

# Atlantic weather regimes and poleward heat transport by transient eddies

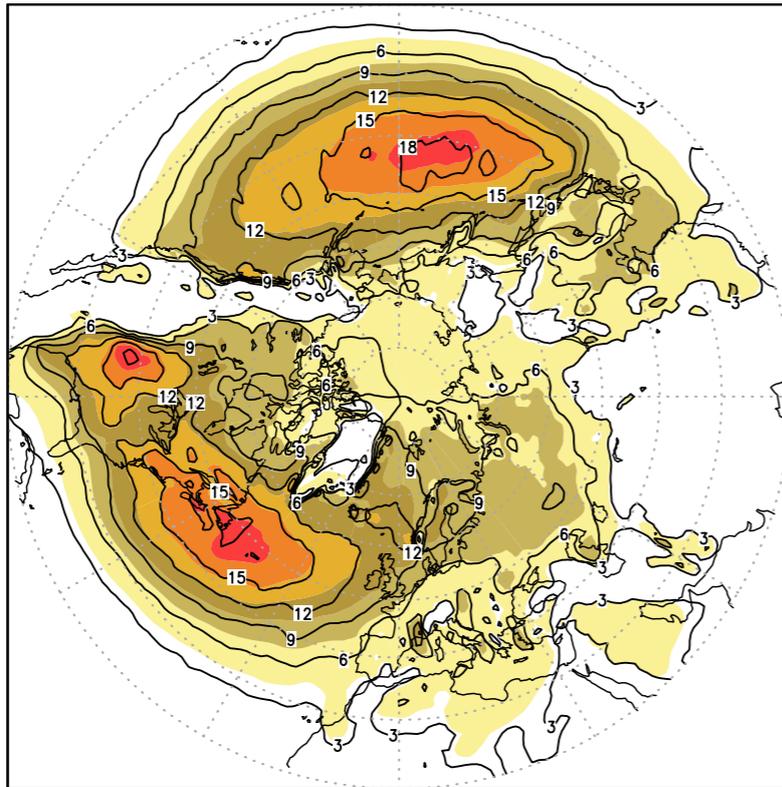
Alessio Bellucci, Paolo Ruggieri, Carmen Alvarez-Castro,  
Panos Athanasiadis, Stefano Materia, Silvio Gualdi

Centro Euro-Mediterraneo sui Cambiamenti Climatici  
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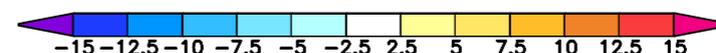
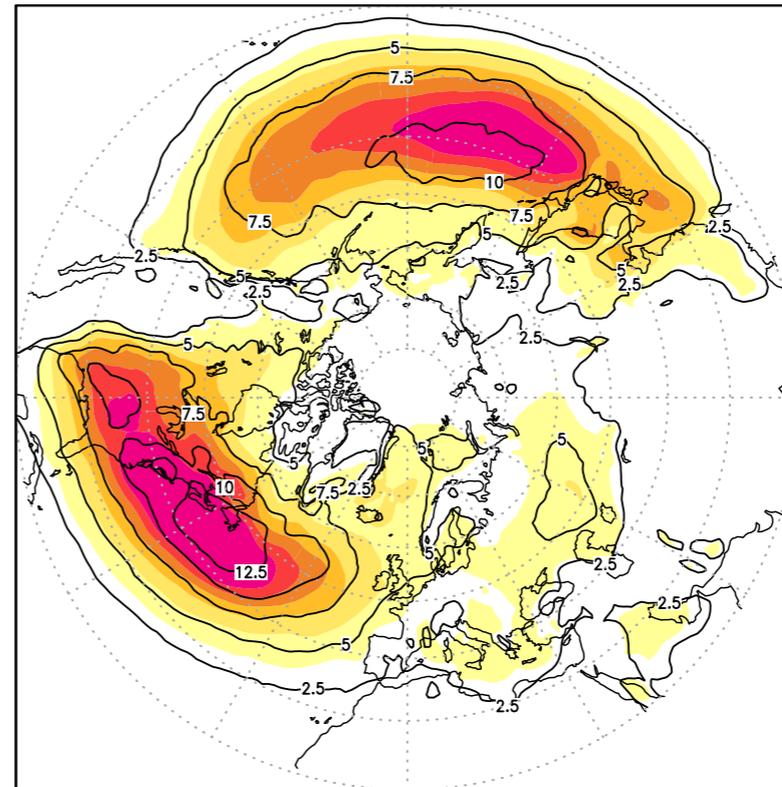
# Overview

## Synoptic

$V^2$  ( $m^2/s^2$ ) mean (shading) and st. dev.



$mV/c_p$  (K m/s) mean (shading) and st. dev.



We investigate links between transient, poleward atmospheric heat transport and the atmospheric circulation in the North Atlantic sector.

We look at the Northern Hemisphere but special emphasis is given to heat transport crossing the 70 °N wall.

Results are based on an atmospheric reanalysis, extended cold season.

# Introduction

- 1) Heat transport performed by eddies is fundamental, transient eddies are a key (if not the dominant) component of AHT in the polar cap and its surroundings (Overland et al. 1990, Miletta et al. 2000).
- 2) Storminess changes accordingly with the low-frequency, large-scale flow of the atmosphere. The variability of the storm track is largely ascribable to the alternating phases of the NAO/AO, but no clear evidence is found of any modulation (by the NAO) of high latitude transient eddy heat transport. (Miletta et al. 2000)
- 3) There is emerging and growing evidence of impulsive and regional injections of heat and moisture into the Arctic. These phenomenon is deemed to be linked with the large scale circulation of the atmosphere and with poleward propagation of storms. (Messori and Czaja 2013, Woods et al. 2013)

# Data

Data used are obtained from ERA-Interim

6-hourly on a  $1^{\circ} \times 1^{\circ}$  regular longitude-latitude grid

10 selected pressure levels (between 100 and 925 hPa)

Period: 1980-2017, extended cold season (NDJFM)

# Methodology

Moist static energy:

$$m = c_p T + L_v Q + \Phi$$

$c_p$  is specific heat of dry air,  $L_v$  latent heat of vaporisation

Transient eddy heat flux is defined as the product  $\mathbf{V}'\mathbf{m}'$

Prime denotes a bandpass filter

Synoptic 2-9 days

Intra-seasonal 10-90 days

Storminess is defined as  $\mathbf{V}'\mathbf{V}'$

A jet latitude index is computed following Woollings et al. (2010)

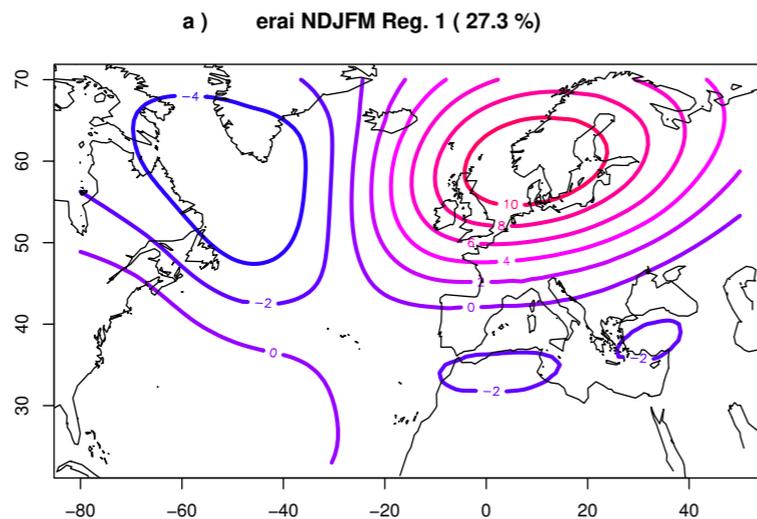
# Methodology: Weather regimes

Weather regimes have been computed following the methods of Michelangeli et al. (1995) and Yiou et al. (2008).

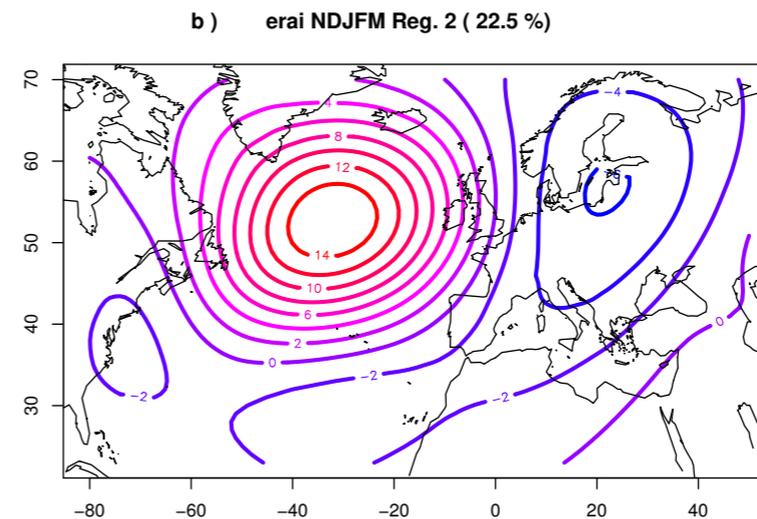
The first ten Empirical Orthogonal Functions (EOFs) of Z at 500 hPa. A k-means algorithm is applied, 4 weather regimes, over the North Atlantic region [80°W - 50°E; 20 - 70°N] on daily data over the period 1980-2017, in NDJFM.

