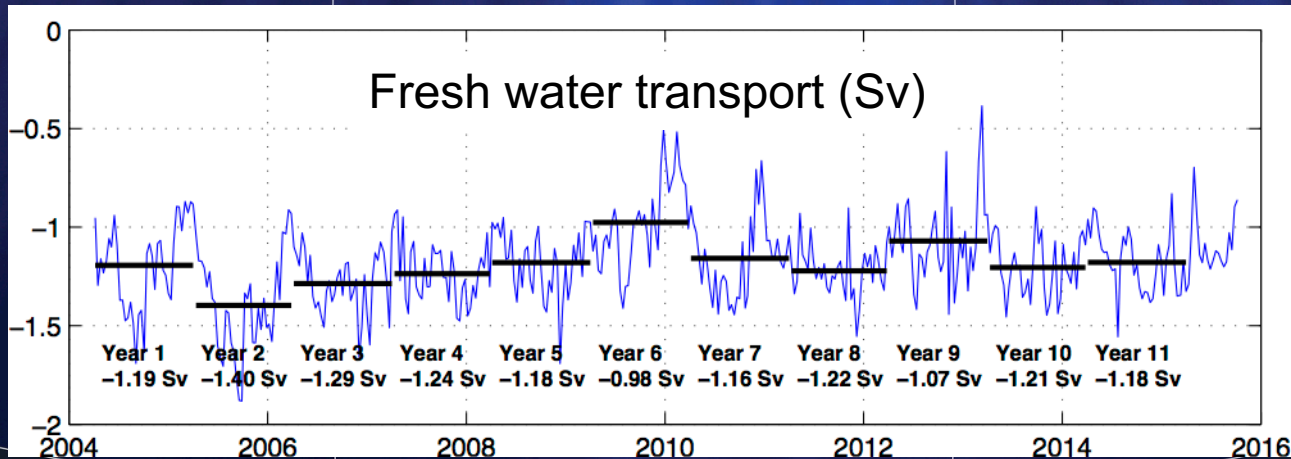


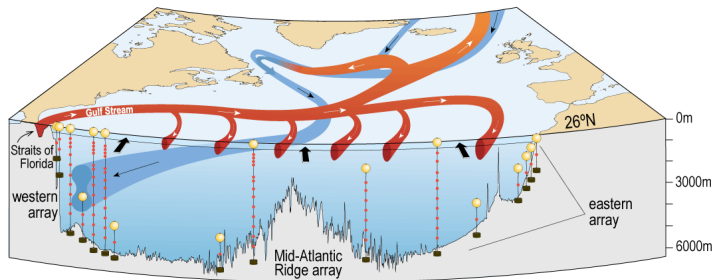
Transports of freshwater and heat in the subtropical North Atlantic

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G. D. McCarthy¹, D. Rayner¹, M. Baringer³, C. S. Meinen³ and Harry Bryden⁴

1) NOC, UK, 2) RSMAS, USA, 3) NOAA, USA., 4) University of Southampton

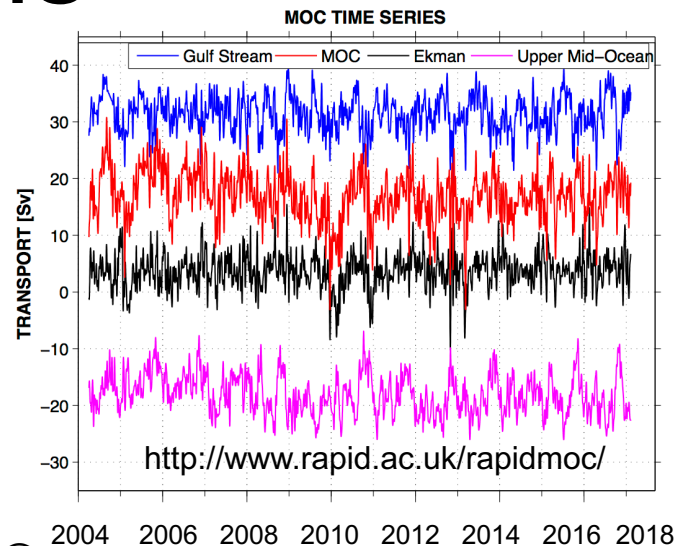
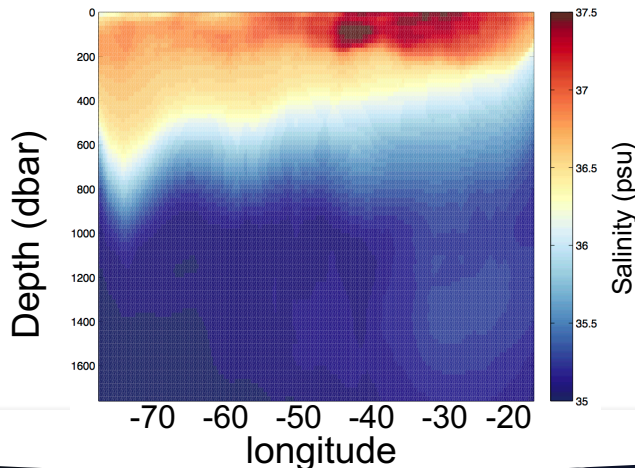


Observations



Temperature and salinity data from the RAPID-AMOC array

Optimally interpolated Argo data (top 1800db)



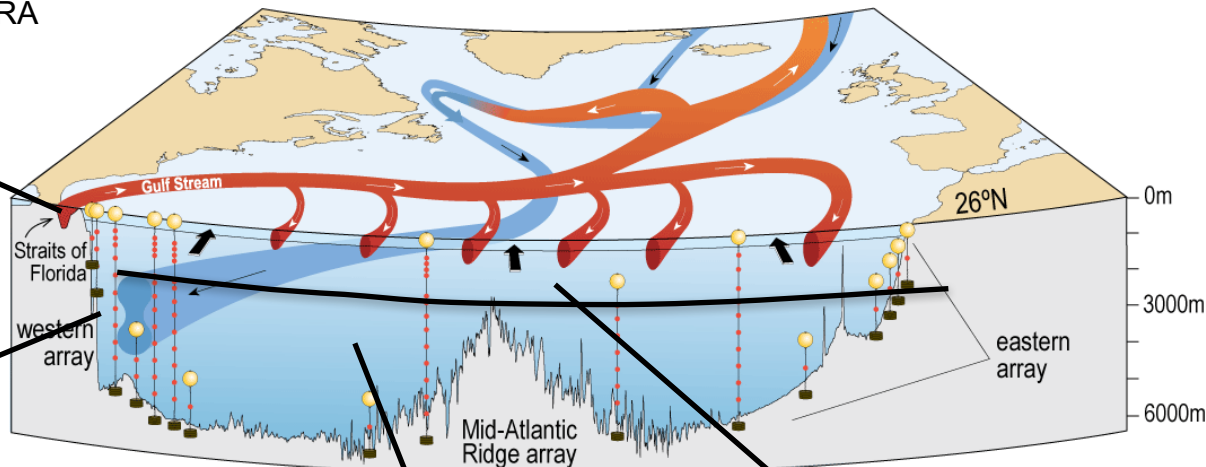
Transport time series from the RAPID-AMOC array

Method

Estimates of temperature, salinity
And velocity fields every 10 days
2004 to 2015

Ekman (uppermost layer)
Salinity from Argo OI
Transport from ERA

Florida Straits
Cable transport +
Transport weighted salinity
(Szuts and Meinen, 2012)



Western Boundary Wedge (west of wb2)
Transport and salinity from
RAPID-AMOC moorings

Interior > 1800db (east of wb2)
Transport and salinity from
RAPID-AMOC moorings

Interior < 1800db (east of wb2)
Salinity from Argo OI
Transport from Argo density field
referenced to RAPID-AMOC
transports

Method

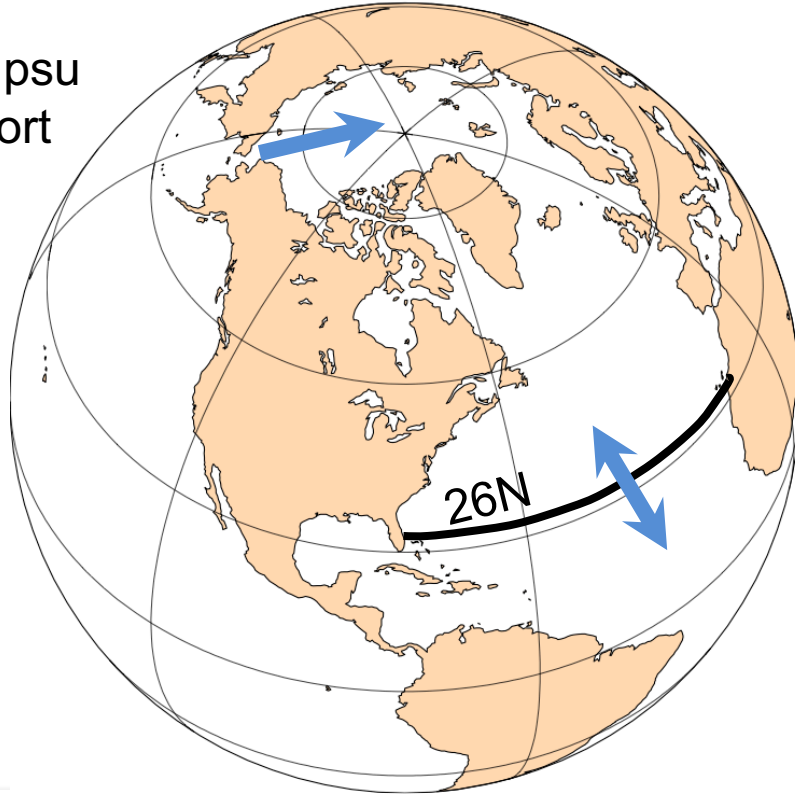
Bering Strait:

0.8 Sv at 32.5 psu

Salinity transport

of 26 Sv psu

(Woodgate et al., 2005)



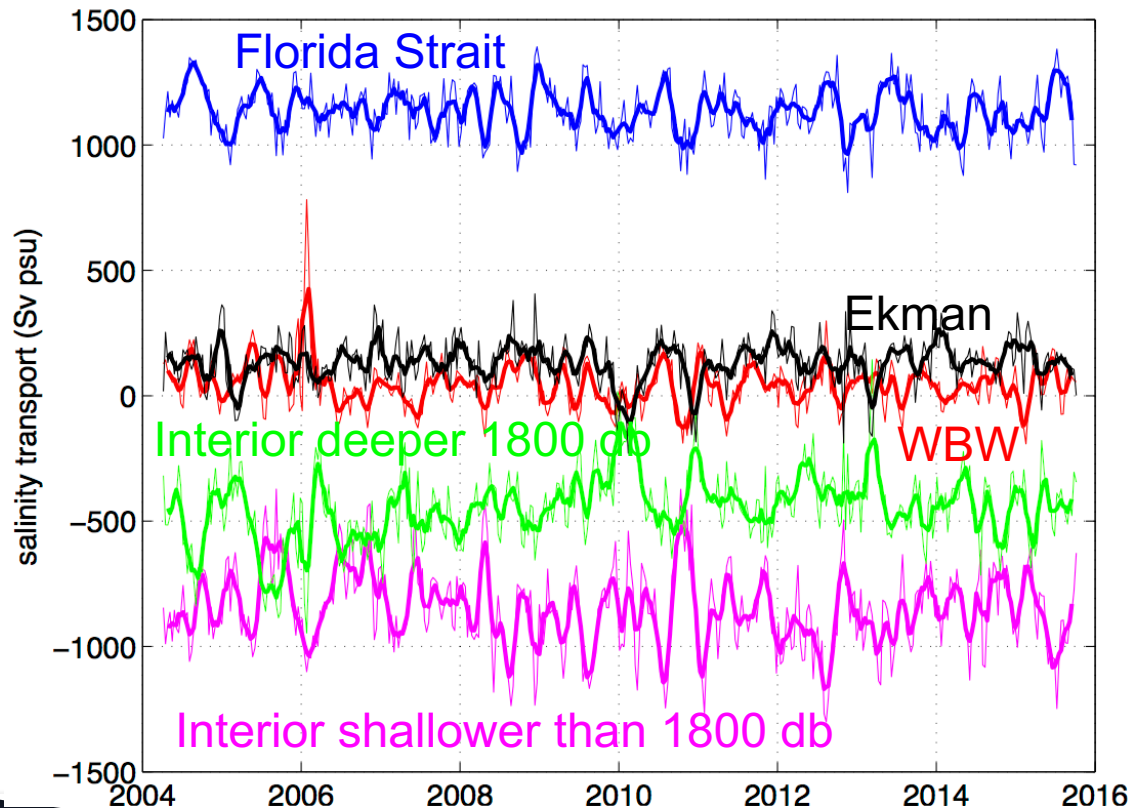
Set the section salinity transport equal to the Bering Strait salinity transport.

Salt is conserved between BS and 26N

The net volume flux across the Section is the freshwater flux.

