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Downscaling of Wind Resources From Mesoscale Tendencies with URANS

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The NEWA Model-Chain



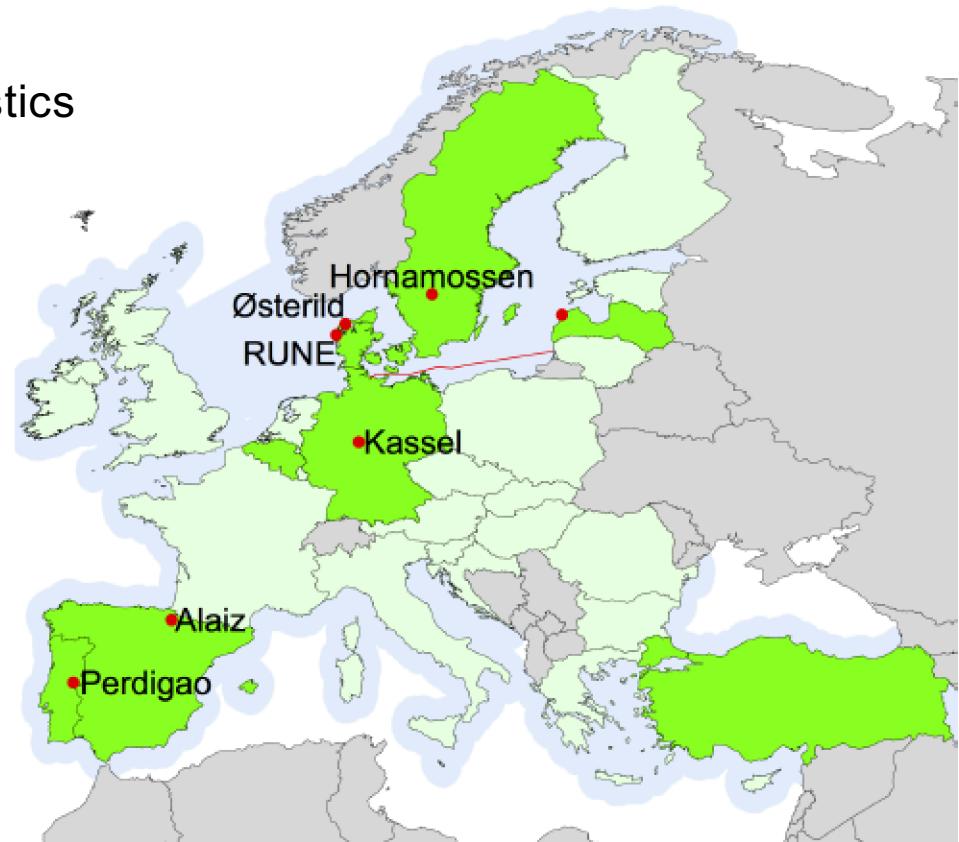
- ✓ Blending of mesoscale and microscale models
- ✓ Systematically validated with high-fidelity experiments over a range of conditions
- ✓ Uncertainty quantification

@mesoscale

- NEWA database of wind resource characteristics
- 30-year WRF down to 3 km resolution
 - + downscaling to 100 m using WAsP (generalized wind methodology)
- Mesoscale multiphysics ensemble (UQ)

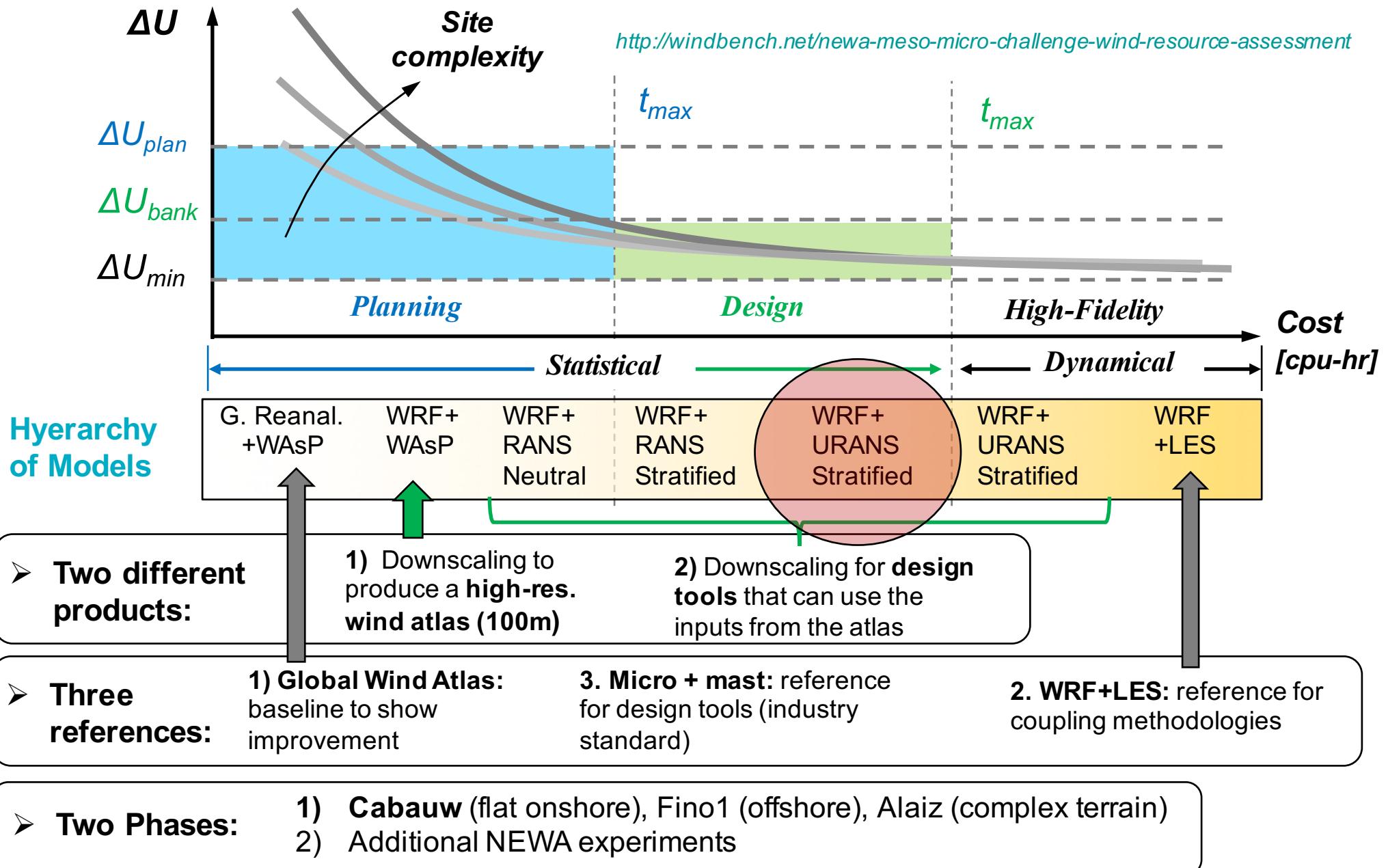
@microscale

- Design tools based on CFD (RANS or LES)
- Driven by (mesoscale) wind atlas inputs
- Include diurnal-cycle thermal stratification
- Surface-layer characterization based on aerial lidar scans of vegetation canopy
- **Open source** model built on OpenFOAM (NEWAfoam)



