



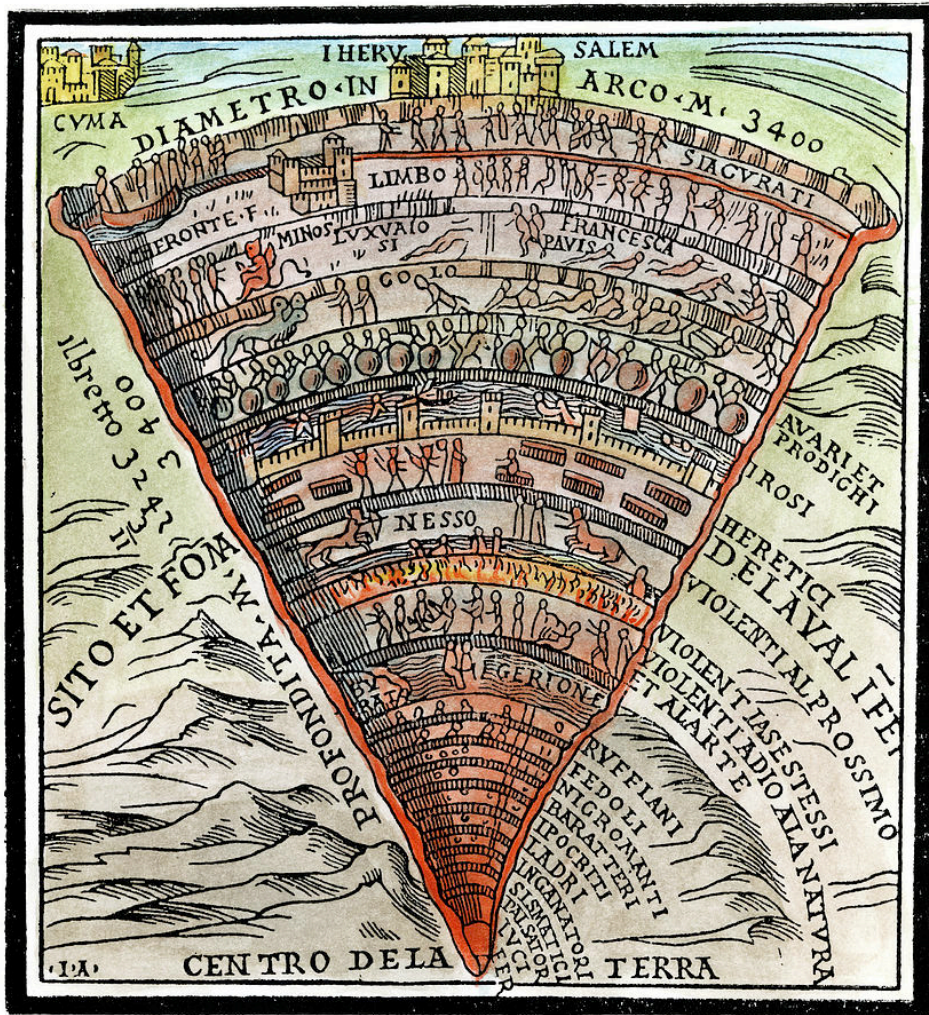
Earth tomography with neutrinos

Andrea Donini (IFIC, Valencia)

Nature Physics 15 (2019) 37

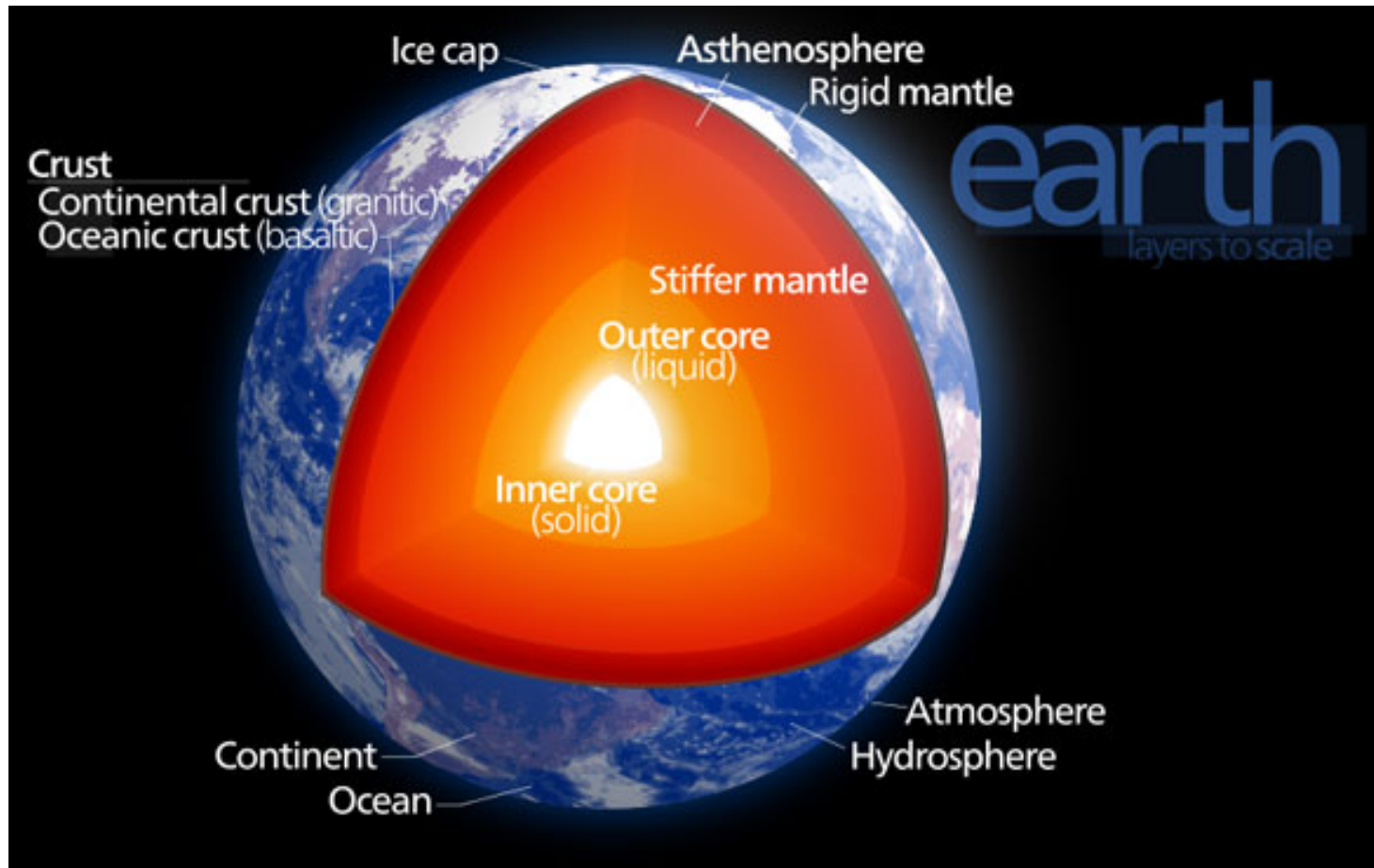
in collaboration with:
S. Palomares-Ruiz
J. Salvadó

An outdated view of Earth's interior



DANTE'S INFERNO, c1520.
Woodcut from a Venetian edition
of the Divine Comedy.

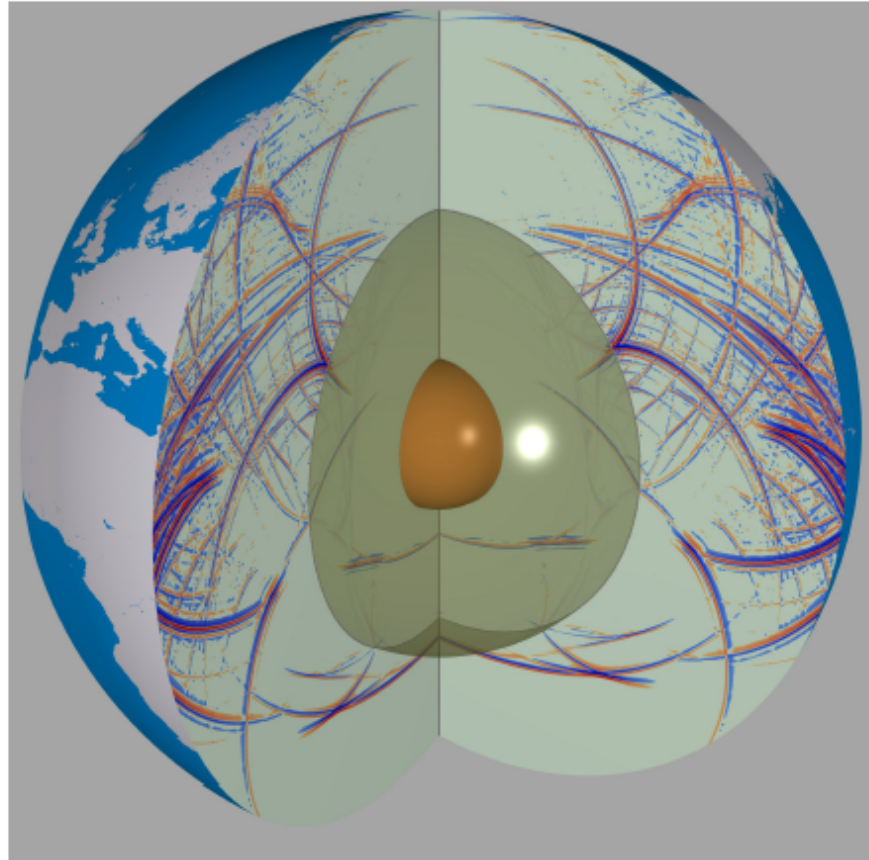
A more recent view



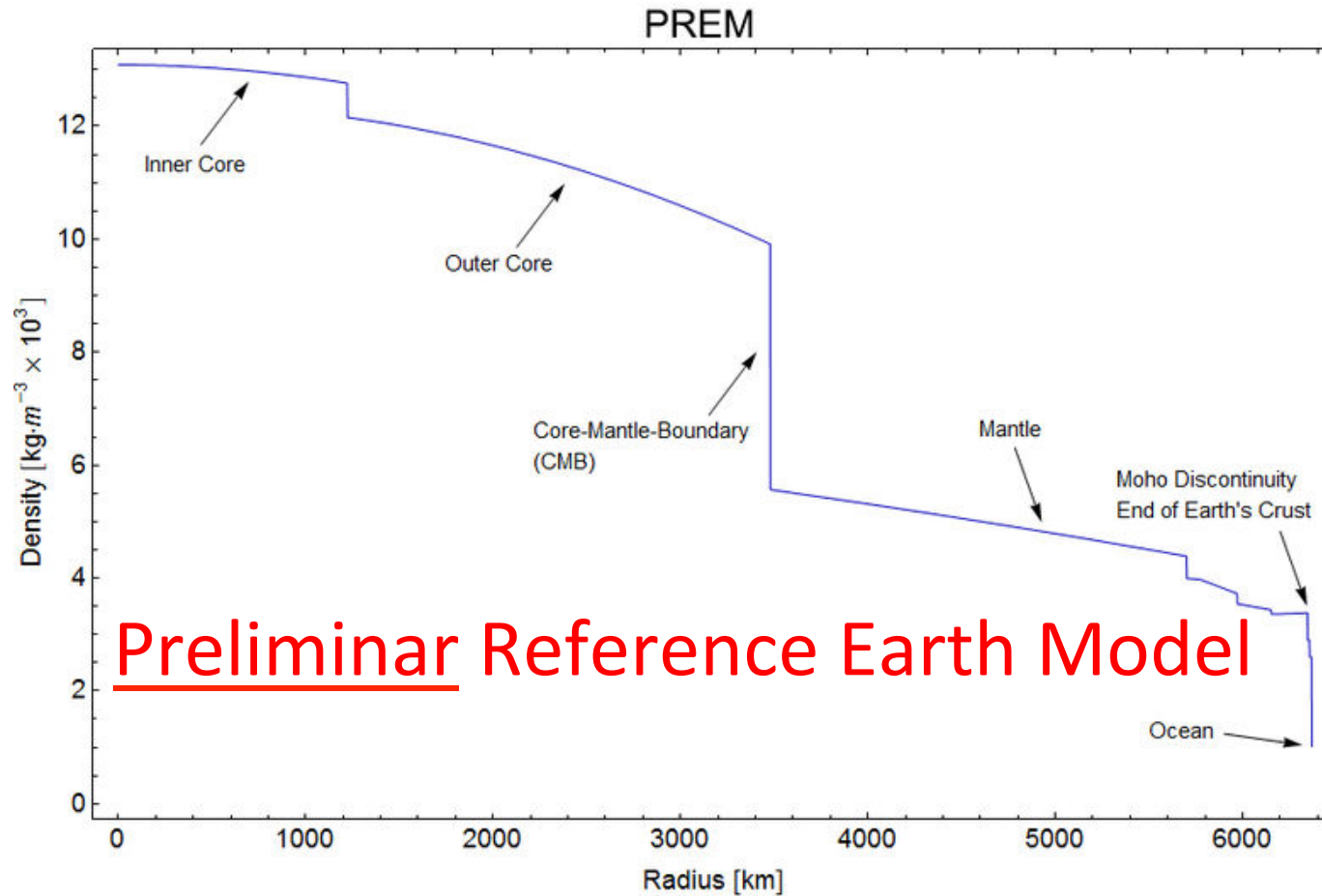
How densities are measured?

seismology

$O(100)$ /year earthquakes
with magnitude
larger than 6

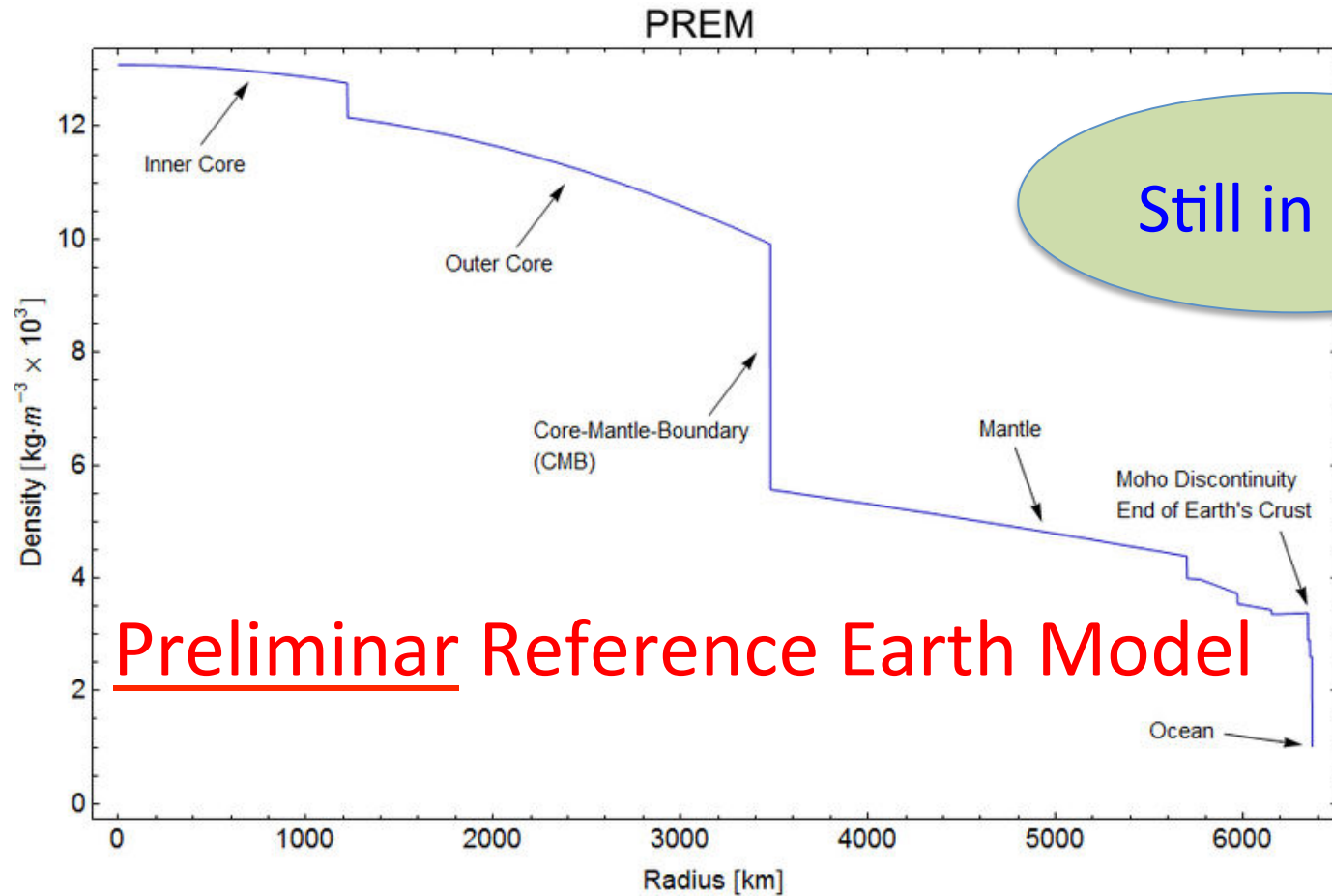


1-dimensional density profile



[Dziewonski and Anderson, Physics of the Earth and Planetary Interiors, 25 (1981)]

1-dimensional density profile



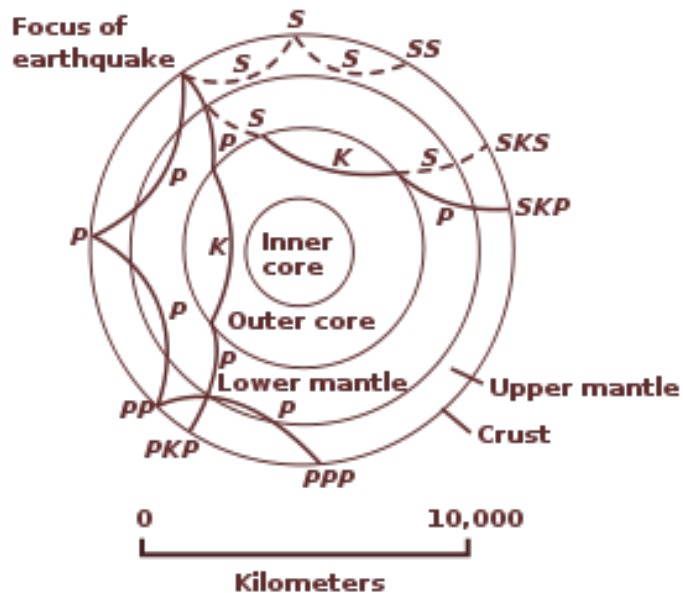
[Dziewonski and Anderson, Physics of the Earth and Planetary Interiors, 25 (1981)]

