

Short communication

# Elevated atmospheric CO<sub>2</sub> concentration and temperature across an urban–rural transect

K. George<sup>a,b,\*</sup>, L.H. Ziska<sup>a</sup>, J.A. Bunce<sup>a</sup>, B. Quebedeaux<sup>b</sup>

<sup>a</sup>*Crop Systems and Global Change Laboratory, USDA-ARS, Building 001, Room 342, 10300 Baltimore Avenue, Beltsville, MD 20705, USA*

<sup>b</sup>*Department of Plant Sciences and Landscape Architecture, University of Maryland, College Park, MD 20742, USA*

Received 2 April 2007; received in revised form 23 July 2007; accepted 10 August 2007

## Abstract

The heat island effect and the high use of fossil fuels in large city centers are well documented, but by how much fossil fuel consumption is elevating atmospheric CO<sub>2</sub> concentrations and whether elevations in both atmospheric CO<sub>2</sub> and air temperature from rural to urban areas are consistently different from year to year are less well known. Our aim was to record atmospheric CO<sub>2</sub> concentrations, air temperature and other environmental variables in an urban area and compare it to suburban and rural sites to see if urban sites are experiencing climates expected globally in the future with climate change. A transect was established from Baltimore city center (Urban site), to the outer suburbs of Baltimore (suburban site) and out to an organic farm (rural site). At each site a weather station was set-up to monitor environmental variables for 5 years. Atmospheric CO<sub>2</sub> was consistently and significantly increased on average by 66 ppm from the rural to the urban site over the 5 years of the study. Air temperature was also consistently and significantly higher at the urban site (14.8 °C) compared to the suburban (13.6 °C) and rural (12.7 °C) sites. Relative humidity was not different between sites whereas the vapor pressure deficit (VPD) was significantly higher at the urban site compared to the suburban and rural sites. An increase in nitrogen deposition at the rural site of 0.6% and 1.0% compared to the suburban and urban sites was small enough not to affect soil nitrogen content. Dense urban areas with large populations and high vehicular traffic have significantly different microclimates compared to outlying suburban and rural areas. The increases in atmospheric CO<sub>2</sub> and air temperature are similar to changes predicted in the short term with global climate change, therefore providing an environment suitable for studying future effects of climate change on terrestrial ecosystems.

Published by Elsevier Ltd.

**Keywords:** Microenvironment; Climate change; Urban ecology; Heat island; Long term

## 1. Introduction

The conversion of rural lands into urban areas with high traffic volumes and dense residential and

commercial buildings greatly affects the local air quality and energy balance. Cities are large consumers of fossil fuels, because in the US as in other countries, the majority of the population resides in urban areas (United Nations, 2004). This results in urban areas being responsible for the largest proportion of anthropogenic emissions such as CO<sub>2</sub> and nitrous oxides (Pataki et al., 2006). Human and vehicle activity has been found to contribute

\*Corresponding author. Crop Systems and Global Change Laboratory, USDA-ARS, Building 001, Room 342, 10300 Baltimore Avenue, Beltsville, MD 20705, USA.  
Tel.: +1 301 504 5527; fax: +1 301 504 5872.

E-mail address: [kate.george@ars.usda.gov](mailto:kate.george@ars.usda.gov) (K. George).

more than 80% of the atmospheric CO<sub>2</sub> in urban areas (Koerner and Klopatek, 2002). Also, the conversion of natural land into roads and buildings changes the albedo and heat capacity of an area resulting in urban areas being significantly warmer than if it remained as a rural landscape (Oke, 1982). If urban environments are elevated in atmospheric CO<sub>2</sub> and temperature for sustained periods this could provide a suitable system for studying the effects of future global climate change on plant population and community dynamics without the high cost of equipment and other resources currently needed to carry out such studies.

The elevation of air temperatures in city centers compared to less built-up areas is a phenomenon that has been recorded since 1833 (Oke, 1982). The main factors contributing to the urban heat island are the high thermal conductivity of buildings and other man-made structures, the low albedo and geometry of city surfaces and low evapotranspiration (discussed in Taha et al., 1991). Heat released directly from building ventilation and vehicular traffic also contribute to city heating and can vary annually and diurnally (Fan and Sailor, 2005). The degree of heating within a city is very variable and is unique to each city based on its location, building layout and traffic (Oke and Maxwell, 1975). A city wide examination of ground level air temperature in Baltimore, USA found that minimum temperatures are closely related to population and the difference between urban and rural minimum temperatures has been increasing as population increases (Brazel et al., 2000). The densely built-up areas in the center of Baltimore had ground level air temperatures 5–10 °C warmer than residential or forested and agricultural areas (Brazel et al., 2000), making it a good model to study climate differences between urban centers and adjacent rural areas.

Near-surface CO<sub>2</sub> concentrations have been documented in several cities across the world (Vancouver, Canada; Kuwait City, Kuwait; Mexico City, Mexico; Basel, Switzerland; Nottingham, UK; Phoenix, USA) to evaluate the dynamics of atmospheric CO<sub>2</sub> over short periods of time (Berry and Colls, 1990; Reid and Steyn, 1997; Idso et al., 2001; Nasrallah et al., 2003; Velasco et al., 2005; Vogt et al., 2006). The majority of these studies analyzed daily and diurnal fluctuations in CO<sub>2</sub> concentrations and concluded that the major source of CO<sub>2</sub> is from vehicular traffic as peak CO<sub>2</sub> concentrations

correlate to high traffic volume during workdays and is significantly reduced at weekends (Idso et al., 1998, 2001, 2002; Nasrallah et al., 2003; Velasco et al., 2005). In Nottingham, UK an 8 month study found a small difference in CO<sub>2</sub> concentration of only 5 ppm between a rural location and an urban city site (Berry and Colls, 1990), although the small difference between sites and large variability in CO<sub>2</sub> concentrations are likely related to the close proximity of power stations to both sites which were only 15 km apart (Berry and Colls, 1990). In Phoenix, USA CO<sub>2</sub> concentration was monitored for nearly a year and values ranged from a daily minimum of 390 ppm rising to a daily maximum of 491 ppm, although a maximum value of 619 ppm was attained (Idso et al., 2002). Over a two week period CO<sub>2</sub> concentration varied significantly from day to day with the highest peak of 650 ppm which was 76% higher than the low of 369 ppm (Idso et al., 2001). Whereas Day et al. (2002) found the urban area was elevated by 19 ppm compared to a suburban area, but their study area was a distance from major streets and less influenced by vehicle emissions. Few studies have concurrently compared urban to rural CO<sub>2</sub> concentrations to determine the amount by which CO<sub>2</sub> concentrations are elevated by urbanization and whether any increases are sustained and consistent from year to year.