NEWA Meso-Micro Challenge

- Determine the applicability range of meso-micro methods for the NEWA validation domain



Design of the ABL model for Wind Energy applications

- Add 1st-order physics to the flow model
 - Turbulence closure through a modified k-eps as proposed by Sogachev et al. 2012
 - Thermal stratification through energy eq.
 - Transient model
 - Inspired in Bass et al. 2010, instead of Idealized forcing, Sanz-Rodrigo et al. 2014 proposed to include large-scale “**tendencies**”
- Avoid adding physical complexity if this is not justified by improved performance
 - No humidity equation
 - No heat transfer by radiation or phase changes
 - No land-surface modeling
- Make best use of onsite measurements to calibrate the model
 - **Nudging at microscale** using typical wind energy measurements

- Bass P, Bosveld FC, Lenderink G, van Meijgaard E and Holtslag AAM. *QJR Meteorol Soc* (2010) 136 671-684
- Sanz-Rodrigo et al. *J. Physics Conf. Ser.* (2014) 753.
- Sogachev A, Kelly MC, Leclerc MY. (2012) *Bound. Layer Meteorol*;145(2):307-327

Momentum Budget at Mesoscale: “Tendencies”

$$\frac{1}{f_c} \frac{\partial U}{\partial t} = -\frac{1}{f_c} \left(U \frac{\partial U}{\partial x} + V \frac{\partial U}{\partial y} + W \frac{\partial U}{\partial z} \right) + V - V_g - \frac{1}{f_c} \frac{\partial \overline{u'w'}}{\partial z}$$
$$\frac{1}{f_c} \frac{\partial V}{\partial t} = -\frac{1}{f_c} \left(U \frac{\partial V}{\partial x} + V \frac{\partial V}{\partial y} + W \frac{\partial V}{\partial z} \right) - U + U_g - \frac{1}{f_c} \frac{\partial \overline{v'w'}}{\partial z}$$

Tendency Advection Coriolis Horizontal Pressure Gradient Turbulence

$$U_{tend} \approx U_{adv} + U_{cor} + U_{pg} + U_{pbl}$$

Coriolis parameter:

$$f_c = 2\Omega \sin \phi$$

SCM with Large-Scale Tendencies and Data Assimilation

- Mesoscale forcing
- Solved at microscale
- Data assimilation at microscale

$$\frac{1}{f_c} \frac{\partial U}{\partial t} = U_{adv} + V + U_{pg} - \frac{1}{f_c} \frac{\partial \bar{uw}}{\partial z} + U_{nud}$$

$$\frac{1}{f_c} \frac{\partial V}{\partial t} = V_{adv} - U + V_{pg} - \frac{1}{f_c} \frac{\partial \bar{vw}}{\partial z} + V_{nud}$$

$$\frac{\partial \Theta}{\partial t} = \Theta_{adv} - \frac{\partial \bar{w\theta}}{\partial z} + \Theta_{nud}$$

$$\delta_{nud} = \frac{\omega_z}{f_c} \frac{(\delta_{obs} - \delta)}{\tau_{nud}}$$