The Earth's core

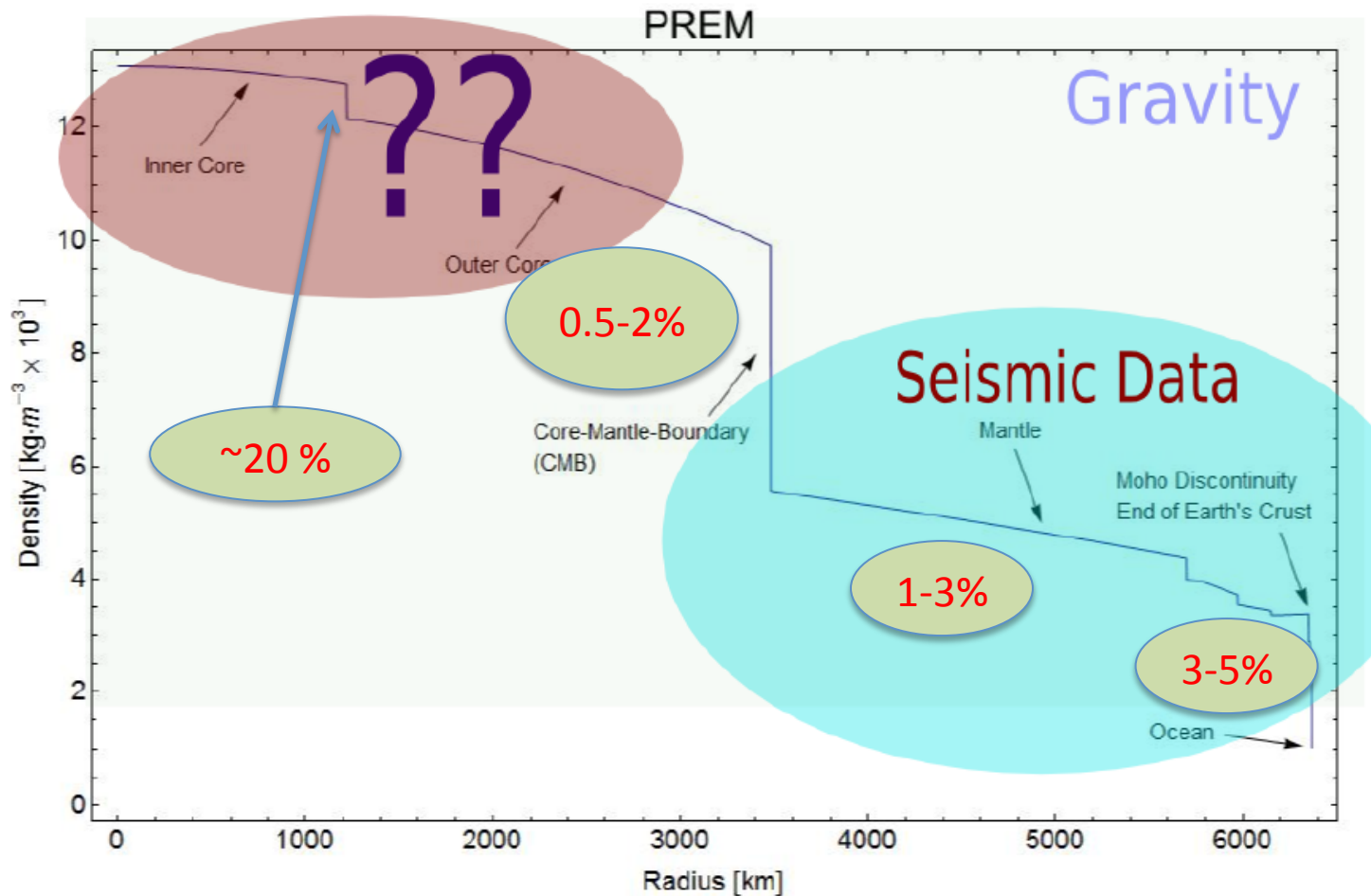
The **OUTER CORE IS LIQUID**,
whereas
the **INNER CORE IS SOLID**
(source of the geodynamo)

IT IS VERY DIFFICULT TO
HAVE DIRECT
INFORMATIONS from the
INNER CORE

Mostly through **global constraints**
and **extrapolations**



Uncertainties from seismology



Neutrinos to study the Earth's interior

An old idea: first mentioned in an unpublished CERN preprint,

A.Placci and E. Zavattini, submitted in Oct 1973 to Nuovo Cimento;
rejected?... never received?....

and in a talk

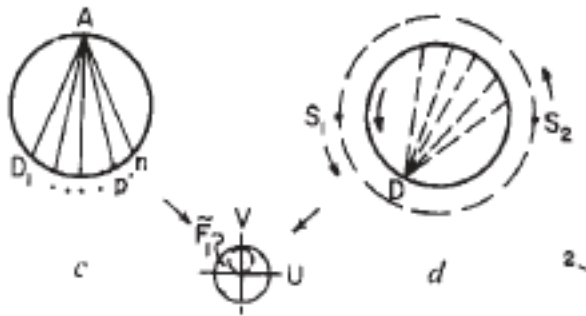
L. V. Volkova and G. T. Zatsepin, Izv. Akad. Nauk. Ser. Fiz. 38N5 (1974)

In modern language, a long-baseline experiment

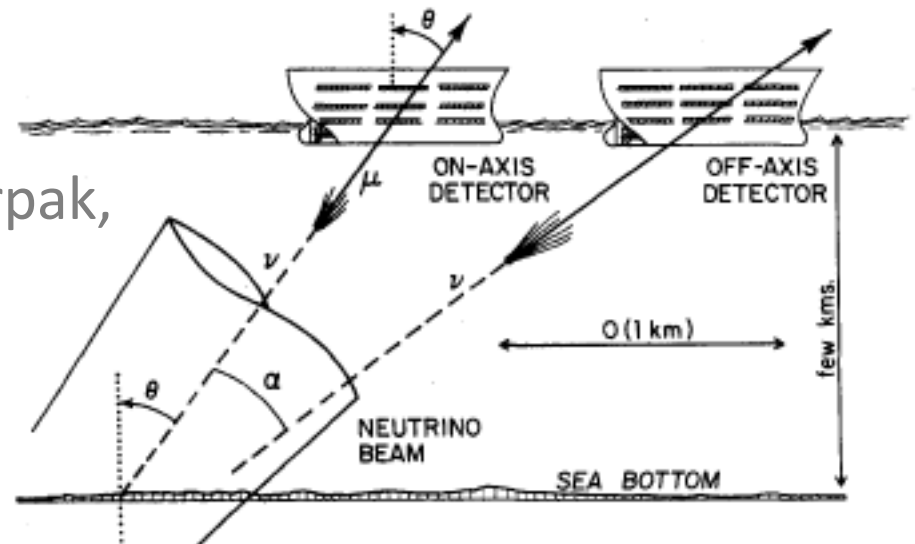
The idea was premature!

Neutrino tomography

Even more premature...

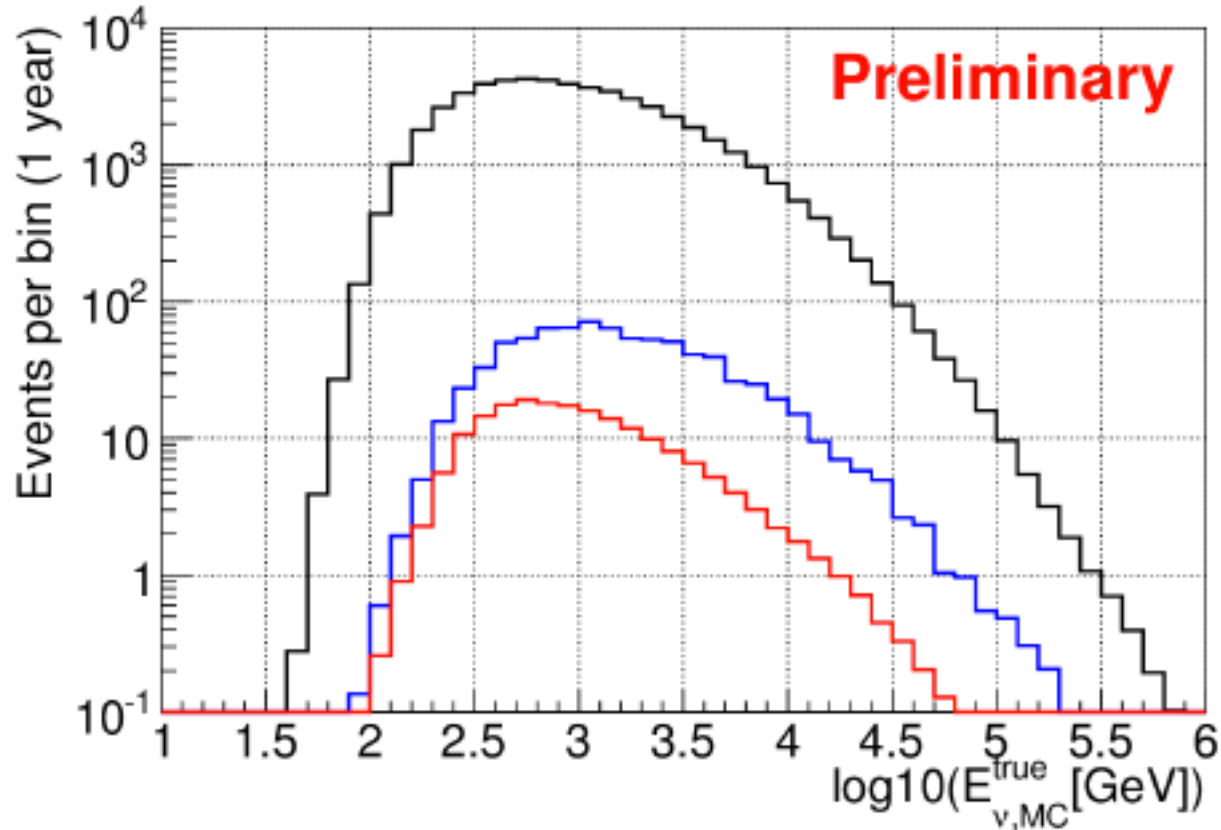


T. Wilson, Nature 309 (1984)



De Rújula, Glashow, Wilson, Charpak,
Phys. Rept. 99 (1983)

An optimal source: atmospheric neutrinos



IceCube contribution to ICRC 2015, arXiv:1510.05223

Two ways to scan the Earth

- Neutrino oscillations (< 1 TeV)

$$P_{ee}^{\pm} = 1 - \left(\frac{\Delta_{23}}{B_{\mp}} \right)^2 \sin^2(2\theta_{13}) \sin^2 \left(\frac{B_{\mp} L}{2} \right) - \left(\frac{\Delta_{12}}{A} \right)^2 \sin^2(2\theta_{12}) \sin^2 \left(\frac{A L}{2} \right)$$

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W. Winter, Nucl. Phys. B 908 (2016) 250; Km3Net, PoS ICRC2017 (2018) 1020

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- Neutrino flux attenuation (> 1 TeV)

$$\frac{d\phi_{\nu}(E, \tau)}{d\tau} = -\sigma_{tot}(E)\phi_{\nu}(E, \tau)$$

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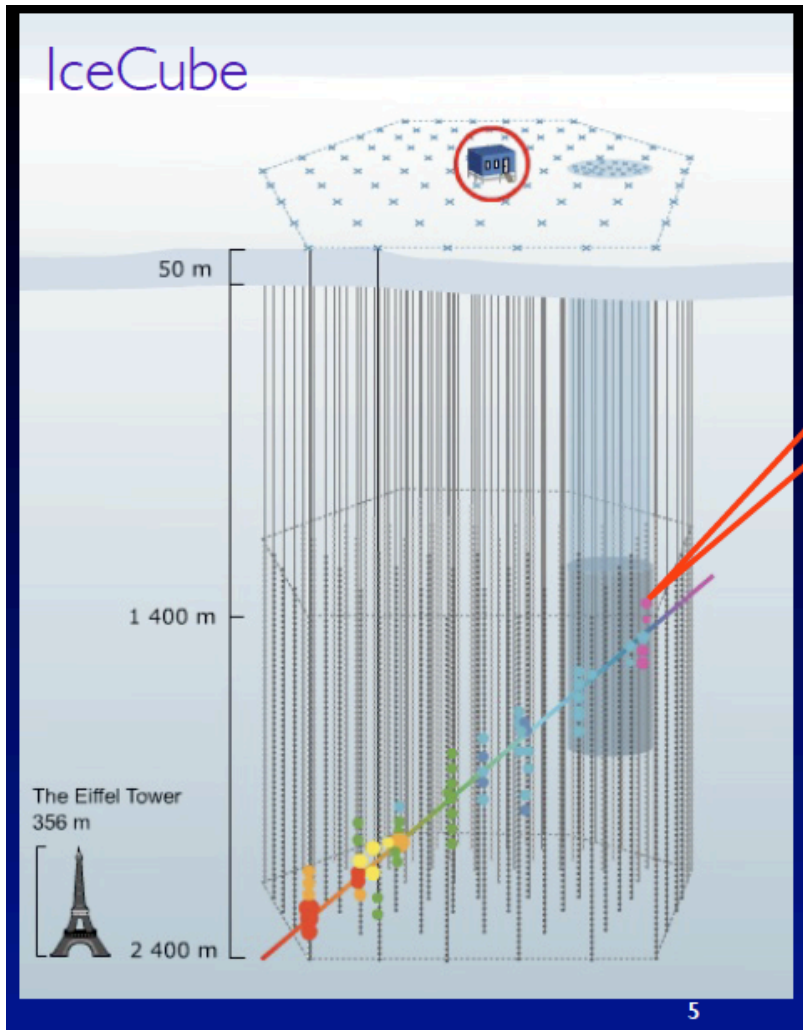
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Gonzalez-García, Halzen, Maltoni, Tanaka, Phys. Rev. Lett. 100 (2008)

The IceCube Experiment

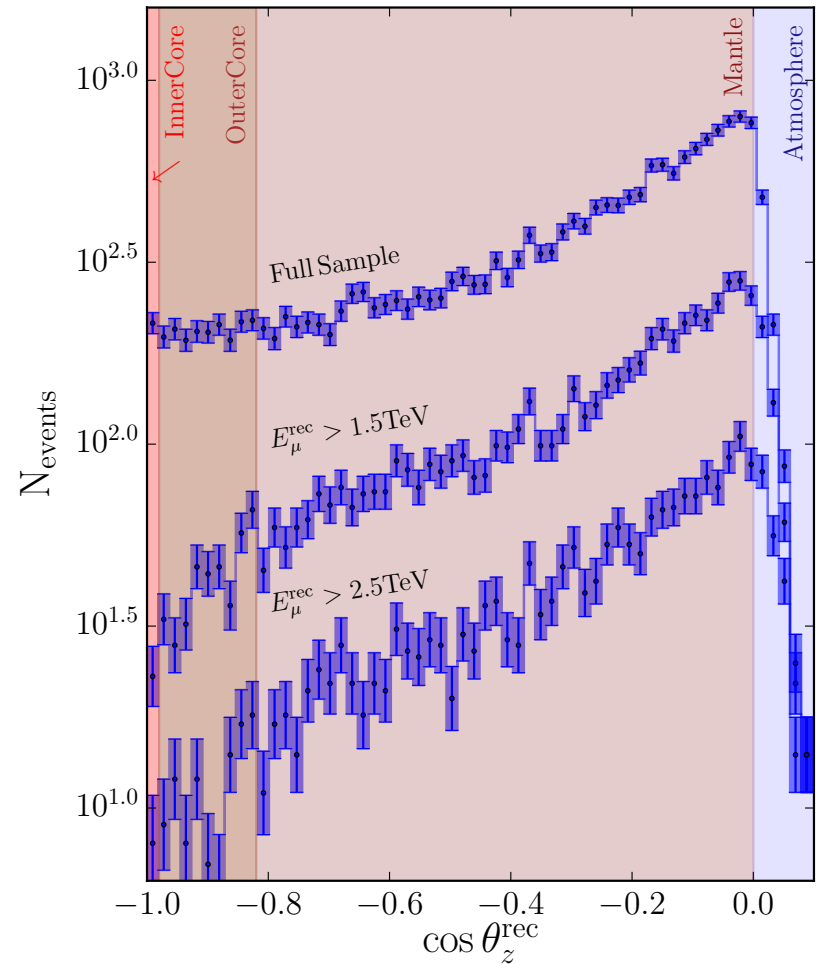
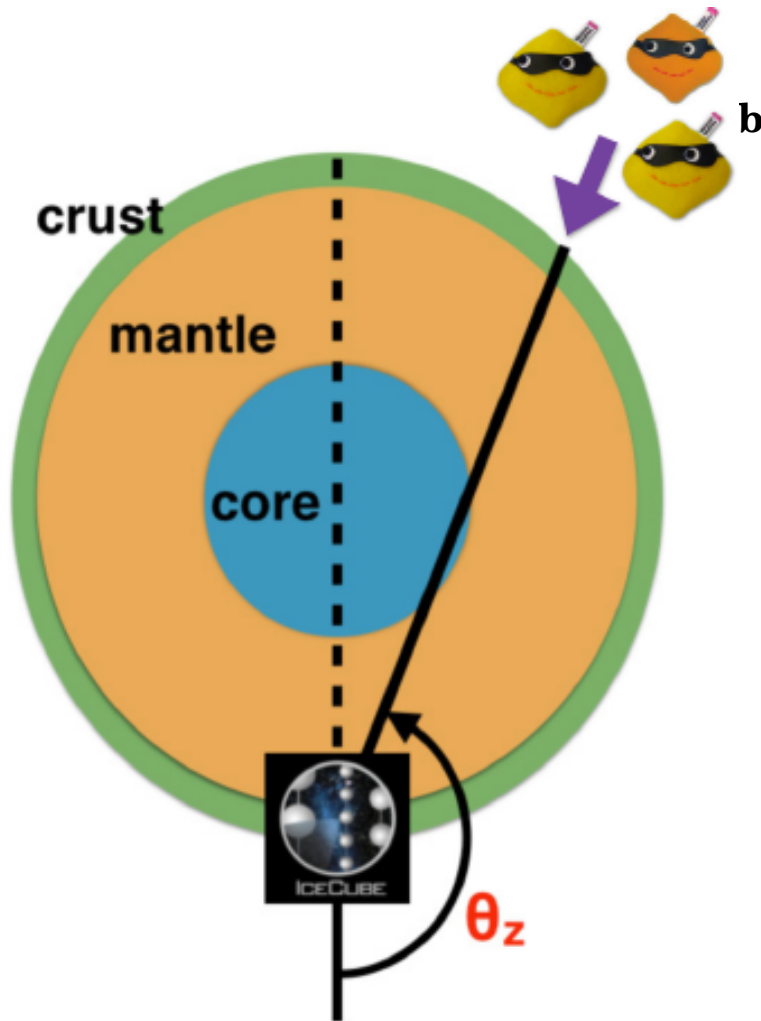


- Deployed in glacial ice at the South Pole
- Array size 1 km^3 , 86 strings, 60 optical sensors (DOMs) per string

The IceCube IC86 data sample

- 1 year of data taking (2011-2012)
 - 20145 muon events over 343.7 days
 - $E_{\mu} = [400 \text{ GeV} \div 20 \text{ TeV}]$
 - The muon direction is a very good proxy of the neutrino direction, with $\Delta\cos\theta < 0.01$
 - **PUBLICLY AVAILABLE!**
- 7 more years of data are not (yet) available.....