Daily data classifications are obtained by the minimum of the Euclidean distances to the centroids

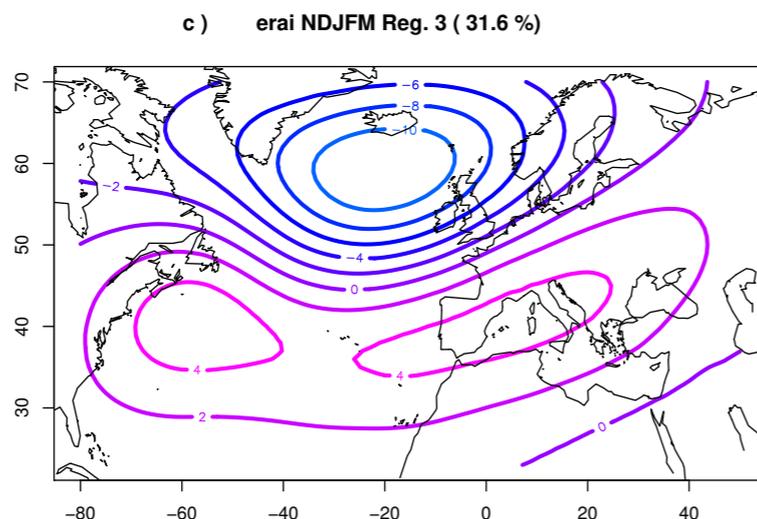
SCAND



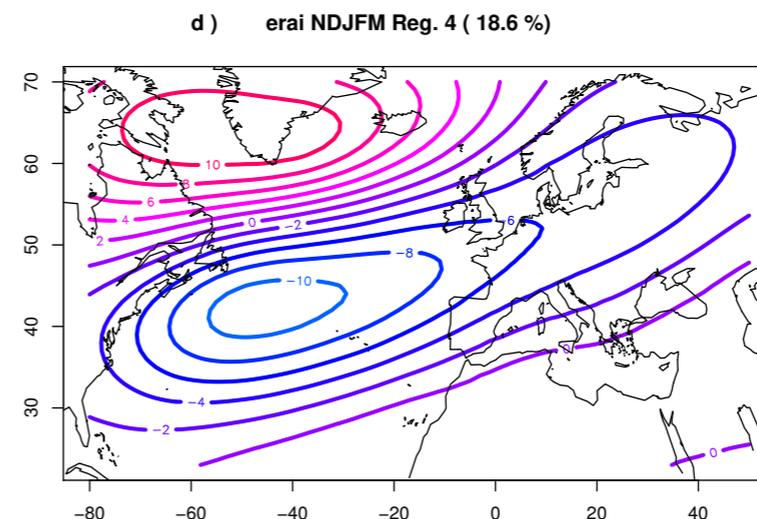
RIDGE



NAO+



NAO-

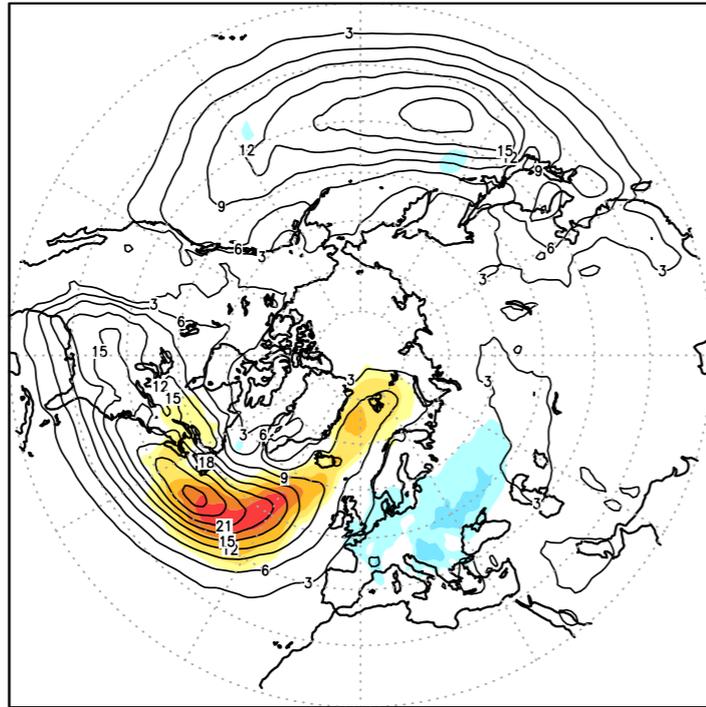


-12 -8 -4 0 4 8 12 NDJFM

# Synoptic eddies heat transport

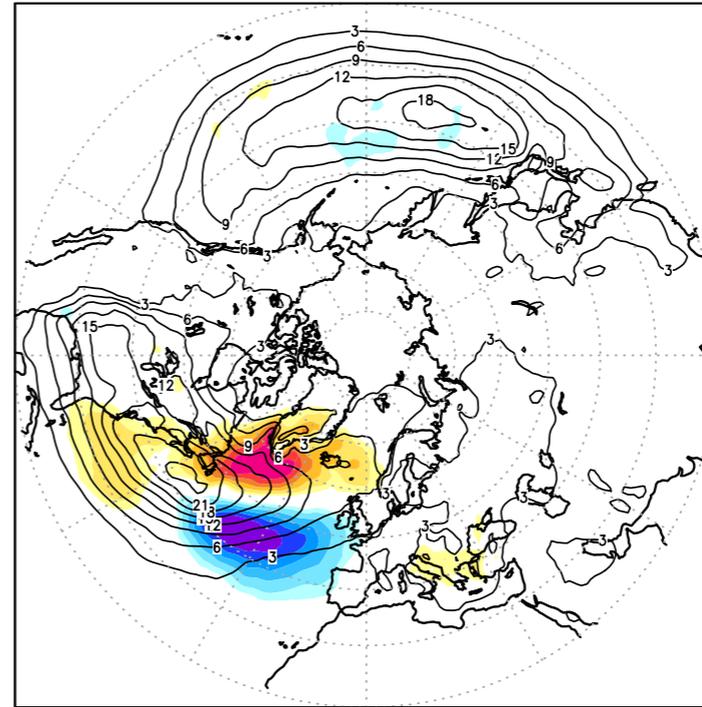
WR=SCAND  $m\sqrt{V}/c_p$  NDJFM 1980-2017 (K m/s)

SCAND



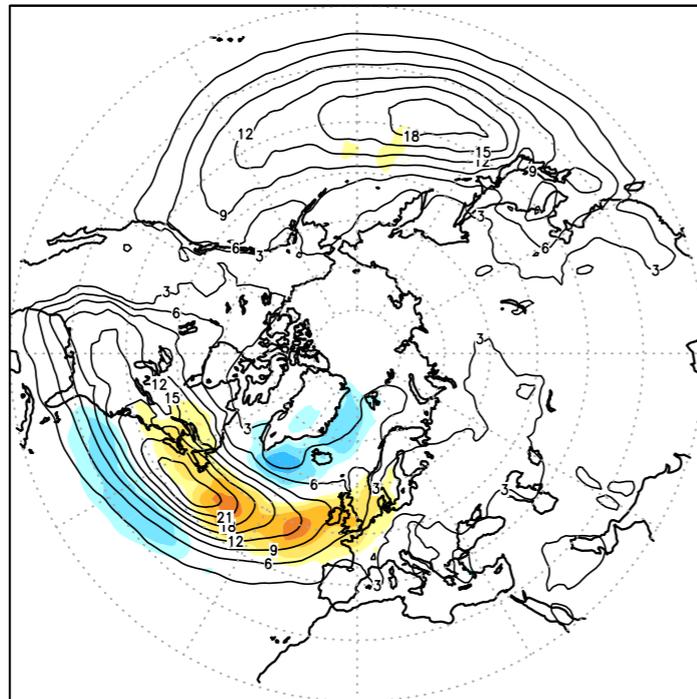
WR=RIDGE  $m\sqrt{V}/c_p$  NDJFM 1980-2017 (K m/s)

RIDGE



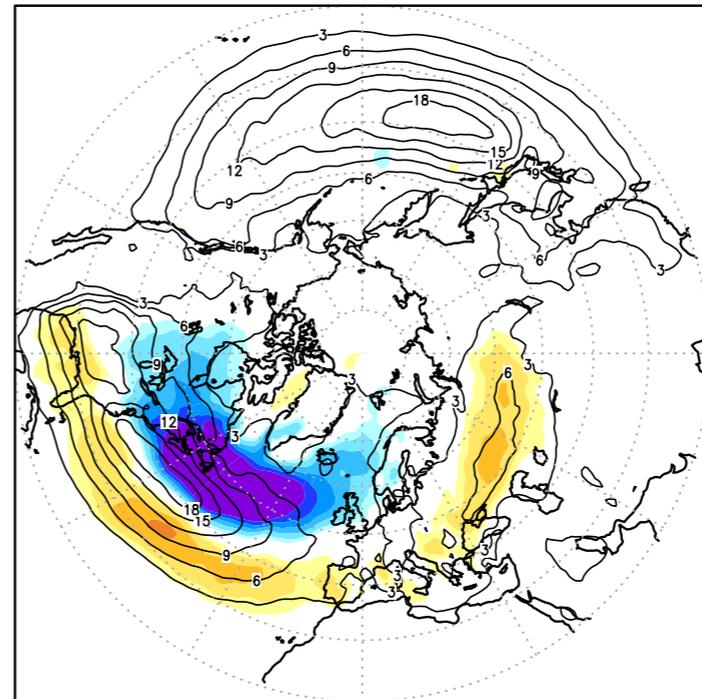
WR=NAO+  $m\sqrt{V}/c_p$  NDJFM 1980-2017 (K m/s)

