PISCES

(Pelagic Interaction Scheme for Carbon and ecosystem studies)



PISCES

Basic information on the model options

PISCES-std vs. PISCES-QUOTA

ln_p4z = .true. PISCES-std (24/25 tracers)



PISCES-std vs. PISCES-QUOTA

ln_p5z = .true. PISCES-QUOTA (39/40 tracers)



PISCES-simple vs. PISCES-STD

ln_p2z = .true. PISCES-Simple (9 tracers)



PISCES-Sediment

Can be used online/offline/standalone



Code structure - Main



Code structure - SMS



Modeling platforms



Parameterizing marine biogeochemistry

Olivier Aumont

LOCEAN

Gp Ζ Ρ E Gd PP Mz Mp N D R S dynamics Conservation:

S(P)+S(Z)+S(D)+S(N)-P(P)-P(Z)-P(D)-P(N)=0

(1) Phyto S(P)=PP P(P)=Gp+Mp(2) Zoo S(Z)=Gp+GdP(Z)=Mz+E(3) Detritus S(D)=Mz+Mp P(D)=Gd+S+R(4) Nutrients S(N)=R+EP(N)=PP



Photosynthesis, growth rate

Photosynthesis

Photosynthesis: Process by which autrophic organism use solar energy to produce organic matter

 $106CO_2 + 16 NO_3^- + HPO_4^{2-} + 122H_2O + 18H^+ + trace elements, light --> C_{106}H_{263}O_{110}N_{16}P + 138O_2$

- The ratio between the different chemical elements is called the Redfield ratio
- The amount of organic matter produced by the photosynthesis is called Gross Primary Production

Growth rate

Growth refers to the increase in biomass (in C units generally)

$$\mu = P^B \frac{Chl}{C} - r$$
 Where r is the respiration rate

From the growth rate, net primary production can be defined as the accumulation of organic matter

Growth rate: General background

- Growth rate is a function of the environmental and biogeochemical conditions and of the species
- It can be expressed as follows:



Growth rate: Temperature

Growth rate increases with temperature until a critical level



A relationship for the enveloppe has been proposed for the first time by Eppley (1972) :

$$f(T) = 1.066^{T}$$

Growth rate increases by 1.9 times every 10°C (Q10).

Since, several alternative expressions have been proposed, but Eppley's relationship remains the most commonly used one.

Growth rate: Light

Growth rate increases with light until a maximum value at which it saturates or even decreases



 α initial slope μ_{max} maximum growth rate $I_{k} = \mu_{max} / \alpha$

$\mu = \mu_{max}$	1–exp	$\left(\frac{-I}{I_k}\right)$	
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I_k is extremely variable between species. For instance, in cyanobacteria, synechococcus spp have a high I_k whereas some prochlorococcus spp have very low I_k.

Ik strongly varies with the average received light (photoacclimation)

Growth rate: Monod model (1942)

Monod model = growth rate is a function of the external concentration of nutrients



K_N increases with size because the S/V ratio decreases. Furthermore, as a result of acclimation processes, K_N decreases with the nutrient concentration



Limitation of the Monod model

- Measured & works best under relatively steady nutrients (or slow change)
- Growth stops when nutrients fall to 0



- Assumes constant stoichiometry
- No luxury uptake of transiently elevated nutrients
- Can be difficult to estimate K_N

Growth rate: Droop model (1968)

- Droop model = growth rate is a function of the internal pool of nutrients (quota)
- The internal pool (quota) is a function of the external concentrations of N



Generally works better, more physiologically grounded, more general



Growth rate: multi-nutrients

Currently, there is no clear consensus on the law which drives growth rate with multiple nutrients. 2 different laws are generally used:

The multiplicative law:

$$\mu = \mu_{max} \frac{N_1}{K_{N_1} + N_1} \frac{N_2}{K_{N_2} + N_2} * \dots$$

The law of the minimum or Liebig's law (1840) :

$$\mu = \mu_{max} \min\left(\frac{N_1}{K_{N_1} + N_1}, \frac{N_2}{K_{N_2} + N_2}, \dots\right)$$

Many other laws do exist but they are not commonly used.