Raw data as a function of E_μ and θ



Comparison with expectations

Flux model

Propagation

Interaction with
nucleons

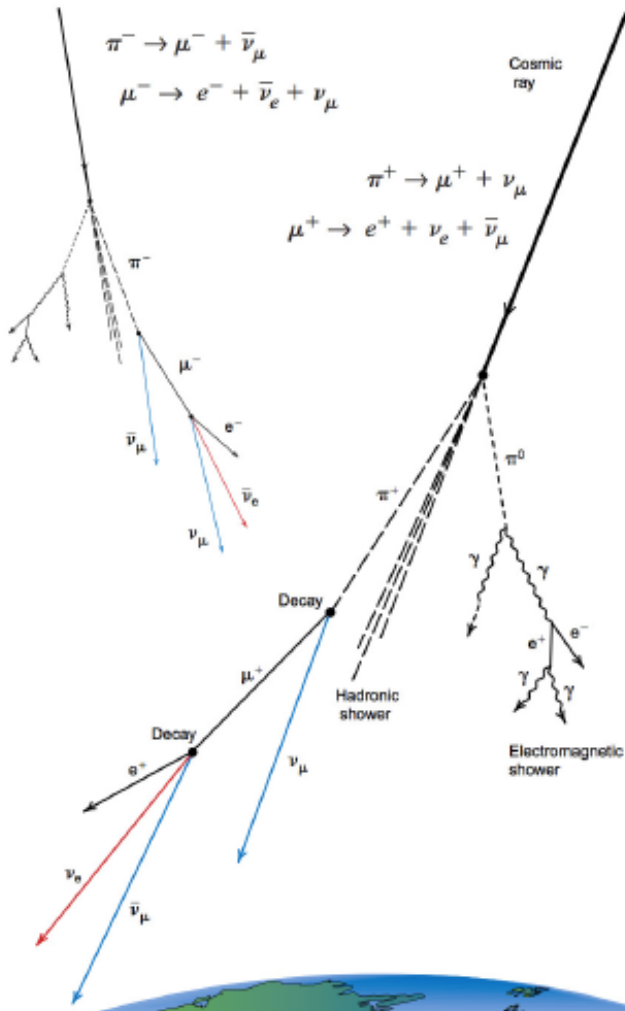
Detector
simulation

Flux model dependence

Primary cosmic ray flux:
Honda-Gaisser model +
Gaisser-Hillas corrections
(HG-GH-H3a)

Hadronic model: QGSJET-II-04

Other options → “discrete”
systematics
(improvements expected, PANE18 talks)



Neutrino propagation

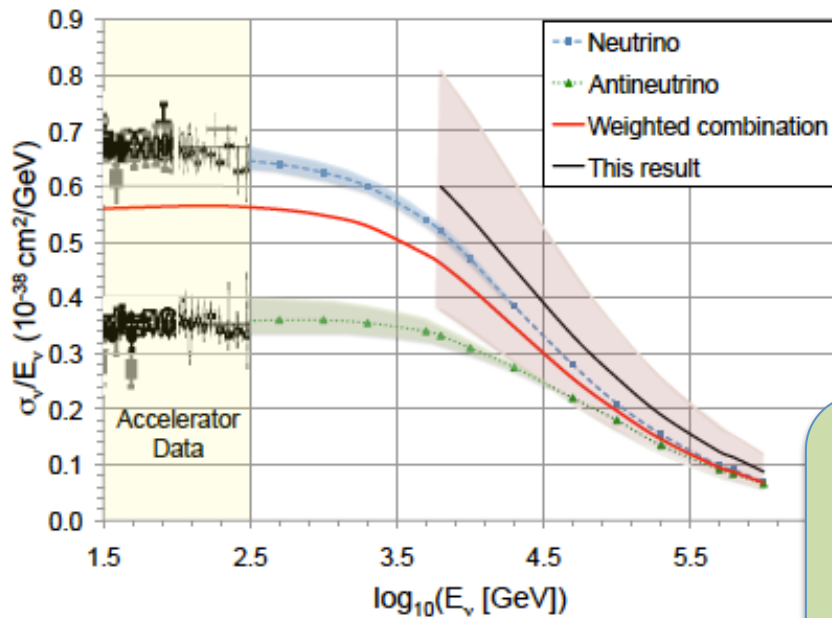
Propagation through the Earth with ν -SQUIDs

Argüelles Delgado, Salvado & Weaver, Comput. Phys. Commun. 196 (2015)

- **Neutrino Oscillations:** evolution Hamiltonian in matter
(dominant below 1 TeV)
- **Neutrino Attenuation:** inelastic CC and NC interactions with matter
(dominant above 1 TeV)
- Neutrino regeneration due to tau decays
- **Migration to lower energy bins** due to NC interactions

Neutrino-nucleon interaction

Parton distribution functions: HERAPDF

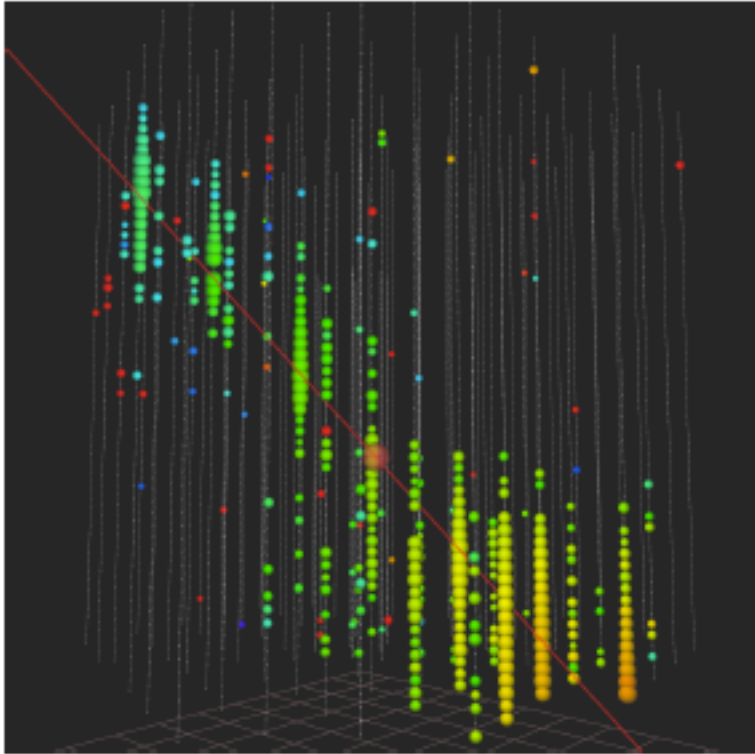


Aarsten et al, Nature 551 (2017)

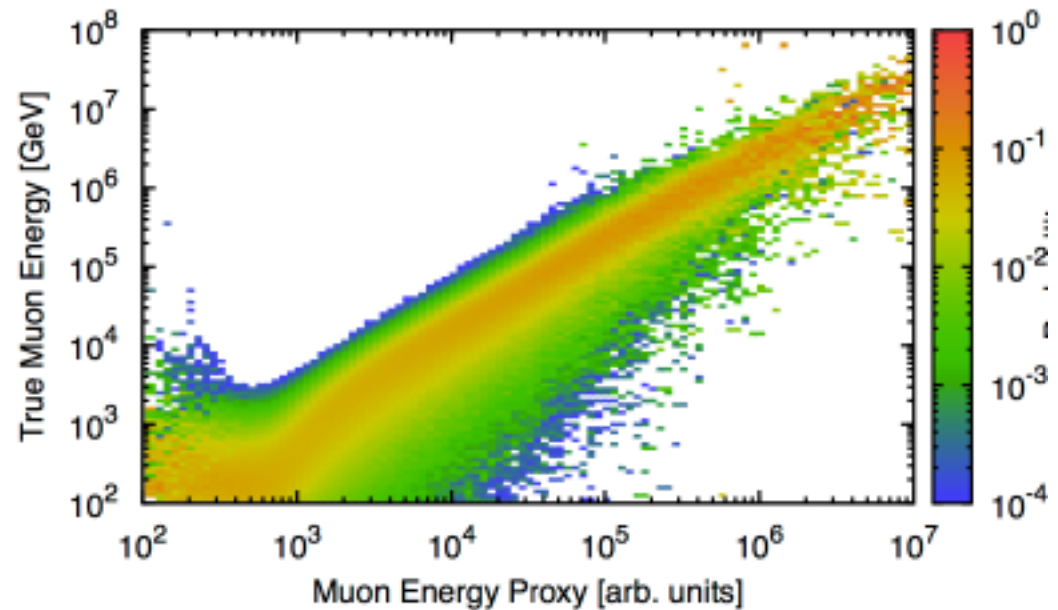
νN ($\bar{\nu} N$) cross-sections
at 2-3% (4-10%) errors

ICECUBE MEASUREMENT
 $1.30^{+0.21}_{-0.19}$ (stat) $^{+0.39}_{-0.43}$ (syst) $\times \sigma_{SM}$

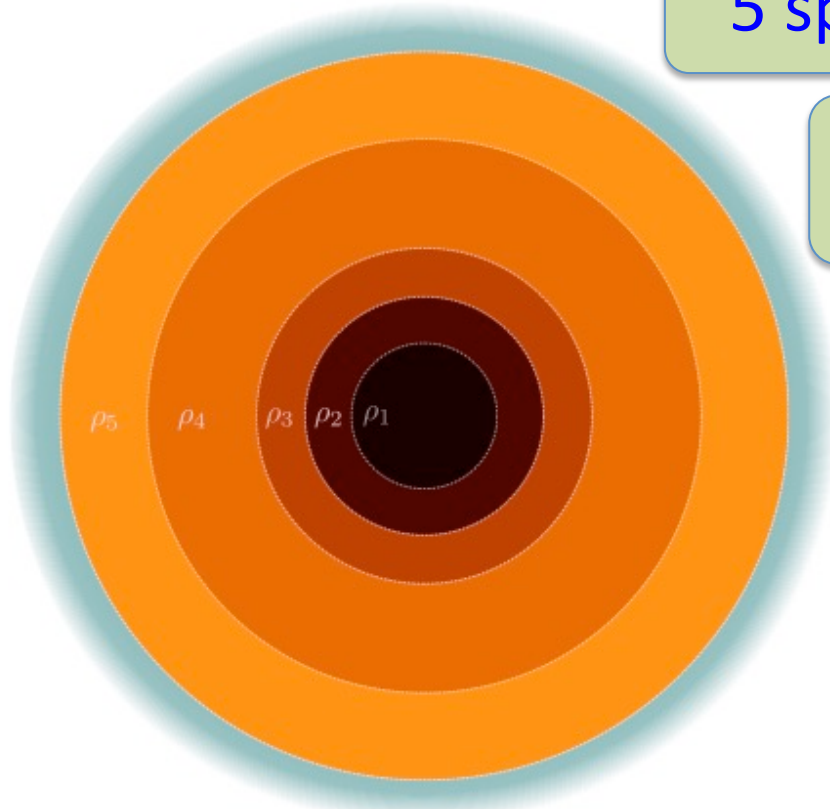
Detector simulation



We use the official
IceCube MC to map
 E_{obs}^{μ} , $\theta_{\text{obs}}^{\mu}$ into E_{rec}^{ν} , $\theta_{\text{rec}}^{\nu}$



Our Earth's model



5 spherical layers

Inner Core, one layer

$L_1 = 1242$ km

Outer Core, two layers

$L_2 = 2373$ km,

$L_3 = 3504$ km

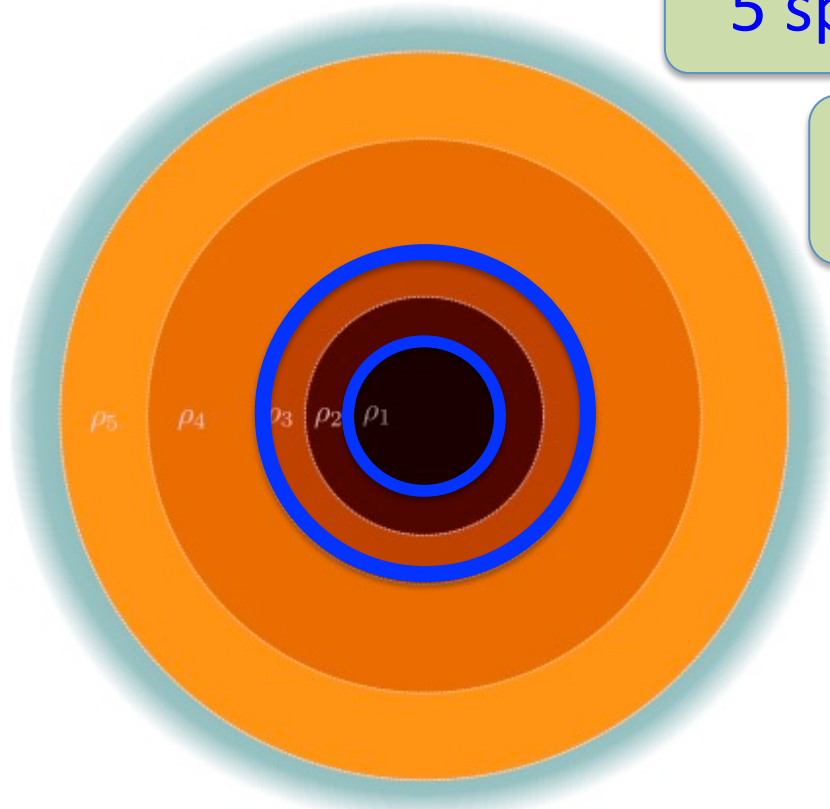
Mantle, two layers

$L_4 = 4938$ km,

$L_5 = 6371$ km

No crust!

Our Earth's model



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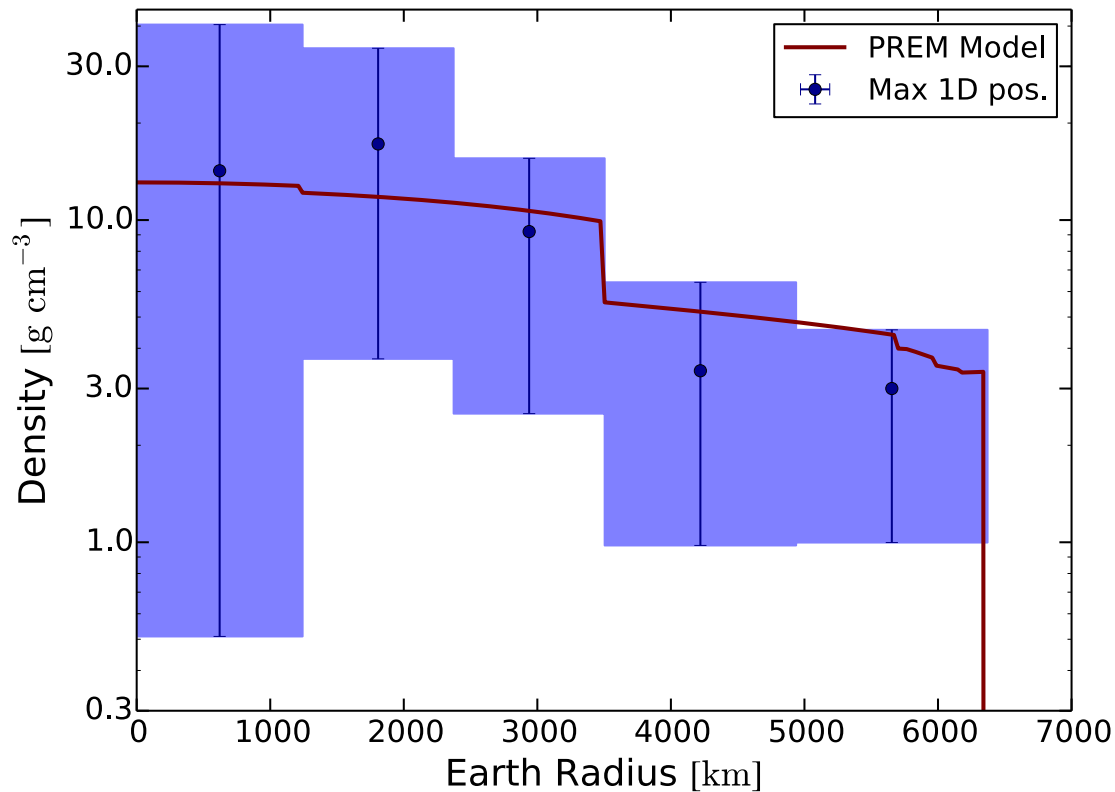
$L_4 = 4938$ km,

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ICB and CMB fixed!

No crust!

First 1-d density profile with neutrinos



Analysis performed
with MultiNest

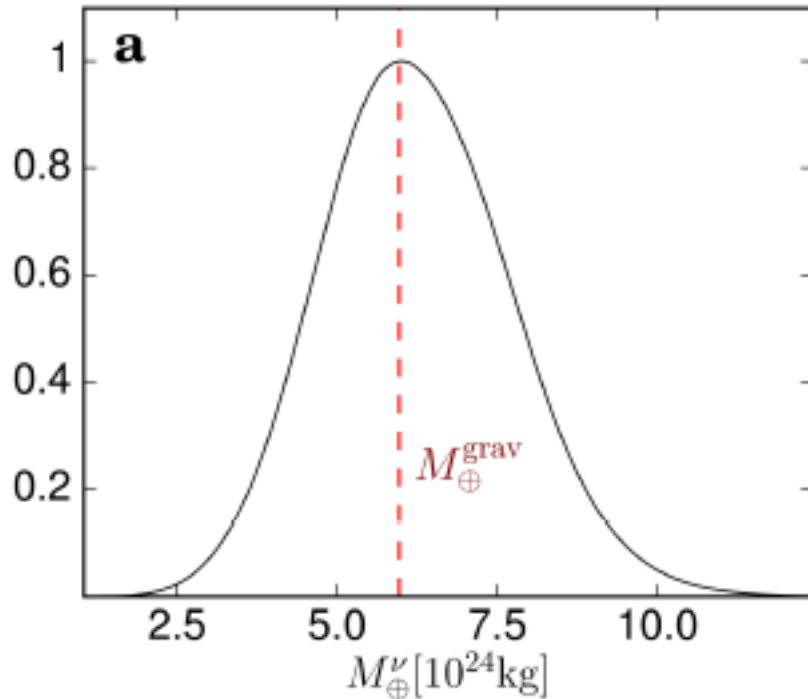
5 Earth layers densities

and

4 systematic errors:

- Flux normalization
- Pion-to-kaon ratio
- Spectral shape
- DOM Efficiency

The Earth's mass



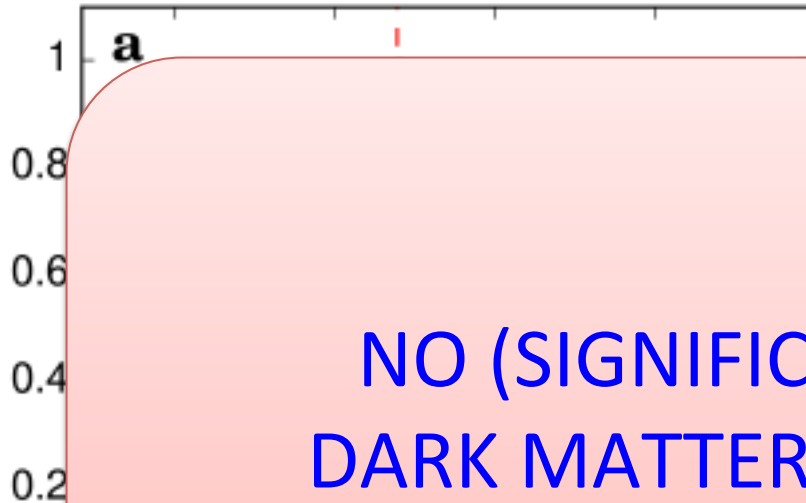
First Electro-weak measurement
of the Earth's mass

$$M_{\text{earth-v}} = (6.0^{+1.6}_{-1.3}) \times 10^{24} \text{ kg}$$

Gravitational measurement of the Earth's mass

$$M_{\text{earth-grav}} = (5.9722 \pm 0.0006) \times 10^{24} \text{ kg}$$

The Earth's mass

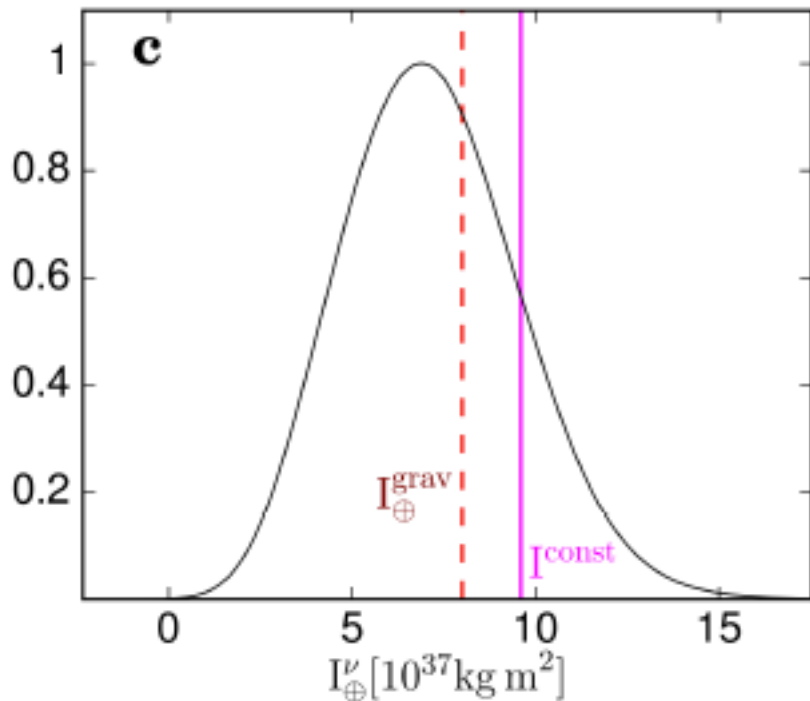


NO (SIGNIFICANT AMOUNT OF)
DARK MATTER INSIDE THE EARTH!