$$\tau_{nud} = 10\text{min} : 1\text{hr}$$

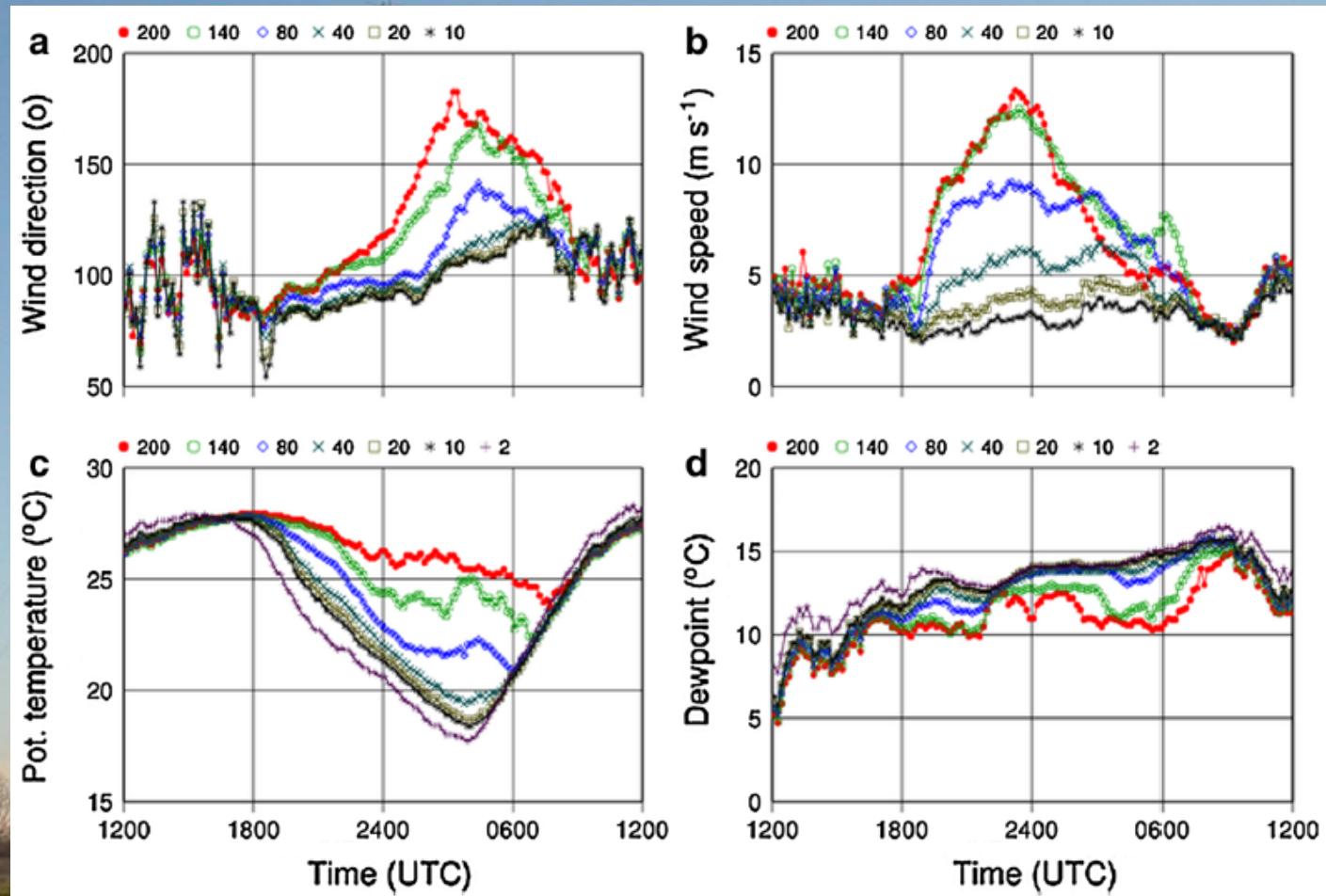
$\omega = 1$ within measurements

- M-O at the surface based on mesoscale inputs

$$\Theta_2 = \Theta_0 + \frac{\theta^*}{K} \left[\lg \left(\frac{2}{z_{0t}} \right) + \Psi_h \left(\frac{2}{L_0} \right) \right]; \quad \theta^* = -\frac{\bar{w'\theta'}}{u^*}$$

GABLs 3 (GEWEX Atmospheric Boundary Layer Studies)

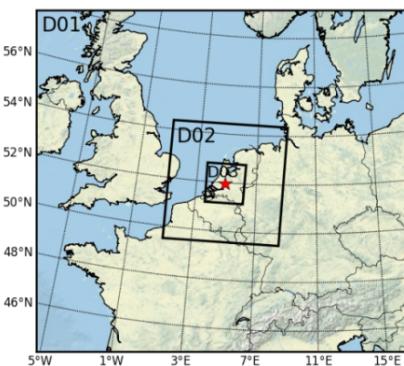
- Stationary synoptic
- Clear skies
- No fog
- Substantial LLJ
- 6 years → 9 days → night of 1–2 July 2006



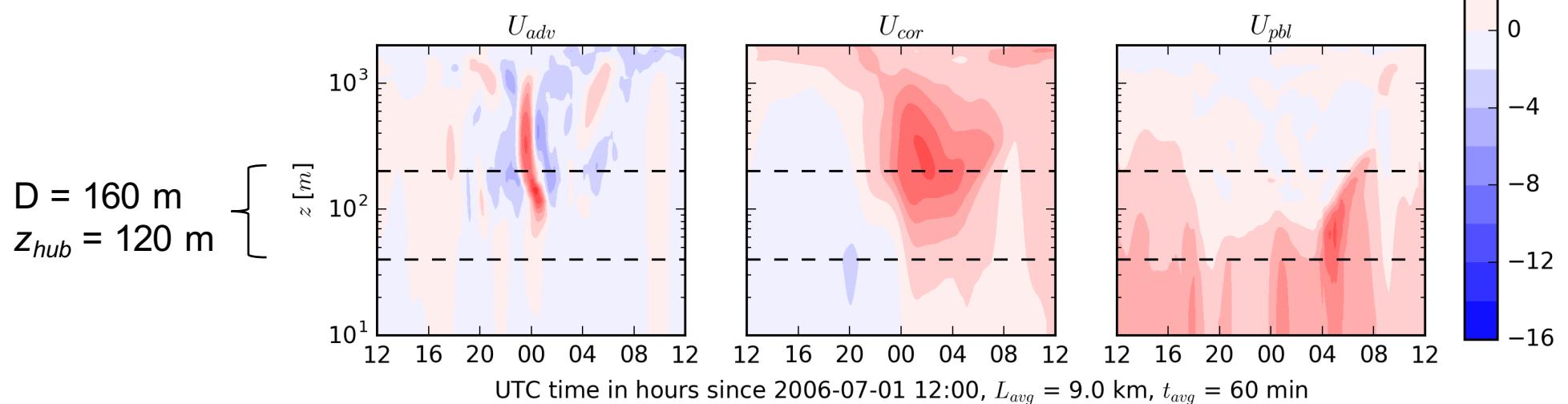
Mesoscale Tendencies from WRF during GABLS3

WRF-YSU

- ERA Interim
- $9 > 3 > 1$ km
- 24 hr spin-up
- Tendencies from d02 (3 km)
- $L_{av} = 3 \times 3 = 9$ km
- $t_{av} = 60$ min (rolling average)



$$U_{tend} \approx U_{adv} + U_{cor} + U_{pg} + U_{pbl}$$



GABLS3 Benchmark revisited for Wind Energy ABL models

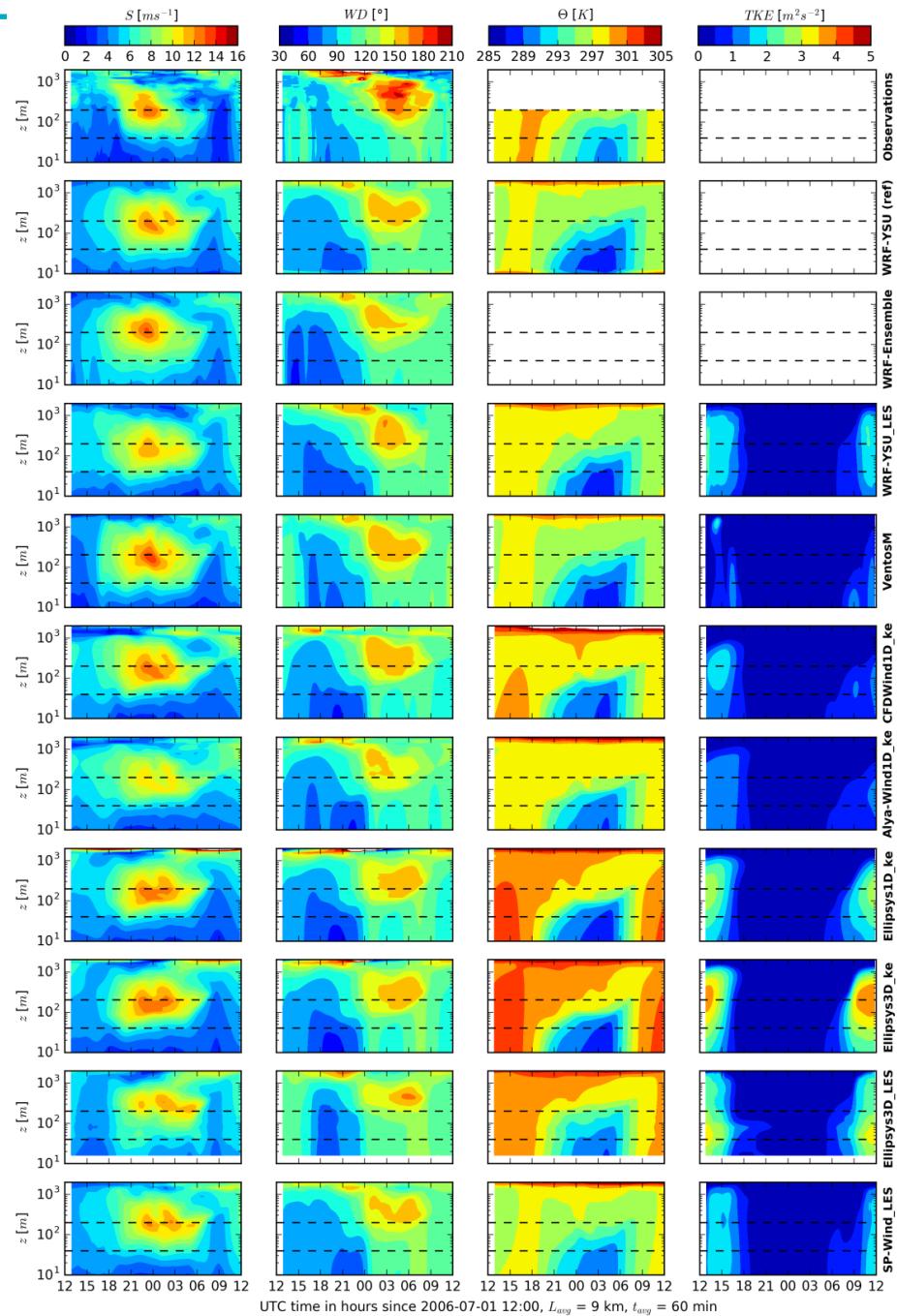
Table 1: Summary of model simulations. Monin-Obukhov similarity theory (MOST) surface boundary conditions use either heat flux (H), 2-m (T_2) or skin temperature (T_{SK}) from WRF