NAO+



WR=NAO-  $m\sqrt{V}/c_p$  NDJFM 1980-2017 (K m/s)

NAO-



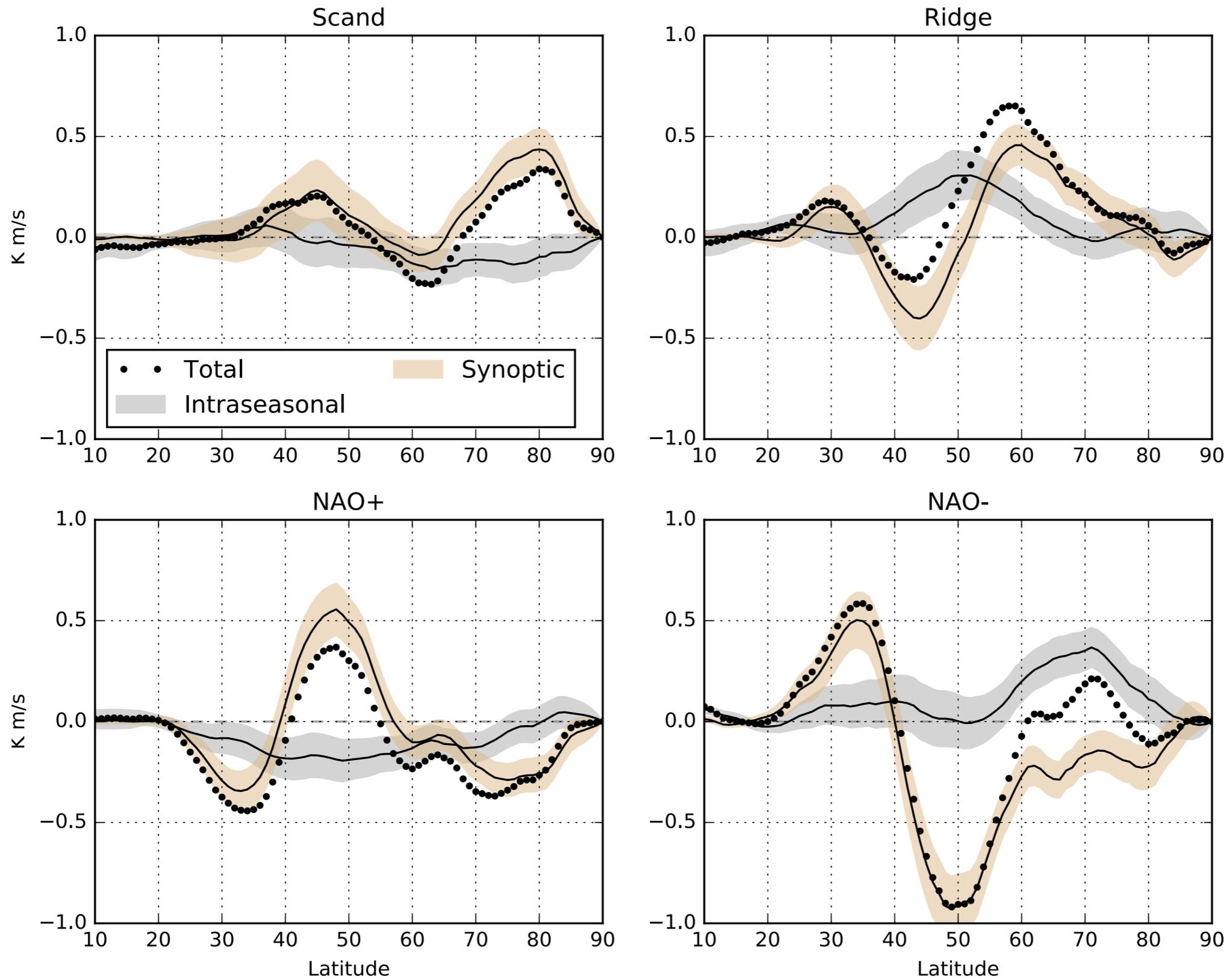
Anomaly



Full field

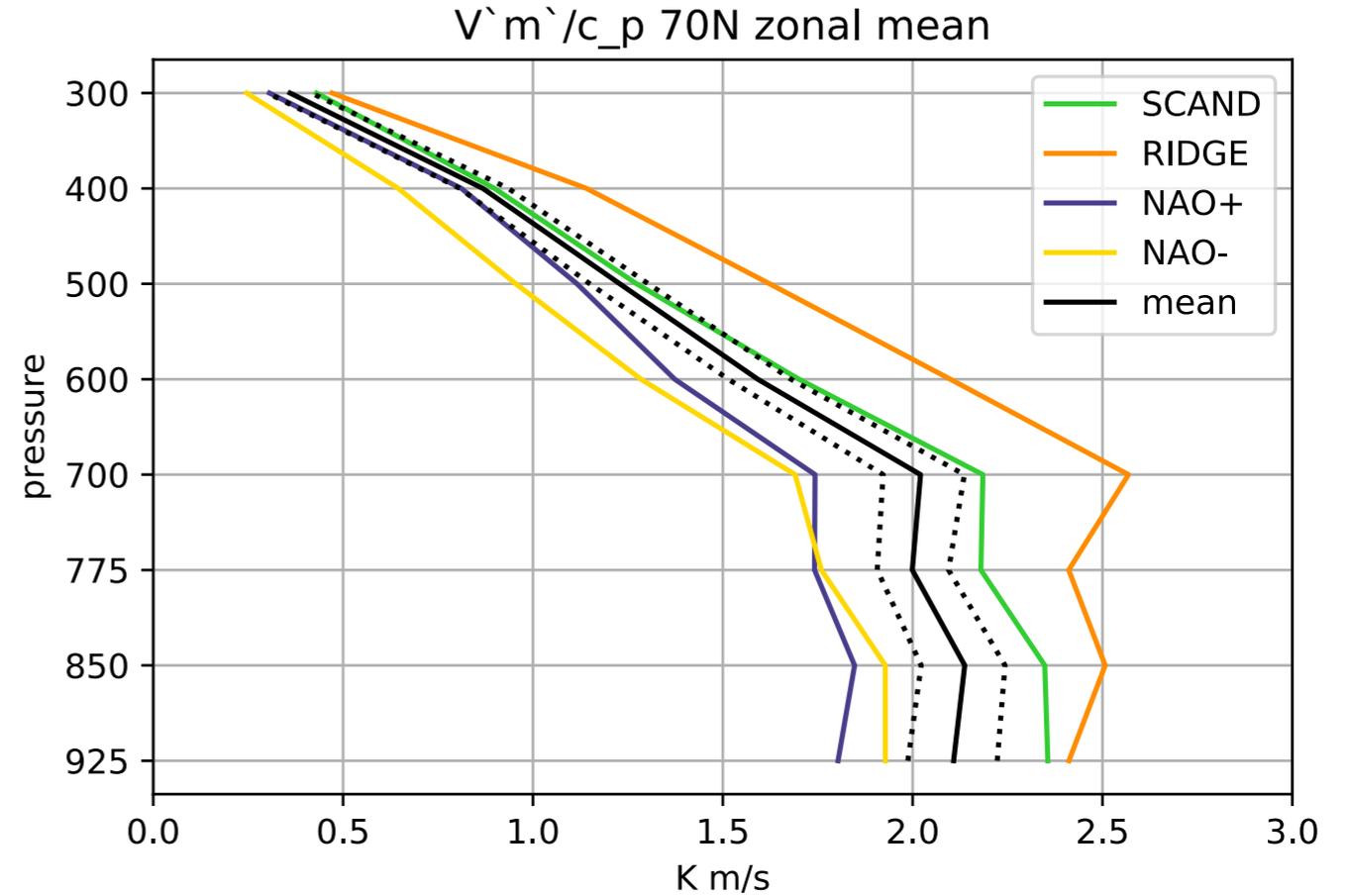
# Zonal mean heat transport

Zonal mean  $V \cdot \overline{m}$  mean and 95% confidence interval

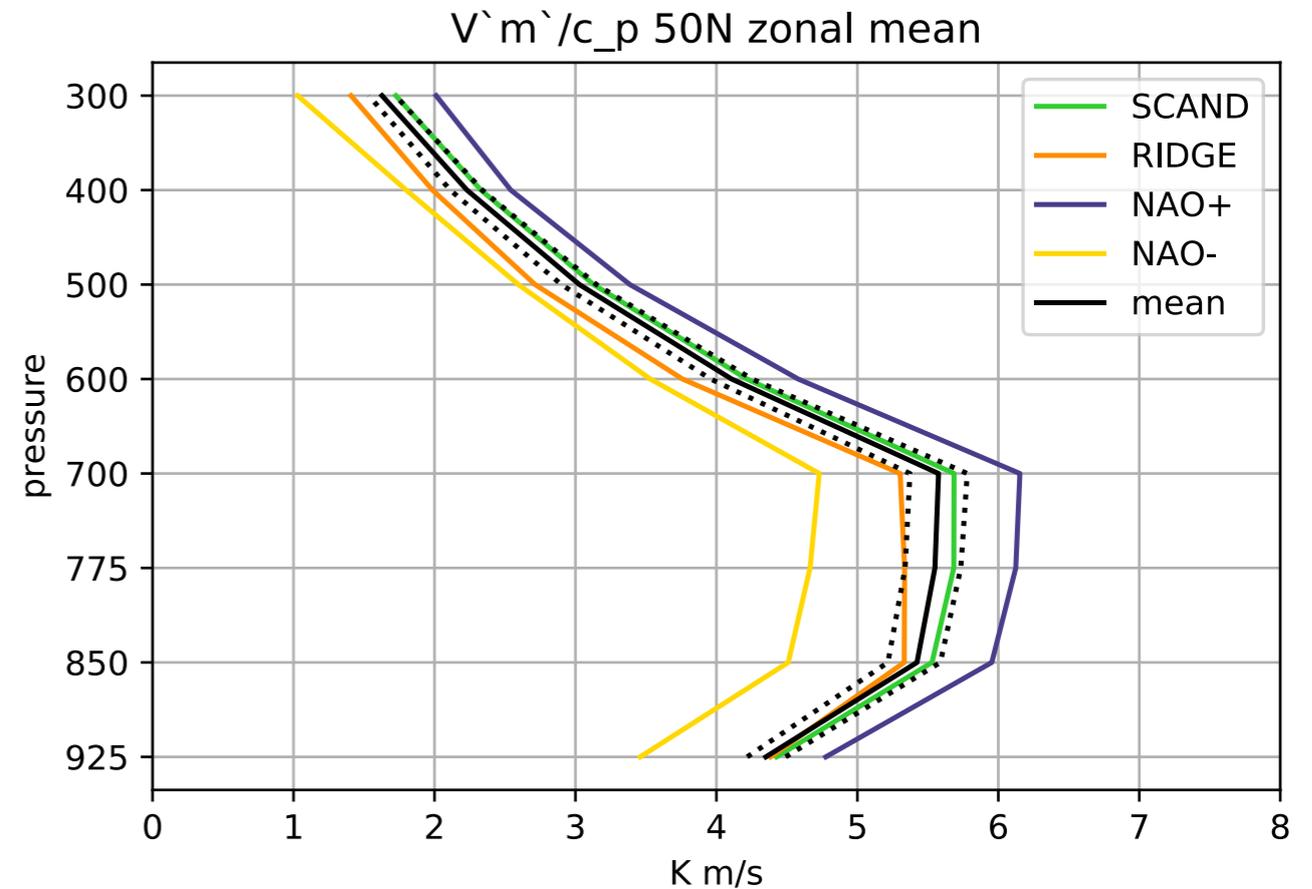