$$FW = -1 \times (26.0 - ST_{26N}) / \overline{Sal}_{26N}$$

Salinity Transports 26N



FS: 1135 ± 102 Sv psu

Ekman: 133.2 ± 94.2 Sv psu

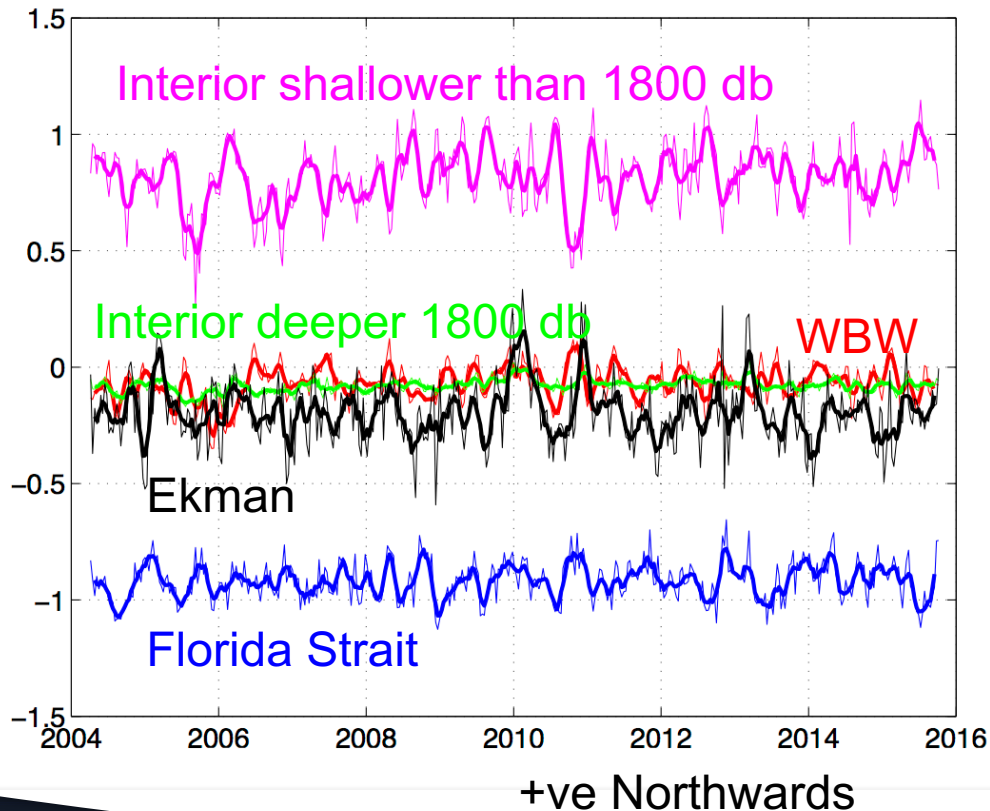
WBW: 45.8 ± 98.5 Sv psu

Int deep: -447.6 ± 147.7 Sv psu

Int shallow: -851.2 ± 165.5 Sv psu

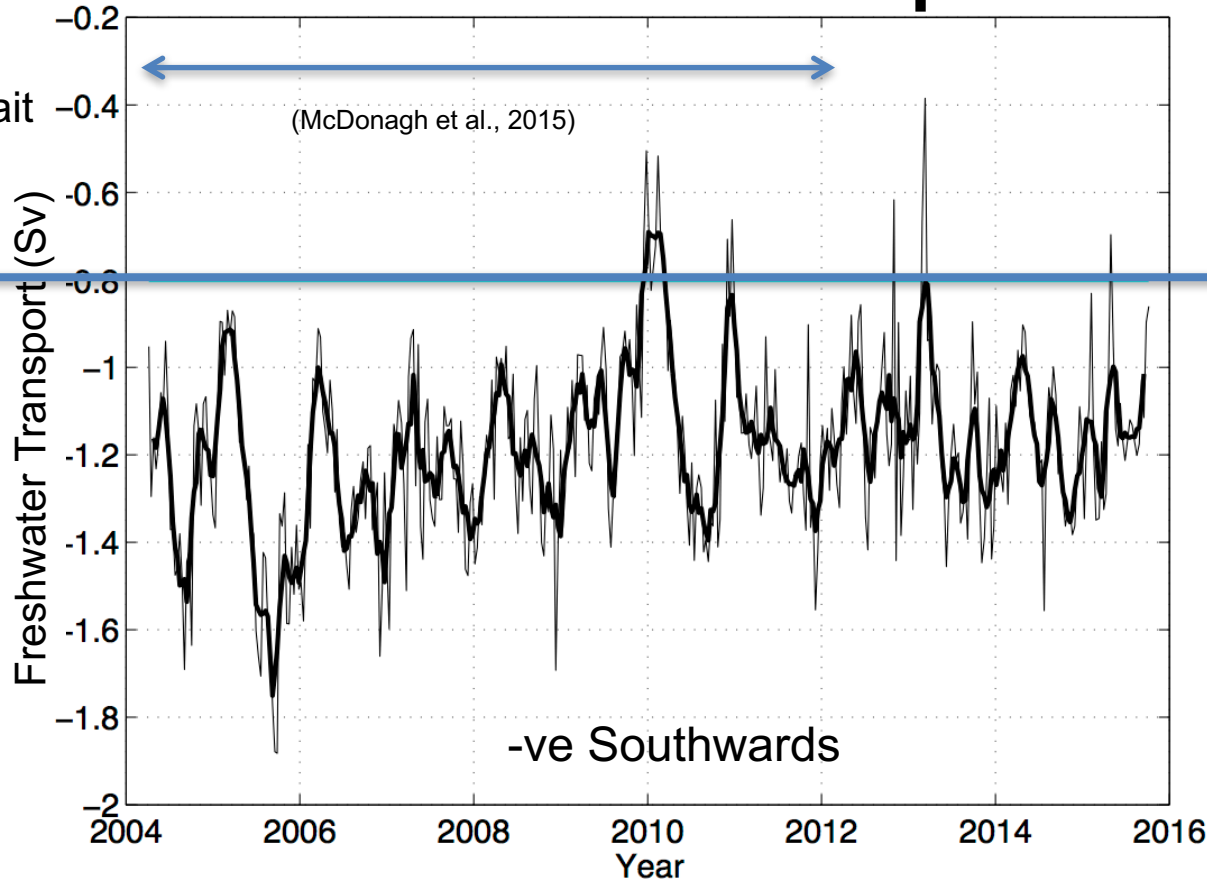
+ve Northwards

Equivalent Fresh water transports



Dominant balance between the Florida Straits (-0.91 Sv) and the upper interior (0.81 Sv)

Freshwater Transport at 26N

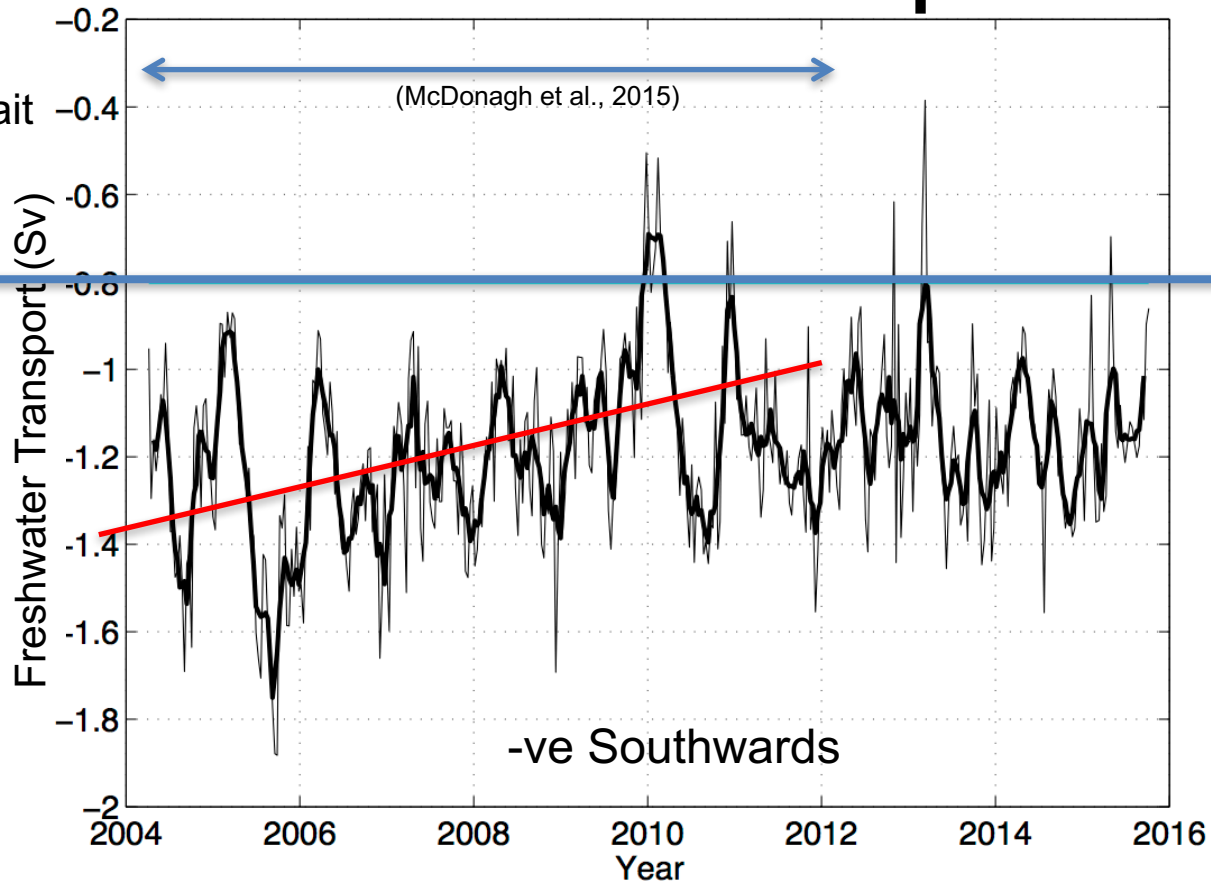


Bering Strait
-0.8 Sv

$E > P$
Net freshwater loss

$E < P$
Input of freshwater
Into the ocean

Freshwater Transport at 26N



Bering Strait
-0.8 Sv

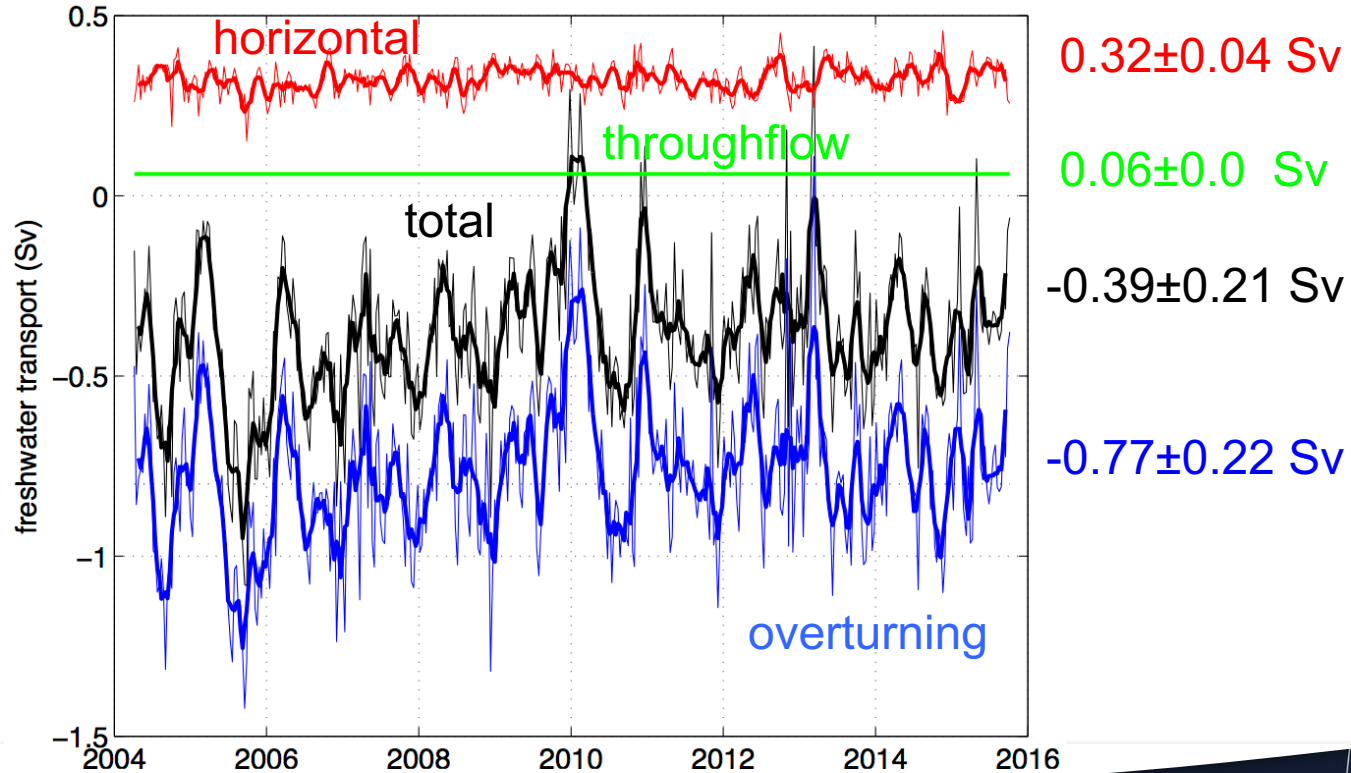
$E > P$
Net freshwater loss

$E < P$
Input of freshwater
Into the ocean

Components freshwater DIVERGENCE

Fresh Water
Divergence Includes:

air-sea fluxes (E-P),
runoff,
ice/melt formation,
oceanic storage of
Fresh water

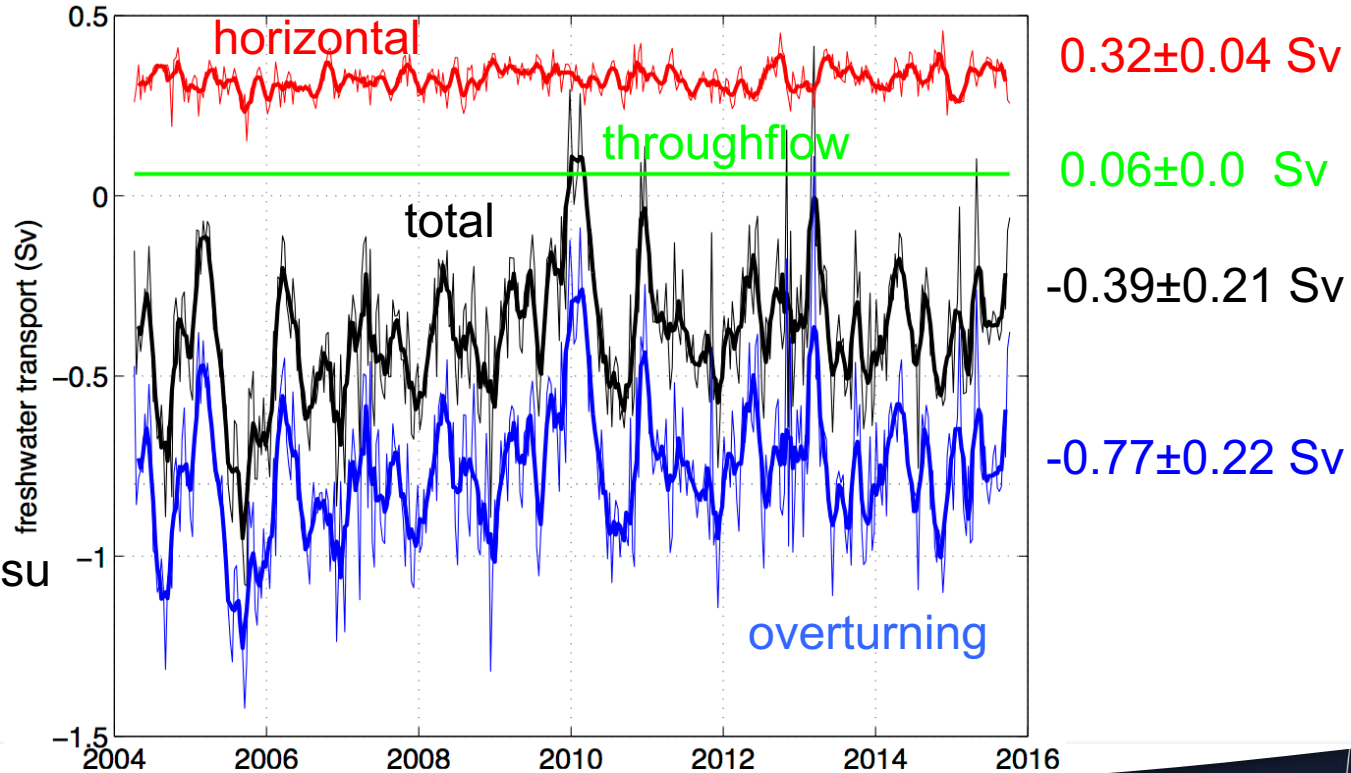
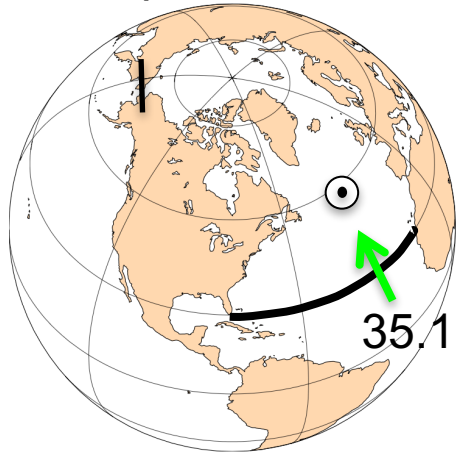


Components freshwater DIVERGENCE

Throughflow

Loss of freshwater
Between BS and 26N.

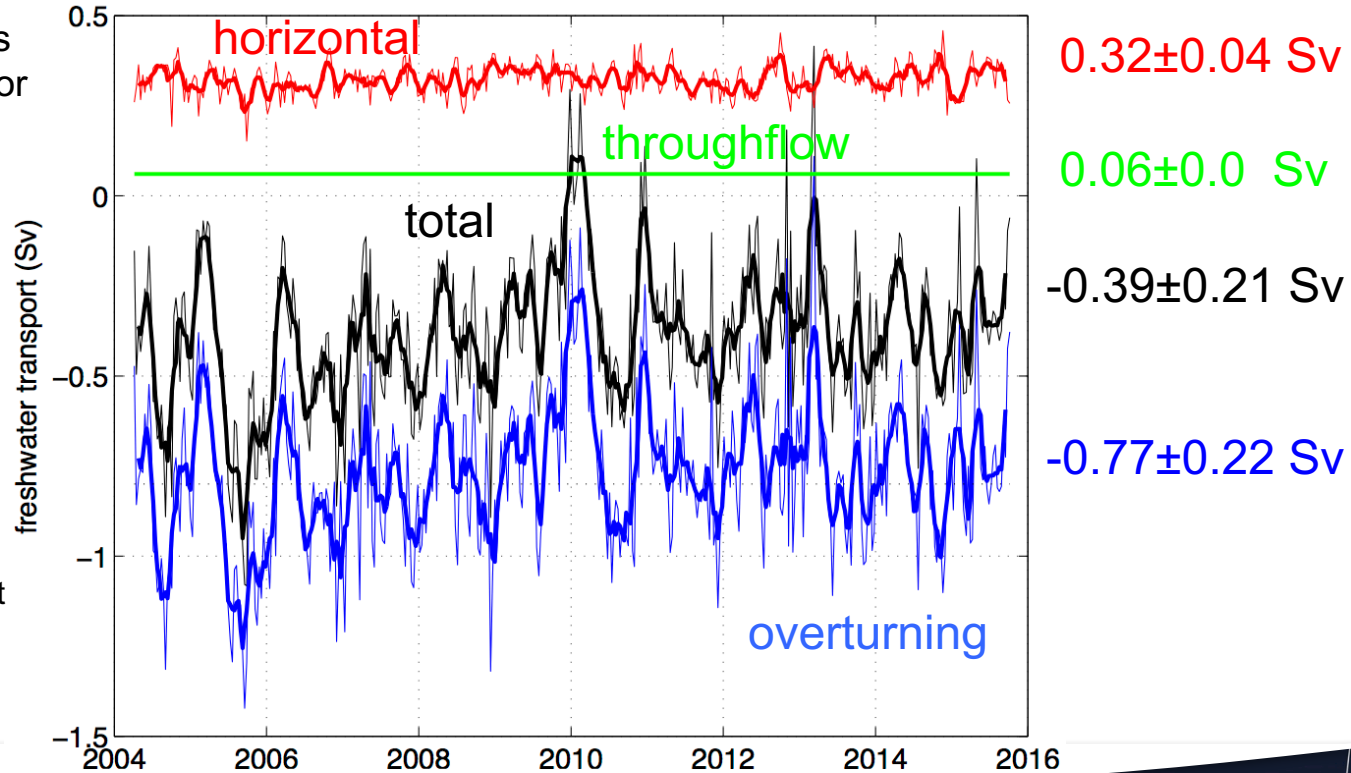
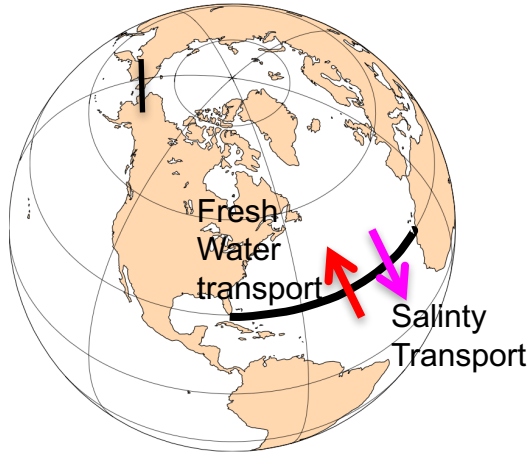
32.5 psu



