# Tropospheric circulation during the early 20th century Arctic warming

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### Background

The early twentieth century Arctic warming (ETCAW) between 1920-1940 is an exceptional feature of climate dynamics in the last century and its warming rate was only recently matched by anthropogenic global warming amplification in the Arctic region.

However, atmospheric warming during the ETCAW was strongest in the winter mid-troposphere of the European sector and is believed to be triggered by an exceptional **case** of natural climate variability.



<sup>1880-2014</sup> with respect to 1951-1980

Beitsch et al. 2014 analysed the climatic conditions of 26 Arctic warming events within an Earth System model. They found a strong increase of stationary atmospheric energy **transport** into the Arctic during the warming event, whereas transient and mean meridional energy transports decrease.

The case of the ETCAW underlines the importance of yearly and decadal internal **variability** in Arctic climate evolution.



How important is internal variability for the formation of this warming event?



What are the atmospheric conditions and dynamics to sustain this warming?



How well do recent centennial reanalysis datasets capture the event compared to general circulation models?



*Is it possible to quantify mid-tropospheric* heat transport into the Arctic?

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## **Results for Arctic winter at 700 hPa**





All four observation-based datasets agree very well in temperature magnitude and correlation, staying within the variability of the models. All datasets show a mid-tropospheric temperature signal during the ETCAW, which is unique in magnitude until the 21st century. In general, GCMs, reanalyses, and reconstructions values match well.

lantic and Pacific domain.

Reanalyses and reconstruction depict a strong positive geopotential height anomaly over the Eurasian part of the Arctic with negative/weaker anomalies over Greenland. This pressure setup enhances meridional winds over the North Atlantic and transports southern airmasses into the Arctic domain.

# **Atmospheric circulation setting** 15 20 25 -25 -20 -15 -10 10

left: 700 hPa geopotential height anomalies for 1920-1939 with reference to 1971-2000 for early winter (DJF) in a) REC1+ERA40, b)20CRv2, c) 20CRv2c, d) ERA20C, e) ERA20CM ensemble mean, f) CCC400 ensemble mean. right) 700 hPa Geopotential height circulation index values for top) Atlantic sector, bottom) Pacific sector

#### **Arctic temperature evolution**

and ERA20CM, as well as the reanalyses 20CRv2, 20CRv2C, ERA20C and the reconstruction REC1.

• We find an increase of **stationary edddy heat transport** at the timing of the warming in the REC1 and 20CRv2 datasets. This peak also fits the increased index values in the 1920s decade over the At-

Moreover, we find a decrease of mean meridional heat flux right before the warming in the independent upper-air reconstruction. The 1920s and 1930s show combined the lowest decadal values of transient eddy heat flux during the whole 20th century.



but only for reanalysis datasets. in all three reanalysis datasets.



datasets, b) the same for stationary heat flux and c) the same for transient eddy heat flux,

Mean meridional flux in reanalyses metric might be unreliable due to surface wind errors

# Conclusions

• Our results support a **meridionalisation of circulation** during the ETCAW. Surface and 700 hPa temperatures in **reanalyses agrees very well with reconstructed temperatures**.

• SST and sea ice forcings are not sufficient to trigger the spatial pattern of the ETCAW. Instead observational input is needed to compute the realistic circulation and heat flux response. Thus it can be concluded that the intrinsic atmospheric variability played a major part in the formation of the ETCAW signal.

• Utilizing the 700 hPa heat transport as surrogate for tropospheric processes, it could be shown that reanalysis and reconstruction datasets show peak values of stationary heat **flux for the ETCAW**. The independent reconstruction shows a decrease of mean meridional heat flux with a negative lag of about 5 years and the analysis of reanalyses exhibits a decrease of transient eddy heat flux into the Arctic domain (agreement with Beitsch et al 2014).

### Setup



### Methods

• We investigate the cold season (DJF) with a focus on the 700 hPa level as a surrogate for mid-tropospheric processes. During boreal winter, temperature differences between polar and subpolar airmasses are strongest and Bekryaev et al. 2010 found that boreal winter together with autumn showed the strongest warming signal. Moreover, Overland and Turet 1994 stated that the poleward energy transport is maximized between 800 and 600 hPa.

We compute different components of northward heat transport at 700 hPa and 60° N in the GCM, reanalysis and reconstruction datasets. The zonal mean northward heat flux at a certain level can be written as



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100	150	Years into the past	
REC1 T+GPH			
20CRV2			
0CRv2c			
RA-20C			
ECHAM 5.4 (CCC400) X	30		
0CM) X 10			
100	150	Years into the past	
constructions (Gries	ser et al. 2010)		

**20CRv2** (Compo et al. 2011) and 20CRv2c (Cram et al. 2015) reanalysis

**ERA-20C** reanalysis (Poli et al. 2015) **and ERA-20CM** (Hersbach et al. 2005) GCM ensemble

• ECHAM5.4 ensemble GCM run (CCC400) (Bhend et al. 2012)

			Deviation from zonal average		
	Mean meridional circulation flux	Stationary Eddy flux	Transient Eddy flux	Deviation from time average	
vT] =	[v][T]+	[v*T*]	+[vT']		