Urban areas are affecting the microclimate, but few studies have recorded these changes for long periods of time to ensure the consistency of data and suitability for investigating effects on plant biological systems, nor monitored other global climate change variables concurrently. The aim of this study was to investigate whether a high population city center has a climate similar to that predicted in the short term (50–100 years) with global climate change. Baltimore was selected as it is one of the largest cities in the USA with dense residential and commercial buildings and high traffic volumes in the city center. The outskirts of Baltimore become more suburban with green areas on the outskirts of the city and becoming rural dominated by agricultural land. This location is ideal for comparing microclimate changes from an urban city center transitioning to a more suburban and rural areas. The objectives of this study were to characterize the microenvironment associated with an urban location relative to a suburban and rural location. A secondary objective was to compare the microenvironmental characteristics to

climatic conditions predicted with global climate change.

## 2. Methods

### 2.1. Site description

A transect was established running west from downtown Baltimore city to a rural agricultural area in western Maryland. Three sites were selected along the urban–rural gradient: an organic farm near Buckeystown, Maryland ( $39^{\circ}18'N$   $77^{\circ}26'W$ , elevation 109.8 m) approximately 87 km west of Baltimore (rural site), a nature center approximately 11 km west of Baltimore ( $39^{\circ}18'N$   $76^{\circ}41'W$ , elevation 98.9 m) on the outer edge of the city (suburban site), and Baltimore city center ( $39^{\circ}16'N$   $76^{\circ}36'W$ , elevation 6.8 m; urban site). Each site is surrounded by grass, which is mowed frequently through the growing season. The urban site is surrounded by large commercial and residential buildings and is very close to a large body of water within the city center (Fig. 1C). The area is bordered by busy city roads less than 0.1 km in distance from the site with annual average daily traffic of 58,083 vehicles (only 2005 and 2006 data available, Maryland State Highway Administration). The suburban site is surrounded by trees as it is part of the Carrie Murray nature center within the Gywnn Falls Park (Fig. 1B). The park area is approximately 480 ha and is surrounded by housing and lawns. The nearest roads are 1.5 km with annual average daily traffic of 33,716 vehicles (2002–2006, Maryland

State Highway Administration). The rural site is located on an organic farm (Fig. 1A), which predominantly grows alfalfa and orchard grass for animal feed. The area is dominated by agricultural land mainly for grazing with a few residences scattered across the landscape. The nearest roads to the site are 1.1 km away and the annual average daily traffic is 5298 vehicles (2002–2006, Maryland State Highway Administration). At each site four plots were established  $2 \times 2 m^2$  in size, each containing the same uniform fallow agricultural soil to a depth of 1.1 m and extant seed bank. The plant communities in each plot were allowed to establish naturally, further site descriptions and results are described in Ziska et al. (2003, 2004).

### 2.2. Site microenvironmental measurements

At each site a weather station was established that monitored the following variables with a 15 min averaging interval using a CR10X data logger (Campbell Scientific, USA): air temperature and relative humidity (CS500, Campbell Scientific, USA), atmospheric  $CO_2$  concentration (S151, Quibit, Canada), soil temperature (Model 107, Campbell Scientific, USA) and moisture (Echo EC-20, Decagon Devices, USA), wind speed and direction (Model 03001, R. M. Young Company, USA), and total and diffuse radiation (Sunshine sensor BF3, Delta-T Devices, UK) and precipitation (Tipping bucket rain gage TE525, Texas Instruments, USA). The sensors had a sky view

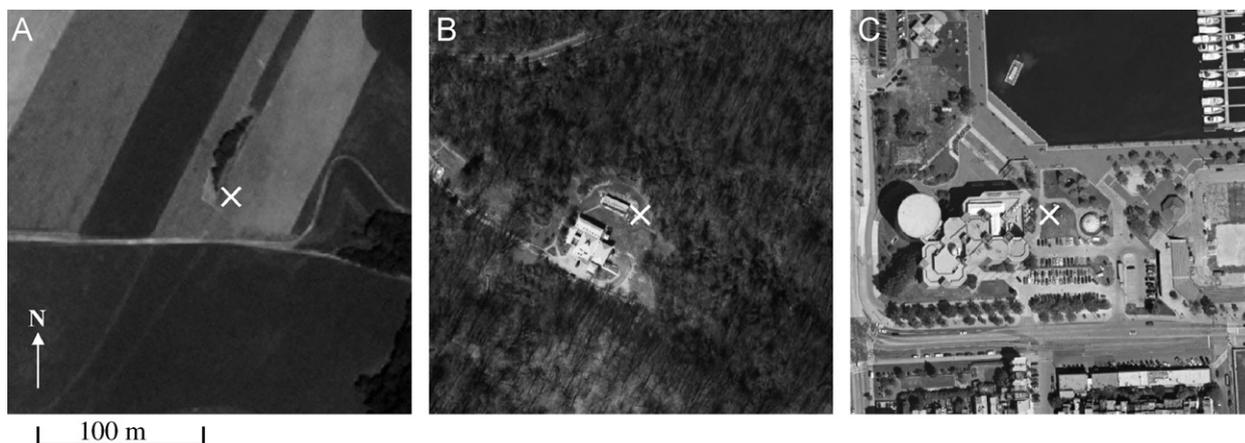


Fig. 1. Aerial view of the three sites: (A) rural; (B) suburban; and (C) urban. Images were obtained from Google Earth Beta (v4.1.7076.4458).

factor of 0.93, 0.83 and 0.87 at the rural, suburban and urban sites respectively, indicating the sensors had an open sky and were not subjected to long periods of shading by trees and buildings. All sensors were factory calibrated before their installation in the field. The CO<sub>2</sub> analyzers were calibrated every two weeks with known CO<sub>2</sub> concentrations. Air temperature sensors were not aspirated. Atmospheric CO<sub>2</sub>, air temperature, wind speed and direction and radiation variables were measured 1.5–2.0 m off the ground. Tipping rain buckets for precipitation were located approximately 1 m off the ground. Soil temperature and moisture were measured at a depth of 10 cm below the soil surface.