Name	Input	Turbulence	z-Levels	Surface B.C.
WRF-YSU (ref)	ERA Interim	YSU	46	Noah
WRF	ERA Interim, GFS	MYJ, MYNN, QNSE, TEMF, YSU	46	Noah
WRF-YSU_LES	ERA Interim	LES-TKE	101	Noah
WRF-VentosM_ke	ERA Interim	YSU/ $k-\epsilon$	70	MOST, H
CFDWind1D_ke	WRF (ref)	$k-\epsilon$	301	MOST, T_2
Alya-CFDWind1D_ke	WRF (ref)	$k-\epsilon$	500	MOST, T_2
Ellipsys1D_ke	WRF (ref)	$k-\epsilon$	512	MOST, T_{SK}
Ellipsys3D_ke	WRF (ref)	$k-\epsilon$	192	MOST, T_{SK}
Ellipsys3D_LES	WRF (ref)	Smagorinsky	128	MOST, T_{SK}
SP-Wind_LES	WRF (ref)	LES-TKE	500	MOST, T_2

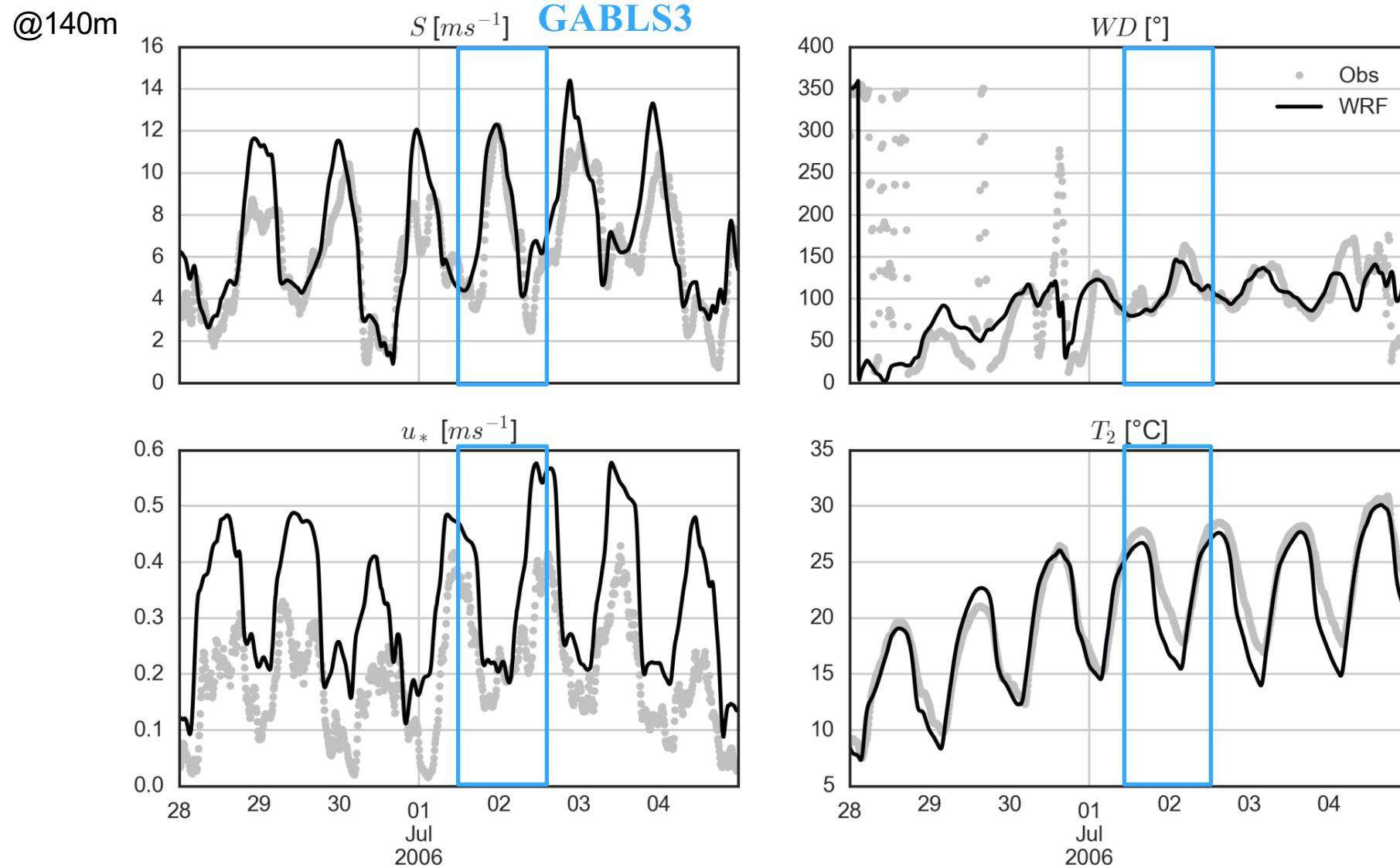
- A fare agreement among all models with the implementation of the tendencies.