# Vertical profile at 70 °N

70 °N Zonal mean  
Both phases of NAO  
correspond to weak transport  
Scand and Ridge to  
intensified transport

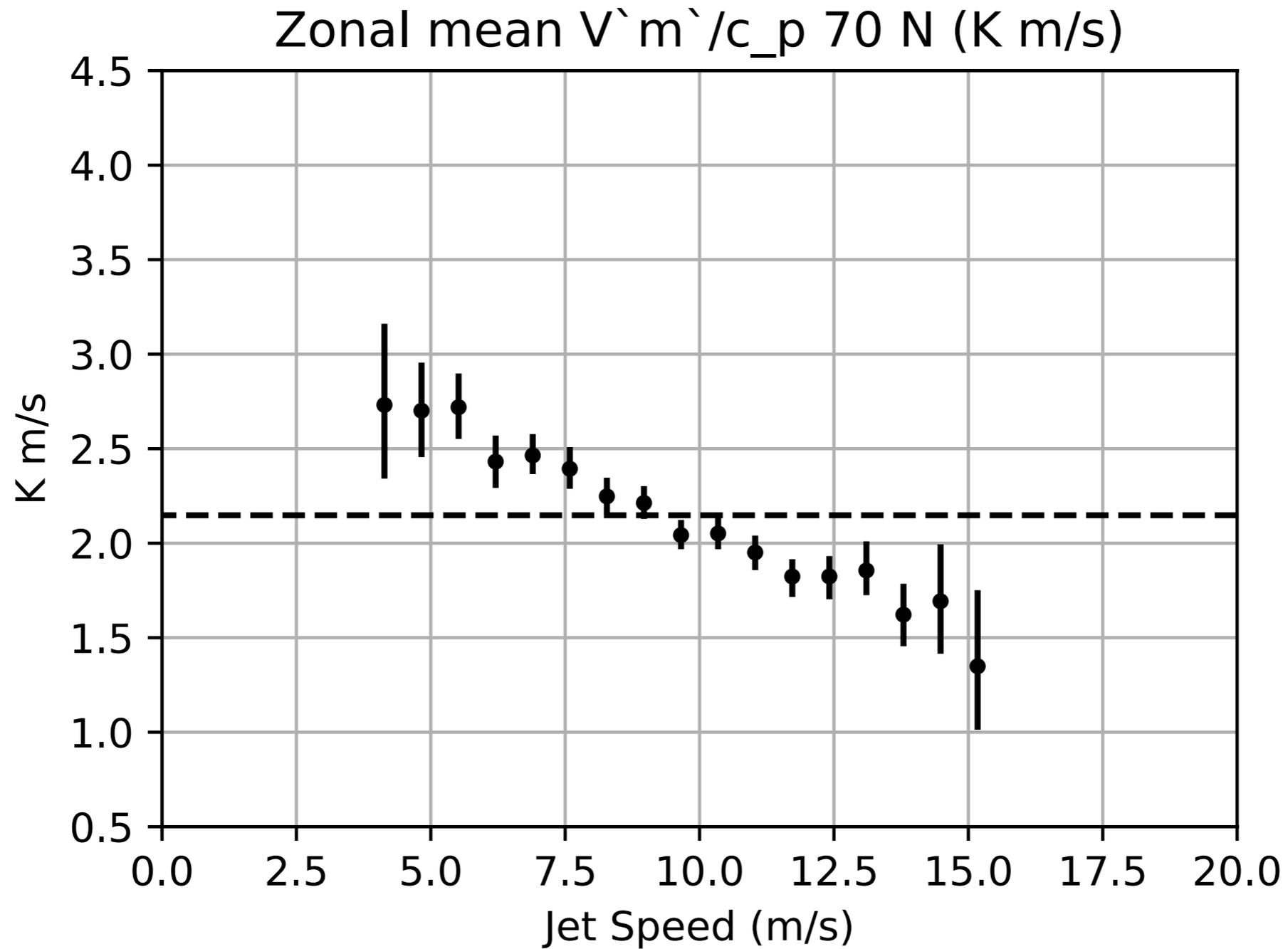


50 °N Zonal mean  
Strong Modulation by  
the NAO with opposite sign  
in the subtropics.