Additionally ozone at a height of approximately 1 m was monitored periodically through the summer months of 2003 and 2004 and more extensively in 2005 and 2006 using chemically sensitive filter paper. Wet deposition of nitrate and nitrite in rain water and dry deposition of nitrate from chemically sensitive filter paper, changed weekly, were monitored throughout the year of 2005. Soil nitrogen content was measured at the start of the growing season to estimate the potential impact of the input of nitrogen to a site from deposition. To remove variability associated with water availability, evaporation at each site was monitored through the growing season using an evaporation pan

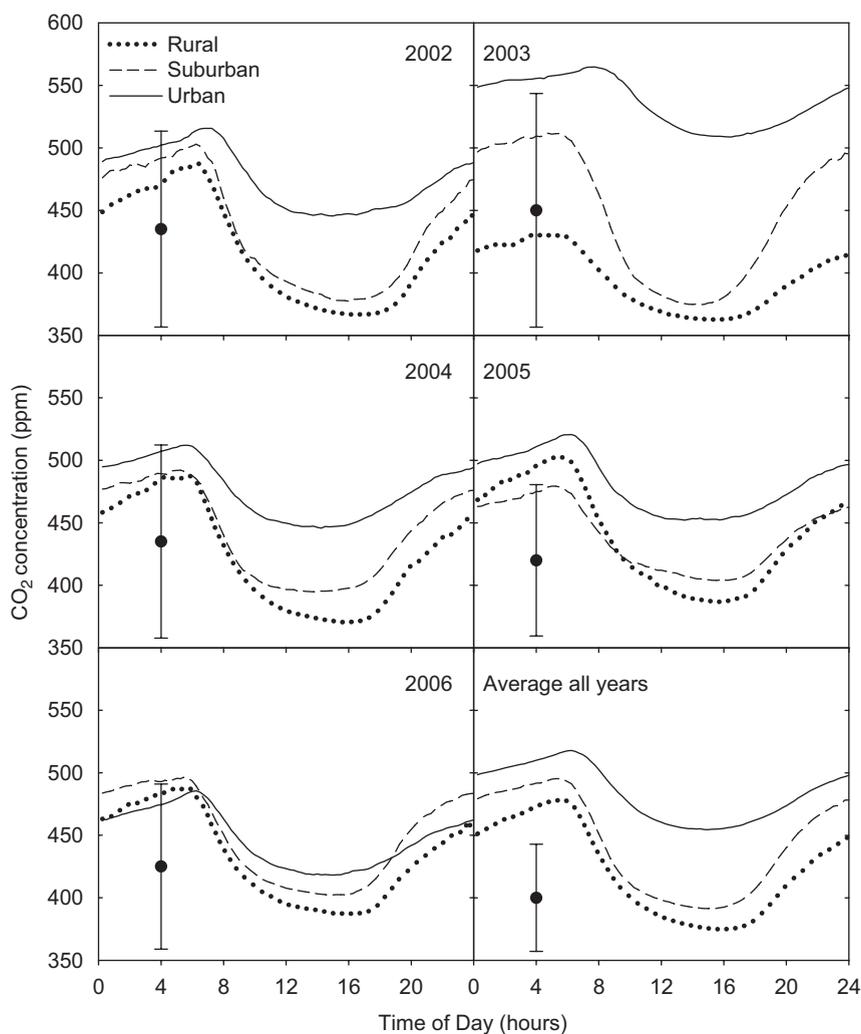


Fig. 2. Near surface atmospheric CO<sub>2</sub> concentration averaged over a 24 h period for each site along the transect for the 5 years of the study. Data were recorded every 15 min resulting in 96 values over a 24 h period consequently data points and error bars are not included but the data are connected by straight lines. The error bars shown in each graph is the maximum standard deviation from all sites.

(EP180, Global Water Instrumentation, USA) and any soil moisture deficit at a site compared to the others was eliminated by watering. This was to remove any variation in water availability to plant communities growing at each site for another experiment being conducted along this transect. The growing season was defined as the period between the last frost preceding winter and the first frost prior to the onset of the following winter. Frost was defined as temperatures reaching below 0 °C for at least an hour.

### 2.3. Statistics

Atmospheric CO<sub>2</sub> was examined to see if there were diurnal differences between sites and years

using analysis of covariance. The covariate was time of day and the independent factors were site and year. Air temperature, soil temperature, precipitation, RH and VPD were also analyzed using analysis of covariance but the covariate was day of year. Differences in ozone and soil nitrogen content between sites were analyzed using analysis of variance. Data were transformed where appropriate to meet the assumptions of normality and equality of variances for ANOVA. Total and diffuse radiation at the urban and rural sites was analyzed using the nonparametric Mann–Whitney *U* test as the data were not normally distributed. The variance of soil moisture variables was not equal so the nonparametric Kruskal Wallis test was performed. All statistics were performed using Statview (SAS Institute, USA).

