generation
779 methods applied in OF (static boundary) and EXP (moving wave paddle), which caused an
780 underestimation of the incident wave energy and an overestimation of the wave setup in OF
781 compared to EXP. Consequently, when applying OF for a design of a hybrid coastal defence system,
782 the incident wave energy is recommended to be calibrated, while the wave setup development for a
783 static boundary condition with active wave absorption in OF is actually closer to the field condition
784 compared to EXP (finite water mass).

785 A detailed comparison of snapshots at key time instants of bore interactions leading up to two
786 selected bore impacts on the vertical wall, revealed that slight errors in wave phases can lead to very
787 different bore interaction patterns on the promenade and finally to different bore impact types on the
788 wall.

789 Future work includes a detailed inter-model comparison between the OF model presented here,
790 a weakly compressible SPH model (DualSPHysics), and a non-hydrostatic wave model (SWASH) for
791 the same case [64].

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794 and L.C.; writing—original draft preparation, V.G.; writing—review and editing, V.G., C.A., T.S., M.S., L.C., A.K.
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805 publish the results.

806 Appendix A. Numerical convergence analysis

807 The OF result is influenced by many of its settings, of which the spatial discretisation of the
808 model domain and time stepping are the most important [44]. Their convergence analysis is
809 presented here. The numerical model convergence analysis is based on η at the experimental wave
810 gauge locations over the wave flume up to the dike toe, since it is the most important driver of model
811 performance of the subsequent processes on the dike. The wave force at the vertical wall is not
812 suitable as reference for the grid convergence analysis, because relatively small differences in wave
813 phase can cause very different types of bore interactions on the promenade and therefore very
814 different resulting bore impacts (section 3.4).

815 A.1. Model convergence statistics

816 For the convergence analysis, four customised statistical error indicators are considered, among
817 which the first three are defined to reflect several aspects of the η time series considered (i.e. wave
818 setup, wave height and wave phase):

- 819 • Freeboard normalised bias, NB :

$$NB = \frac{B}{R_c}, \quad (A1)$$

820 in which R_c is the freeboard, and B is the bias defined by (7). The bias or difference in the wave
 821 setup is normalised with the freeboard which is one of the governing parameters for waves
 822 overtopping a dike [78].

823 • Residual error of the normalised standard deviation, $RNSD$:

$$RNSD = 1 - \sigma^*, \quad (A2)$$

824 in which σ^* is given by (6) and in which the observed time series is the reference time series and
 825 the predicted time series is the considered time series. A positive $RNSD$ signifies a higher wave
 826 height and a negative $RNSD$ signifies a lower wave height compared to the reference.

827 • Residual error of the correlation coefficient, RCC :

$$RCC = 1 - R, \quad (A3)$$

828 in which R is the correlation coefficient, given by (9), between the reference time series and time
 829 series of interest. Lower RCC values indicate better phase correspondence of the considered time
 830 series to the reference.

831 • Normalised mean-absolute-error, $NMAE$, given by:

$$NMAE = \frac{MAE}{O_{max} - O_{min}} \times 100\%, \quad (A4)$$

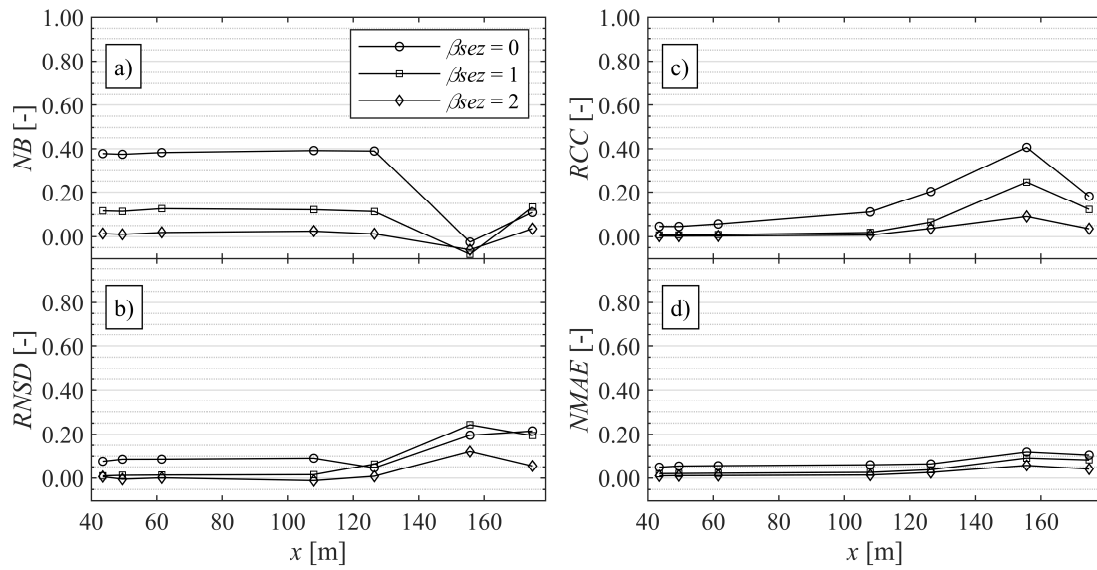
832 in which MAE is the mean absolute error, given by (3), and O_{max} and O_{min} are the maximum and
 833 minimum value of the reference time series.