# Heat transport and Jet speed

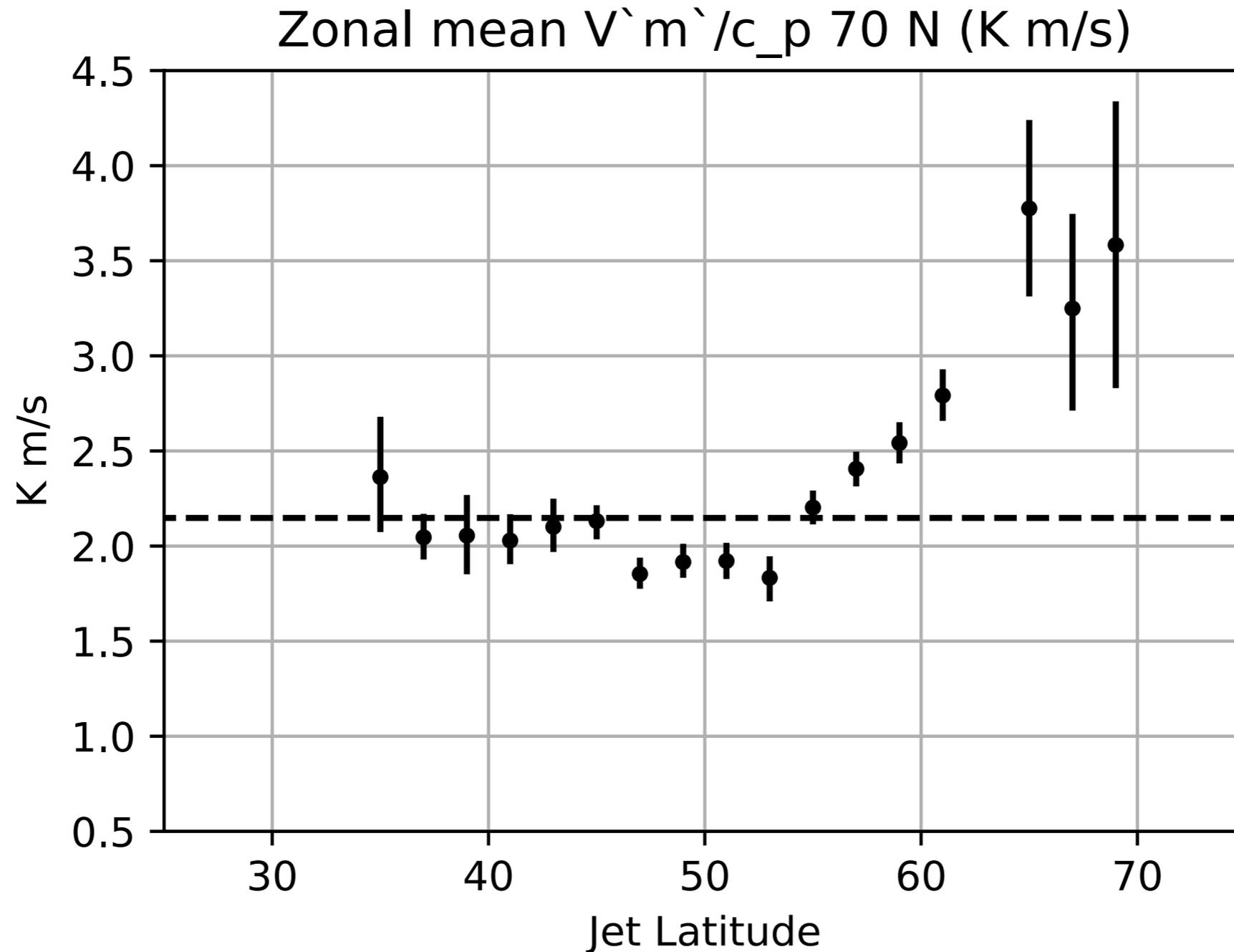
Jet speed measured in the sector 60W-0 15-75N at 850 hPa



Heat transport decreases linearly with jet speed

# Heat transport and Jet latitude

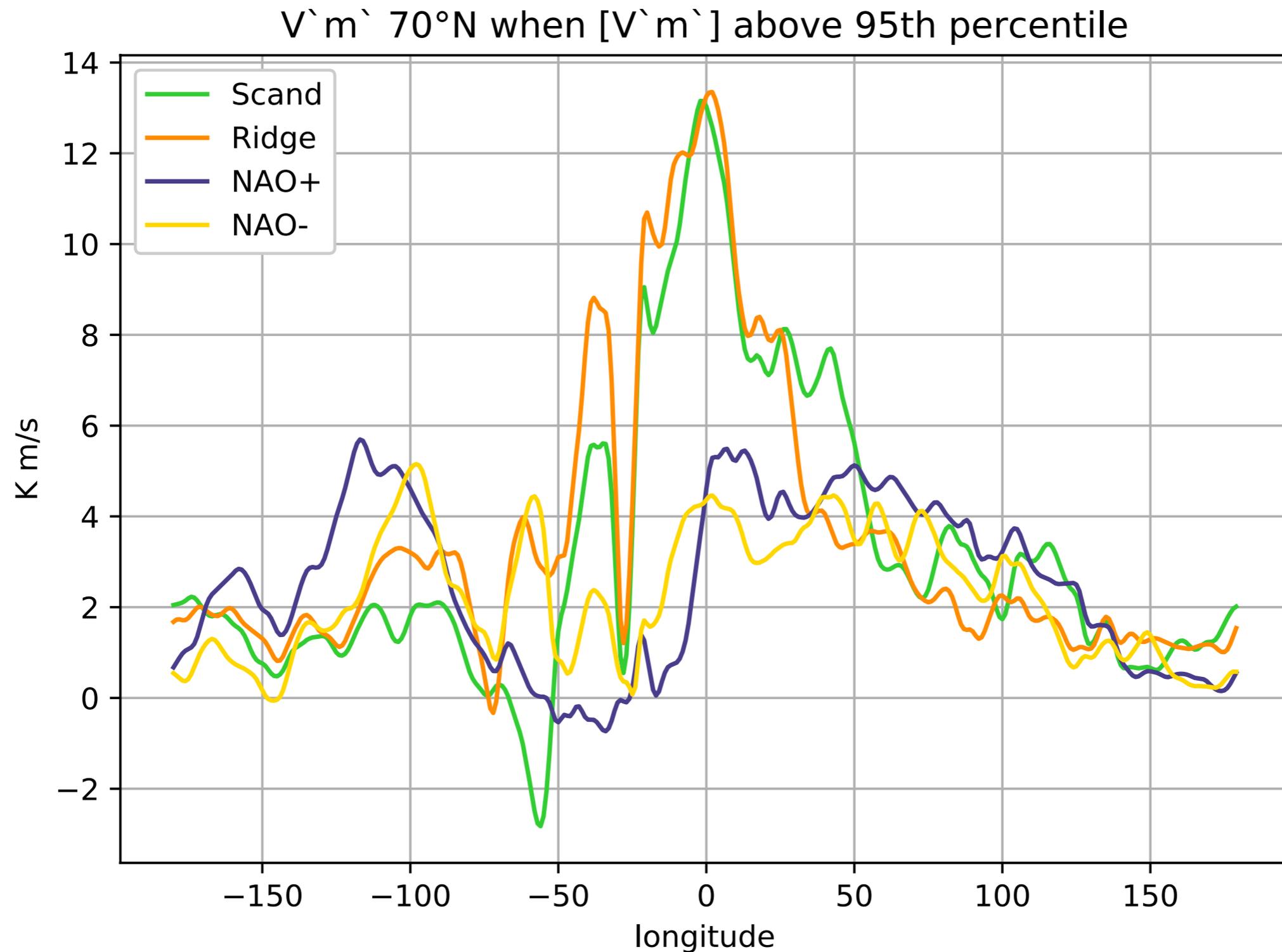
Jet latitude measured in the sector 60W-0 15-75N at 850 hPa  
Trimodal distribution



“Threshold” behaviour, strong HT with Northern regime

# Extremes

Longitudinal profile of synoptic eddy heat flux anomalies at 70° N in very strong zonally averaged



# Summary

- 1) Transient eddy heat transport (TEHT) is substantially modulated by WR on a regional scale.
- 2) Zonally average TEHT at  $70^{\circ}\text{N}$  is not significantly modulated by the phases of the NAO. It is on average stronger during Scand and Ridge and weaker in both phases of the NAO
- 3) TEHT at  $70^{\circ}\text{N}$  depends linearly on the North Atlantic jet speed. It also depends on jet latitude, being significantly large in the case of northern jet.