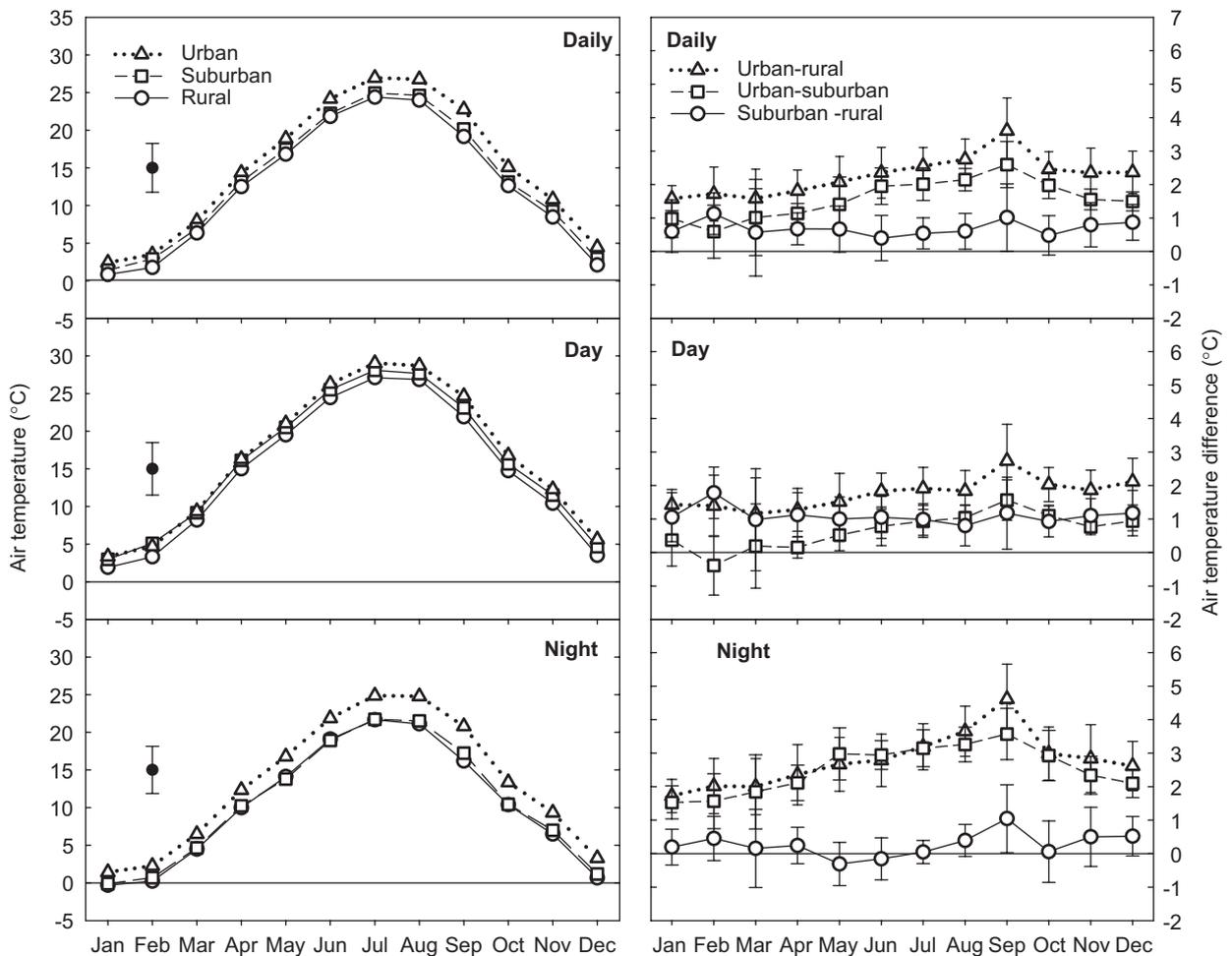


Fig. 3. Daily (24h), daytime and nighttime air temperature for each site along the transect averaged over the 5 years of the study. The error bars for the graphs on the left side are the maximum standard deviation. The graphs on the right represent the deviations in temperature between sites, the error bars are one standard deviation.

### 3. Results

#### 3.1. Atmospheric CO<sub>2</sub>

Atmospheric CO<sub>2</sub> concentration was significantly different between the three sites ( $P < 0.01$ ; Fig. 2). The highest concentration on average across the 5 years of the study was at the urban site (488 ppm) the lowest at the rural site (422 ppm) and the suburban site intermediate to the other two sites (442 ppm). CO<sub>2</sub> concentration at the urban site compared to the rural was increased on average by

16% from 2002 to 2006, which is low compared to 2000 and 2001 where CO<sub>2</sub> concentration was elevated by 21% (Ziska et al., 2004) and 31% (Ziska et al., 2003), respectively. This variation is consistent with our finding that CO<sub>2</sub> concentration differs significantly between years ( $P = 0.01$ ), although the average range in CO<sub>2</sub> concentrations between years for this study was small 443–459 ppm. Time of day, as expected, also significantly affected CO<sub>2</sub> concentration ( $P < 0.01$ ) with the lowest CO<sub>2</sub> concentration in the early afternoon and peaking in the early hours of the

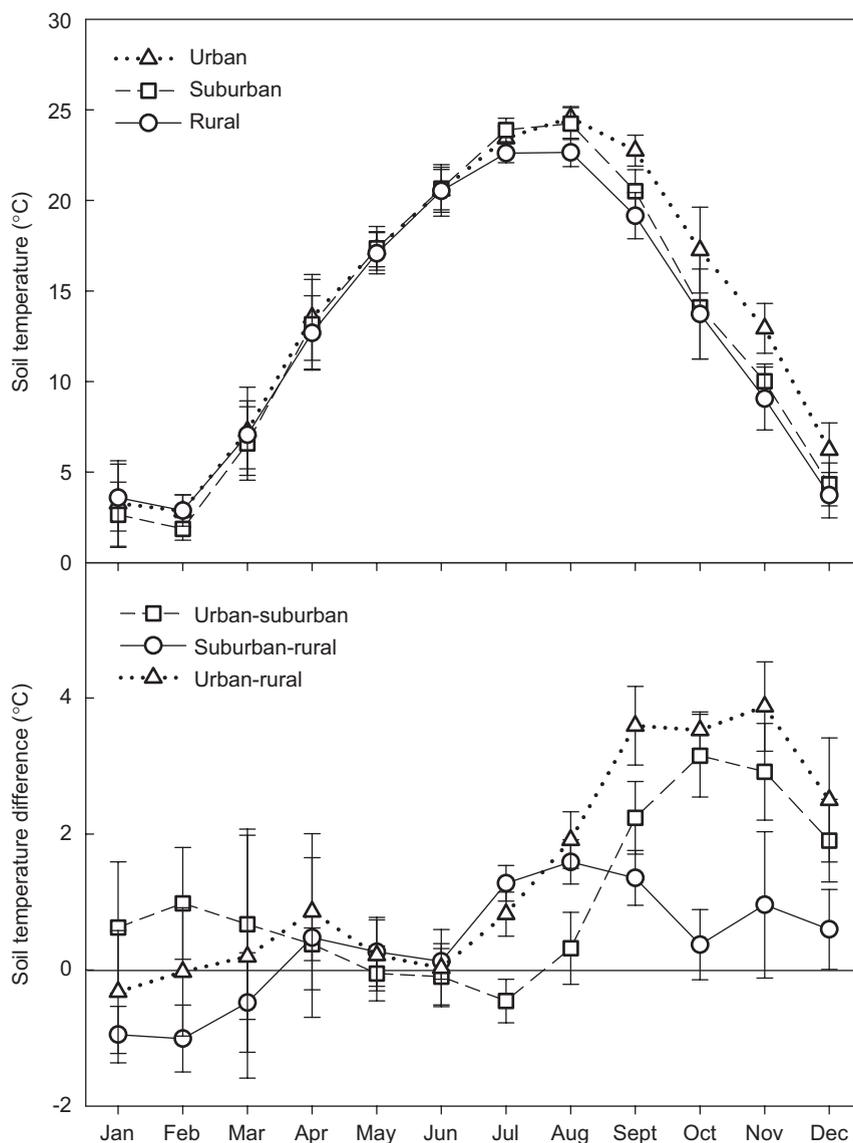


Fig. 4. Annual soil temperature (10 cm depth) averaged over the 5 years of the study and differences between sites on a monthly basis. The error bars are one standard deviation.

morning (Fig. 2). There was no difference between CO<sub>2</sub> concentrations on the weekend days (448.6 ppm) compared to the week days (448.7 ppm) at any site.