Almost all RANS models in the benchmark are based on the modified k -eps proposed by Sogachev et al. 2012

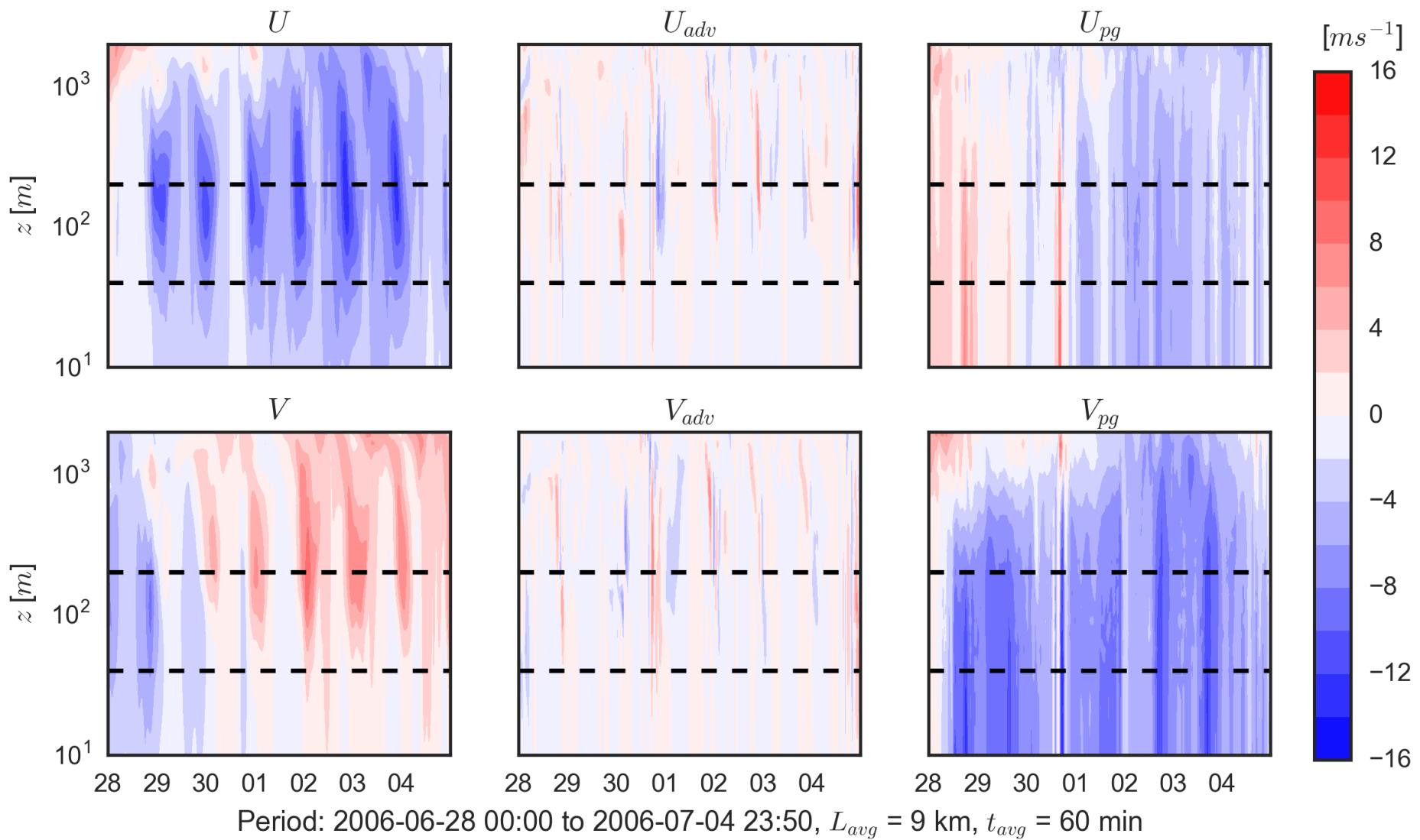
- *Sanz Rodrigo J, et al. (2017). J. Phys: Conf. Ser. 854 012037*
- *Sogachev A, Kelly MC, Leclerc MY. (2012) Bound. Layer Meteorol ;145(2):307-327*