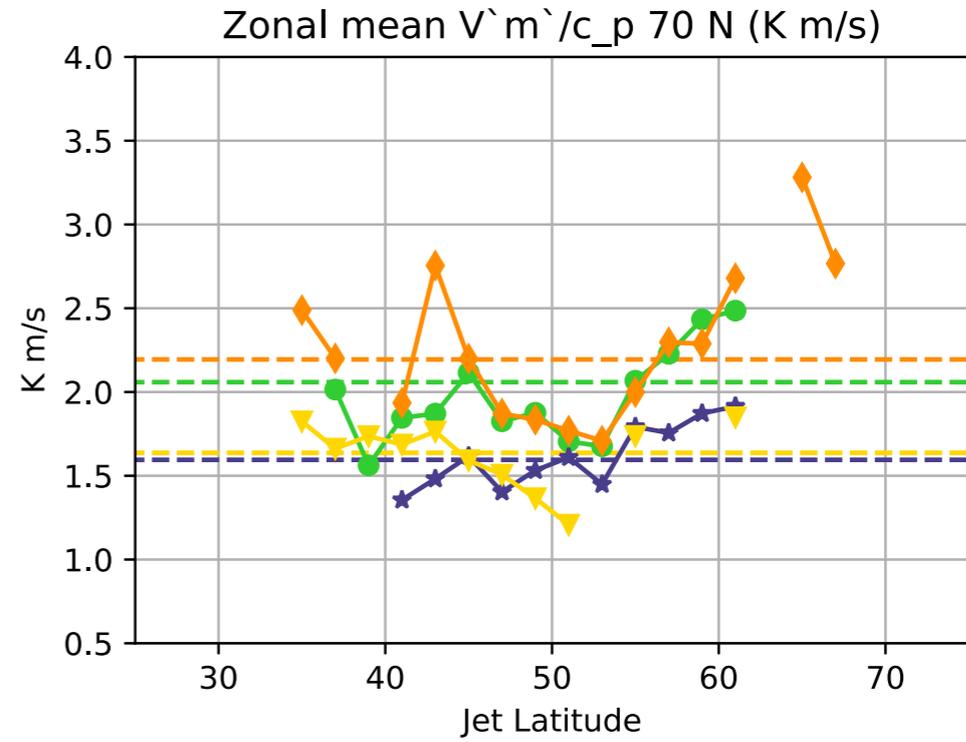
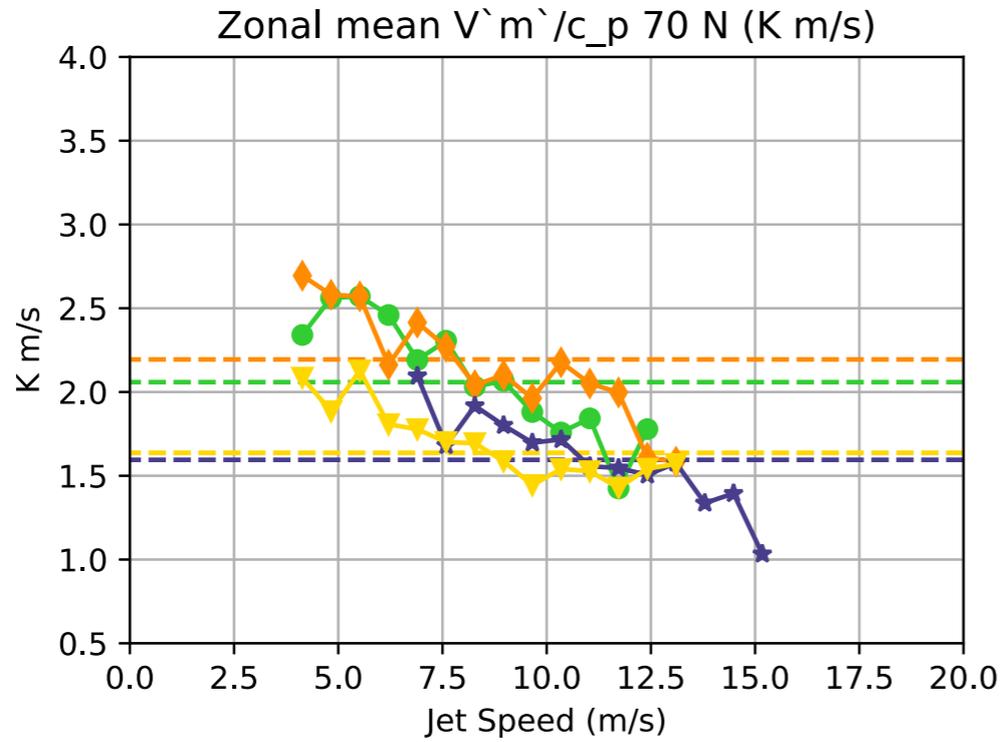
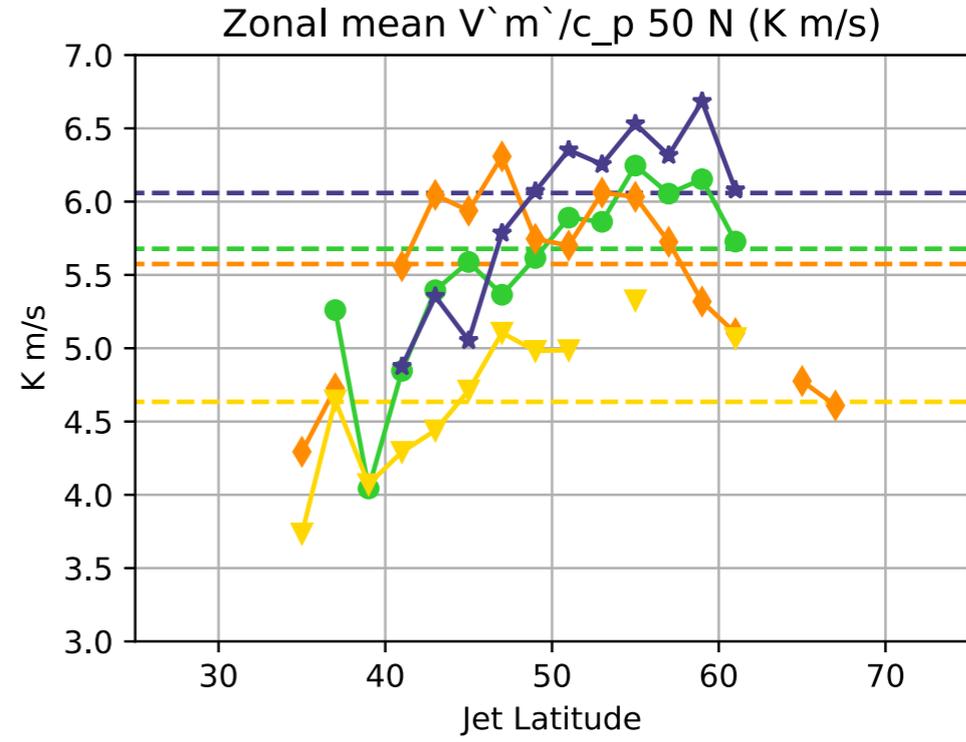
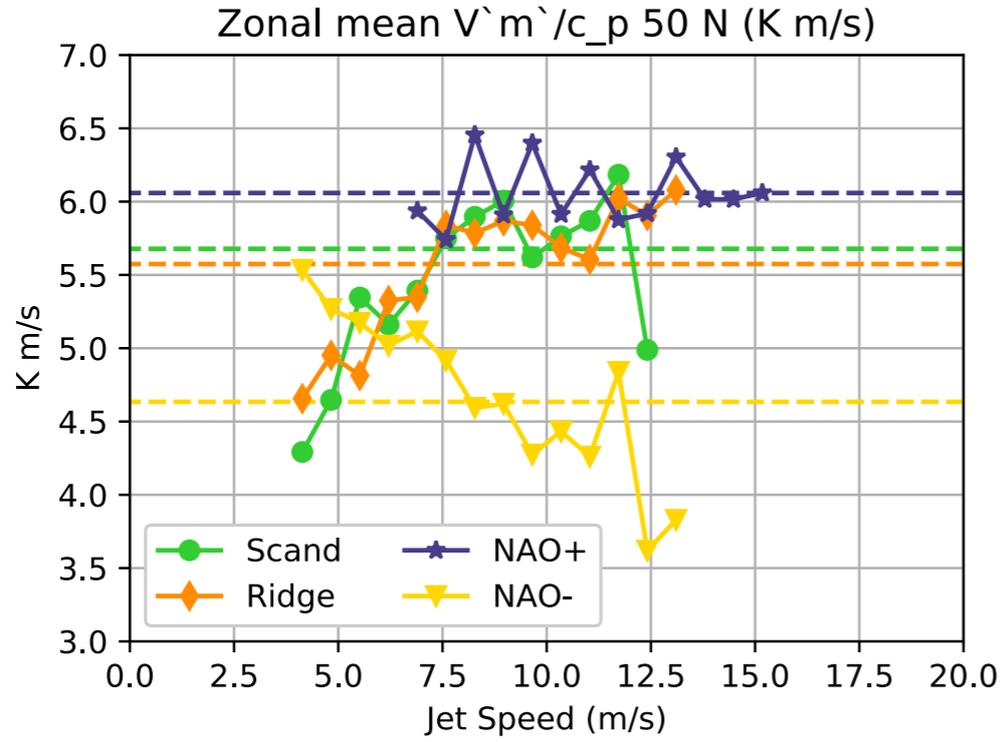
The Blue-Action project has received funding from the European Union's Horizon 2020 research and innovation programme under grant agreement No 727852