### 3.2. Air and soil temperature

Daily air temperature differed significantly between the sites ( $P = 0.01$ ; Fig. 3) with the highest air temperature at the urban site (14.8 °C), the lowest at the rural site (12.7 °C) and the suburban site falling in between the two extremes (13.6 °C). Across the State of Maryland average annual air temperature from 1901 to 2000 was 12.1 °C (<http://www.ncdc.noaa.gov/oa/climate/research/cag3/md.html>). Daily air temperature was also significantly different between years ( $P < 0.01$ ) with the highest average temperature in 2006 (14.3 °C) and the lowest in 2003 (13.2 °C). The same significant differences were apparent when temperature was calculated for night time and day time ( $P < 0.01$ ; Fig. 3). The temperature differences at night time were: urban (13.1 °C), suburban (10.9 °C) and rural (10.4 °C) and at day time: urban (16.5 °C), suburban (16.2 °C) and rural (14.8 °C). The biggest difference in temperature occurring at night between urban and rural sites (2.7 °C). Although these differences in temperature between sites are small, compared to the annual variation in temperature, they are consistent throughout the year (Fig. 3). The increase in air temperature at the urban site resulted in the growing season being longer compared to the suburban and rural sites. On average the growing season was 258 days at the urban site and 215 and 210 days, respectively, at the suburban and rural sites (Table 1).

Soil temperature (10 cm depth) was significantly different between sites ( $P < 0.01$ ; Fig. 4). The highest temperature was at the urban site (14.8 °C) and the lowest at the rural site (14.1 °C). The greatest

difference in soil temperature was between the urban and rural sites (0.7 °C) and the urban and suburban sites (0.5 °C) there was no significant difference in soil temperature between the rural and suburban site (0.1 °C). The major difference between sites occurred in the last 4 months of the year (Fig. 4) as the urban site remained warmer longer while temperatures dropped quicker at the other two sites.

### 3.3. Moisture variables

Precipitation was not significantly different between sites ( $P = 0.17$ ) or between years ( $P = 0.54$ ), although there was considerable variation in precipitation over the 5 years of the study (Table 2). The highest precipitation was in 2003 (1196 mm), which was also the wettest year on record for the State of Maryland. This resulted in relative humidity being significantly greater ( $P < 0.01$ ) and vapor pressure deficit (VPD) to be significantly reduced ( $P < 0.01$ ) in 2003 compared to the other years (Table 2). Relative humidity was not significantly different between sites ( $P = 0.12$ ) but VPD was significantly greater at the urban site ( $P < 0.01$ ) as air temperature was also higher.

Soil moisture was significantly different between site and year. The wettest site on average was the urban site this was because soil moisture in 2004 at this site was higher than previous years (Table 1). This may have been because of additional watering to meet evaporative demand. Apart from the high soil moisture at the urban site in 2004 all other sites and years had very similar average soil moisture.

### 3.4. Solar radiation

Total and diffuse radiation was measured at the urban and rural sites during 2005 and 2006. On a daily basis there was no significant difference

Table 1

Growing season length in days at each site over 5 years based on the last day that a frost occurred after winter and the first day that a frost appeared before winter

Year	Rural			Suburban			Urban		
	Last frost day	First frost day	Growing season length	Last frost day	First frost day	Growing season length	Last frost day	First frost day	Growing season length
2002	4 April	2 November	213	8 April	1 November	208	24 March	27 November	249
2003	16 March	23 October	222	2 April	23 October	205	15 March	2 December	263
2004	8 April	10 November	217	8 April	10 November	217	25 March	18 December	269
2005	16 April	11 November	210	22 March	17 November	241	15 March	18 November	249
2006	9 April	15 October	190	23 March	13 October	205	21 March	5 December	260

Table 2  
Moisture variables from each site across the transect

Variable	Site	2002	2003	2004	2005	2006	Average
Precipitation (mm)	Rural	1484	1288	906	978	767	1112 (4)
	Suburban	778	1150	970	1065	1052	1027 (4)
	Urban	785	1151	818	679	809	867 (3)
RH (%)	Rural	61.5 (20.7)	77.1 (14.9)	72.9 (15.2)	70.9 (13.9)	69.1 (14.5)	70.3 (8.9)
	Suburban	63.8 (18.1)	78.6 (16.1)	69.5 (15.7)	68.6 (14.2)	68.1 (15.9)	68.9 (10.2)
	Urban	60.2 (15.6)	68.4 (15.8)	66.3 (15.5)	64.9 (15.8)	60.7 (16.1)	64.1 (7.5)
VPD (kPa)	Rural	0.62 (0.38)	0.35 (0.25)	0.42 (0.26)	0.47 (0.31)	0.50 (0.30)	0.47 (0.20)
	Suburban	0.60 (0.38)	0.37 (0.26)	0.48 (0.28)	0.52 (0.32)	0.55 (0.35)	0.51 (0.22)
	Urban	0.72 (0.47)	0.54 (0.38)	0.59 (0.36)	0.67 (0.46)	0.72 (0.46)	0.65 (0.33)
Soil moisture (%)	Rural			13.8 (1.8)	10.1 (5.3)	9.7 (4.7)	12.0 (3.4)
	Suburban			10.1 (1.4)	9.3 (3.0)	12.5 (2.3)	10.8 (1.6)
	Urban			19.3 (2.2)	8.7 (3.9)	12.7 (3.5)	14.4 (2.7)

Precipitation is summed over the year. RH, VPD and soil moisture are a daily average over the year. Average is the over all 5 years and the values in brackets are the daily standard deviations. Soil moisture was not recorded in 2002 and 2003.

between total radiation ( $P = 0.08$ ) at the urban (average  $347.9 \mu\text{mol m}^{-2} \text{s}^{-1}$ ) and rural (average  $398.5 \mu\text{mol m}^{-2} \text{s}^{-1}$ ) sites. Diffuse radiation although more than 50% lower than total radiation, was not significantly different ( $P = 0.19$ ) between the urban (average  $159.2 \mu\text{mol m}^{-2} \text{s}^{-1}$ ) and rural (average  $189.3 \mu\text{mol m}^{-2} \text{s}^{-1}$ ) sites. Although total radiation is slightly reduced in the urban compared to the rural site, diffuse radiation is also lower suggesting that the potential for increased air pollution in the urban environment is not increasing diffuse radiation.