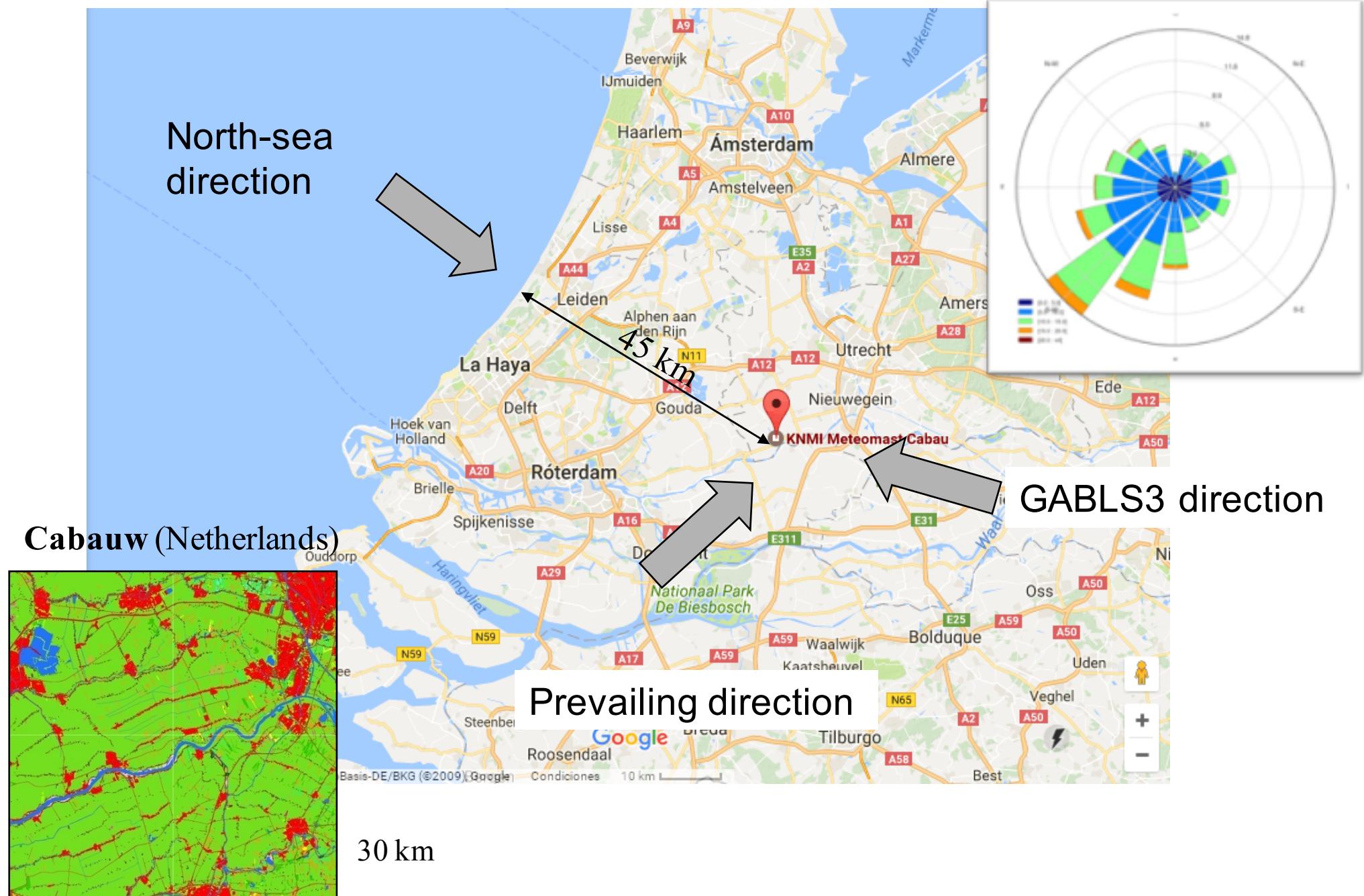
A typical week: GABLS3 direction



A typical week: GABLS3 direction tendencies

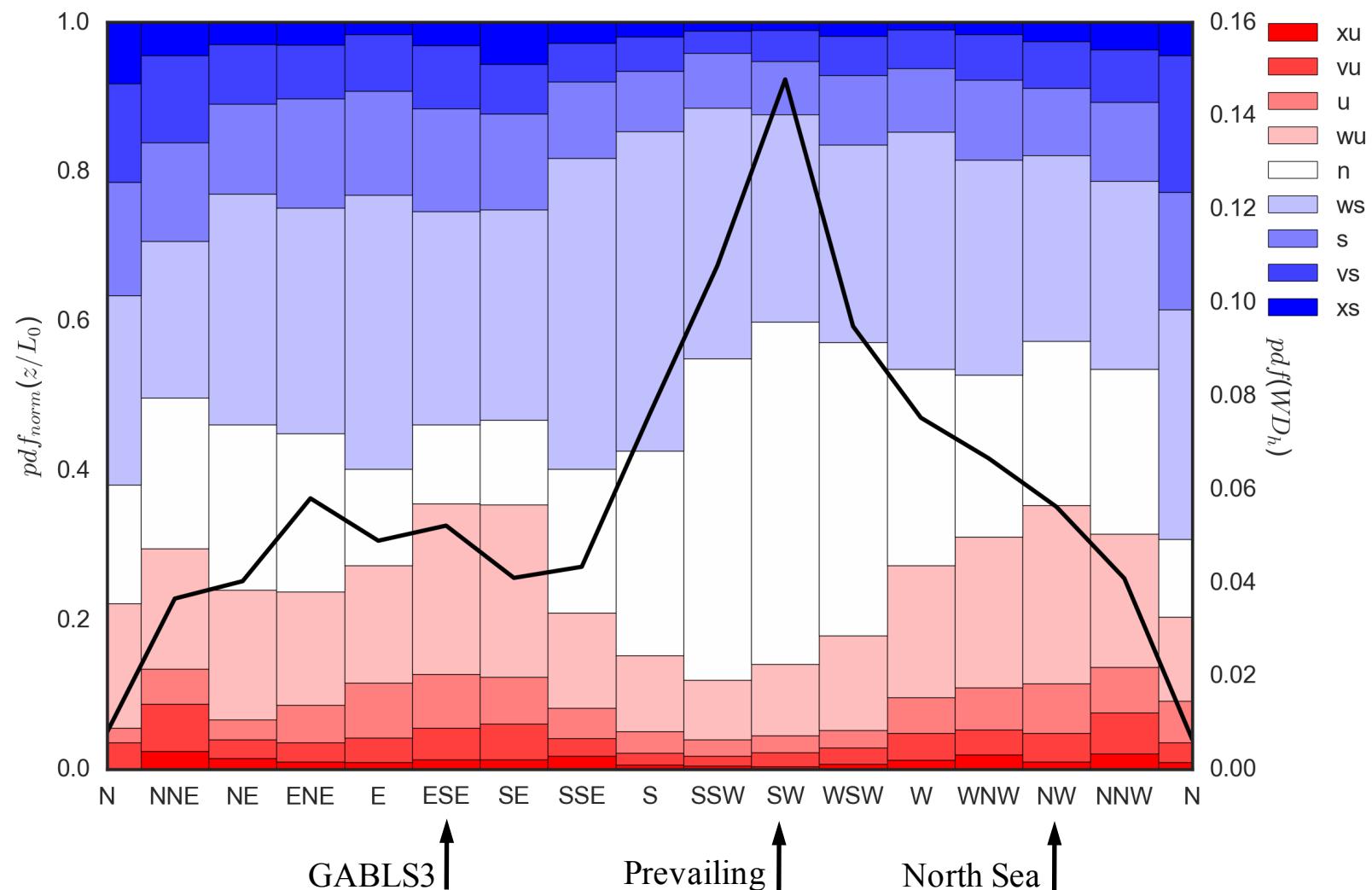


Cabauw year 2006: From flow case to wind climate characterization



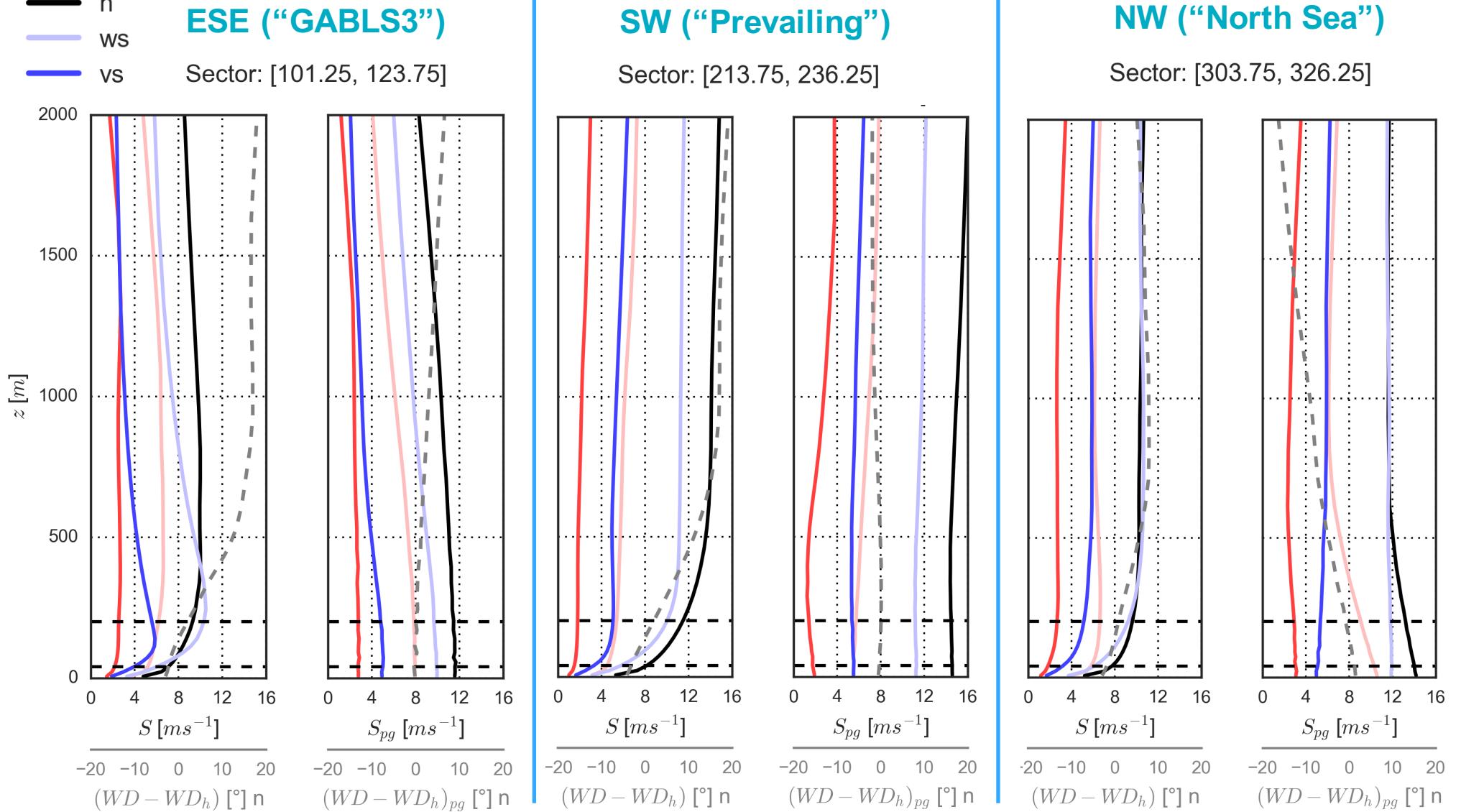
Wind Climate Classification: the common approach