Components freshwater DIVERGENCE

Horizontal

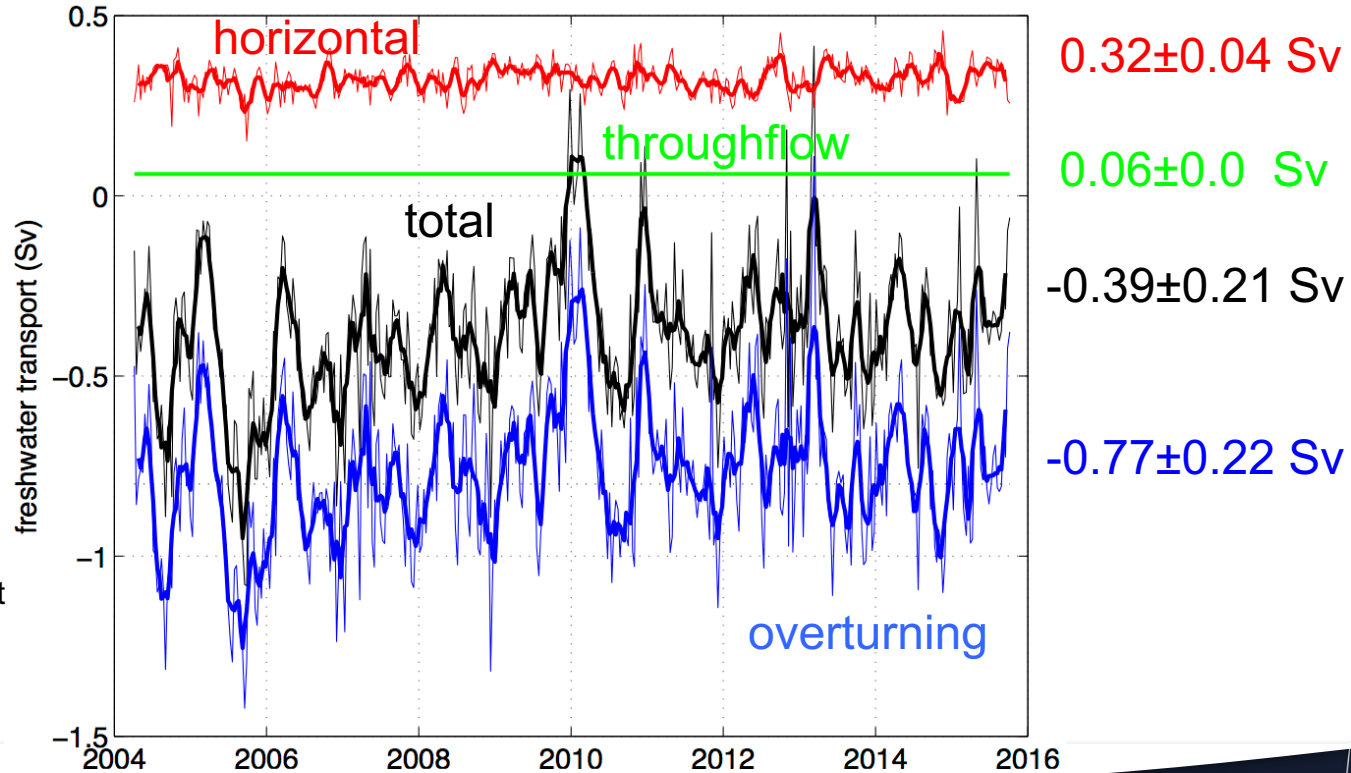
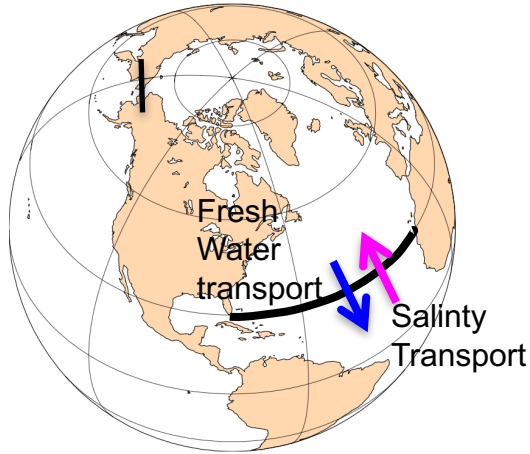
Balance between the northwards flowing FS current and the interior upper ocean return flow



Components freshwater DIVERGENCE

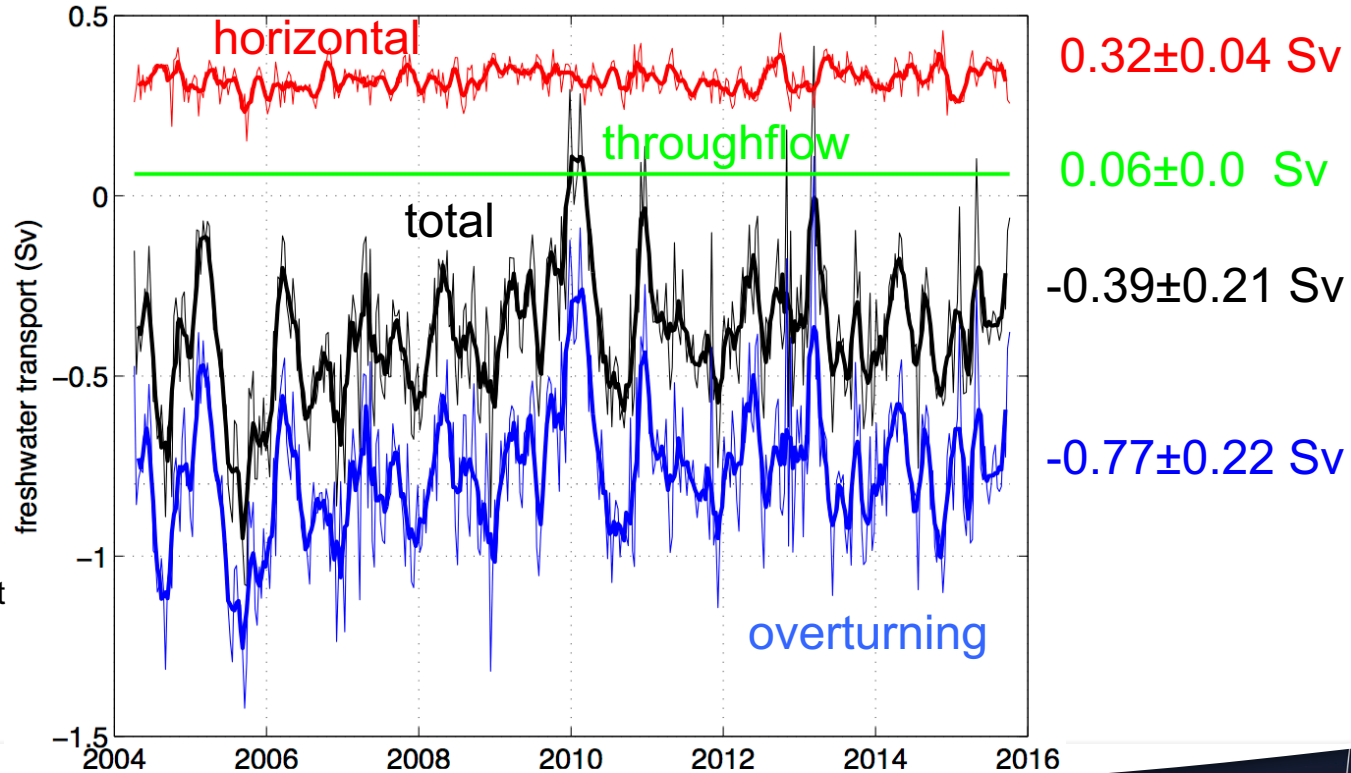
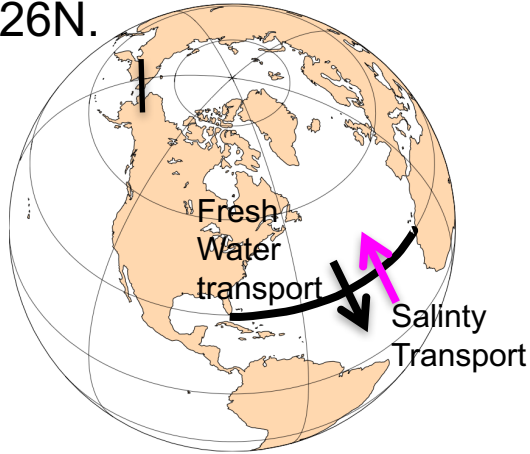
Overtuning

Largest and most variable component.