### 3.5. Ozone and nitrogen deposition

Ozone concentrations did not differ significantly between the sites along the transect ( $P = 0.10$ ; Fig. 5). There was a significant difference between years ( $P < 0.01$ ); on average the ozone concentration for 2005 was  $45 \text{ ppb} \pm 21$  (one standard deviation) and for 2006  $35 \text{ ppb} \pm 13$  which were not high enough to affect plant physiology and growth. Ozone was also measured for short time periods in 2003 (day 248–273, September) and 2004 (day 78–174, March–June) and the average ozone concentrations respectively were  $24 \text{ ppb} \pm 7$  and  $34 \text{ ppb} \pm 11$ , which during these short time periods is similar to the ozone concentrations shown for 2005 and 2006 (Fig. 5).

Wet and dry deposition was highest at the rural and lowest at the urban site (Table 3). Soil total nitrogen content was not significantly different

between sites ( $P = 0.59$ ). The addition of wet and dry deposition to soil contributed only 2.1%, 1.8% and 1.2% nitrogen annually to the rural, suburban and urban sites, respectively.

## 4. Discussion

Across the transect, atmospheric  $\text{CO}_2$  and temperature were elevated at the urban site and gradually decreased out to the suburban and rural sites. This is consistent with other studies that have found air temperature and atmospheric  $\text{CO}_2$  are closely related to population and associated high traffic volume in urban city centers (Idso et al., 1998, 2001, 2002; Brazel et al., 2000; Nasrallah et al., 2003; Velasco et al., 2005). Increased air temperature at the urban site significantly increased the VPD. Nitrogen deposition although highest at the rural site was not great enough to increase soil nitrogen content compared to the other sites. Along the transect changes in the microclimate (IPCC, 2007) and deposition of nutrients (Denman et al., 2007) are consistent with predictions of modifications in the environment expected with global climate change. It appears that densely populated urban areas could provide a setting that is suitable for studying the effects of future global climate change on terrestrial ecosystems.

Globally averaged surface atmospheric  $\text{CO}_2$  concentrations are 379 ppm (2005; IPCC, 2007) and are estimated to increase by 50–100 ppm by 2100 (Friedlingstein et al., 2006). Along the transect on average the lowest  $\text{CO}_2$  concentration was at the

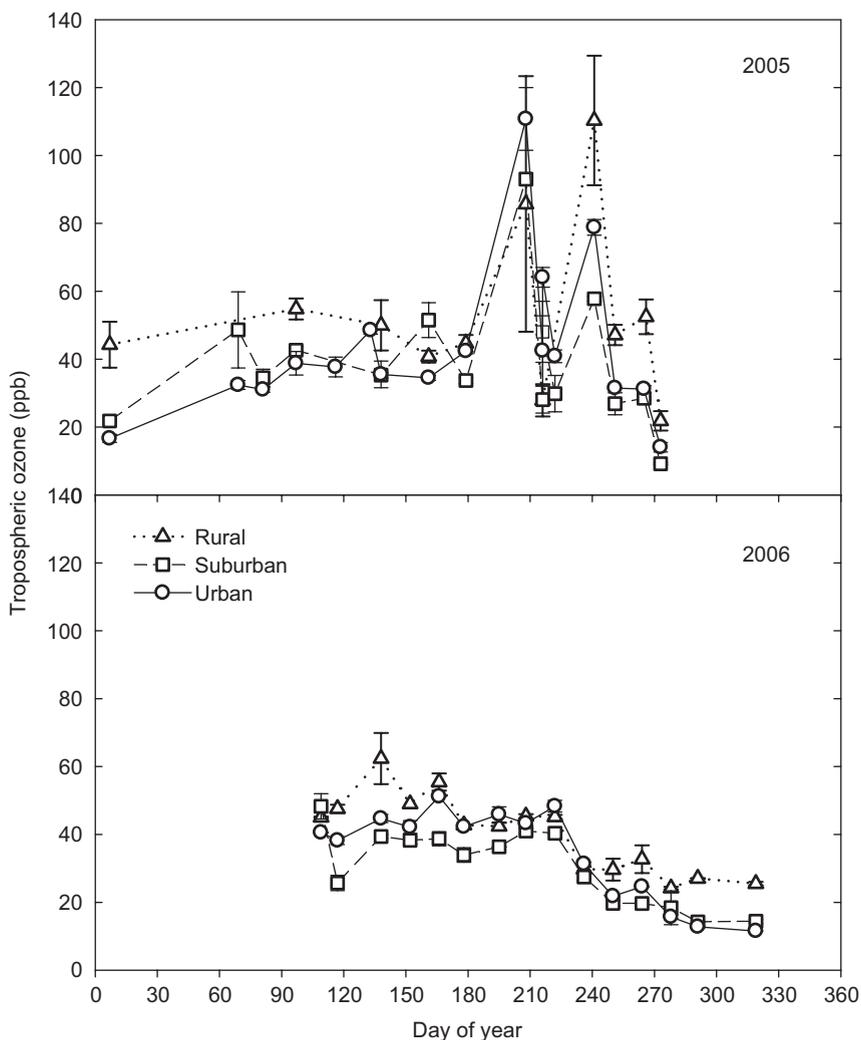


Fig. 5. Tropospheric ozone measured at each site through the growing season of 2005 and 2006. The error bars are one standard deviation.

rural site, 422 ppm, which increased by 66 ppm at the urban site. This difference in  $\text{CO}_2$  concentration between sites was maintained over the 5 year period with the average annual  $\text{CO}_2$  concentration at the rural site ranging from 395–439 ppm and the urban site 448–537 ppm. This difference between the urban and rural sites is similar to the increase in  $\text{CO}_2$  concentration found in Phoenix, USA, where an increase of 111–185 ppm was reported from a pristine rural site to the city center (Idso et al., 1998, 2001). The sustained increase in  $\text{CO}_2$  concentration over 5 years between an urban and rural site, is within the range expected with global climate change predictions (IPCC, 2007), and is expected to significantly impact biological systems.