834 The closer these statistics are to zero, the lower the difference is between the considered and reference
 835 time series.

836 A.2. Convergence analyses

837 The grid convergence analysis varies the refinement level in the surface elevation zone β_{sez} up to
 838 the dike toe (i.e. $\beta_{sez} = 0, 1, 2, 3$; Figure 3) and uses the mesh with the highest level (i.e. $\beta_{sez} = 3$ or $\Delta z =$
 839 $\Delta z = 0.0225$ m) as the reference to which the other – coarser – resolution simulations are compared to.
 840 Convergence is achieved when no significant changes are observed anymore compared to a finer grid
 841 resolution model. The time stepping convergence analysis uses the run with the lowest $maxCo$
 842 number (i.e. $maxCo = 0.15$) as the reference to which other temporally coarser simulations (i.e. $maxCo$
 843 $= 0.45, 0.25$) are compared to. The statistical error indicators from section A.1 are provided in Figure
 844 A1 and Figure A2, respectively. All errors stay close to or less than 5% at the toe of the dike for $\beta_{sez} =$
 845 2 (i.e. $\Delta x = \Delta z = 0.045$ m) and $maxCo = 0.25$. Even though $maxCo = 0.45$ does not show much higher
 846 errors than a value of 0.25, still $maxCo = 0.25$ was preferred, because higher $maxCo$ simulations were
 847 found to be prone to numerical instabilities. In any case, as long as the $maxCo$ number cannot be
 848 defined separately for the air and water phases, the time stepping is mostly determined by the high
 849 spurious velocities that occur at the water-air interface. Because these spurious velocities are much
 850 higher (2-3 times) than the velocities in the water phase, much lower Courant numbers are actually
 851 obtained in the water phase [45]. This also explains why only limited differences between the tested
 852 $maxCo$ values are observed here.

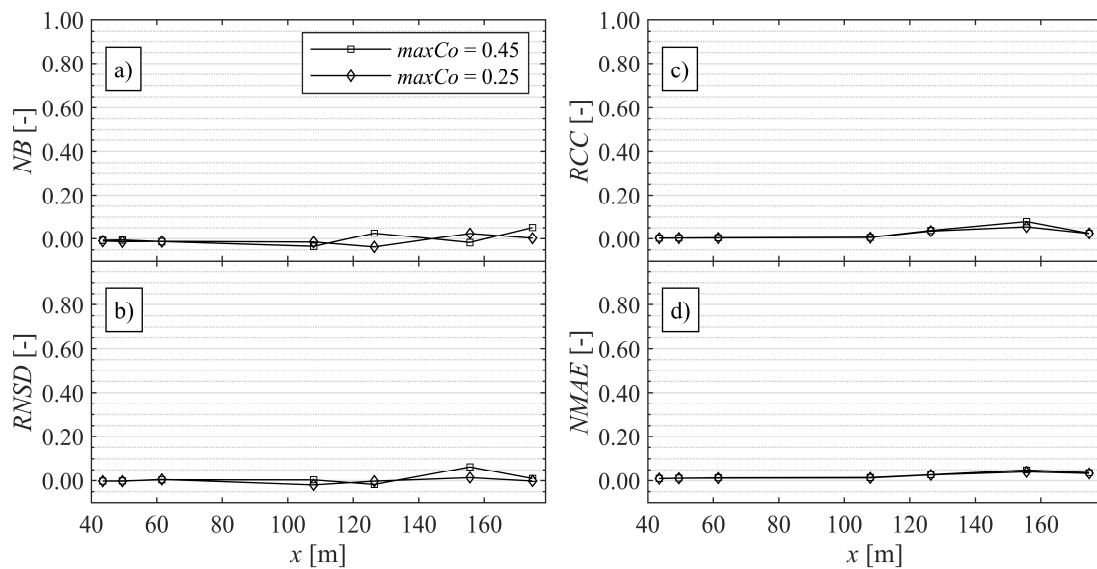
853 Moreover, the $NMAE$ shows in both cases a similar value at the toe of the dike (WG14) to that of
 854 the ~3% obtained between EXP and REXP. The remaining numerical error is therefore assumed
 855 acceptable, and the mesh resolution and time stepping can be considered sufficiently converged for
 856 those settings ($\beta_{sez} = 2$; $maxCo = 0.25$).



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Figure A1. OF model grid resolution convergence analysis of the η time series at the WG locations along the flume up to the dike toe (WG14), based on: a) the normalised bias, b) the residual normalised standard deviation, c) the residual correlation coefficient and d) the normalised mean-absolute-error. The reference is the finest mesh with a refinement level in the surface elevation zones β_{sez} of 3.



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Figure A2. Same as Figure A-1, but a time stepping convergence analysis instead for the mesh with $\beta_{sez} = 2$. The reference is the lowest maximum Courant number applied ($maxCo = 0.15$).

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