Light in the ocean



Light decreases with depth as (Loi de Beer): $I(z) = I(0)e^{-kz}$

- In pure water, the attenuation length is 37m. for a chlorophyll concentation of 0.2 mg Chla/m³, its value is about 20m.
- Blue light penetrates much deeper than red light which remains trapped in the top 10 to 20m of the ocean.

Light in the ocean

Euphotic zone

l > 1% l_o





Chlorophyll



Predation by zooplankton



Holling Type I

$$G_p = gP$$

Holling Type II



Holling Type III



Predation on several preys

Very complex to properly model!



Mitra et al., 2006



Mortality in models does not necessarily represent senescence. It may model :

- senescence
- viral attacks
- aggregation/sinking
- predation by unresolved higher trophic levels

Numerous formulations exist but the two most common expressions are:

$$M = m_p P$$
 $M = m_p P^2$

Representing particles in models



Representing particles in models: POC

If bacterial degradation and sinking are the only active processes



Remineralization

Specific cases :

• If v and λ are constant : F=vD=

$$F = vD = F(\widetilde{z})e^{-\frac{\lambda}{v}(z-\widetilde{z})}$$

• If
$$\lambda$$
 is constant and $v = Az$ $F = vD = F(\widetilde{z}) \left(\frac{z}{\widetilde{z}}\right)^{(-b)}$ where $b = \frac{\lambda + A}{A}$

Observed fluxes vs. modeled fluxes



Fluxes seem to be well approximated by a power law function

But far from being that simple



From Marsay et al. (2015)

A simple set of 4 equations

$$(S-P)(P) = \mu_{max} L_{NUT} L_I P - g \frac{pP}{K_z + pP + (1-p)D} Z - m_p P$$

$$(S-P)(Z) = ga \frac{pP + (1-p)D}{K_z + pP + (1-p)D} Z - m_z Z^2$$

$$(S-P)(D) = g(1-a) \frac{pP + (1-p)D}{K_z + pP + (1-p)D} Z + m_z Z^2 + m_p P$$

$$(B - I)(D) = g(I - u)K_{z} + pP + (1 - p)D$$

$$-g\frac{(1 - p)D}{K_{z} + pP + (1 - p)D}Z - v\frac{\partial}{\partial z}D - t_{d}D$$

$$(G - D)(M) = K - L - D$$

 $(S-P)(N) = -m_{\max} L_{NUT} L_I P + t_d D$

Non-linear equations

Parameters needed to constrain the model

$$\begin{array}{ll} \mu_{\max}, K_N, K_I, m_p & \text{phytoplancton} \\ g, K_z, p, a, m_z & \text{zooplancton} \\ v, t_d & \text{détritus} \end{array}$$

Difficult estimation

- lab experiments, species dependant, equilibrium state
- large variability: non constant
- agregate many processeses
- inverse methods (data assimilation)
- empirical estimation

Evaluation of a model





Tuning the models



- Can be very painful
- The more complex the model is, the more difficult it is





We learn a lot on the model dynamics/behavior

Optimizing the parameters : Assimilation



Variational assimilation



Example: Simultaneous assimilation at 5 1-D stations: Chla, NO3, POC et Si

45 optimized parameters



Kane et al., 2011

Optimizing parameters



Kane et al., 2011

Can be difficult (not the universal cure) !

- Are stations representative of the system?
- What should be done with non assimilated variables?
- Some parameters may not be well constrained by available data

Models Intercomparison Projects (MIPs)

Motivation of these exercises

- Evaluate and intercompare the participating models
- Identify the converging and diverging behaviors and stimulate the model developments
- Estimate the uncertainties, for instance in projections

Ocean biogeochemistry: 1st exercise started in 1995

• 4 participating groups : IPSL, GFDL, MPIM, Hadley



- Carbon cycle (ocean models only) : OCMIP-2, OCMIP-3, OMIP
- Iron cycle : FeMIP
- Carbon cycle (Earth System Models) : CMIP, C4MIP
- Marine ecosystems : FISHMIP