Gravitational measurement of the Earth's mass

$$M_{\text{earth-grav}} = (5.9722 \pm 0.0006) \times 10^{24} \text{ kg}$$

The Earth's moment of inertia



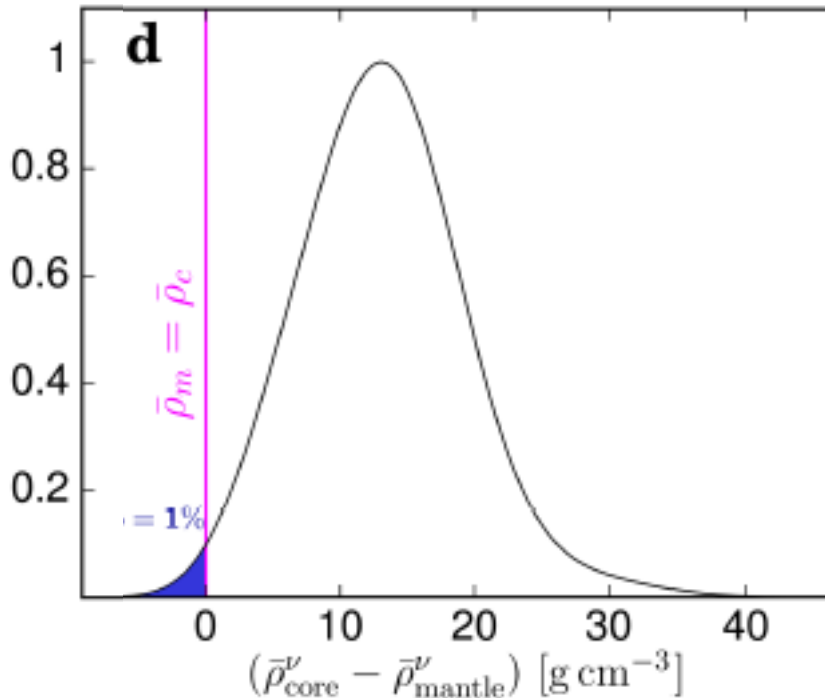
Electro-weak measurement of
the Earth's moment of inertia

$$I_{\text{earth-}\nu} = (6.9 \pm 2.4) \times 10^{37} \text{ kg m}^2$$

Gravitational measurement of the Earth's moment of inertia

$$I_{\text{earth-grav}} = (8.01736 \pm 0.00097) \times 10^{37} \text{ kg m}^2$$

Earth's non-homogeneity



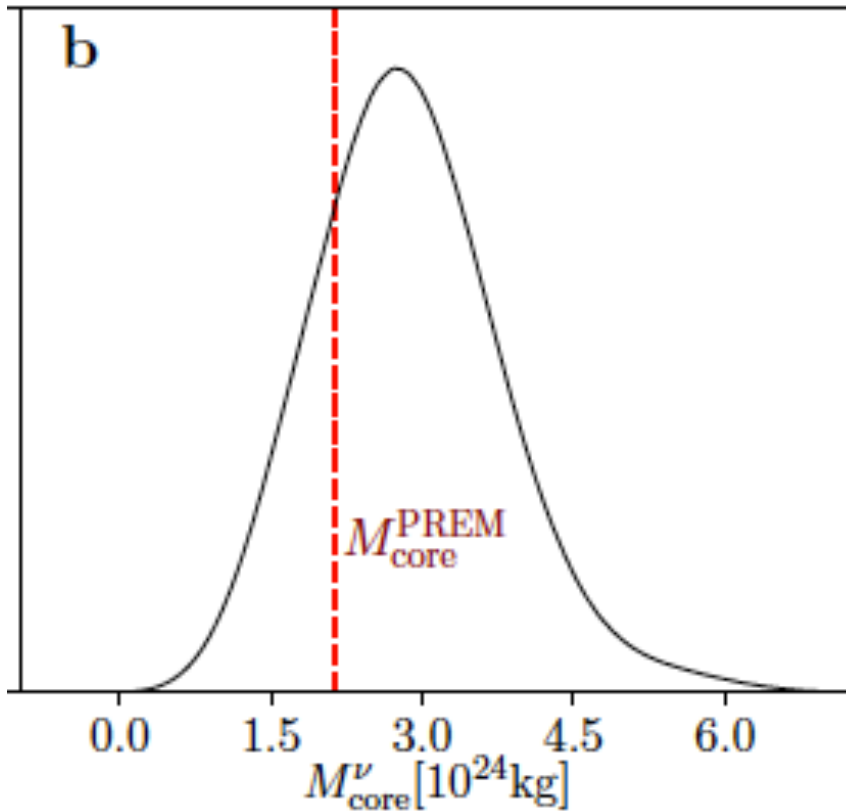
Electro-weak measurement of the Core-Mantle discontinuity

$$\Delta\rho_{\text{CMB-}\nu} = (13^{+5.8}_{-6.3}) \text{ g/cm}^3$$

A homogenous Earth has a p-value $p = 0.01$!!!

2008 Claim: IceCube could reject a homogeneous Earth at 5σ in ten years

The Earth's core mass



Electro-weak measurement of
the Earth's core mass

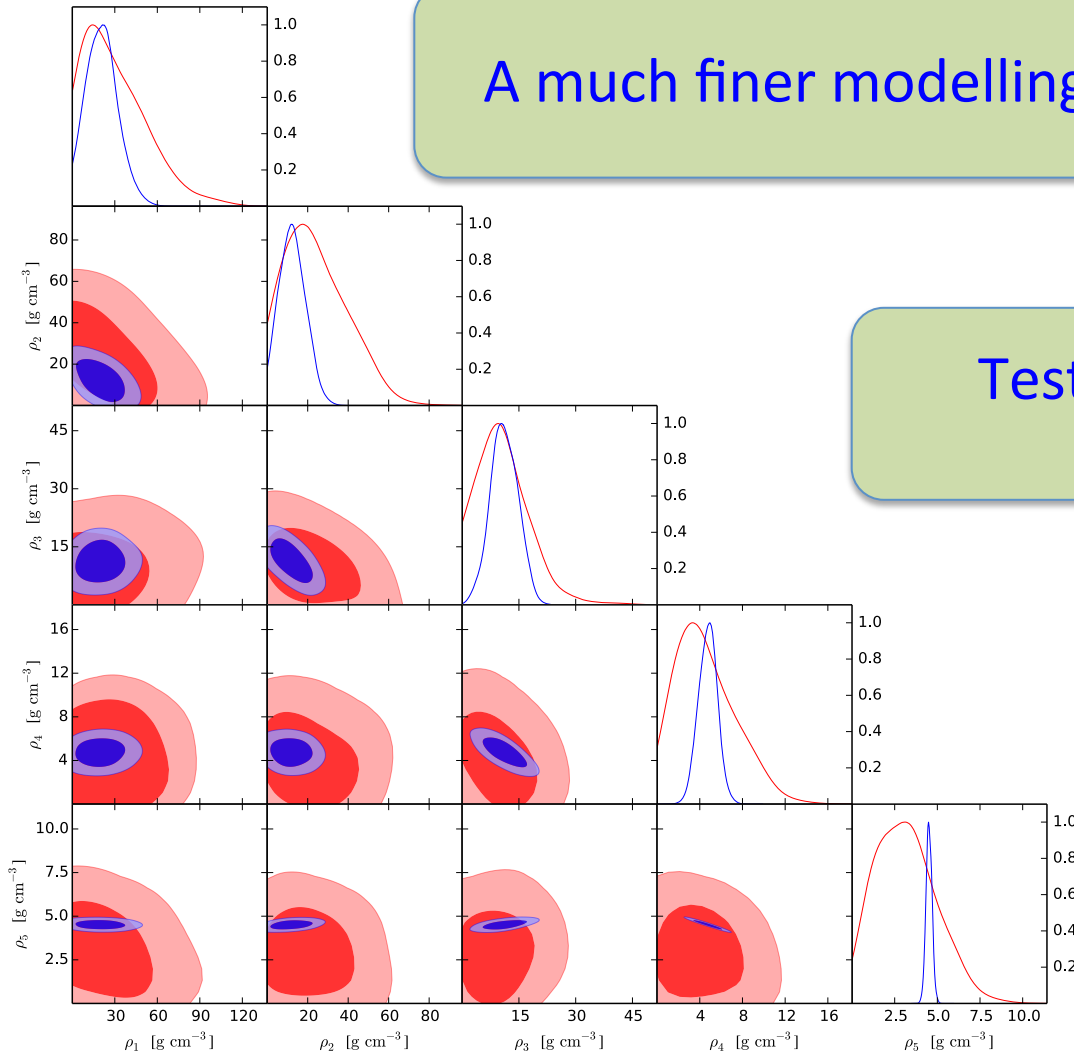
$$M_{\text{core-}\nu} = (2.7^{+1.0}_{-0.9}) \times 10^{24} \text{ kg}$$

Forecast with 10 years of data

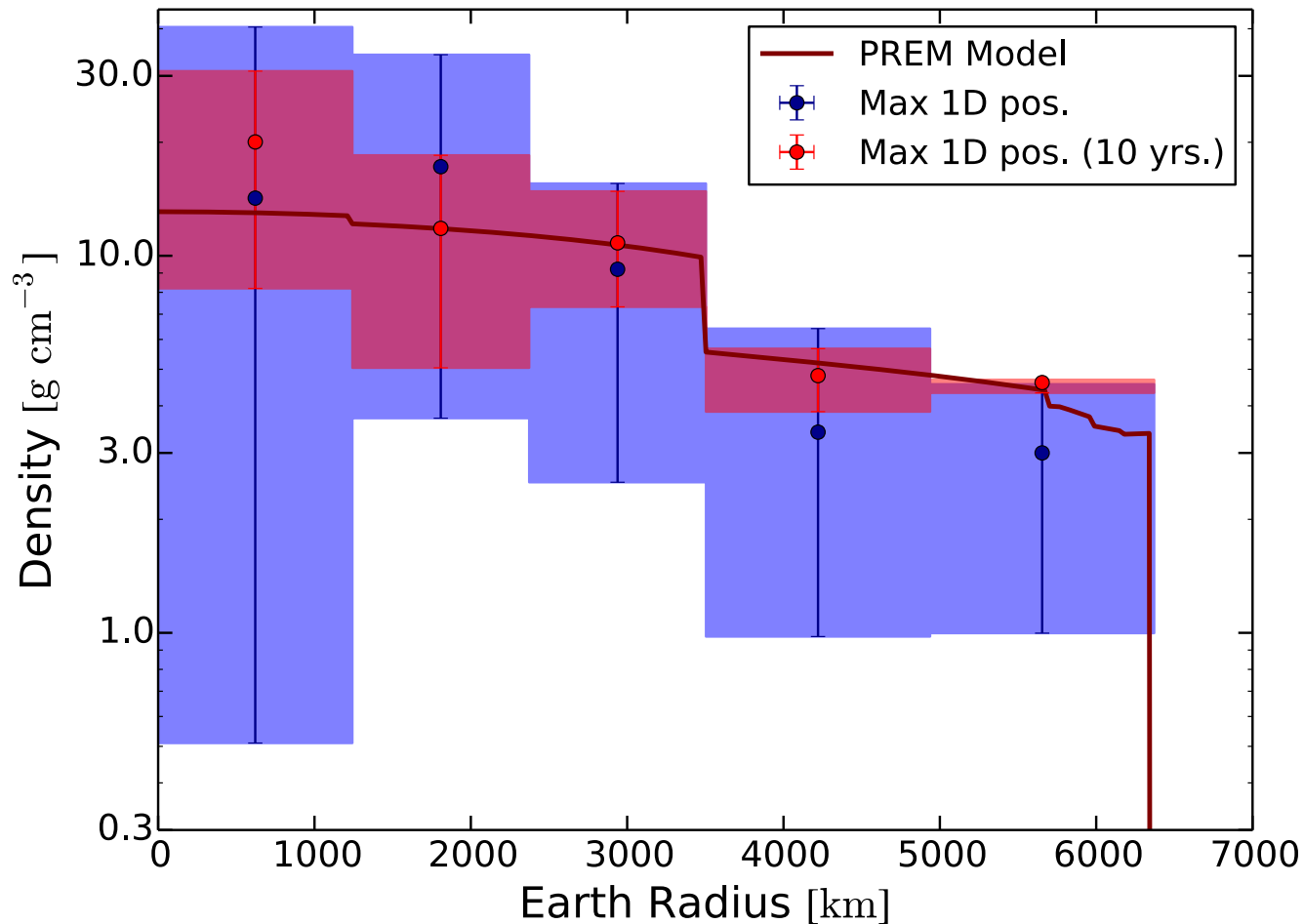
A much finer modelling of the Earth could be done

Test of the Inner-Outer Core discontinuity

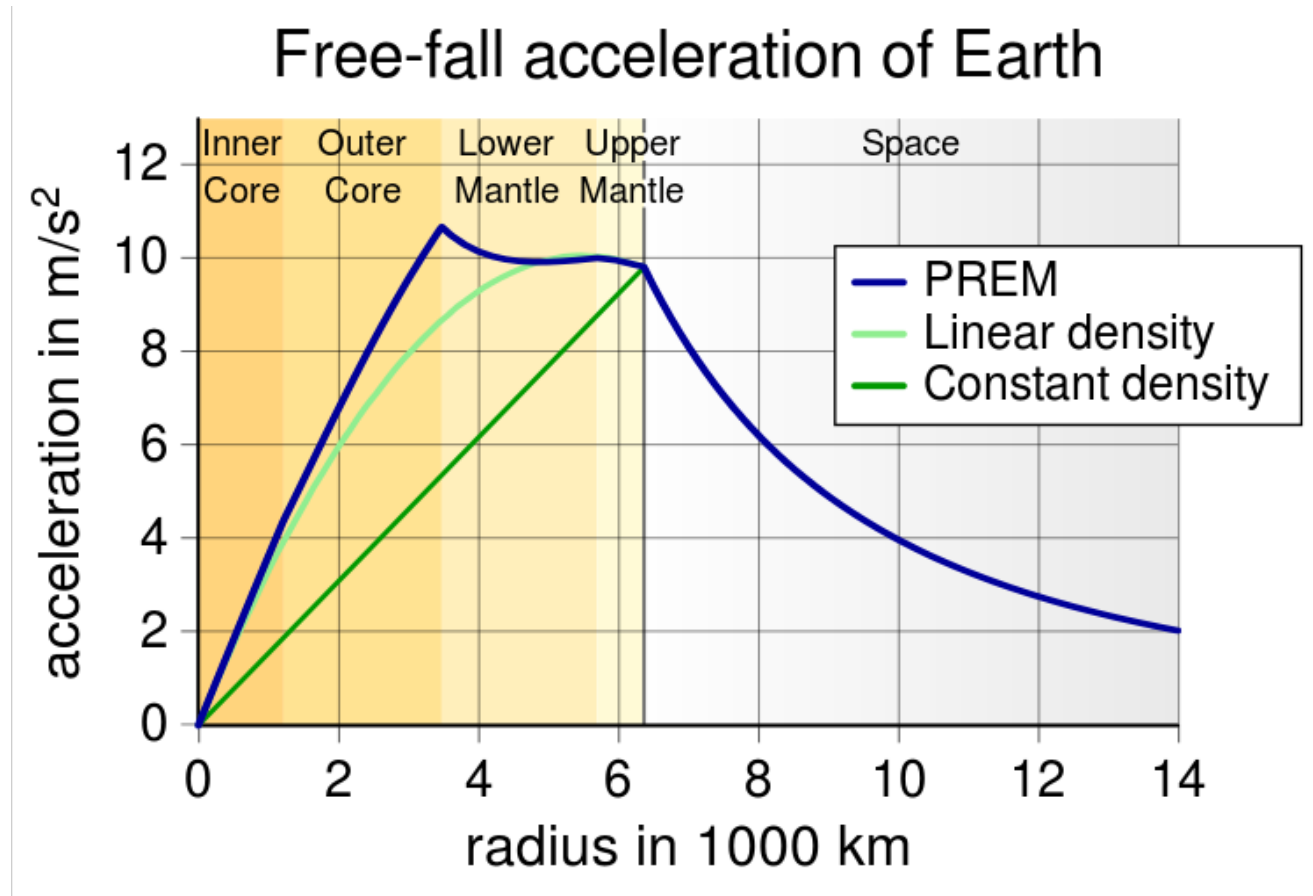
Independent localization of the Core-Mantle Boundary