Air temperature globally varies greatly but on average it is predicted to increase by 1.1–6.4 °C by 2100 and this increase will be greatest on land in the northern hemisphere (IPCC, 2007). Along the transect temperature was significantly different between sites with an average daily difference between the urban and rural sites of 2.1 °C. This difference was greatest in September and at night and is likely a consequence of anthropogenic heating, such as building and industrial energy consumption, and vehicle fuel combustion, which can contribute significantly to the urban heat island in winter (Fan and Sailor, 2005). This resulted in the growing season at the urban site being 36–70 days longer over 5 years than the other two sites as

Table 3

Dry and wet deposition nitrogen measured in 2005 at each site and compared to US-EPA data (<http://cfpub.epa.gov/gdm/index.cfm>)

	Rural	Suburban	Urban	US-EPA, Beltsville, MD
Dry deposition ( $\mu\text{g m}^{-3}$ )				
Nitric acid	3.96	3.27	5.24	2.04
Nitrate				1.33
Ammonium				1.43
Wet deposition ( $\text{kg ha}^{-1}$ )				
Nitrate	7.12	5.08	4.01	2.82
Nitrite	0.19	0.12	0.08	
Ammonium				1.76
Soil N ( $\text{kg ha}^{-1}$ )	771.8	735.8	783.9	
Total wet deposition ( $\text{kg ha}^{-1}$ )	11.75	8.37	6.59	4.38
Total dry deposition ( $\text{kg ha}^{-1}$ )	6.11	4.35	3.43	2.38

Soil nitrogen content was quantified at each site and addition from total wet and dry deposition was estimated across the transect. Total wet deposition is nitrate and nitrite summed from each site and ammonium estimated as 38% of nitrate based on US-EPA values. Total dry deposition is estimated as 52% of wet deposition based on US-EPA data (<http://cfpub.epa.gov/gdm/index.cfm>).

freezing temperatures did not occur until later in the year. Soil temperature was also higher by  $0.7\text{ }^{\circ}\text{C}$  at the urban compared to the rural site. The soil temperature difference between the sites was largest in the last four months of the year, similar to the air temperature patterns. Air temperature was consistently higher at the urban site and similar to predictions of global climate change resulted in fewer frost days and warmer night time temperatures (IPCC, 2007).

Precipitation was above average in Maryland for 4 of the 5 years of the study and 2003 was the wettest year recorded over more than 100 years (<http://www.ncdc.noaa.gov/oa/climate/research/cag3/md.html>). Although precipitation measured in our study was not significantly different between sites and years, there is a great amount of variation across the region and between years (Table 2). Studies have found that precipitation increases over urban areas (Burian and Shepherd, 2005) and is predicted to increase in North America with climate change (Diffenbaugh et al., 2005). Although we saw no differences in precipitation in our study, in the past decade the State of Maryland as a whole is seeing more extreme and higher annual precipitation than experienced in the last 100 years (<http://www.ncdc.noaa.gov/oa/climate/research/cag3/md.html>). A consequence of urbanization on moisture variables observed in our study, was the increase in air temperature at the urban site significantly increased VPD, whereas, during wet years relative humidity was significantly increased

and VPD significantly reduced. This is consistent with a previous study that found air temperature modified moisture variables more than relative humidity except during wet periods (Barradas, 1991). VPD directly affects plant physiology (Aphalo and Jarvis, 1991), influencing gas exchange and growth rates of plants, which can impact the urban climate.

Urban environments can impact air quality variables other than atmospheric  $\text{CO}_2$  such as tropospheric ozone and nitrogen deposition. In 2006 Baltimore experienced 17 days where ozone levels on average were above 100 ppb for 8 h and western Maryland experienced 2 days (considered unhealthy for sensitive groups), all occurring between May and August (<http://www.mde.state.md.us/Programs/AirPrograms/Monitoring/aqsummaries/index.asp>). Our measurements at each site indicated peak values between June and September but on average across the year ozone concentration was below levels that would affect human or plant physiology (McKee, 1994). Wet and dry deposition added a small percent of nitrogen compared to soil nitrogen content. Nitrogen deposition appeared to be highest at the rural site compared to the urban although soil nitrogen values were not different between sites. Total solar radiation appears to be slightly lower at the urban compared to the rural site, which is likely indicative of the urban site experiencing some shading from surrounding buildings during part of the day. Diffuse radiation, which can be increased by particulate matter such as from vehicular emissions, was not different between the sites and on average was lower at the urban compared to the rural site. It is

evident from our data that urban environments provide a microclimate that is representative of changes predicted in the future with global climate change (IPCC, 2007), consequently vegetation within urban areas are currently experiencing elevated atmospheric CO<sub>2</sub> and temperature levels that can significantly affect plant growth compared to rural areas.

Over the 5 year period of this study it appears that a densely populated, urban center, is affecting the microclimate so that it is similar to climate predicted globally in the next 100 years. Over the 5 years of the study CO<sub>2</sub> concentrations and air temperature were consistently higher at the urban site compared to the suburban and rural sites. In the USA 80% of the population resides in urban areas and urbanization is continuing to increase (United Nations, 2004). Increasing land use change from rural agricultural areas to dense populated urban areas will have significant impacts on carbon emissions and cycling (Imhoff et al., 2004).

### Acknowledgments

We would like to thank Ernie Goins, Danielle Reid and Jonathan Clark for assistance with data collection and site maintenance. This research was supported by the Biological and Environmental Research Program (BER), US Department of Energy, Interagency Agreement no. DE-AI02-02ER63360 to L.H.Z. and J.A.B.