Based on z/L at 10 m & velocity @ 140m (classification as used in Sanz-Rodrigo et al. 2014)



- vu
- wu
- n
- ws
- vs

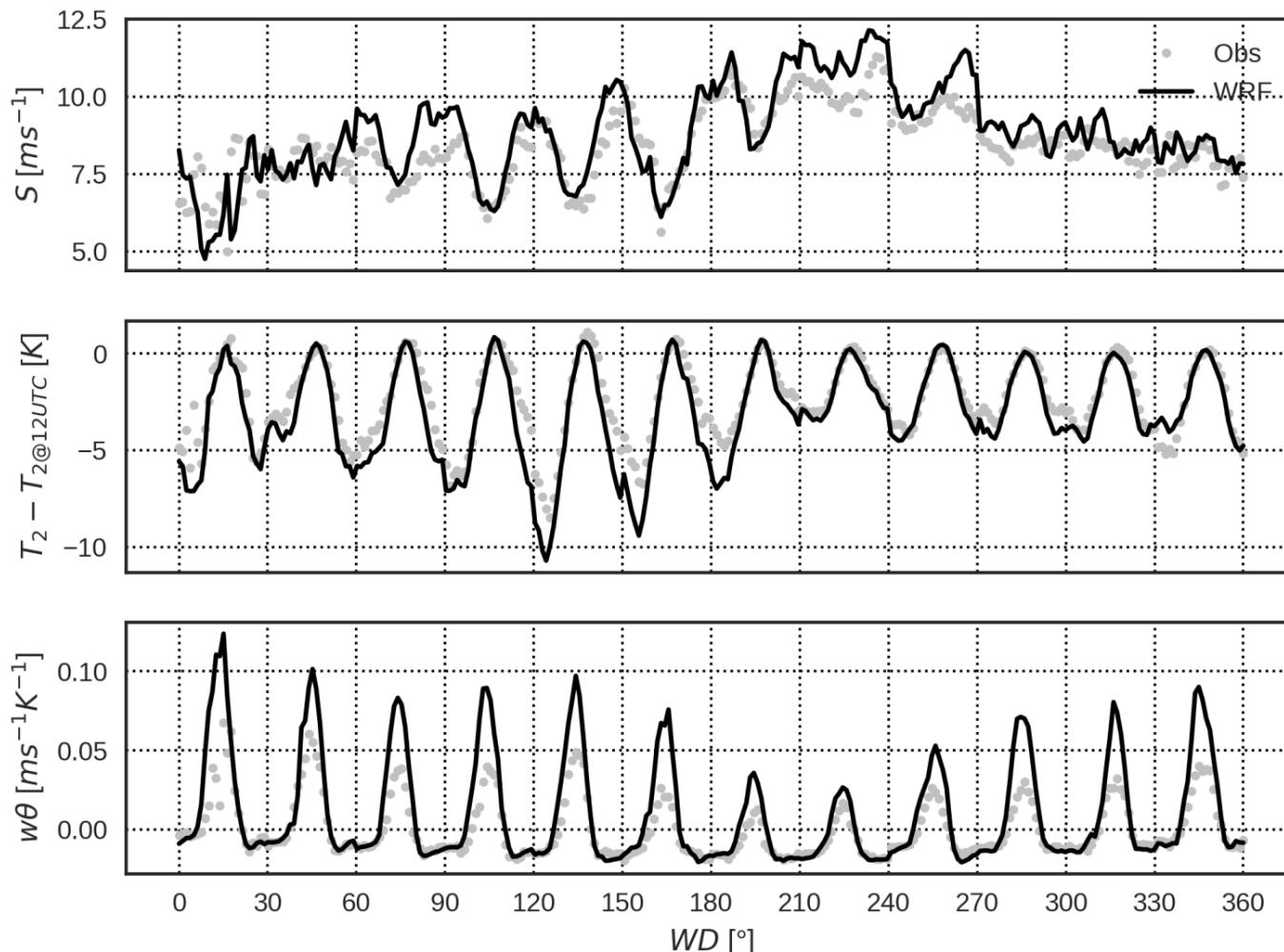
Annual Mean Profiles vs Stability



- As expected, mesoscale pressure gradient significantly change with direction.
- Its value changes significantly with stability and height.
- Therefore, it can provide important information missed by the microscale.