1-d density profile with 10 years

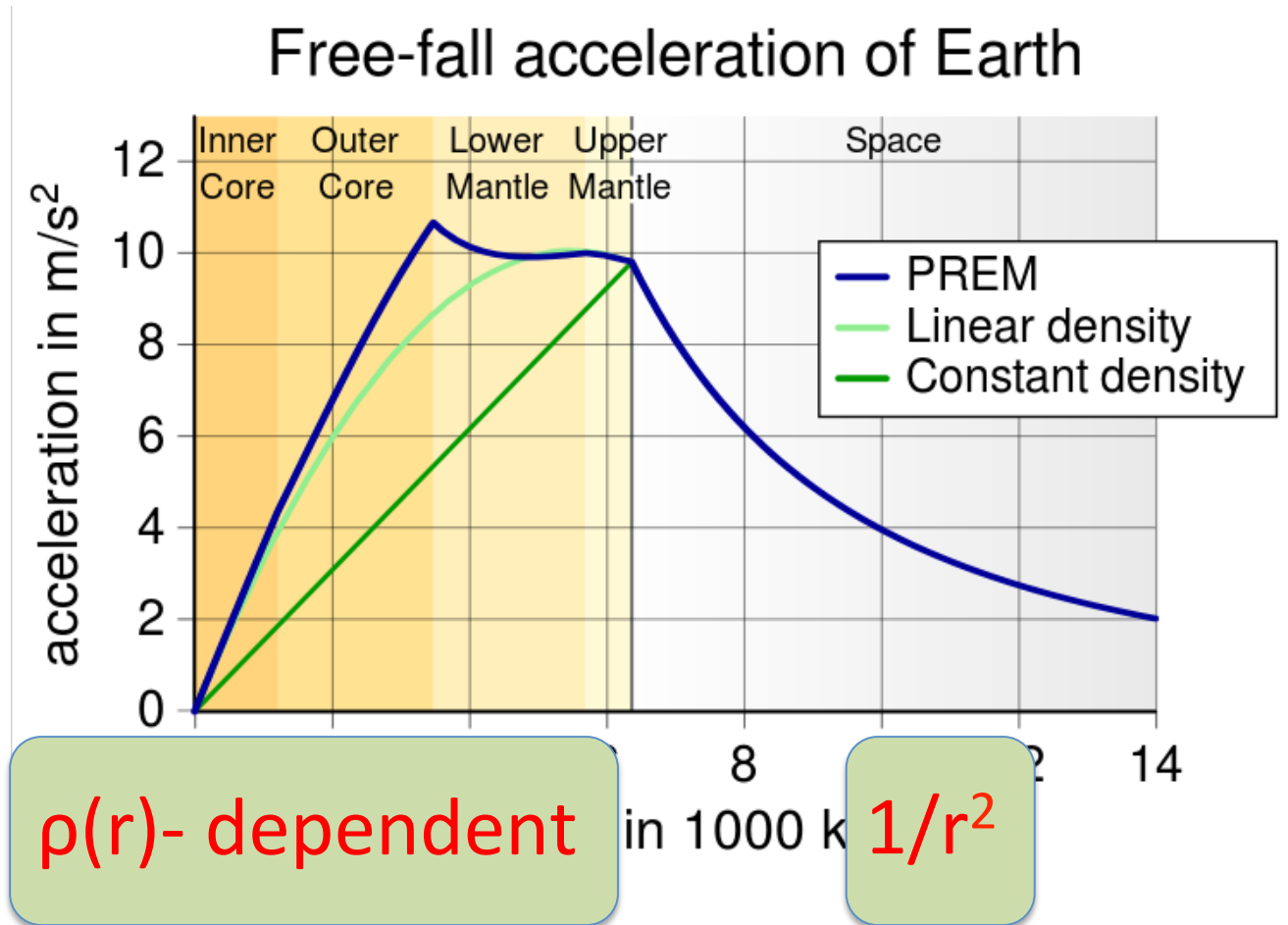


An input to geophysics: $g(r)$



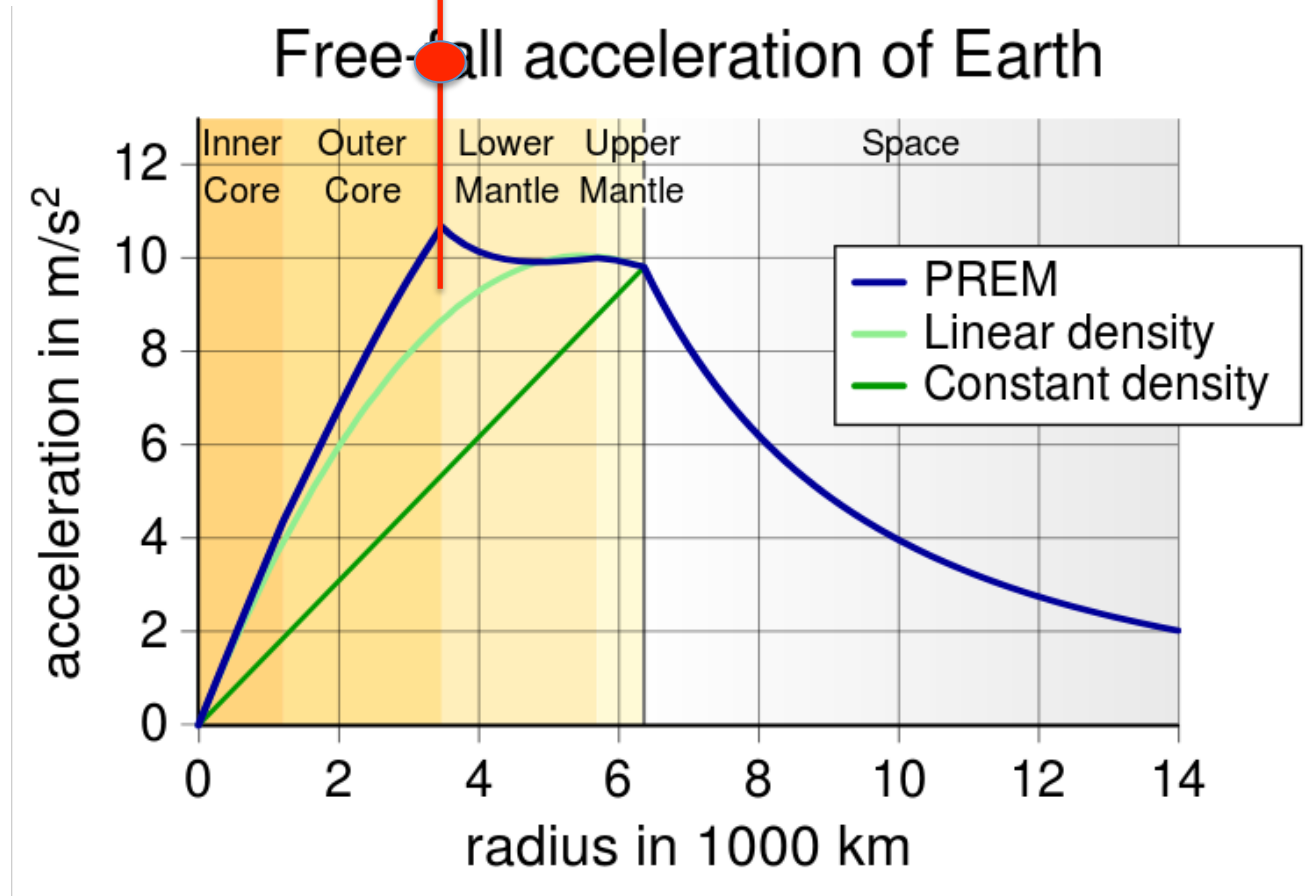
The Earth's gravitational profile is needed to compute $\rho(r)$ from earthquake waves velocities

An input to geophysics: $g(r)$



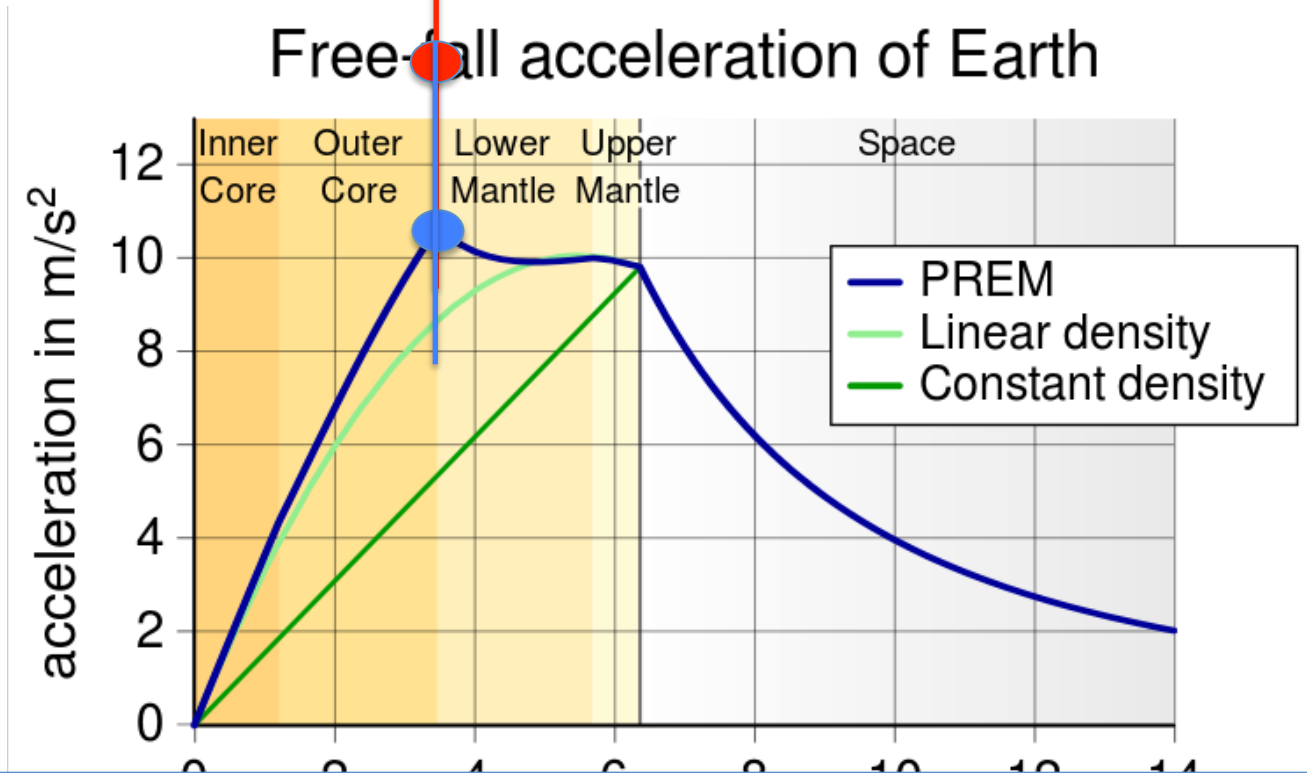
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The Earth's gravitational profile is needed to compute $\rho(r)$ from earthquake waves velocities

An input to geophysics: $g(r)$



A **GOOD** NEUTRINO MEASUREMENT OF $g(r)$ COULD BE ADDED TO SEISMOLOGY AS A PRIOR TO REDUCE ERRORS (STILL TO TEST THE IMPACT)

The Earth's gravitational profile is needed to compute $\rho(r)$ from earthquake waves velocities

Conclusions

It is eventually possible to make a neutrino tomography of the Earth: **first 1-dimensional density profile** (with just one year of IceCube data)! M_{earth} , I_{earth} , $\Delta\rho_{\text{CMB}}$, M_{core}

Precision will hugely increase as soon as 7 other years of IceCube data will become accessible (percent level in the mantle)!

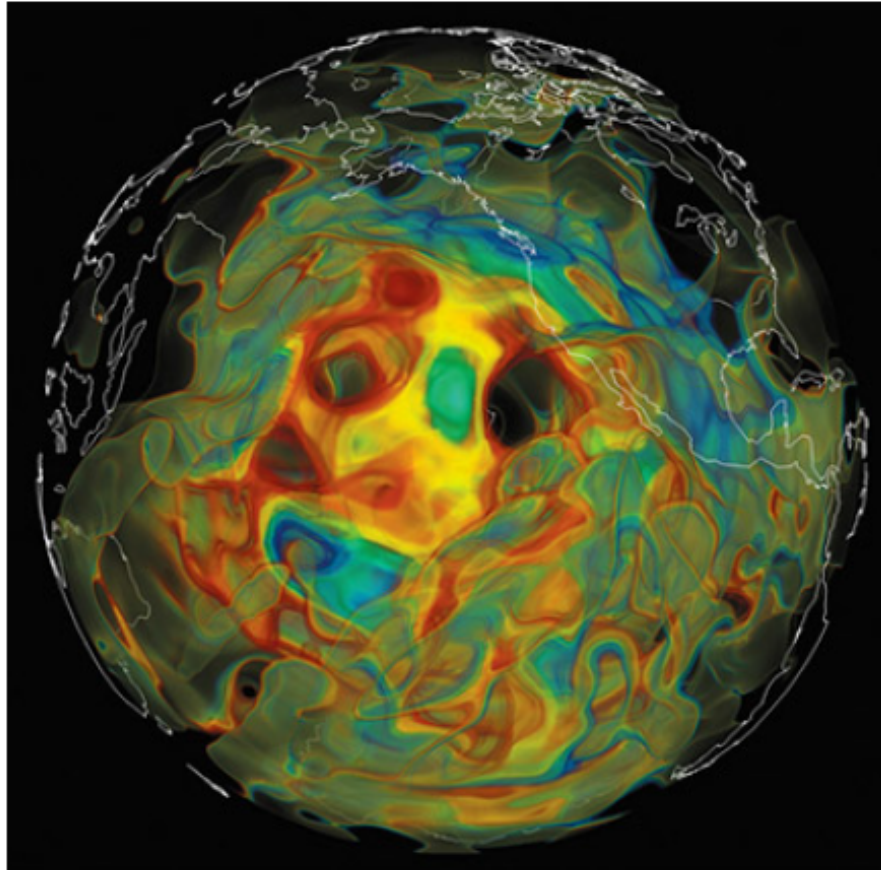
Outlook

Waiting for Km3Net, we made contact with the **Antares Collaboration** to include their data into our analysis (still to be done)

After ARCA is completed, we will look at the Earth's interior from both emispheres (**test of anisotropies**)

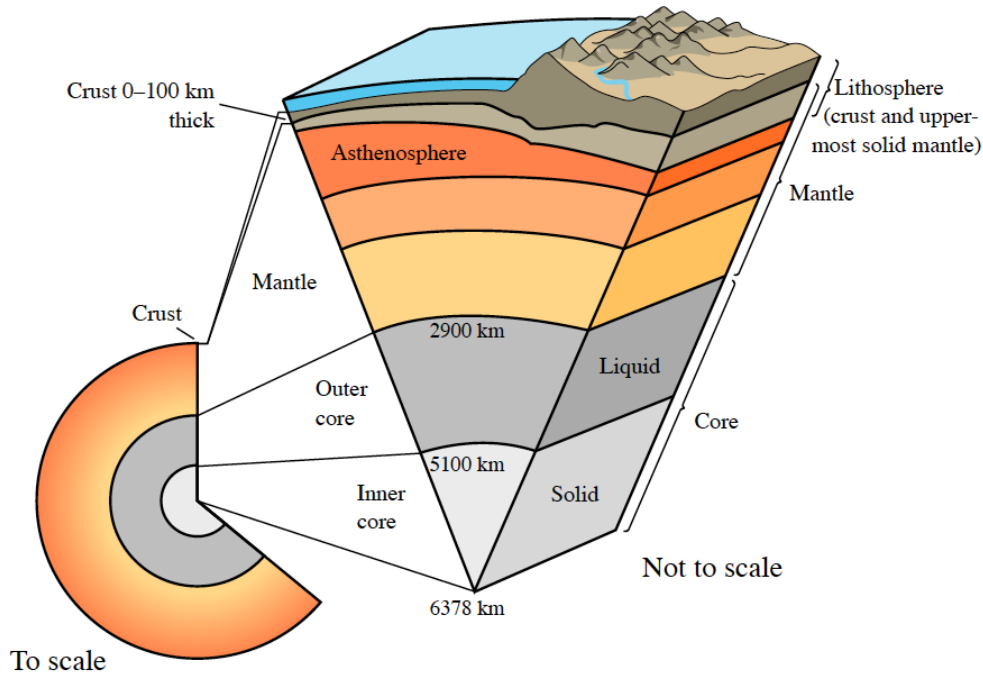
Together with **medical imaging techniques** (started discussions with experts in the field) we could perform **EARTH'S THREE-DIMENSIONAL ν -TOMOGRAPHY**

State-of-art three-dimensional picture from seismography



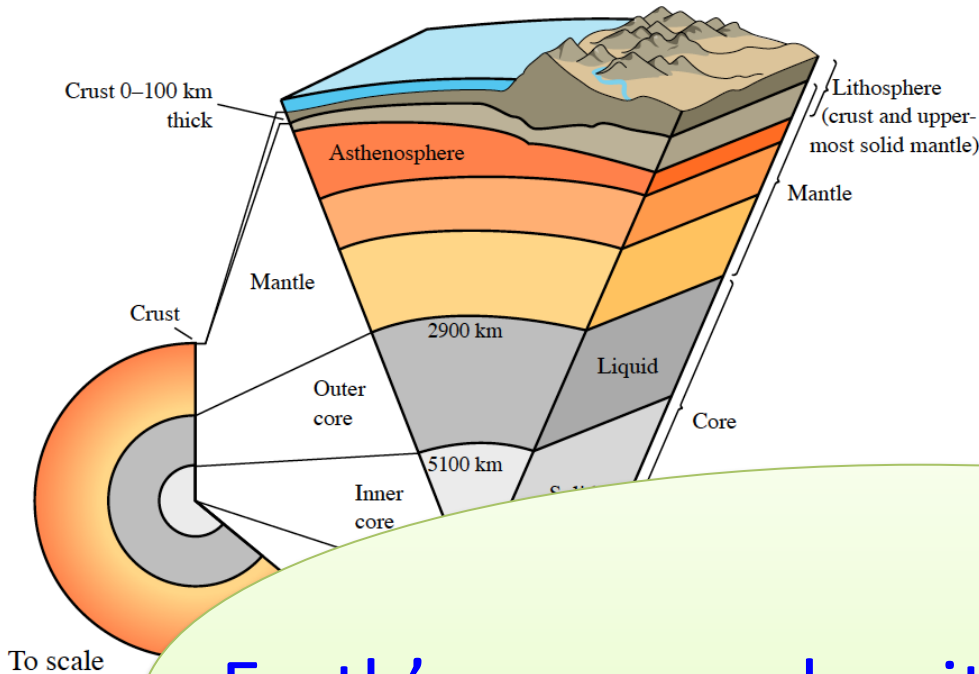
Backup slides

Density at different depths



Depth ^[103] km	Component layer	Density g/cm ³
0-60	Lithosphere ^[n 14]	—
0-35	Crust ^[n 15]	2.2-2.9
35-60	Upper mantle	3.4-4.4
35-2890	Mantle	3.4-5.6
100-700	Asthenosphere	—
2890-5100	Outer core	9.9-12.2
5100-6378	Inner core	12.8-13.1

Density at different depths



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35-2890	Mantle	3.4-5.6
	Asthenosphere	—
		9.9-12.2
		8-13.1

Earth's average density: $\rho = 5.5148 \text{ g/cm}^3$
(granite density is 2.7 g/cm^3)

How densities are measured?

seismology

propagation of earthquake waves
through the Earth: p-waves and s-waves
(v_p and v_s)
composition dependence!