### References

- Aphalo, P.J., Jarvis, P.G., 1991. Do stomata respond to relative humidity? *Plant, Cell and Environment* 14, 127–132.
- Barradas, V.L., 1991. Air temperature and humidity and human comfort index of some city parks of Mexico City. *International Journal of Biometeorology* 35, 24–28.
- Berry, R.D., Colls, J.J., 1990. Atmospheric carbon dioxide and sulphur dioxide on an urban/rural transect-I. Continuous measurements at the transect ends. *Atmospheric Environment* 24A, 2681–2688.
- Brazel, A., Selover, N., Vose, R., Heisler, G., 2000. The tale of two climates—Baltimore and Phoenix urban LTER sites. *Climate Research* 15, 123–135.
- Burian, S.J., Shepherd, J.M., 2005. Effect of urbanization on the diurnal rainfall pattern in Houston. *Hydrological Processes* 19, 1089–1103.
- Day, T.A., Gober, P., Xiong, F.S., Wentz, E.A., 2002. Temporal patterns in near-surface CO<sub>2</sub> concentrations over contrasting vegetation types in the Phoenix metropolitan area. *Agricultural and Forest Meteorology* 110, 229–245.
- Denman, K.L., Brasseur, G., Chidthaisong, A., Ciais, P., Cox, P.M., Dickinson, R.E., Hauglustaine, D., Heinze, C., Holland, E., Jacob, D., Lohmann, U., Ramachandran, S., da Silva Dias, P.L., Wofsy, S.C., Zhang, X., 2007. Couplings between changes in the climate system and biogeochemistry. In: Solomon, S., et al. (Eds.), *Climate Change 2007: The Physical Science Basis. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change*. Cambridge University Press, Cambridge, UK and New York, USA, pp. 500–587.
- Diffenbaugh, N.S., Pal, J.S., Trapp, R.J., Giorgi, F., 2005. Fine-scale processes regulate the response of extreme events to global climate change. *Proceedings of the National Academy of Sciences* 102 (44), 15774–15778.
- Fan, H., Sailor, D.J., 2005. Modeling the impacts of anthropogenic heating on the urban climate of Philadelphia: a comparison of implementations in the two PBL schemes. *Atmospheric Environment* 39, 73–84.
- Friedlingstein, P., Cox, P., Betts, R., Bopp, L., von Bloh, W., Brovkin, V., Cadule, P., Doney, S., Eby, M., Bala, I., John, J., Jones, C., Joos, F., Kato, T., Kawamiya, M., Knorr, W., Lindsay, K., Matthews, H.D., Rayner, T., Reick, C., Roeckner, E., Schnitzler, K.-G., Schnur, R., Strassmann, K., Weaver, A.J., Yoshikawa Zeng, C., 2006. Climate–carbon cycle feedback analysis: results from the C4MIP model intercomparison. *Journal of Climate* 19, 3337–3353.
- Idso, C.D., Idso, S.B., Balling, R.C., 1998. The urban CO<sub>2</sub> dome of Phoenix, Arizona. *Physical Geography* 19, 95–108.
- Idso, C.D., Idso, S.B., Balling, R.C., 2001. An intensive two-week study of an urban CO<sub>2</sub> dome in Phoenix, Arizona, USA. *Atmospheric Environment* 35, 995–1000.
- Idso, S.B., Idso, C.D., Balling, R.C., 2002. Seasonal and diurnal variations of near-surface atmospheric CO<sub>2</sub> concentration within a residential sector of the urban CO<sub>2</sub> dome of Phoenix, AZ, USA. *Atmospheric Environment* 36, 1655–1660.
- Imhoff, M.L., Bounoua, L., DeFries, R., Lawrence, W.T., Stutzer, D., Tucker, C.J., Ricketts, T., 2004. The consequences of urban land transformation on net primary productivity in the United States. *Remote Sensing of Environment* 89, 434–443.
- IPCC, 2007. Summary for Policymakers. In: Solomon, S., Qin, D., Manning, M., Chen, Z., Marquis, M., Averyt, K.B., Tignor, M., H.L. Miller, M. (Eds.), *Climate Change 2007: The Physical Science Basis. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change*. Cambridge University Press, Cambridge, United Kingdom and New York, USA, pp. 18.
- Koerner, B., Klopatek, J., 2002. Anthropogenic and natural CO<sub>2</sub> emission sources in an arid urban environment. *Environmental Pollution* 116, S45–S51.
- McKee, D.J., 1994. *Tropospheric Ozone: Human Health and Agricultural Impacts*. Lewis Publishers, Boca Raton, pp. 333.
- Nasrallah, H.A., Balling, R.C., Madi, S.M., Al-Ansari, L., 2003. Temporal variations in atmospheric CO<sub>2</sub> concentrations in Kuwait City, Kuwait with comparisons to Phoenix, Arizona, USA. *Environmental Pollution* 121, 301–305.
- Oke, T.R., 1982. The energetic basis of the urban heat island. *Quarterly Journal of the Royal Meteorological Society* 108, 1–24.
- Oke, T.R., Maxwell, G.B., 1975. Urban heat island dynamics in Montreal and Vancouver. *Atmospheric Environment* 9, 191–200.
- Pataki, D.E., Alig, R.J., Fung, A.S., Golubiewski, N.E., Kennedy, C.A., McPherson, E.G., Nowak, D.J., Pouyat, R.V., Romero Lankao, P., 2006. Urban ecosystems and the North American carbon cycle. *Global Change Biology* 12, 2092–2102.

- Reid, K.H., Steyn, D.G., 1997. Diurnal variations of boundary layer carbon dioxide in a coastal city—observations and comparison with model. *Atmospheric Environment* 31, 3101–3114.
- Taha, H., Akbari, H., Rosenfeld, A., 1991. Heat Island and Oasis effects of vegetative canopies: micro-meteorological field-measurements 44, 123–138.
- United Nations, 2004. *World Urbanization Prospects: The 2003 Revision*. United Nations: New York.
- Velasco, E., Pressley, S., Allwine, E., Westberg, H., Lamb, B., 2005. Measurements of CO<sub>2</sub> fluxes from the Mexico City urban landscape. *Atmospheric Environment* 39, 7433–7446.
- Vogt, R., Christen, A., Rotach, M.W., Roth, M., Satyanarayana, A.N.V., 2006. Temporal dynamics of CO<sub>2</sub> fluxes and profiles over a central European city. *Theoretical and Applied Climatology* 84, 117–126.
- Ziska, L.H., Gebhard, D.E., Frenz, D.A., Faulkner, S., Singer, B.D., Straka, J.G., 2003. Cities as harbingers of climate change: Common ragweed, urbanization and public health. *Journal of Allergy and Clinical Immunology* 111, 290–295.
- Ziska, L.H., Bunce, J.A., Goins, E.W., 2004. Characterization of an urban/rural CO<sub>2</sub>/temperature gradient and associated changes in initial plant productivity during secondary succession. *Oecologia* 139, 454–458.