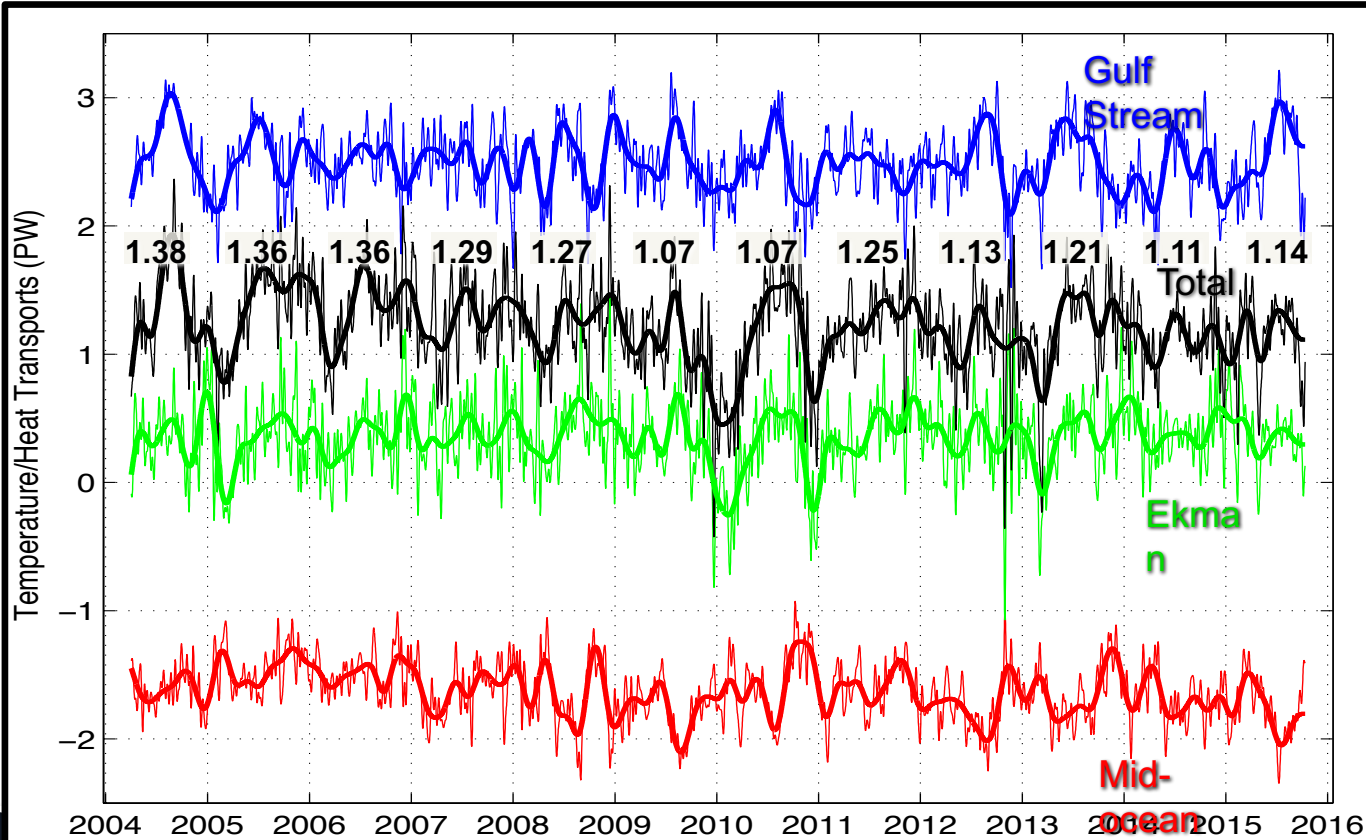


Components freshwater DIVERGENCE

Total:
Freshwater INPUT
Into the North Atlantic
ocean between BS and
26N.



Heat Transport at 26N (time series updated to 2015)



2004-2015
mean heat
transport:
 1.22 ± 0.11 PW

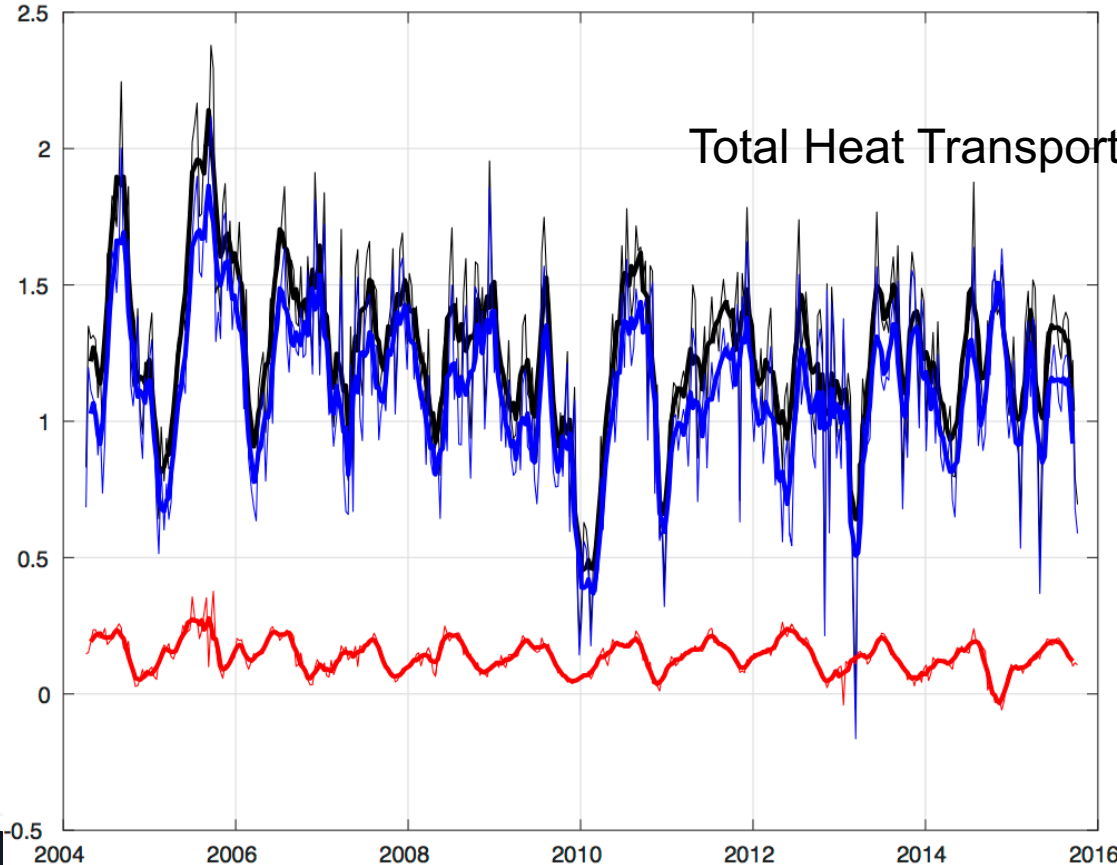
Multi-year
averages:
2004-2008
 1.33 ± 0.05 PW