How densities are measured?

seismology

propagation of earthquake waves
through the Earth: p-waves and s-waves
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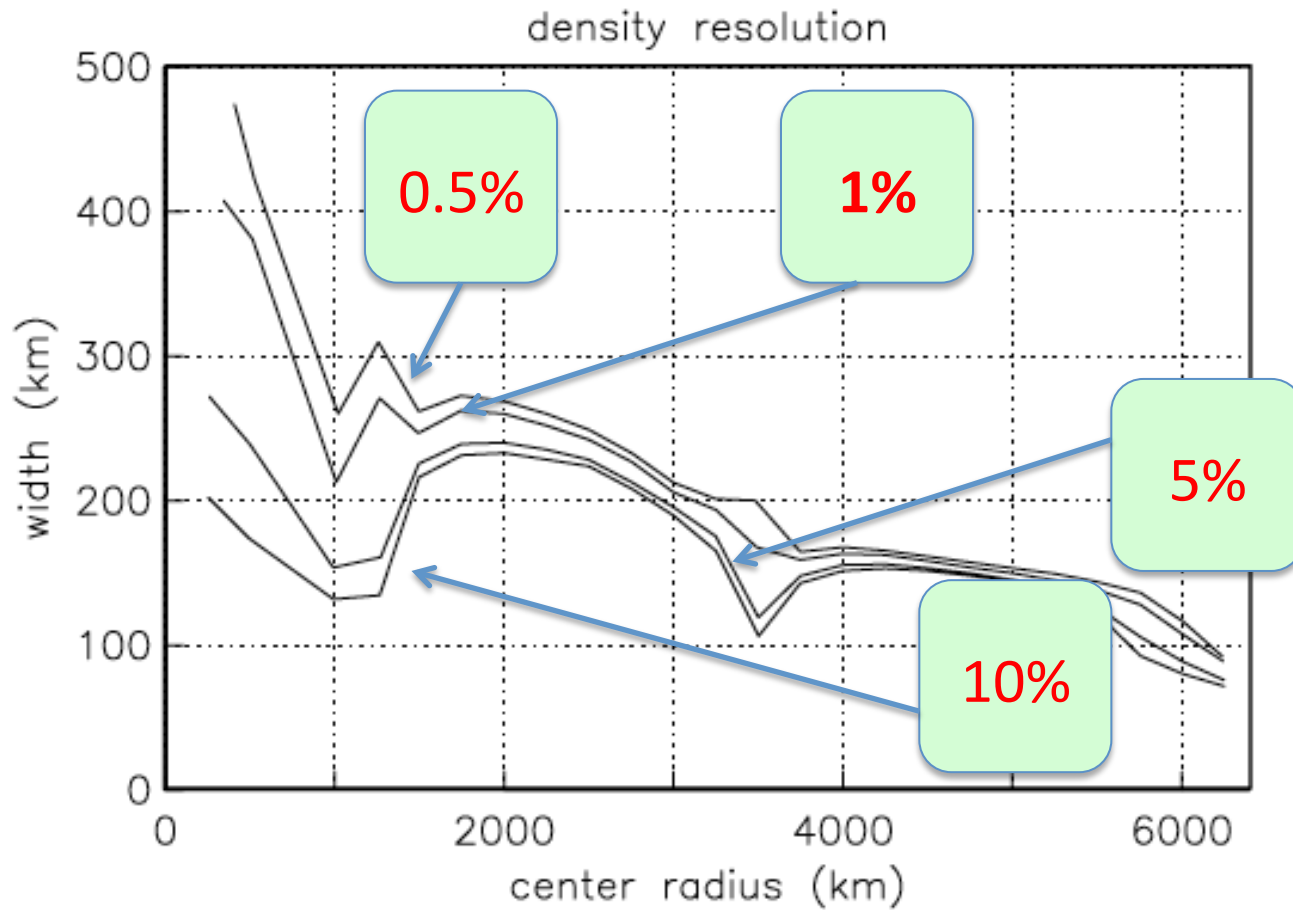
Adams-Williamson equation (1924)

$$\frac{d\rho}{dr} = -\rho(r) \frac{g(r)}{\Phi(r)}$$

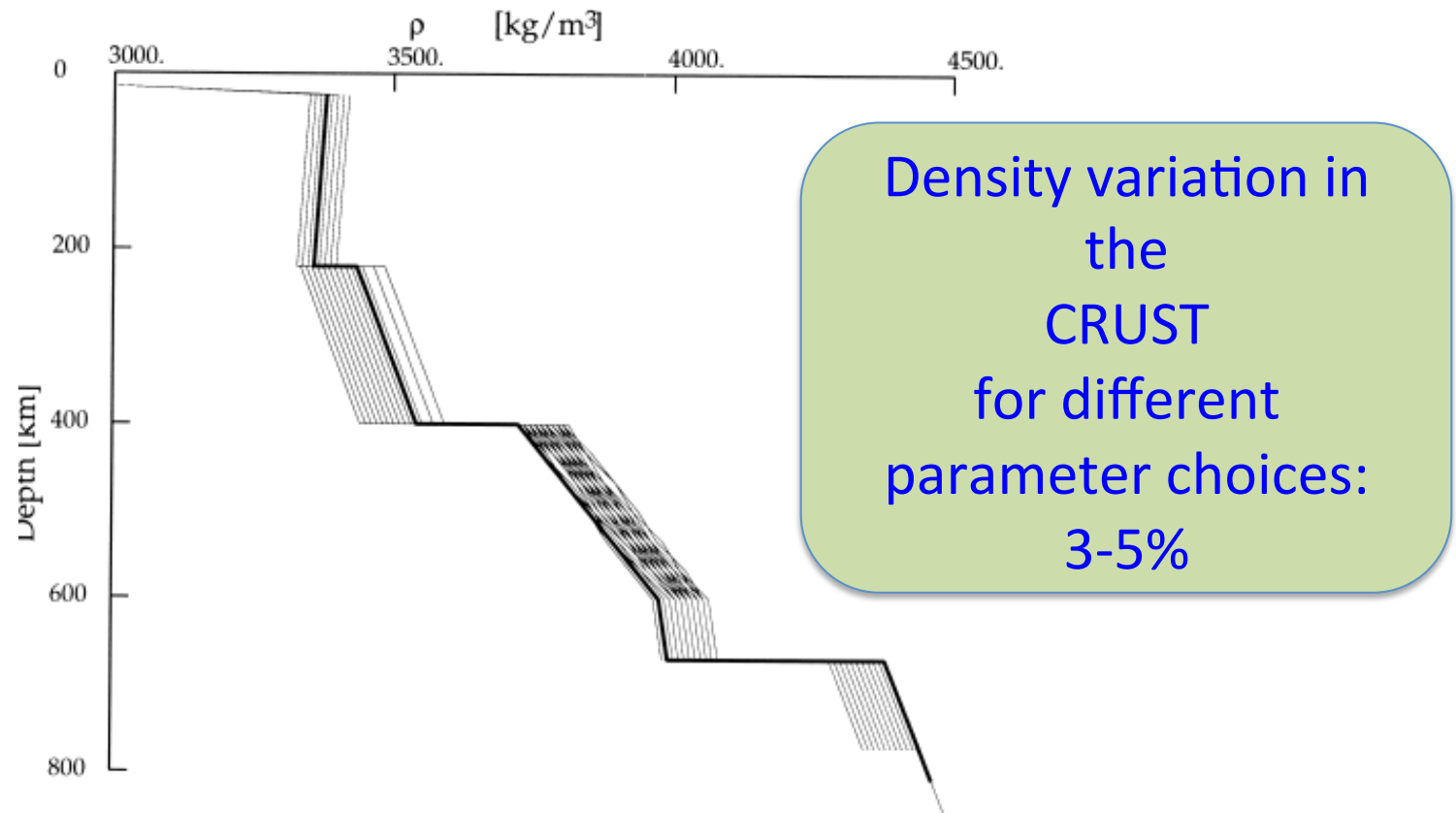
$$\Phi(r) = v_p^2 - \frac{4}{3}v_s^2$$

Composition dependence! Gravitational profile dependence!

Uncertainties in the core density



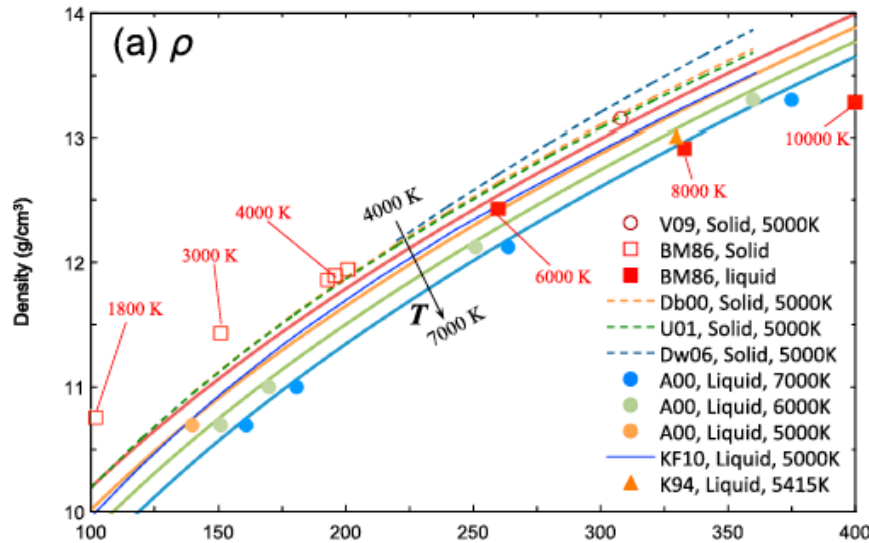
Model dependence of the profile...



[Kennett, Geophysical Journal International, 132 (1998)]

Inner core uncertainties

Strong dependence of the IC density on temperature, pressure and composition



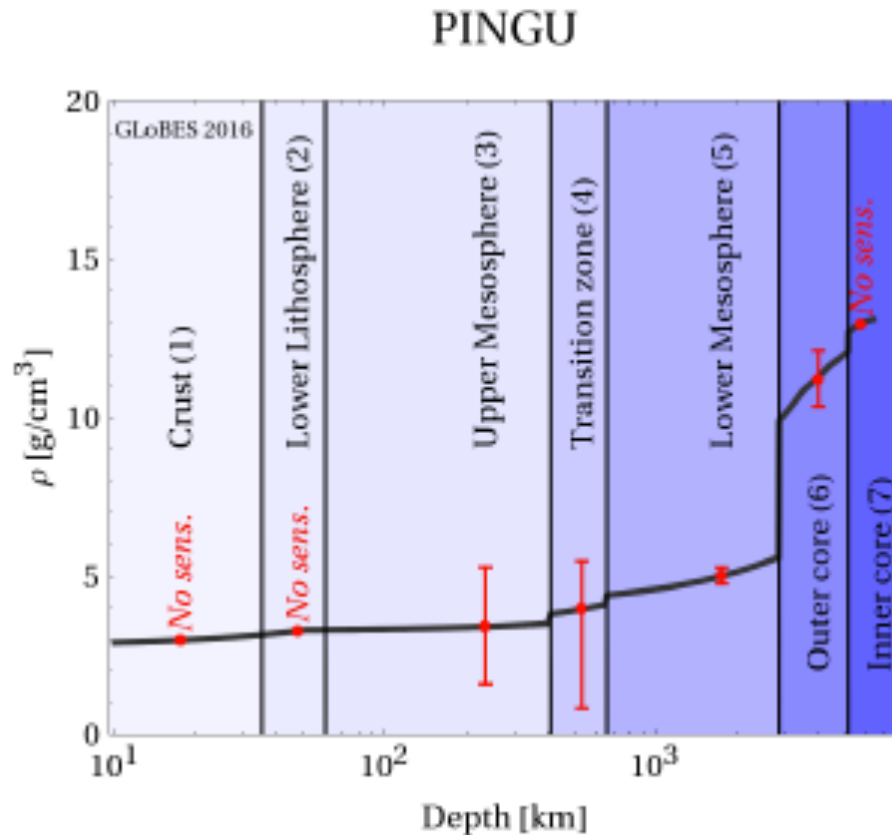
Estimated temperature range still very large: 4000-10000 K

Composition guessed (iron-nickel?)

Missing Xenon problem

Ishikawa, Tsuchiya, Tange, J. GeoPhys. Res. (Solid Earth) 119 (2014)

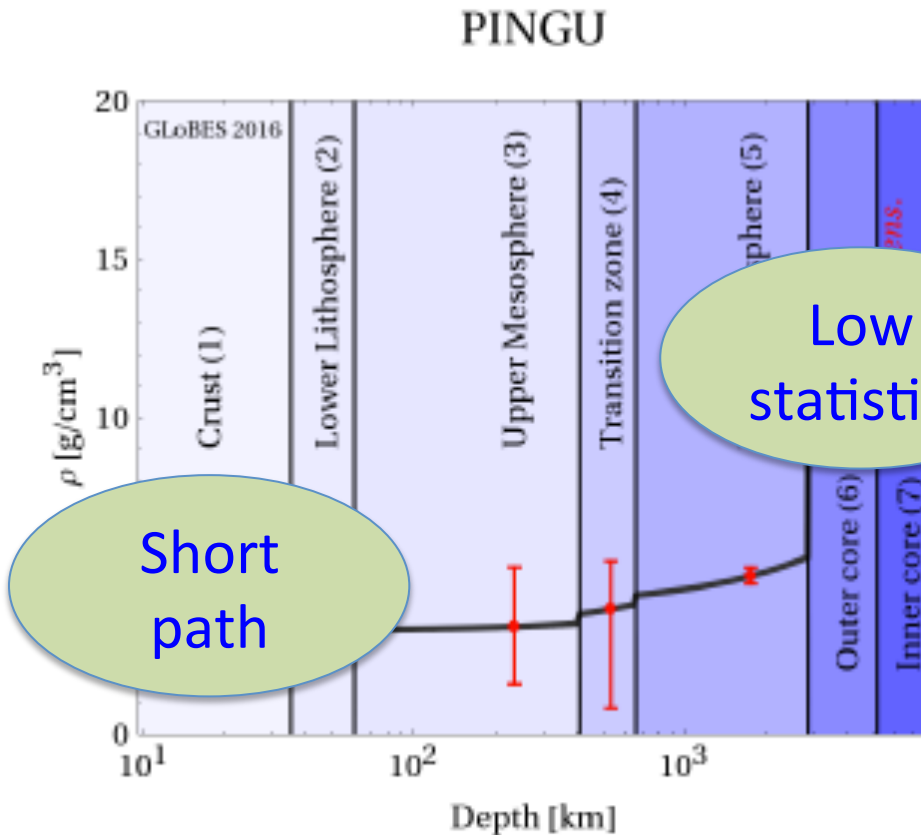
“Recent” forecasts, 1



After 10 years of data taking
at PINGU or ORCA using
neutrino oscillations

Winter, Nucl. Phys B908 (2016)

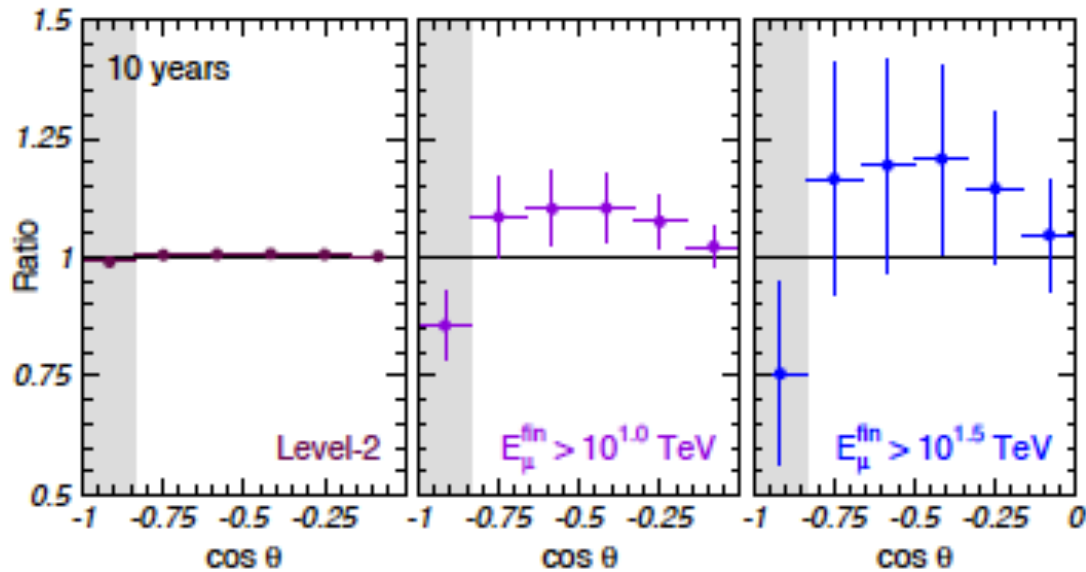
“Recent” forecasts, 1



After 10 years of data taking at PINGU or ORCA using neutrino oscillations

Winter, Nucl. Phys B908 (2016)

“Recent” forecasts, 2

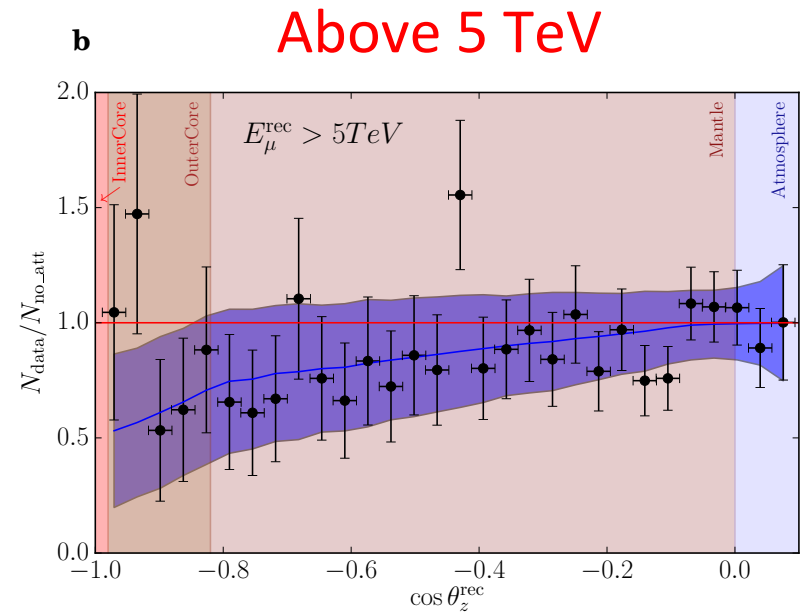
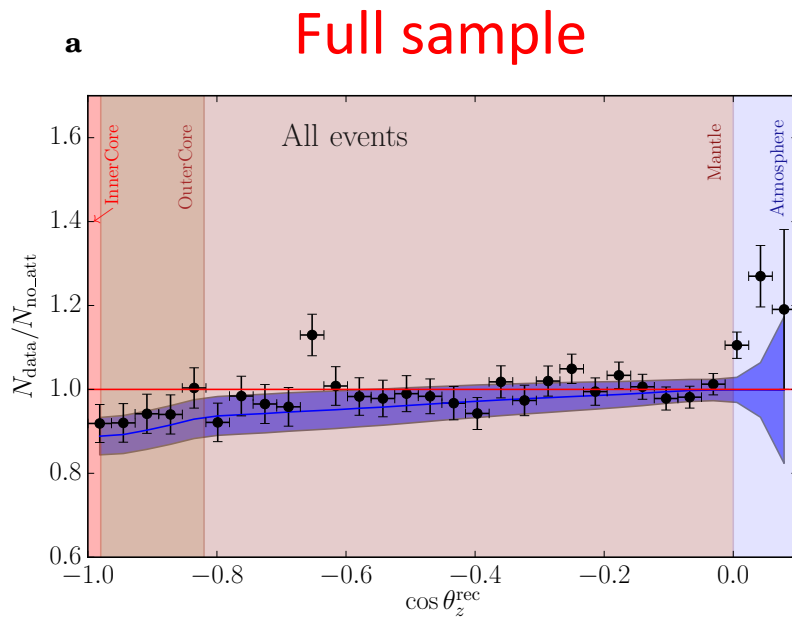