Sector-wise Annual Mean Diurnal Cycle

- However, binning has several drawbacks
 - We should be able to produce pseudo-stationary profiles consistent with the bined tendencies
 - Loss of time-dependent information from previous cycles
- Alternative to binning by stability (steady approach) to retain diurnal cycle dynamics (unsteady approach)



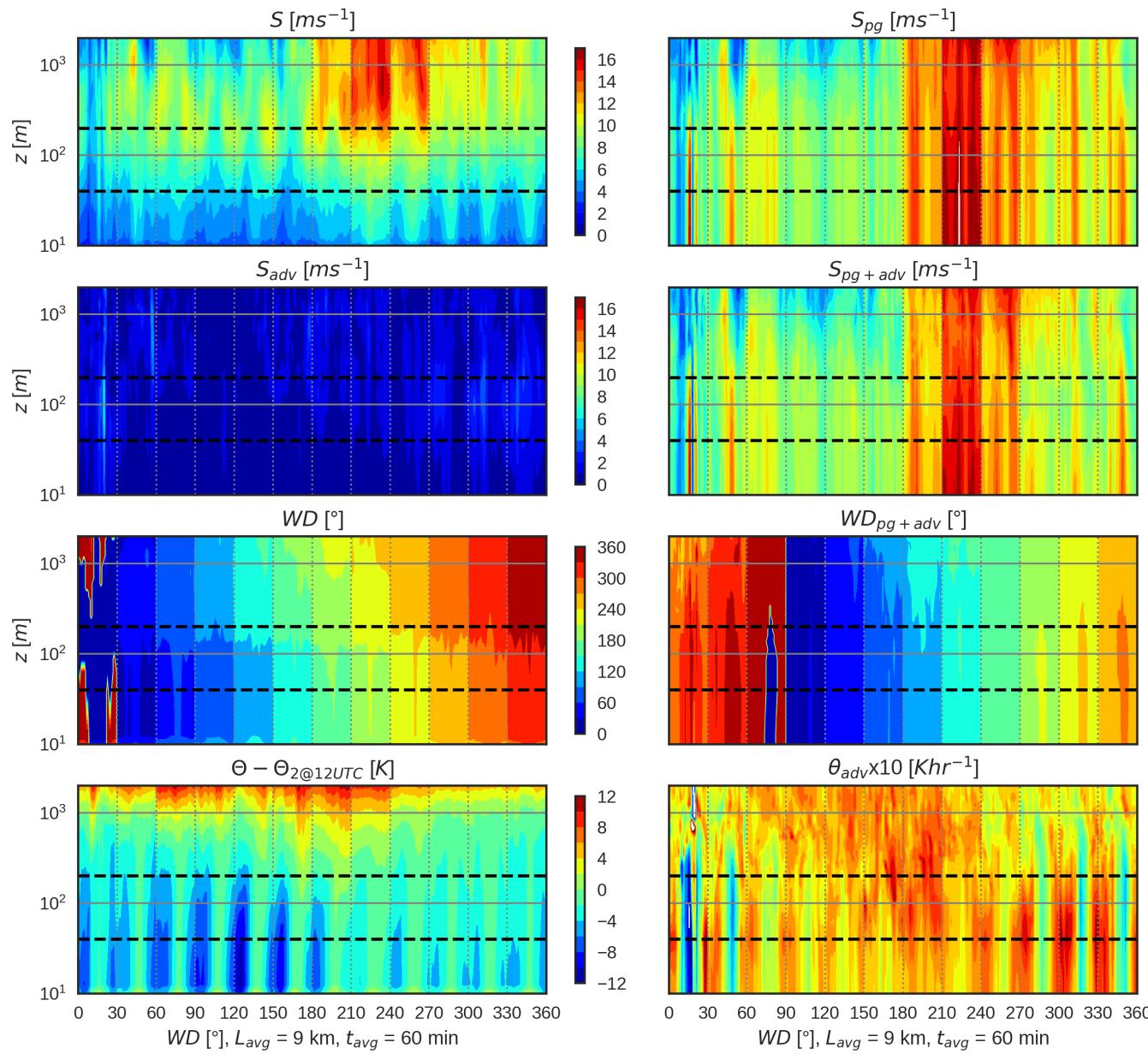
$Z_{ref} = 140 \text{ m}$

30° bins

$S_{140} > 4 \text{ m/s}$

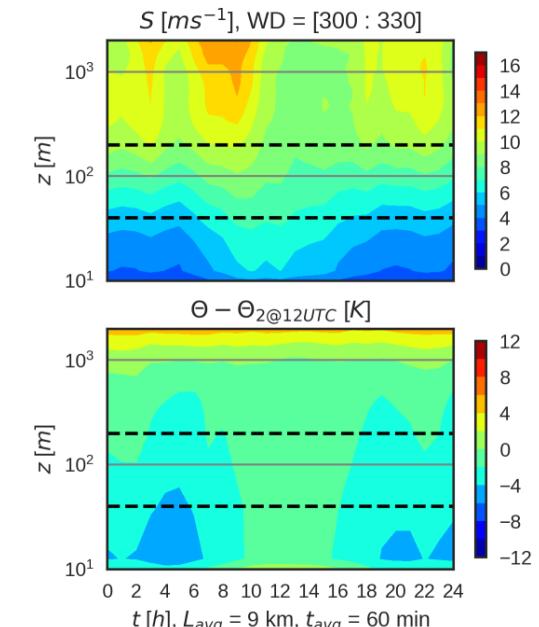
$T = 12 \times 24 = 288 \text{ h}$

Sector-wise Annual Mean Diurnal Cycle



$z_{ref} = 140 \text{ m}$
 30° bins
 $S_{140} > 4 \text{ m/s}$

“North Sea”



Conclusions & Outlook

So far...

Under NEWA framework a :

- Meso-micro offline coupling method verified with GABLS3
- Methodology for annual wind resources based on 1D sector-wise mean diurnal cycles now under development for the “NEWA Meso-Micro Challenge”

Outstanding Challenges

- From statistics to “suitable” model forcings (smoothing, nudging, etc)
- 3D methodology for complex terrain: nesting 3D tendencies and boundary conditions
- Optimization of 360 cycle using surrogate model and other statistical methods
- $k-\varepsilon$ model of Sogachev et al. (2012) seems to require a calibration (I_{max} in complex terrain, etc)

Acknowledgements

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- **MesoWake**, European Comission FP7-PEOPLE-2013-IOF, 624562, <http://www.windbench.net/mesowake-2014-2017>
- **NEWA**, European Comission FP7-ENERGY-2013.10.1.2, 618122, <http://www.neweuropeanwindatlas.eu/>; MINECO, Spain, PCIN-2014-011-C07-02
- **IEA Task 31 Wakebench Phase 2**, International Energy Agency
- **PRACE-MesoWake**, PRACE 13th Call, MareNostrum, Barcelona



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