2009-2015
 1.14 ± 0.07 PW

2011-2015
 1.17 ± 0.06 PW

+ve Northwards

Heat Transport at 26N

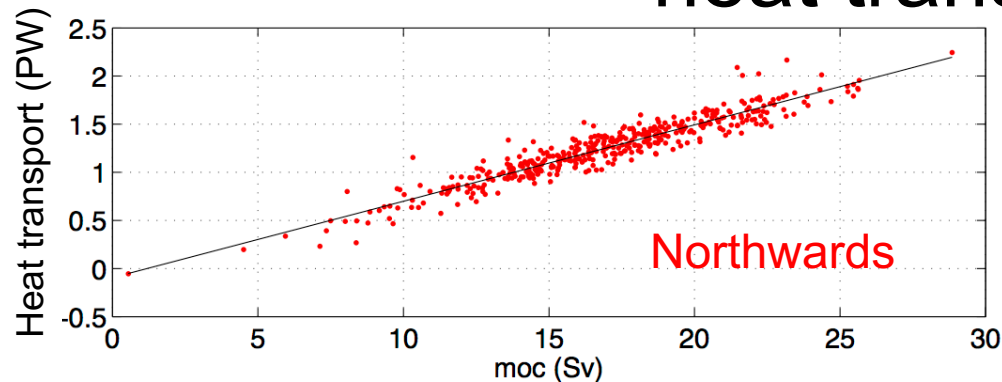


Total Heat Transport = 1.25 ± 0.34 PW

Overturning : 1.11 ± 0.33 PW

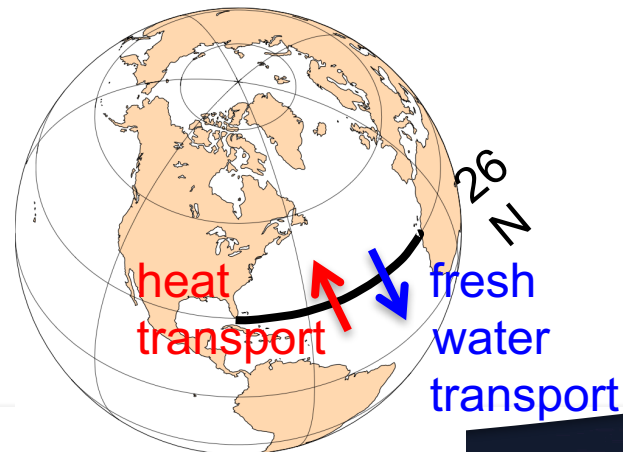
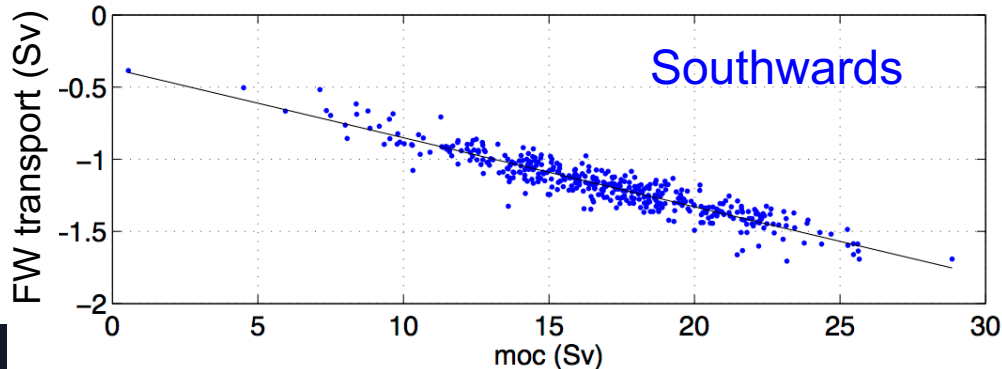
Horizontal = 0.14 ± 0.06 PW

Relationship between AMOC, freshwater and heat transports

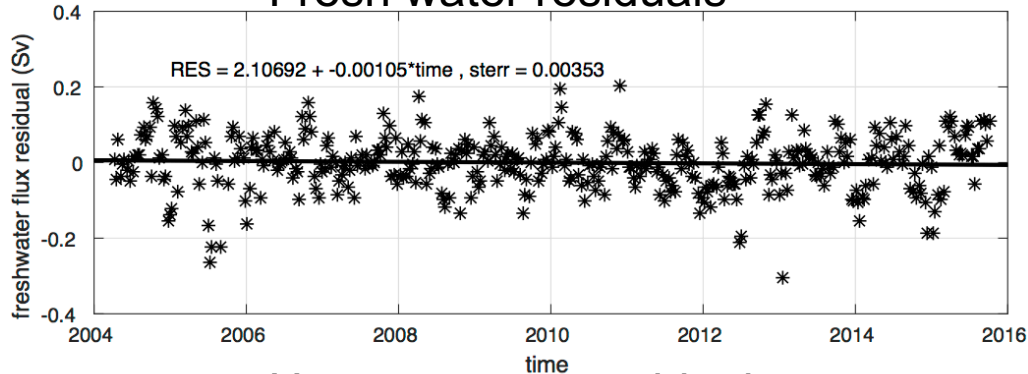


$$\text{Heat transport} = -0.10 + 0.79\text{MOC}, R^2 = 0.89$$

$$\text{FW transport} = -0.37 - 0.048\text{MOC}, R^2 = 0.87$$



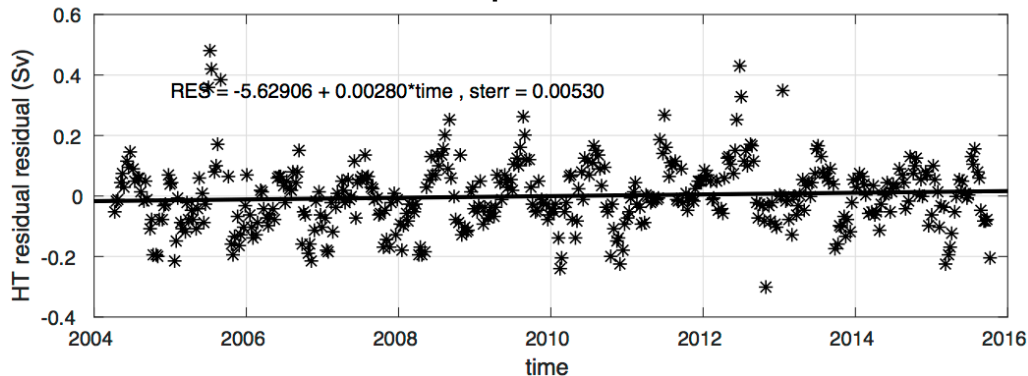
Fresh water residuals



2004 to 2012 (McDonagh et al., 2015)
Increasing -ve residuals over time.

Consistent with an increase in upper ocean salinity at **26N**.

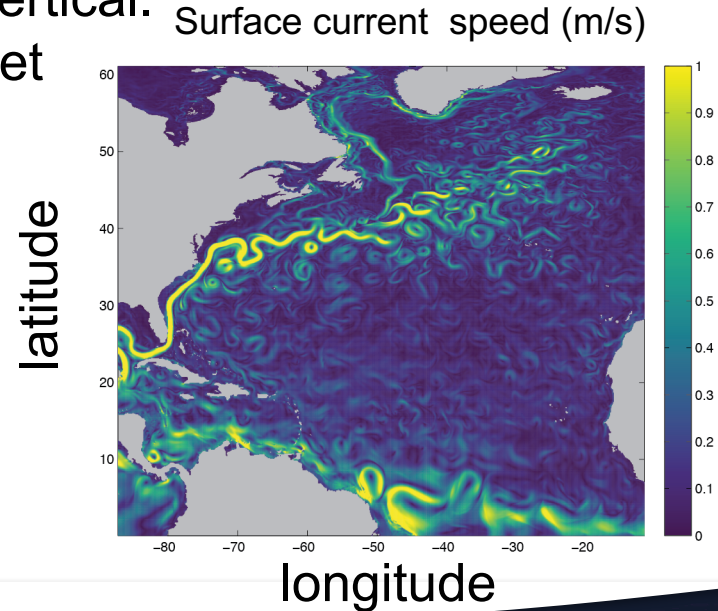
Heat transport residuals



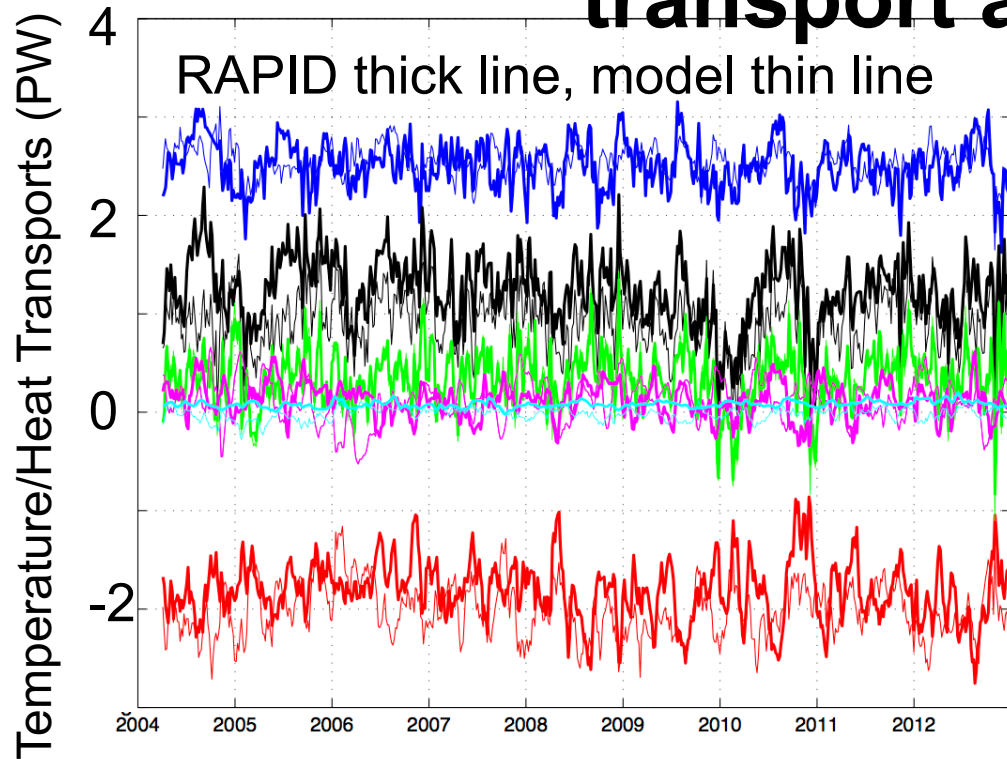
Consistent with an increase in upper ocean temperature at **26N**.