After 10 years of
data taking
at IceCube using
neutrino attenuation

Claim: IceCube could reject a homogeneous Earth at 5σ in ten years

Gonzalez-García, Halzen, Maltoni, Tanaka, Phys. Rev. Lett. 100 (2008)

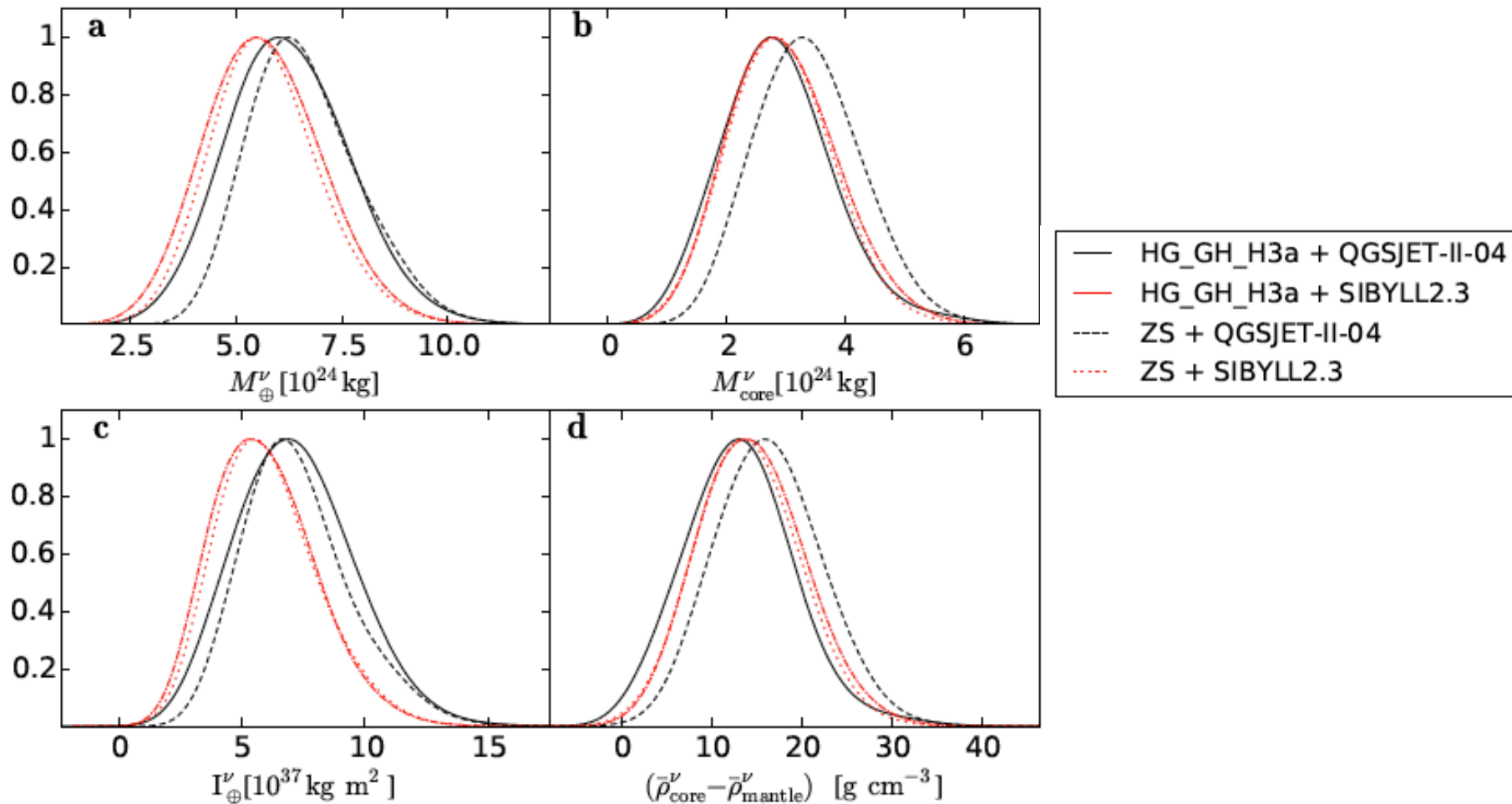
Including energy cuts: $N_{\text{data}}/N_{\text{noatt}}$



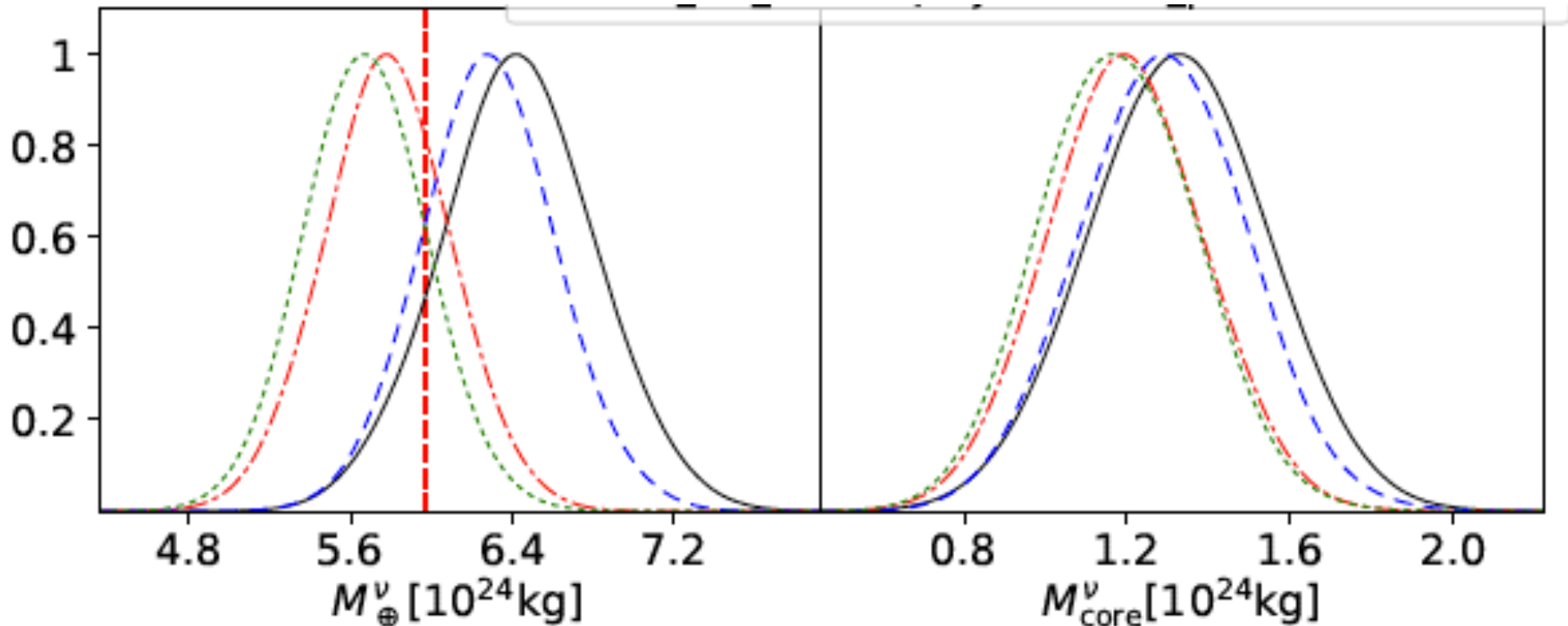
Full sample
useful for
normalization

For $\cos \theta > -0.6$,
attenuation can be
as large as 50%

Flux and hadronic model dependence



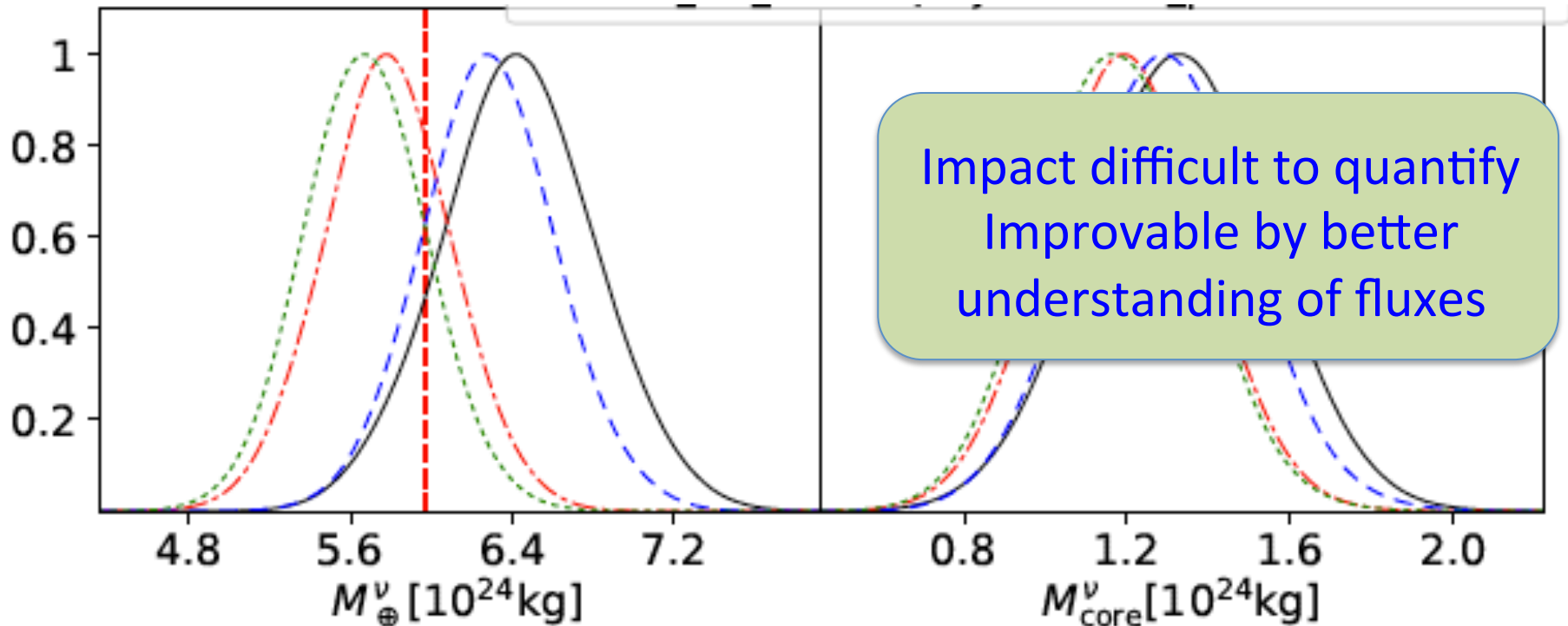
Flux and hadronic model dependence, 2



- GH_HG_H3a + QGSJET -> GH_HG_H3a + QGSJET
- - - GH_HG_H3a + QGSJET -> GH_HG_H3a + SIBYLL2.3
- - - GH_HG_H3a + QGSJET -> ZS_pamela + QGSJET
- - - GH_HG_H3a + QGSJET -> ZSa_pamela + SIBYLL2.3

Forecast for
10 years of MC data

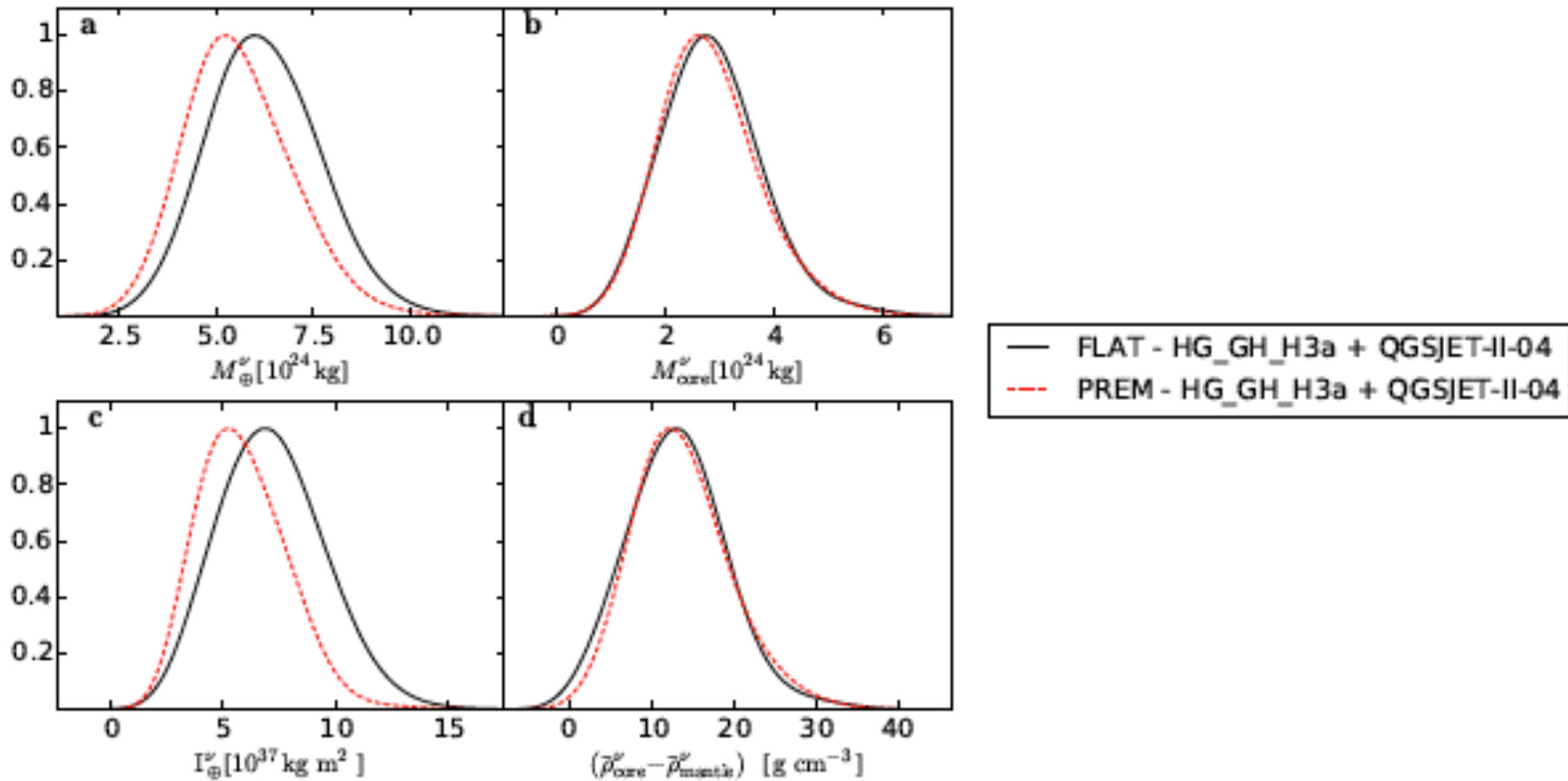
Flux and hadronic model dependence, 2



- GH_HG_H3a + QGSJET -> GH_HG_H3a + QGSJET
- - - GH_HG_H3a + QGSJET -> GH_HG_H3a + SIBYLL2.3
- - - GH_HG_H3a + QGSJET -> ZS_pamela + QGSJET
- - - GH_HG_H3a + QGSJET -> ZSa_pamela + SIBYLL2.3

Forecast for
10 years of MC data

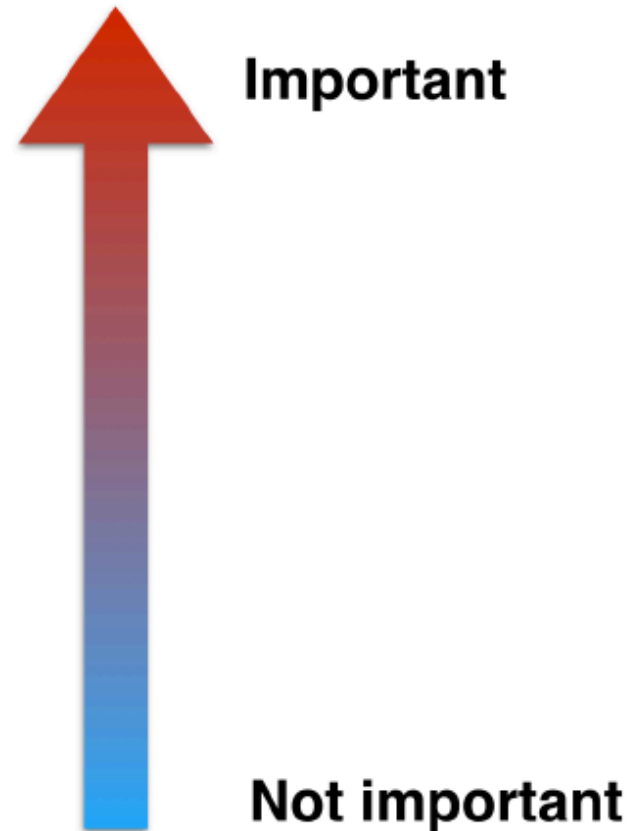
Earth's profile dependence



Systematics importance

- ▶ DOM efficiency
- ▶ Flux continuous parameters
 - ▶ spectral index
 - ▶ π/K ratio
 - ▶ $\nu/\bar{\nu}$ ratio Full Implementation
- ▶ Air shower hadronic models Marginally irrelevant precise check
- ▶ Primary cosmic ray fluxes Marginally irrelevant precise check
- ▶ Hole Ice Irrelevant
- ▶ Neutrino cross sections Irrelevant
- ▶ Bulk ice scattering/absorption Irrelevant

continuous systematics
discrete systematic



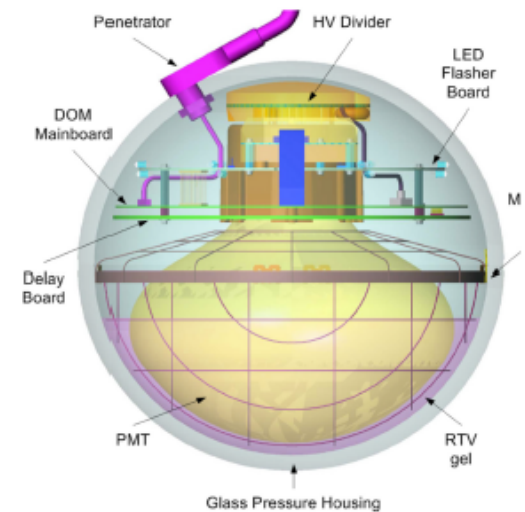
Systematics importance

- ▶ DOM efficiency
- ▶ Flux continuous parameters
 - ▶ spectral index
 - ▶ π/K ratio
 - ▶ $\nu/\bar{\nu}$ ratio Full Implementation
- ▶ Air shower hadronic models Marginally irrelevant precise check
- ▶ Primary cosmic ray fluxes Marginally irrelevant precise check
- ▶ Hole Ice Irrelevant
- ▶ Neutrino cross sections Irrelevant
- ▶ Bulk ice scattering/absorption Irrelevant

continuous systematics
discrete systematic



Important



D.O.M.