Thank you

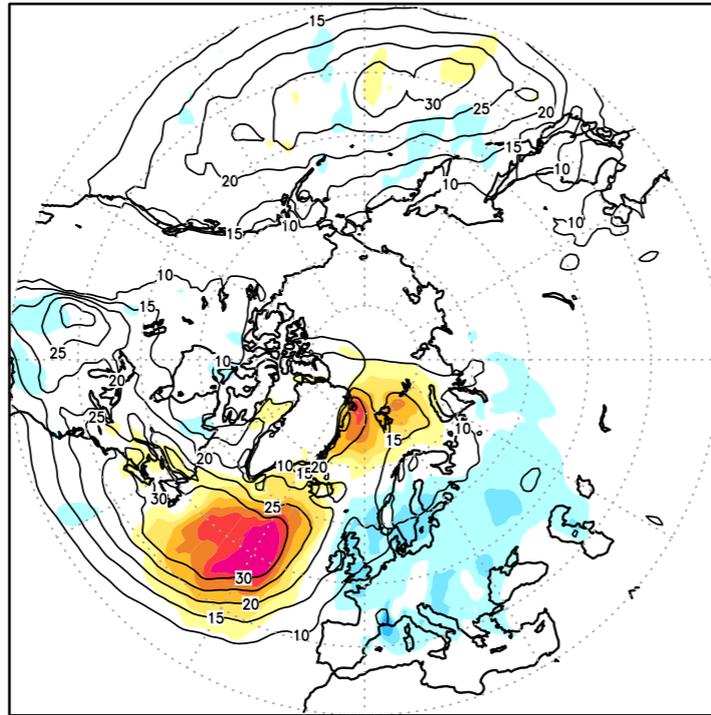
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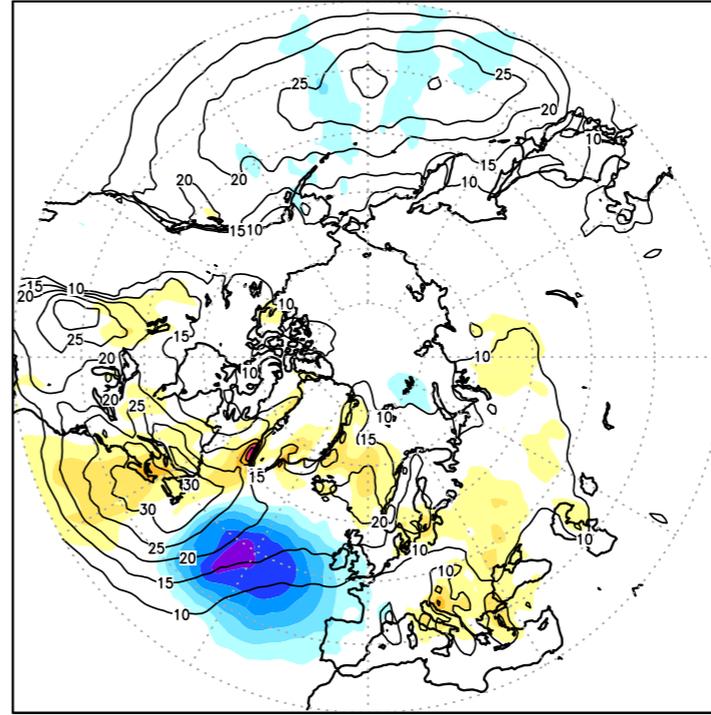
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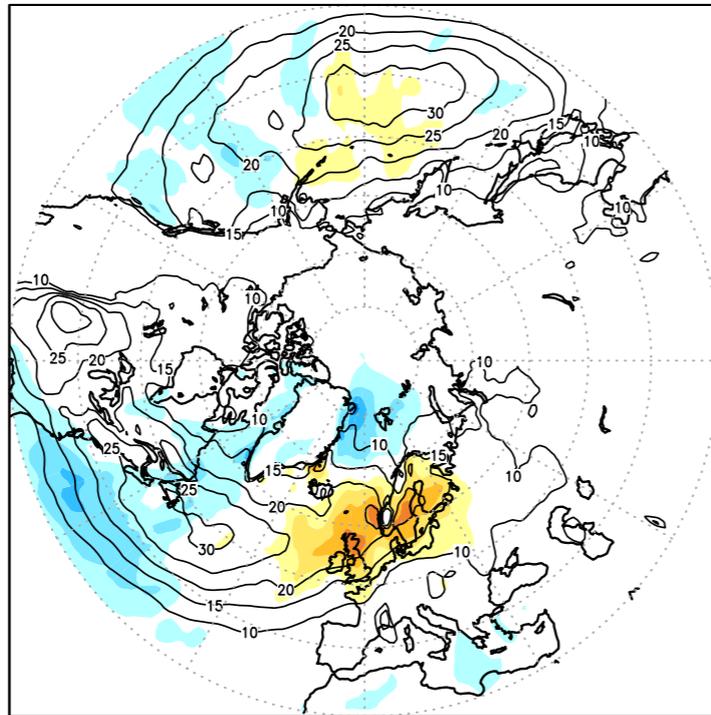
WR=SCAND V'V' NDJFM 1980-2017 (K m/s)



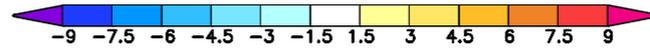
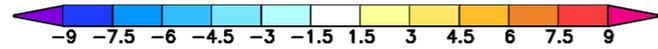
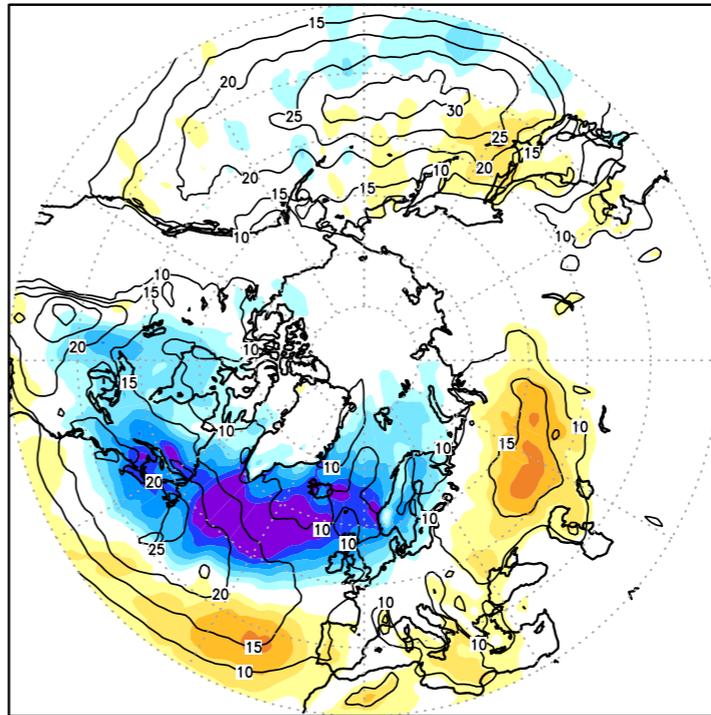
WR=RIDGE V'V' NDJFM 1980-2017 (K m/s)



WR=NAO+ V'V' NDJFM 1980-2017 (K m/s)

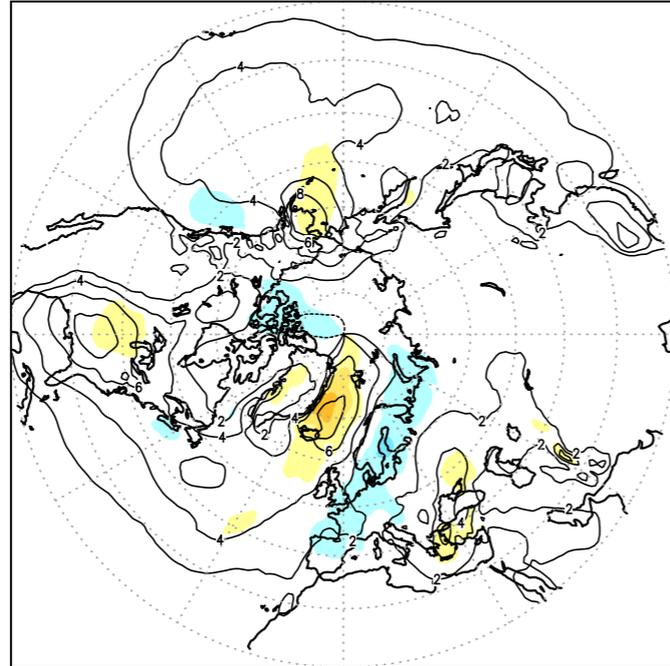


WR=NAO- V'V' NDJFM 1980-2017 (K m/s)

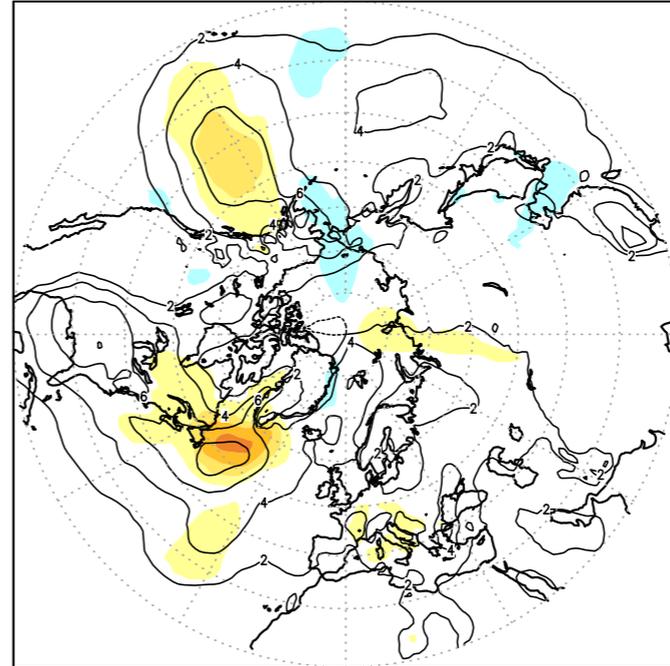


INTRA-SEASONAL EDDIES TRANSIENT  $V^{\prime}M^{\prime}$  850 hPa

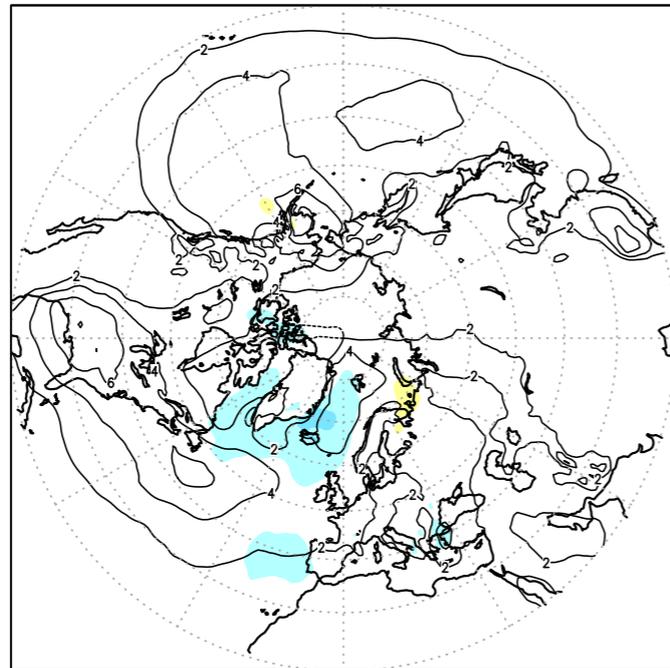
WR=SCAND  $m^{\prime}V^{\prime}/c_p$  NDJFM 1980-2017 (K m/s)