Can an ocean only model reproduce the Heat transport at 26N?

- NEMO-LIM2 ocean circulation/sea ice model (Marzocchi, 2015; Moat, 2016)
- 1/12 degree Horizontal with 75 levels in the vertical.
- Forced by the Drakkar Surface Forcing dataset (air temperature, winds, humidity, surface radiative heat fluxes and precipitation)
- Eddy resolving (in Mid latitudes)
- Run from 1958 to 2015



Can an ocean only model reproduce the Heat transport at 26N?

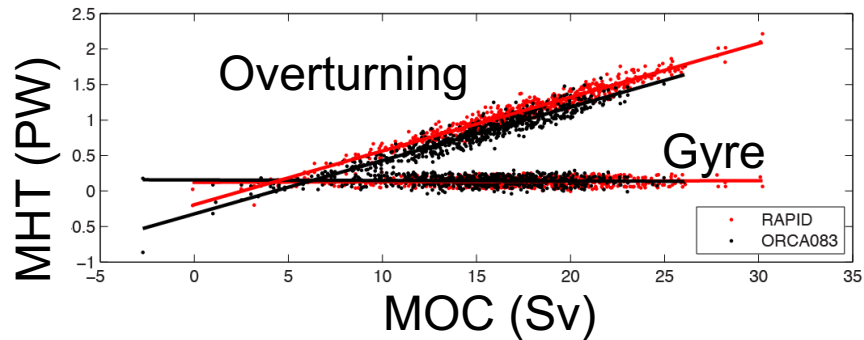
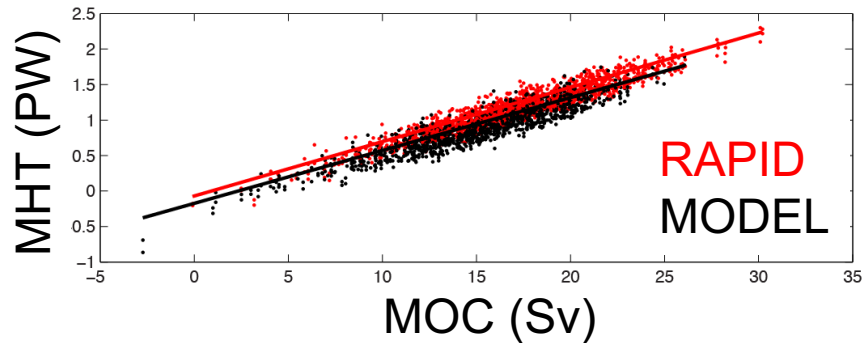


	ORCA083 (PW)	RAPID (PW)
FC	2.54 ± 0.16	2.50 ± 0.25
EKMAN	0.34 ± 0.31	0.35 ± 0.30
Mid-Ocean	-2.02 ± 0.27	-1.81 ± 0.31
WBW	0.12 ± 0.21	0.12 ± 0.18
EDDY	-0.02 ± 0.06	0.08 ± 0.03
Total heat transport	0.93 ± 0.32	1.24 ± 0.36

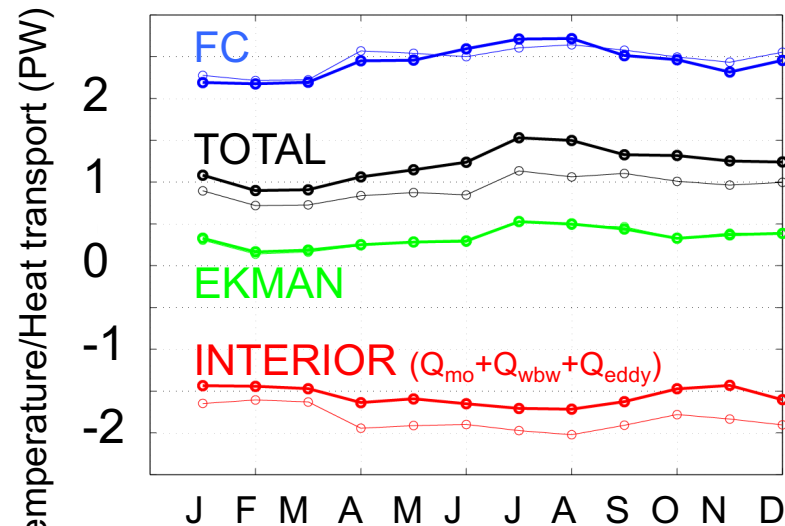
ORCA12 HT 0.3 PW lower than RAPID, but good agreement in the variability.

(Moat et al., JGR-Oceans, 2016)

RAPID MODEL comparison at 26.5°N



Climatological Seasonal Cycle



Model underestimates the MOC
(2004 to 2012)

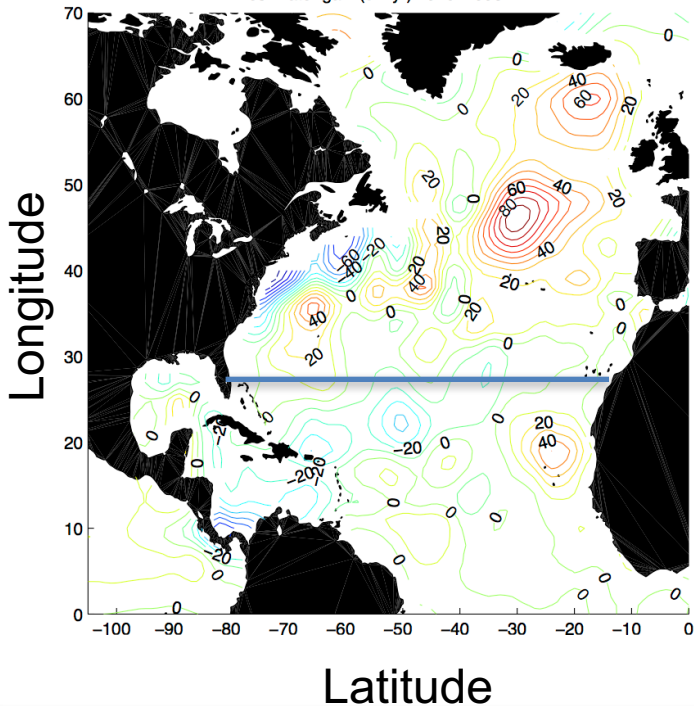
RAPID: 17.20 ± 4.60 Sv

ORCA083: 15.18 ± 4.03 Sv

Freshwater and heat storage (upper 1000 m)

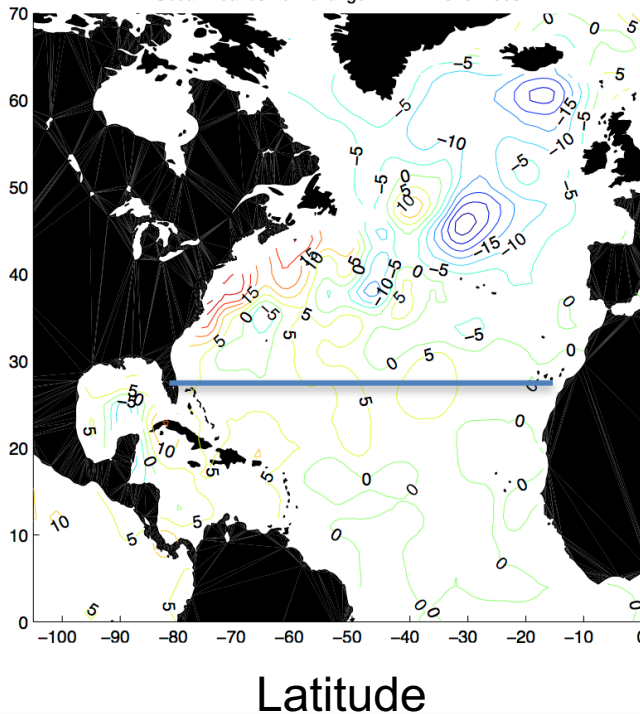
Freshwater Gain

Freshwater gain (cm/yr) 2016–2008 EN4



Ocean heat content change

Ocean heat content change W/m² 2016–2008



2008 to 2016

AMOC 26N lower by
2.5 Sv

**Freshening and cooling
In the region of the
North Atlantic current**

Why?

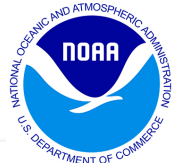
Slow down of the MOC.
Ocean current position
changing.
Surface fluxes.

Summary

Extended the fresh water time series at 26N to October 2015.
data will be made available soon.

Heat transport time series at 26N April 2004 to October 2015
available: https://www.rsmas.miami.edu/users/mocha/mocha_results.htm

Variability in the heat and freshwater transports is dominated by variability in the MOC



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