Not important

Systematics importance

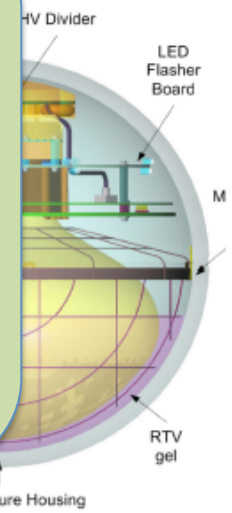
(as in Salvado talk on sterile neutrinos)

- ▶ DOM efficiency
- ▶ Flux continuous parameters

- ▶ S
- ▶
- ▶

- ▶ Air sh
irrelev
- ▶ Prima
irrelev
- ▶ Hole
- ▶ Neutr
- ▶ Bulk ic
Irrelevant

Not taken into account:
OPTICAL PROPERTIES OF THE ICE

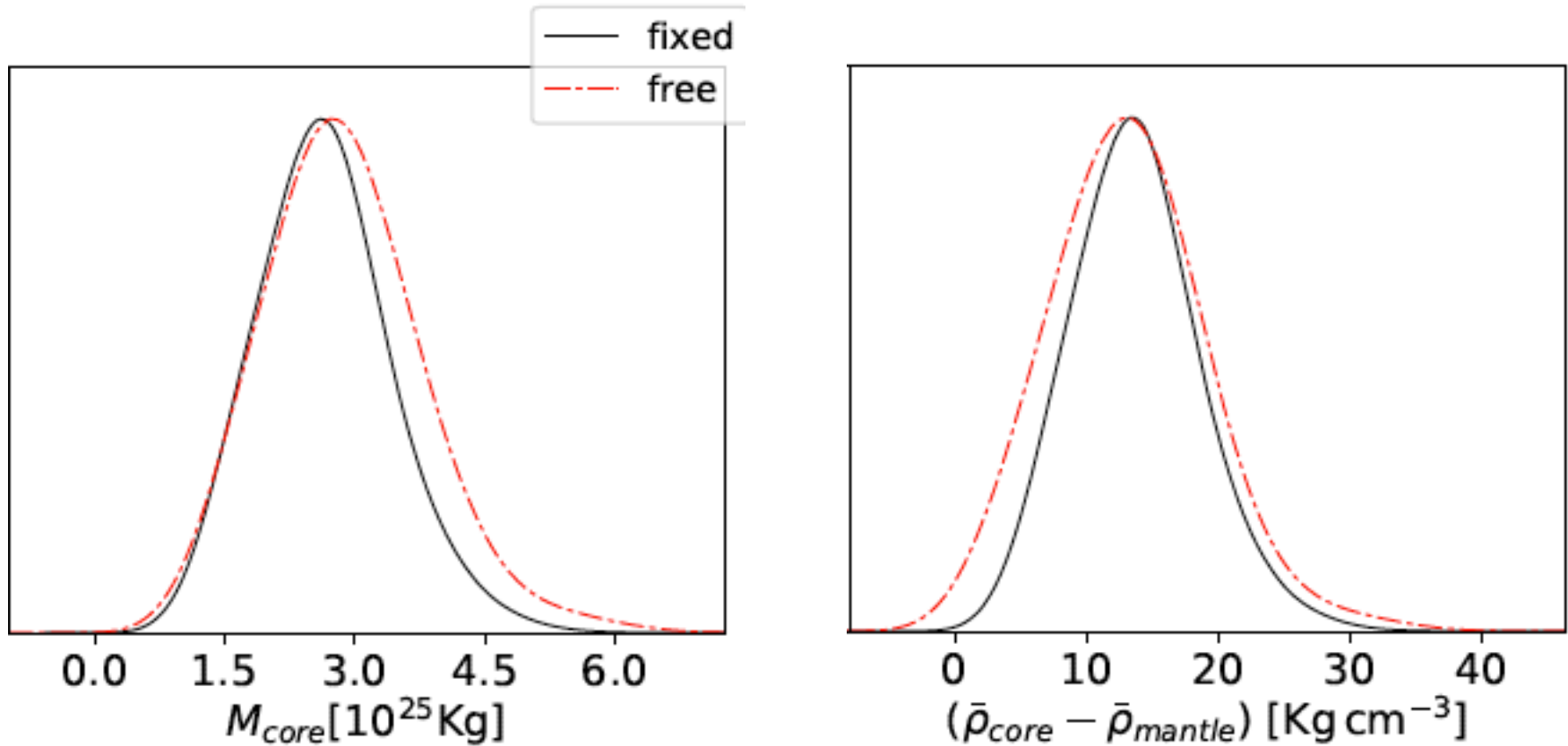


D.O.M.

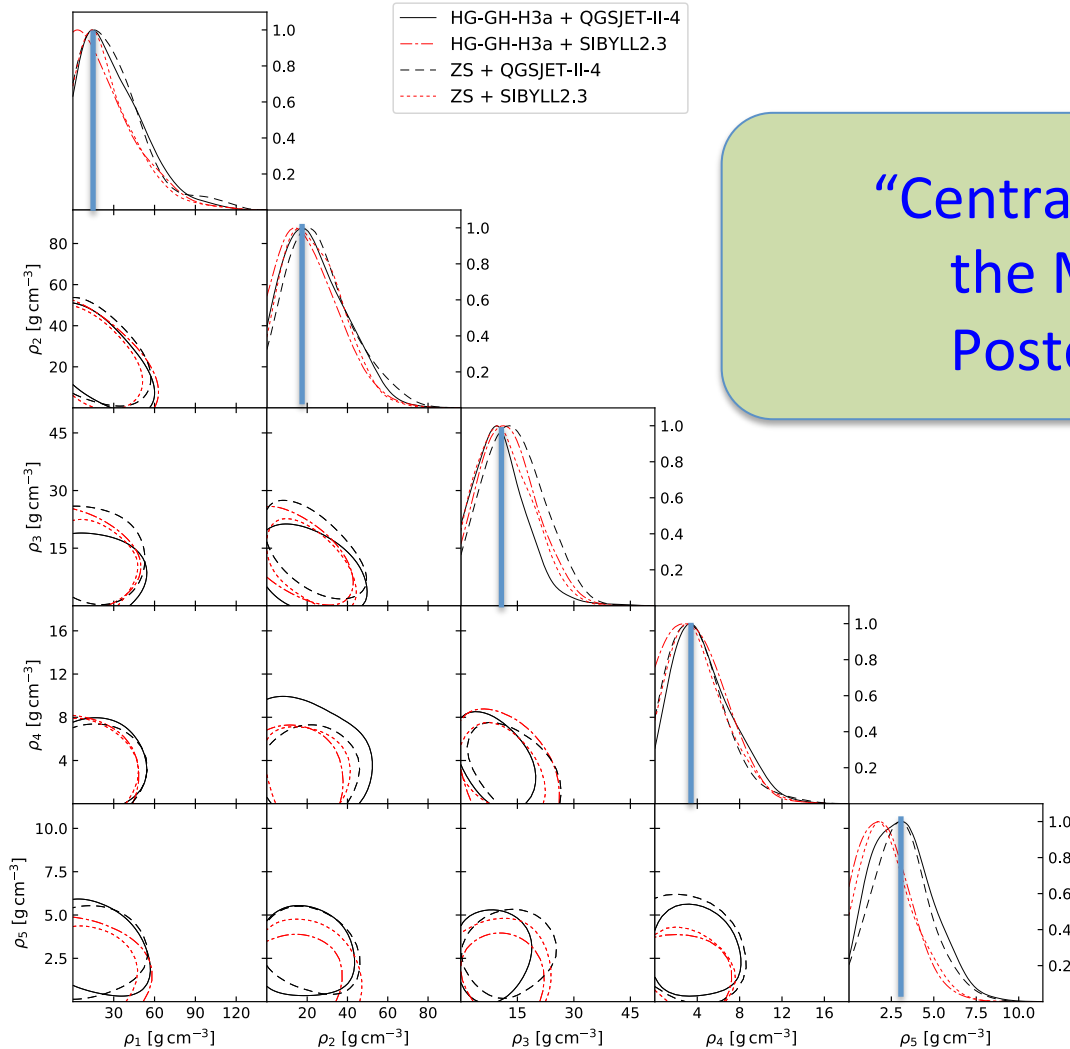
Not important

continuous systematics
discrete systematic

Impact of systematics on the error

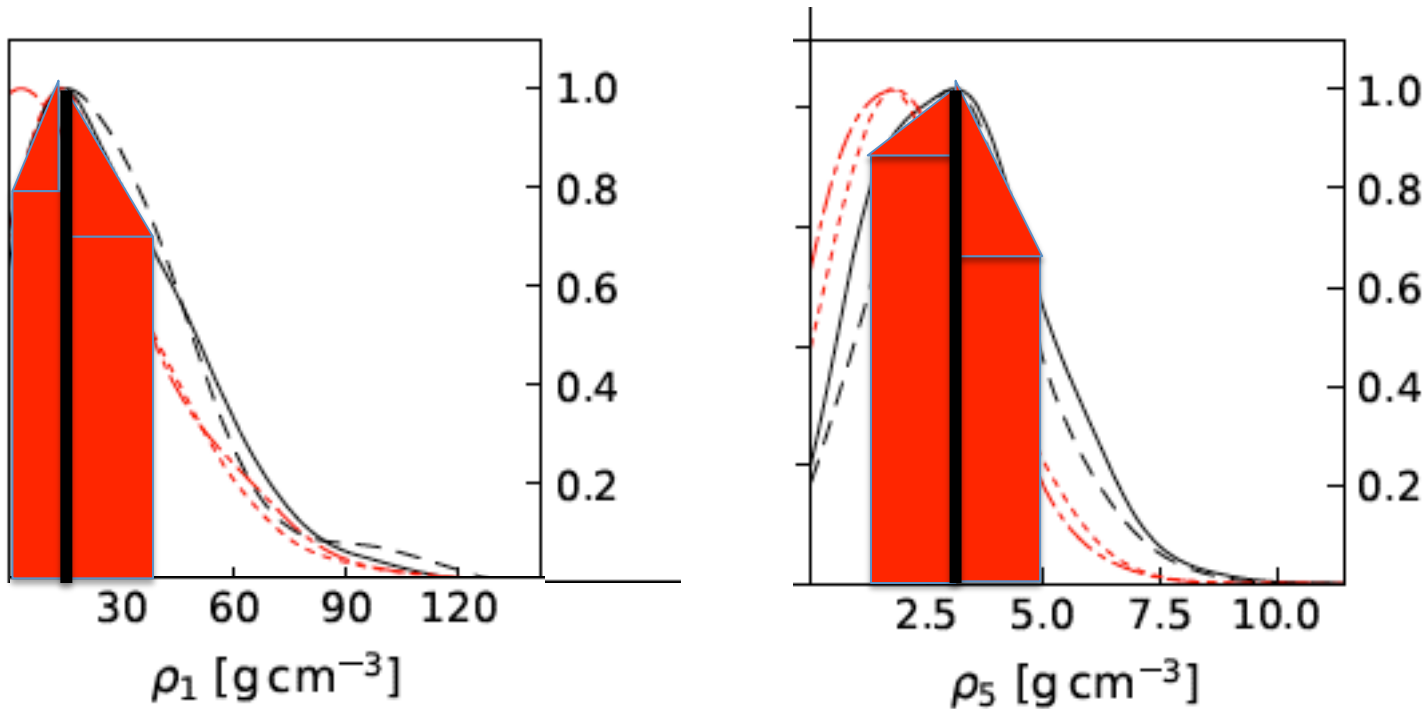


What are the dots?

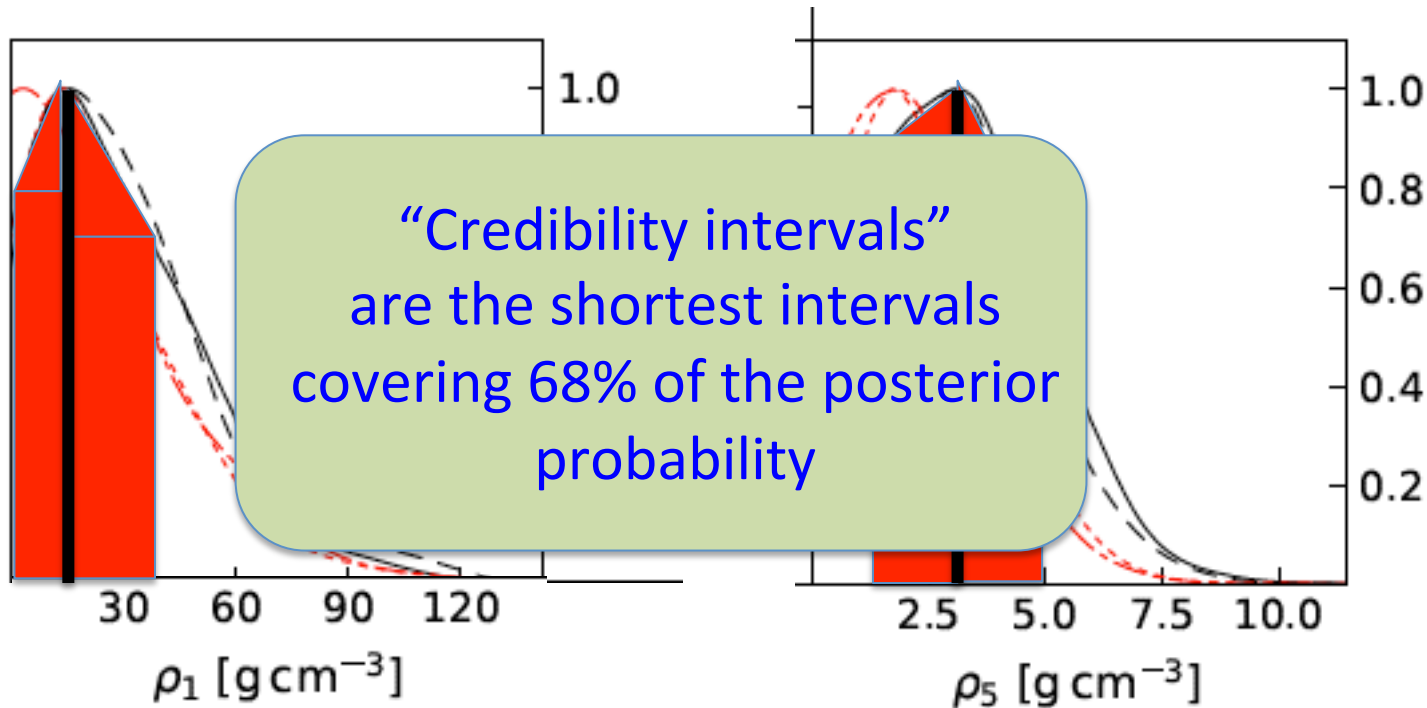


“Central values” represent the Maximum of the Posterior Probability

How to get asymmetric (bayesian) credibility intervals?



How to get asymmetric (bayesian) intervals?



Results for different models

	Piecewise flat Earth's profile				PREM Earth's profile
	HG-GH-H3a + QGSJET-II-04	HG-GH-H3a + SIBYLL2.3	ZS + QGSJET-II-04	ZS + SIBYLL2.3	HG-GH-H3a + QGSJET-II-04
M_{\oplus}^{ν} [10^{24} kg]	$6.0^{+1.6}_{-1.3}$	$5.5^{+1.5}_{-1.3}$	$6.2^{+1.4}_{-1.2}$	$5.5^{+1.3}_{-1.2}$	$5.3^{+1.5}_{-1.3}$
M_{core}^{ν} [10^{24} kg]	$2.72^{+0.97}_{-0.89}$	$2.79^{+0.98}_{-0.85}$	$3.27^{+0.92}_{-0.89}$	$2.84^{+0.89}_{-0.88}$	$2.62^{+0.97}_{-0.84}$
I_{\oplus}^{ν} [10^{37} kg cm ²]	6.9 ± 2.4	$5.4^{+2.3}_{-1.9}$	$6.7^{+2.3}_{-2.0}$	$5.5^{+2.2}_{-1.9}$	$5.3^{+2.3}_{-1.7}$
$\bar{\rho}_{\text{core}}^{\nu} - \bar{\rho}_{\text{mantle}}^{\nu}$ [g/cm ³]	$13.1^{+5.8}_{-6.3}$	$14.0^{+6.0}_{-5.9}$	$15.9^{+6.0}_{-5.9}$	$13.5^{+6.1}_{-5.5}$	$12.3^{+6.3}_{-5.4}$
p - value mantle denser than core	1.1×10^{-2}	2.4×10^{-3}	9.4×10^{-4}	4.6×10^{-3}	3.8×10^{-3}

Constraints and derived quantities

- Gravitational measurement of the Earth's mass

$$M_{\oplus} = \frac{4\pi}{3} \int_0^{R_{\oplus}} dr r^2 \rho(r) = 5.972 \times 10^{24} \text{kg}$$

- A derived quantity: Earth's mean moment of inertia

$$I_{\oplus} = \frac{8\pi}{3} \int_0^{R_{\oplus}} dr r^4 \rho(r) = 0.3307144 M_{\oplus} R_{\oplus}^2$$

A constant density would give $I_{\oplus}(\rho(r) = \rho_0) = 0.4 M_{\oplus} R_{\oplus}^2$.

Impact of constraints

Adding the gravitational Earth's mass as an external constraint, results in fixing the mantle density:

$$\rho_5 = [1.22-4.78] \text{ g/cm}^3 \rightarrow [4.43-4.79] \text{ g/cm}^3$$

Rather small impact on the core density, instead:

$$\rho_{\text{core}} = [10.2-20.8] \text{ g/cm}^3 \rightarrow [9.7-18.6] \text{ g/cm}^3$$

Tomography

noun to·mog·ra·phy \ tō-'mä-grə-fē \

imaging by sections or sectioning, through the use of any kind of penetrating wave ([Wikipedia](#))

a method of producing a three-dimensional image of the internal structures of a solid object (such as the [human body](#) or the [earth](#)) by the observation and recording of the differences in the effects on the passage of waves of energy impinging on those structures ([Merriam-Webster](#))