WR=RIDGE  $m^{\prime}V^{\prime}/c_p$  NDJFM 1980-2017 (K m/s)



WR=NAO+  $m^{\prime}V^{\prime}/c_p$  NDJFM 1980-2017 (K m/s)



WR=NAO-  $m^{\prime}V^{\prime}/c_p$  NDJFM 1980-2017 (K m/s)

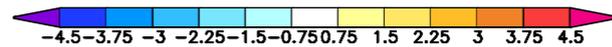
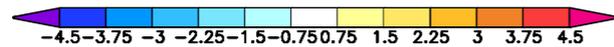
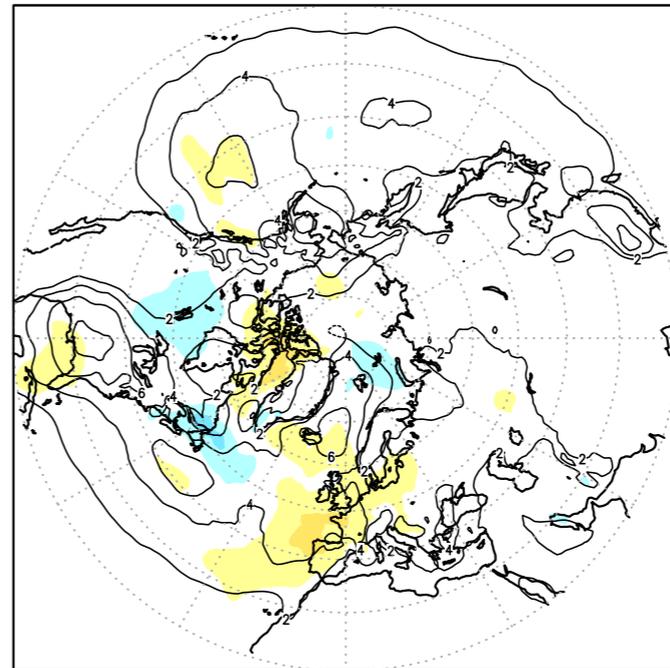
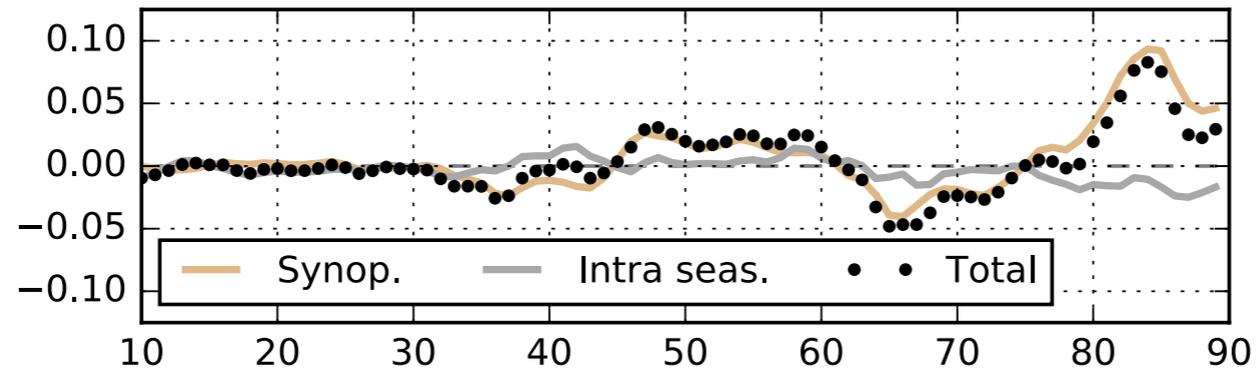


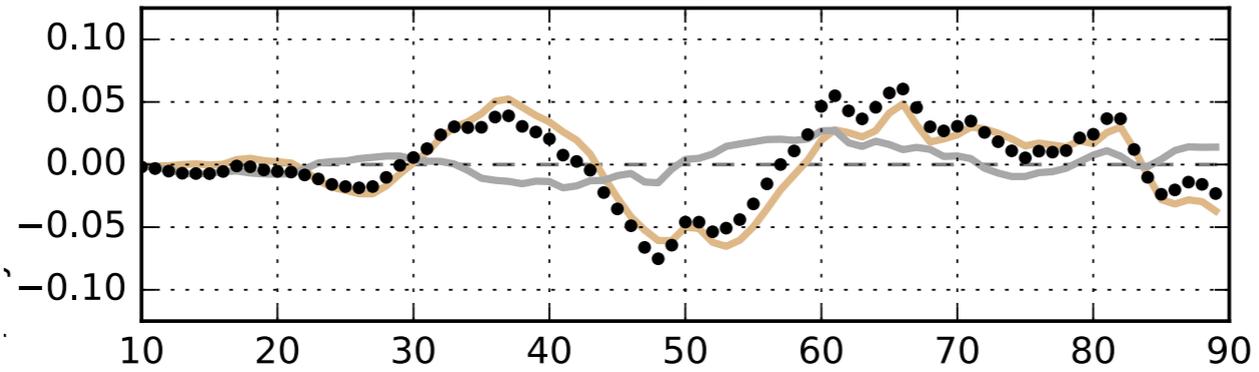
FIG. 7. As in figure 6 but for transient intra-seasonal eddies.

Temperature eddy tendency 80W-50E (K/day)

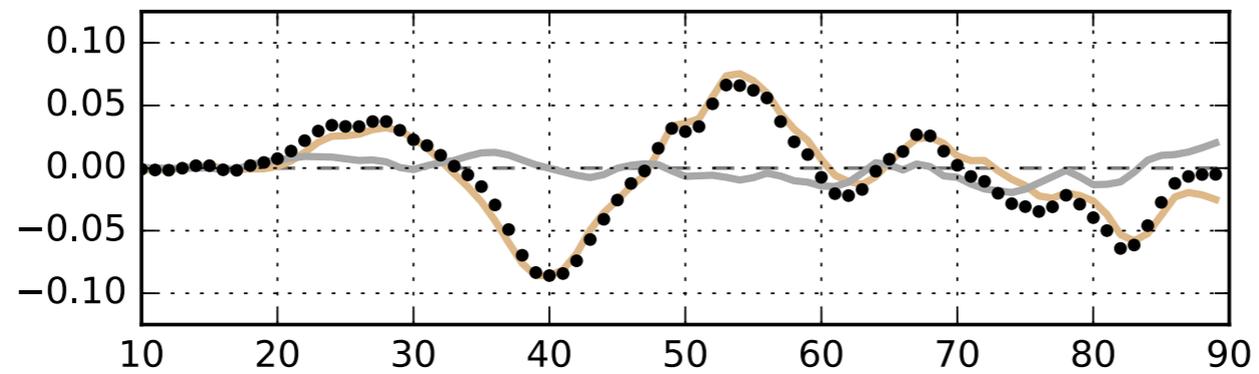
Scand



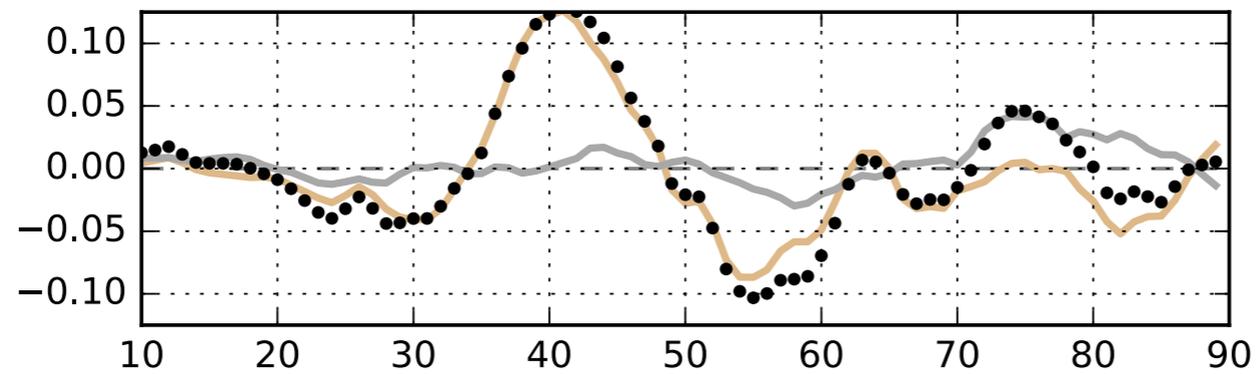
Ridge



NAO-



NAO+



Latitude

